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Program and Reports

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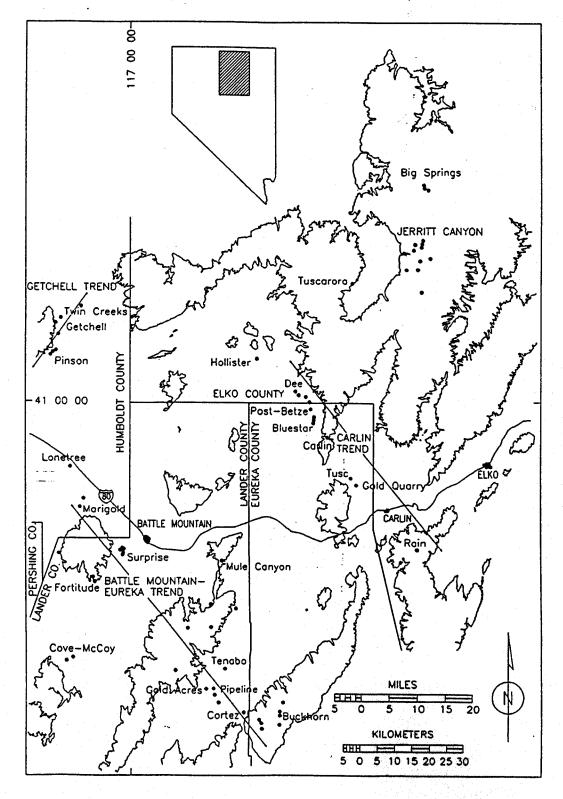


Figure 1: Location map for the Jerritt Canyon District (adapted from Phinisey, et al., 1996)

Geology, Alteration, and Geochemical Dispersion in the Paleohydrologic System at the SSX mine, Jerritt Canyon District, Elko County, Nevada

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Introduction

The SSX deposit (South Saval Extension) is one of several sedimentary rock-hosted gold deposits in the Jerritt Canyon District. The district is located in the north-central part of the Independence Mountains, approximately 40 miles north-northwest of Elko, Nevada (fig. 1). The location of the SSX deposit is shown in figure 2, and it was discovered by exploration drilling along structural features observed by geologists in the Saval open-pit. The SSX mine is located 500 to 1000 feet below the surface, and is being mined by underground methods (underhand stope/fill) at an average grade of approximately 0.4 oz/ton with a cutoff of 0.2 oz/ton. The goal of this research is to characterize the ore system in terms of fluid flow paths using alteration/geologic mapping and geochemical data, supported by petrographic and instrumental analyses.

Geologic Units

The geology of the district is characterized by faulted and folded Paleozoic sedimentary rocks of the lower and upper plates of the Roberts Mountains thrust (Birak, 1986). The upper plate rocks are pelagic, siliciclastic and volcanic rocks of the Upper Cambrian to Middle Ordovician Snow Canyon Formation (Leslie, 1990). The lower plate rocks include portions of the Ordovician Pogonip Group, the Ordovician Eureka Quartzite, the Ordovician-Silurian Hanson Creek Formation, and a Silurian portion of the Roberts Mountains Formation (Hofstra et al., 1991). There are also four distinct types of intrusive rock recognized in the

district: a Pennsylvanian basalt dike and three Tertiary types ranging in composition from basalt to quartz monzonite (Phinisey, 1995).

The units exposed in the SSX mine area are the Hanson Creek Formation, the Roberts Mountains Formation and the Snow Canyon Formation, all of which are intruded by a basalt dike (likely to be one of the Pennsylvanian basalt dikes described by Phinisey). The Hanson Creek Formation has been divided into five distinct units by IMC staff and labeled numerically from top to bottom (order of appearance in surface drilling). The following descriptions are in order from oldest to youngest. Hanson 5 is described as fine-grained, laminated limestone with sparse chert nodules. Hanson 4 is a banded, fine to mediumgrained, laminated, thinly bedded limestone with abundant chert nodules. Hanson 3, the primary host for gold mineralization, is a rhythmically interbedded, locally fossiliferous micrite and calcareous siltstone, which grades upward from nodular and undulatory bedding surfaces to tabular and planar beds. Hanson 2 is a fine-grained, locally dolomitic limestone with the least amount of silty particles. Hanson 1 is a unit consisting of tabular, rhythmically interbedded chert, limestone, and dolostone (Hohbach and Roth, 1983). The Saval Discontinuity separates the Hanson Creek Formation from the overlying Roberts Mountains Formation which is best described as a locally carbonaceous, finely laminated calcareous siltstone.

Orebody Controls

There appear to be two primary factors controlling ore-grade gold mineralization at SSX: 1) stratigraphy, and 2) structural controls. The ore is hosted mainly within the Hanson Creek Formation in unit 3, where the enhanced permeability of silty interbeds appears to have conducted fluids. Samples of micritic beds and silty interbeds were collected and assayed by I.M.C. mine geologists. Silty interbeds were found to contain significantly higher gold values than micrite beds. Silty interbeds also exhibit alteration features that will be described in the following section.

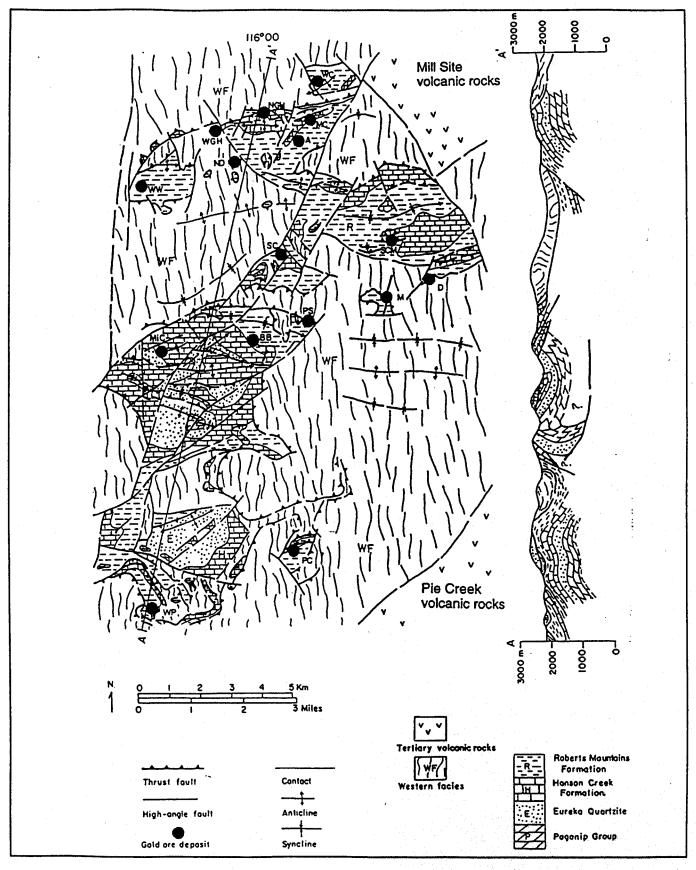


Figure 2: Geologic map and cross section of the Jerritt Canyon District showing locations of gold deposits. SC-Saval Canyon, BB-Burns Basin, PS-Patani Springs and South Saval Extension SSX (adapted from Phinisey, et al., 1996)

The second factor controlling ore-grade mineralization is structure. There are two important sets of faults. The older set is partially occupied by the basalt dike (labeled by mine geologists as the South Boundary Fault), and strikes N70W to N50W with nearly vertical dips. These older faults are 6 to 36-inch, gouge-filled shears characterized by high carbon, arsenic and iron sulfides, and silica content. The dike ranges from 10 to 25 feet in thickness and occupies one of the southernmost of these faults.

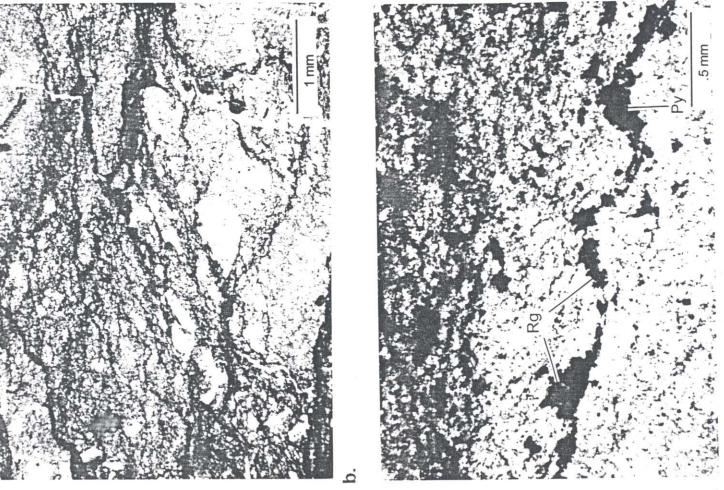
There are at least two faults in the mapped workings comprising the younger set of structures. One strikes N16E and dips 60 to 80 NW, and the other strikes approximately N40E and dips 62 NW. Both faults truncate the basaltic dike, have normal and possibly left lateral offset, and share similar features: dolomitized wallrocks, carbon enrichment, arsenic and iron sulfides, silicification, and brecciation (both tectonic and decarbonatization/collapse). The ore mineralization/grade is more substantial with proximity to these structures and is highly localized in the Hanson 3 both in the hanging wall (mapped area) and footwall (not exposed).

Alteration

Four different types of wallrock alteration were mapped in the SSX workings: 1) decarbonatized wallrock; 2) arsenic minerals, 3) carbon enrichment, and 4) silicified rock. The degree of decarbonatization was determined using 10% HCl and categorized by amount of effervescence in terms of weak, moderate, and strong. Decarbonatized rock locally exhibits macroscopic dissolution features such as collapse breccias and slump blocks, but the most consistent evidence for decarbonatization is the development of stylolites. The stylolites are 0.01 mm to 0.8mm thick, crenulated structures, consisting of insoluble residue/carbonaceous material, and locally pyrite, arsenic sulfides, and silica (fig. 3). In weakly decarbonatized Hanson 3, stylolite development is concentrated along and adjacent to siltstone interbeds. Stronger decarbonatization results in a higher concentration of stylolites and a resulting decrease in the thickness of micrite beds. Intensely decarbonatized intervals



- Stylolites along siltstone interbeds of Hanson Creek Formation, unit 3, plane light, 40X
- b. Stylolites in fossiliferous Hanson 3 micrite, plane light, 40X
 - c. Realgar, pyrite and silica filled stylolite in Hanson 3 micrite, plane light, 100X



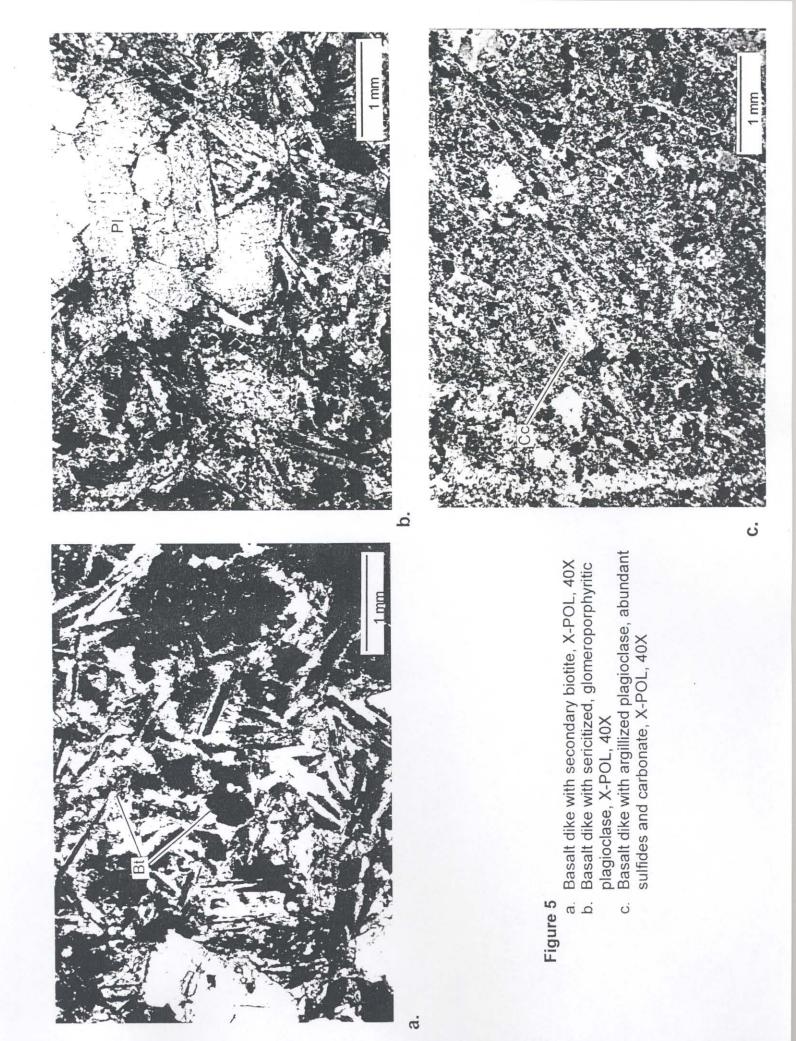
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without breccia result in a rock with the appearance of equally thick dark and light colored beds. Decarbonatization to the point of sanding is evident in the Hanson 2, and locally in Hanson 3 micrite, although very strongly decarbonatized Hanson 3 results mainly in collapse breccia formation. The strongly decarbonatized rock is localized near structures and is commonly accompanied by other alteration types.

Some of the altered rock that was mapped as decarbonatized is now known to be dolomitized. Petrographic analysis of samples taken near the northeasterly faults revealed rhombic carbonate overgrowths on anhedral "dusty" carbonate grains (fig. 4). Thin-section billets stained with alizarin red indicate dolomite instead of abundant calcite. Dolomitization is localized in northeasterly-trending fault-related breccias. Dolomite is also present in veinlets and as breccia filling between dolomitized fragments.

Carbon enrichment is apparent throughout the mapped workings, but is localized in siltstone interbeds adjacent to northeasterly and most northwesterly structures. Outside of the ore zone, carbon flooding is apparent along calcite veins. Arsenic minerals tend to coincide with carbon-enriched rock as well as in silicified rock. Silica is present in veinlets and grains of jigsaw-textured quartz and some chalcedonic quartz veinlets in the basaltic dike, but no true jasperoid was apparent in the mapped area.

Alteration of the basalt dike increases in intensity from east to west approaching the northeasterly faults. The least altered of the dike samples is identical to deuterically altered basaltic dike described by Phinisey (1995) and is characterized by potassium metasomatism resulting in secondary biotite (fig. 5a) and k-feldspar overgrowths on plagioclase (not observed). The next most intense dike-hosted alteration is phyllic alteration characterized by sericite, quartz, pyrite, and leucoxene (fig. 5b). The most intense alteration of the dike is argillic, with abundant kaolinite, sericite, smectite, quartz, carbonate and marcasite (fig. 5c).



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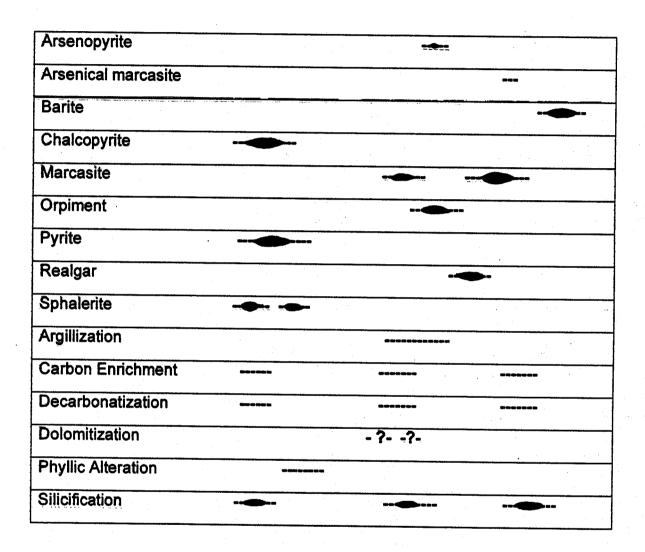
Paragenetic Relationships

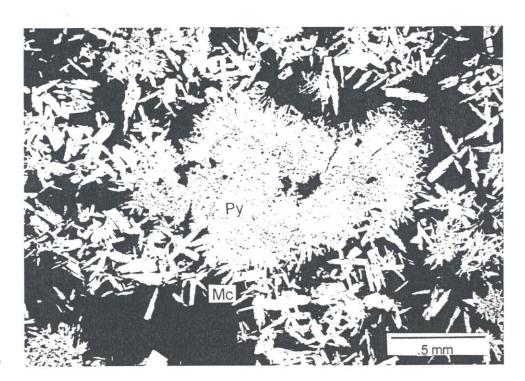
Petrographic analysis of 80 thin sections and 8 polished-thin sections has permited a preliminary interpretation of some paragenetic relationships (fig. 6). Early base-metal sulfides present in the basalt dike are replaced by marcasite and arsenic sulfides (fig. 7a). In the arsenic-rich assemblage, a progression from oxidized to more reduced fluids is evidenced by early orpiment replaced and rimmed by realgar (fig. 7b). In the basalt dike, anhedral pyrite is rimmed and replaced by euhedral marcasite, and marcasite veinlets cut earlier realgar (fig. 8).

Away from northeasterly structures and spatially associated arsenic sulfides, arsenopyrite rims on marcasite may represent arsenic mineralization. A backscatter electron image of a rimmed marcasite crystal (fig. 9) reveals a progression from marcasite to arsenopyrite to porous marcasite to arsenical marcasite. This image was taken from a sample of basalt dike in an arsenic sulfide-poor, iron sulfide-rich zone away from a northeasterly fault. Within the ore zone, marcasite is the dominant iron sulfide and realgar is the dominant arsenic sulfide. In the basalt dike, pyrite is rimmed and replaced by marcasite overgrowths. This replacement is progressively more complete with proximity to northeasterly faults, and coincides with more intense alteration of the dike.

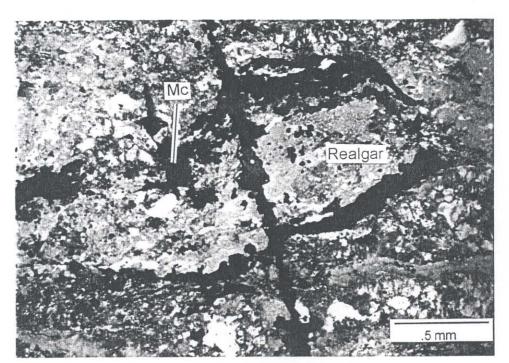
Chalcopyrite and sphalerite were found in the basalt dike near northeast-trending faults but outside of the ore zone. Sphalerite commonly has chalcopyrite "disease" and pyrite and marcasite replace these sulfides. Sulfides in the basalt dike are found mainly in ferromagnesian mineral sites and in veinlets. Barite is localized in veinlet cores and breccia vugs.

Paragenetic Diagram





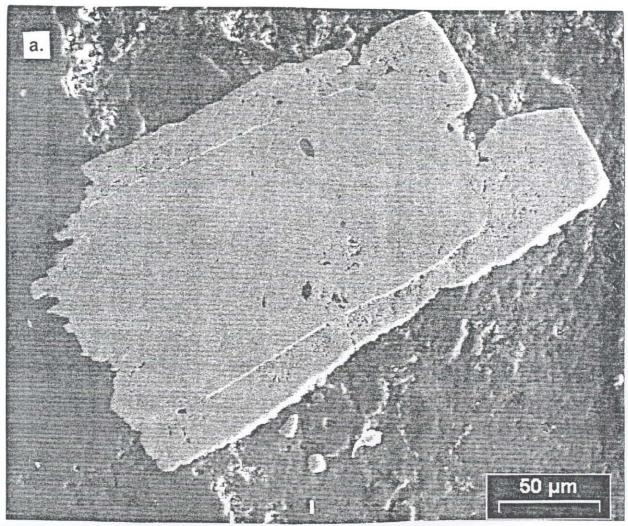
a.

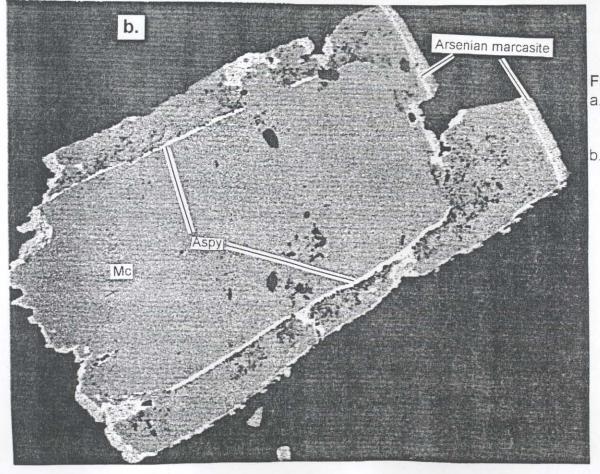


b.

Figure 8

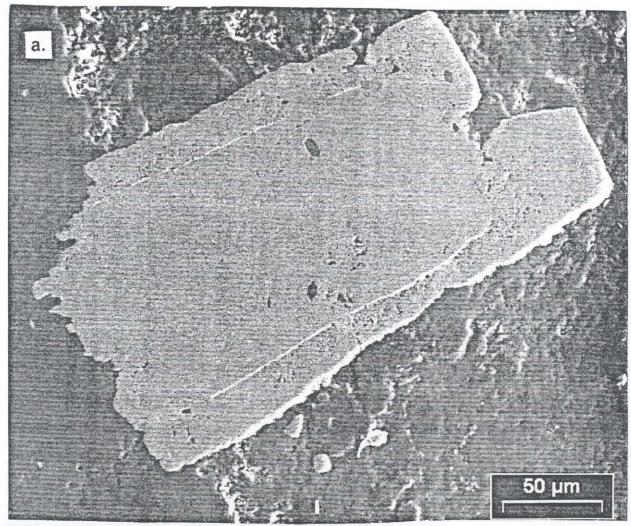
- a. Pyrite rimmed by marcasite, reflected light, 100X
 b. Marcasite veinlet cutting realgar, reflected light, X-POL, 100X

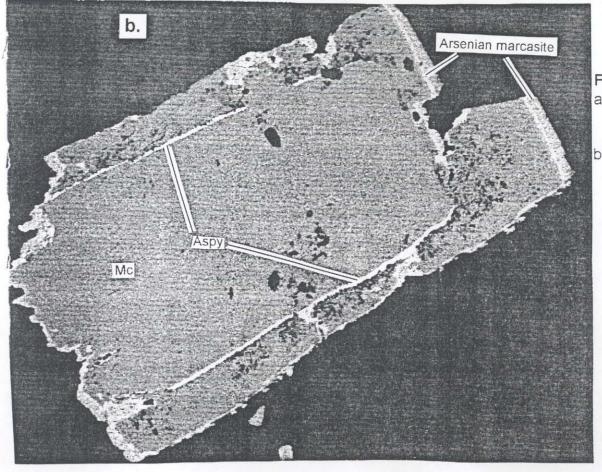




- Figure 9

 a. Secondary electron image of zoned marcasite, 400X
- b. Backscatter electron Image, 400X





- Figure 9
 a. Secondary electron image of zoned marcasite, 400X
- b. Backscatter electron Image, 400X

Geochemistry

The geochemical data for the SSX project area were obtained from 164 surface downhole rotary and core drill-holes covering approximately 1 square mile. The drill-hole depths average approximately 1000 feet and composite samples were taken at average intervals of 50 feet. 3146 samples were analyzed by Chemex Labs for 33 elements: Au (ppb), Ag (ppm), Al (%), As (ppm), Ba (ppm), Be (ppm), Bi (ppm), Ca (%), Cd (ppm), Co (ppm), Cr (ppm), Cu (ppm), Fe (%), Ga (ppm), Hg (ppb), K (%), La (ppm), Mg (%), Mn (ppm), Mo (ppm), Na (%), Ni (ppm), P (ppm), Pb (ppm), Sb (ppm), Sc (ppm), Sr (ppm), Ti (%), Tl (ppm), U (ppm), V (ppm), W (ppm), and Zn (ppm). Of these elements, Be, Bi, Ga, La, Na, Ti, and U were not variable enough for study.

A spreadsheet program-was used to organize and analyze the geochemical data. A statistical summary of each element is shown in table 1. A correlation matrix is given in table 2, with significant values (greater than 0.5) in bold. There are six drill-holes within the mapped area that were used to characterize the geochemical profile of major structures.

Correlation

The data were analyzed using a computer and spreadsheet program to generate a correlation matrix (table 2). There are three distinct groups of elements that have a significant correlation. 1) The highest degree of correlation is between cadmium, vanadium and zinc. These elements also have a lesser degree of correlation with silver, molybdenum, nickel, and copper. 2) Scandium, cobalt, iron, aluminum and copper correlate well with each other, and potassium, nickel, chromium and manganese weakly correlate with these elements. 3) Finally, another group of elements that has a significant degree of correlation is gold, mercury, arsenic, and thallium.

Table 1 Statistical Summary for Downhole Geochemistry, SSX Project Area

Element	Mean	Standard Error	Median	Mode	Standard Deviation	Sample Variance	Kurtosis	Skewness	Range	Minimum	Maximum
Au (ppb)	186.6969793	13.65990794	10	0	766.0517293	586835.252	81.42543782	8.128751455	10000	0	10000
Ag (ppm)	0.407408585	0.013105396	0.2	0.1	0.734954556	0.5401582	110.7335725	7.910543447	15.7	0.1	15.8
AI %	0.510864865	0.007560712	0.44	0.45	0.424007015	0.179781949	22.26455153	3.809249708	4.27	0.02	4.29
As (ppm)	92.16947536	6.151963685	24	1	345.0039664	119027.7369	203.7577795	12.35748373	8189	1	8190
Ba (ppm)	426.7230525	6.438755978	350	170	361.0873641	130384.0845	10.81987192	2.629198158	3060	10	3070
Ca %	9.370346582	0.07646663	9.18	15	4.288271497	18.38927243	-0.885936138	-0.236083834	14.85	0.15	15
Cd (ppm)	1.441033386	0.052641837	0.5	0.25	2.952169984	8.715307615	43.95230863	5.829968718	36.75	0.25	37
Co (ppm)	3.888871224	0.087502502	3	3	4.907166518	24.08028324	28.27515863	4.581960924	64.5	0.5	65
Cr (ppm)	78.53481717	1.367946493	57	4	76.71484914	5885.168078	8.291673536	2.433661436	656	0	656
Cu (ppm)	30.07281399	0.502024881	18	6	28.15370574	792.6311468	2.929805855	1.627113965	239	0	239
Fe %	1.129812401	0.013411221	0.96	0.99	0.75210527	0.565662338	12.48492196	2.888084207	7.59	0.09	7.68
Hg (ppb)	2273.045787	129.6693573	340	110	7271.896405	52880477.33	81.30013375	7.840627399	99990	10	100000
K %	0.224798092	0.001898761	0.22	0.21	0.106483069	0.011338644	2.03208005	0.755443397	0.84	0.01	0.85
Mg %	3.462521463	0.03271443	3.83	1.38	1.834635037	3.36588572	-0.406984116	0.177762739	10.32	0.03	10.35
Mn (ppm)	24.26995231	0.351202369	19	16	19.69553406	387.9140619	21.51420584	3.596655988	260	2	262
Mo (ppm)	7.589189189	0.250840171	3	0	14.06719191	197.8858882	19.33749409	3.914505085	125	0	125
Ni (ppm)	31.5036566	0.609276551	21	6	34.16841152	1167.480346	10.36648721	2.827949797	269	1	270
P (ppm)	1276.220986	20.90053046	1000	280	1172.108009	1373837.186	12.4283097	2.762607355	10000	0	10000
Pb (ppm)	6.644833068	0.171422578	6	6	9.61342954	92.41802752	657.9602428	20.91320354	343	1	344
Sb (ppm)	22.31707472	1.887240383	2.6	1.8	105.837006	11201.47184	62.84281496	7.715552863	999.9	0.1	1000
Sc (ppm)	2.568839428	0.050488764	2	2	2.83142502	8.016967641	25.03946068	4.484313995	30	0	30
Sr (ppm)	260.0944356	5.925467857	112	95	332.3020125	110424.6275	2.871438229	2.000156731	1600	0	1600
TI (ppm)	5.289348172	0.077921407	5	5	4.369855869	19.09564032	1563.651122	35.51905834	205	5	210
V (ppm)	83.98314785	3.906315281	34	3	219.0673312	47990.49558	59.60932696	7.003627511	2749	1	2750
W (ppm)	5.73608903	0.116226302	5	5	6.518005837	42.48440009	688.8647622	24.27947159	225	5	230
Zn (ppm)	143.4015898	4.326822394	86	24	242.649496	58878.77792	40.97561781	5.632716953	2998	2	3000

Table 2 Correlation Matrix: SSX Project Area

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Mo				Ī												-	8.0	0.393 0.486	140.0	0.005	0.101	-0.17	0.01	0.826	0.002	0.853	QW W
Mn														Γ	-	6 F.O	0.142	013	0.213	0.11	0.426 0.101	-0.17	0.04	90.0	0.012	0.0	Mn
Ma														F	Q T	0.243	0.184		90.0	-0.16	0.075	0.41	89.0 P	,	드	0.181	Ma
×													-	0.02	0.249	-0.06 0.022 0.243	0.276 0.184 0.142	0 0.168 0.144	-0.04 0.142 -0.06 0.213 0.041 0.086 0.042	-0.17	-0.06 0.413 0.075	-0.33	90.0	0.491 0.047 -0.06 0.099 0.188	-0.03	0.05 0.164 0.562 0.099 -0.07 0.135 0.181	×
ЬH												F	0.14	-0.19	-0.11 0.249	90.0	-0.07	Q	0.04	0.329	90.0	0.0 9		90.0-	0.041	-0.07	Ha
Fe											-	0.02	0.547	-0.12	0.601	90.0	0.47	0.07		90.0	0.846	-0.36	0.054	0.047	-0 0.035 0.041	88	Fe
Cn										-	355 0.647	60.0- 280.	142 0.498 0.547	-0.13	0.371	189 0.496 0.036	0.705	0.42	0.268 0.189	0.15 -0.08	0.463	-0.36 -0.36	013 0.039 0.054 0.503	0.491	P	0.552	3
Š									-	0.445	0.385	0.085	0.142	-0.31	0.253	0.189	0.276	0.175	0.161	0.15	0.171	-0.42	0.013	131	0.024	0.164	১
တ								-	0.286	0.503	0.89	0.005	0.404	-0.14	0.625	-0.01	0.466	600.0	0.114	-0.07	0.884	-0.21	0.019	0.022	0.12	0.05	ප
ρ							1	0.02	-0.7 0.122 0.286	-0.57 0.475 0.503	-0.6 0.015	-0.02 0.005	-0.53 0.088 0.404	0.2	-0.09 0.525	0.83	-0.24 0.763 0.466	-0.08 0.431 0.009	-0.25 0.024 0.114		-0.32 0.094 0.884	-0.12	-0.05 0.005 4E-04 0.019	0.015 0.945 0.022 0	900	0.957	S
Ca						-	0.016	0.43	-0.7	-0.57	9.0-	0.052	-0.53	0.182	-0.37	-0.03	-0.24	-0.08	-0.25	5E-04	-0.32		0.005	0.015	-0.03 0.039 0.00 8	尀	င္မ
Ва					-	0.021	-0.12 0.016	90.0	-0.03	-0.18	-0.11	-0.06 0.052	-0.12	0.168 0.182	φ	-0.12	-0.15	-0.1	0.033	0.007	90.0 9	-0.15 0.667	900	<u>۲</u>	89	কা	Ba
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Ag		1	960.0	0.074	0.13	-0.14	-0.03 0.532	-0.01 0.749	0.284	0.369	-0.02 0.035 0.812		0.05 0.675	-0.02	Ģ Ģ	-0.07 0.445 0.034	-0.08 0.432 0.441						1		7 0 0 0 0	0.07 0.504 0.129	Ag
Au	1	0.254	-0.11	0.639 0.074	6 8	0.047	9.03	-0.01	0.05	-0.07 0.3 69	-0.02	0.722 0.242	-0.13	-0.13	-0.07	-0.07	-0.08	-0.04 0.417	-0.05 0.031	0.216 0.185	-0.06 0.012	8	0.556 0.181	200	9	8	Au ,
	Au (ppb)	Ag (ppm)	Al %	As (ppm)	Ba (ppm)	Ca %	Cd (ppm)	Co (ppm)	Cr (ppm)	Cn (bbm)	Fe %	Hg (ppb)	8	Mg %	Mn (ppm)	Mo (ppm)	Ni (ppm)	P (ppm)	┪		ᆿ	Sr (ppm)		(mdd) A	\dashv	Zu (ppm)	

Relationships to Geology

The correlation of certain elements is most likely related to lithology. Most notably, Sc, Co, Fe, Al, and Cu (± K, Ni, Cr, Mn) are most elevated in dike samples (Appendix 1). Other elements seem to be associated with structural features. Higher values of Zn, V, Cd and other base metals generally occur in the Roberts Mountains Formation and Snow Canyon Formation above the ore deposit, but high Zn and V values also coincide with northeasterly structures (fig. 10). The observed relationship between early base-metal sulfides and later marcasite may indicate a remobilization of Zn and other elements, with redeposition above the ore deposit along structures and within sedimentary rocks. The geochemical profiles of two drill holes show a typical relationship between Au and As within the ore zone (fig. 11).

Flow Paths

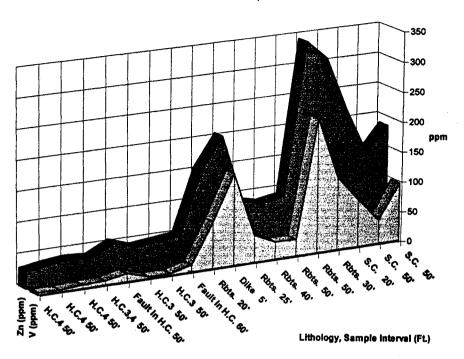
The direction of alteration intensity combined with drill-hole geochemistry may indicate direction of fluid flow. Specific drill-hole intervals were compared at the 7040' level to see if differences could be found with distance from northeasterly structures. The concentrations of Au, As, and Hg steadily increase with proximity to the northeast-trending faults (fig. 12). There is also a decrease in Ca and an increase in Mg toward the faults, as well as a depletion in Al concentration.

Discussion

So far, there is some evidence to suggest that hydrothermal fluids were transmitted primarily along northeast-trending faults. The best indication of this is the fact that ore zones and gold grades increase significantly with proximity to these faults. The older northwest-trending faults also have high grades, but the extent of the ore is much more constrained. Alteration effects are more prominent along the northeast-trending faults, and alteration intensity in the basalt dike increases with proximity to them. On a larger scale, ore grade-

Figure 10





SP-339C Zn, V

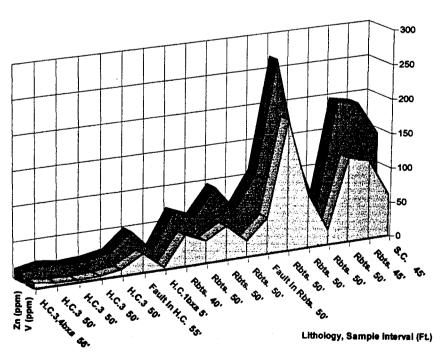
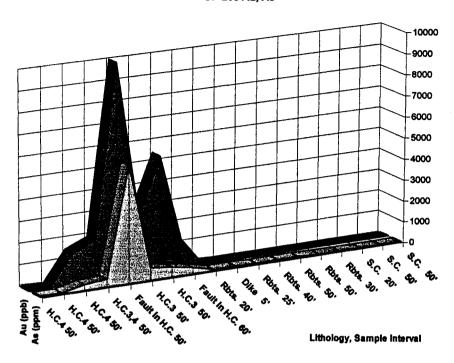
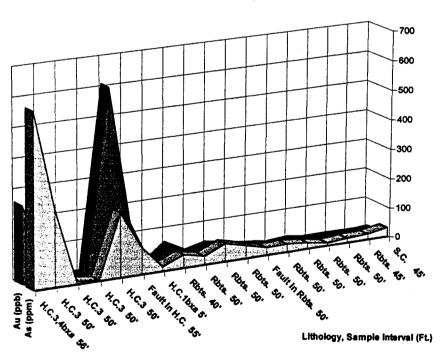


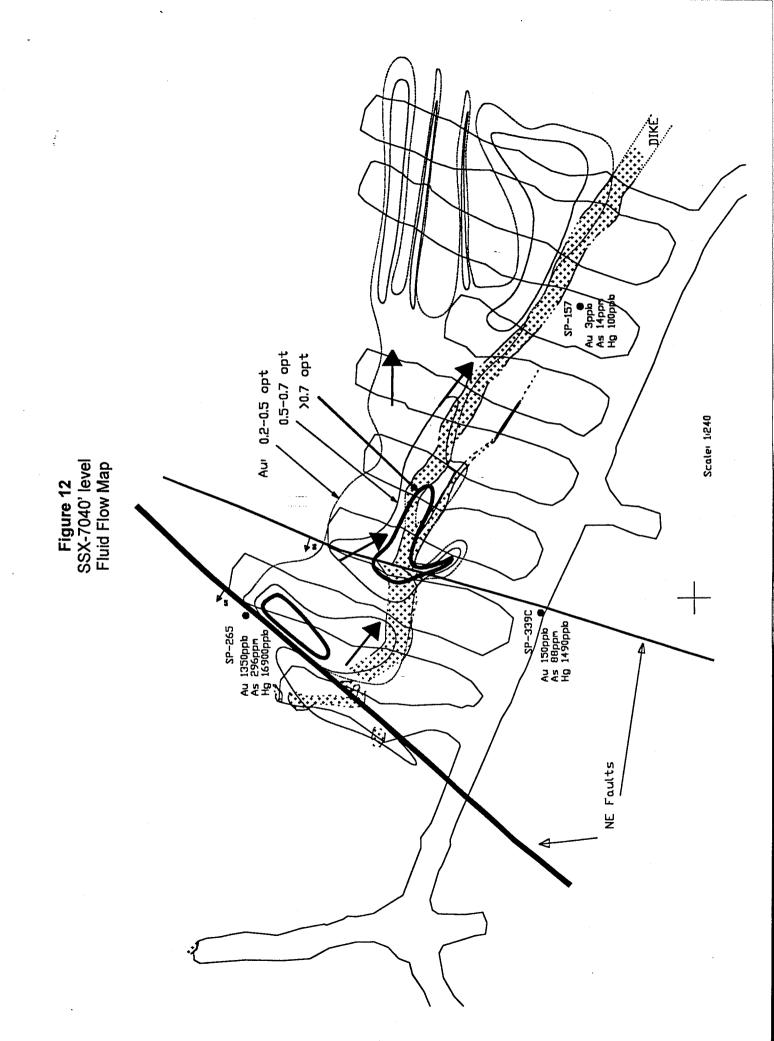
Figure 11

SP-265 Au, As



SP-339C Au, As





thickness maps for the SSX project area show elongate, northeast-trending ore pods that may lie along these structures where they cut favorable stratigraphy.

The hydrothermal fluids were acidic, as evidenced by the widespread occurrence of marcasite in the ore zone. Paragenetic relationships suggest a fluid that evolved from a relatively oxidized to a more reduced and acidic fluid. Correlation of certain elements may imply that the composition of the ore fluid was enriched in As, Hg, and possibly Tl. Also, the ore fluid was probably siliceous, as petrographic analyses indicate a close relationship between arsenic minerals and silica.

The conspicuous concentration of base-metals in the rocks overlying the ore deposit and adjacent to the northeast-trending faults may indicate remobilization and precipitation of these elements higher up in the hydrothermal plume. The presence of chalcopyrite and sphalerite outside of the ore zone in the basalt dike and observed replacement of these by later iron sulfides may support this idea.

Acknowledgments

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Appendix 1: SSX Geochemical Data For Surface Drill Holes Intercepting Mapped Workings

Zn (ppm)	122	154	108	230	12			1	1005	178		84	120	200	24	24		Zu (bbm)	32	280	44	22	48	24	18	4 0	20	20		(mod) u7	74	102	176	172	172	8 6	142	112	72	32	100	26	26	24	6.7
W (ppm)	5	5	2	2	0	0	0	0	0	9	0	9	0	0 4	מע	5		(mdd) M	0 4	10	20	10	20	230	30	0 0	10	10		(mod) vv	2	5	5	2	0	0 4	2	5	5	2	0	0 40	2	4	0
(mdd) A	37	42	24	62	13	200	207	720	60	78	900	28	70	38	1	9			0	18	108	21	100	12	1 02	-	101	9	_		22	30	90	85	31	40	107	53	90	= 1	1	2 6	4	a	9
(mod)	2	9	0	n	2	2	2 4	2 4	0 4	0 4	0	0	0	מיים	2	2	\rightarrow	-	200	5	5	5	2	2	0 4	0 4	2	S	1	V (1984)	2	5	5	0	0 4	2 4	2	5	2	0	0 0	2	2	¥	0
or (ppm)	61	72	200	120	000	404	440	188	000	200	0 4	000	800	311	875	648		1 June	00	559	428	1000	248	895	1386	1145	152	370	/man/	78	104	109	88	104	000	78	94	84	69	900	1120	1330	1265	1275	16/01
(mod)	2	2 0	2 0	26	2 5	4 6	2 6	1	-	26	2 6	3 6	20	0	2	2		ac (pam) ac	0	4	21	2	22	4 0	1	-	-	0	So (moon) Co		2	2	2	200	3 6	2	2	2	-	0	1	-	-	-	1
SD (ppm) Sc	1.8	7	7 0	0 8	-	-	6	18.5	28	1 4 6	-	- 1	2 0	12	5.2	30			20	12	32	48	12.5	4.0	200	12.5	80	11.5	Sh (mon) AS	_	1.4	1.4	2.4	7	28	3.2	3.4	4	5.4	4.1	0.8	0.4	0.4	0.8	200
(MATH) SC	4 0	20 0	000	0 0	4	00	8	0	9	0 0	100	100	7 0	14	10	89	Ob /month Ob	+	2	2	4	-	- 6	7	200	2	-	2	Ph (mm) Ph	+-	14	8	9	* '	100	4	9	4	4	4 0	2	-	-	-	
DIAM) FO	510	1380	1780	530	890	1140	1090	2750	730	780	800	1380	2880	270	270	370	/nom/	+	830	370	970	270	0801	210	180	230	2540	1620) HO (mout)	10	840	480	650	200	270	020	2120	000	070	280	200	210	220	250	2000
1		32								1		1	1		8	6	A (mon) IV	1	L		107			2 0			21	3	a) d (moon)		17	32	1	\perp	L	L				0 7	L		Ш	5	2
115/41/1		0 4	12	2 -	3	7	8	90	10	1	100	100	13.0	-	0	0	1 man	-	9	-			0				П	0	2		2	=	12	2 2	7	2	3	2	0 0	200	0	0	0	0	
	98	200	25	17	20	16	14	18	22	23	18	23	8	32	23	20	(max) (Mo (max)	15	3	23	90	35	35	28	19	20	15	89	m) Mo (nom)	-	34	47	30	29	20	18	14	13	4 6	21	22	20	23	20	-
2 2	00	1 25	020	16	60	37	84	52	81	12	4.4	54	79	5.87	3	1.12	18	31	93	24	2.97	48	26.0	38	28	38	82		(Whi (Dom)																
0	0 6	_	_	_	_							1	_	\perp			+	1			_	_	_		_			1.7	Mg °		3.21		-		1	- 1	- 1	- 1		1					
		0	0	0.22	0.2	0.2	0.2	0.2	0.2	0.2	0.35	0.32	0.13	-	Ц	0.2	% ×			0.21	_	0.2	-	-	Ľ		_	0.03	× ×	0.24	0.29	0.28	0.29	0 22	0.36	0.27	0.38	0.27	000	0.12		0.19	0.21	0.18	
1		180	ı				06	420	280	150	380	3150	9850	15100	37000	34800	Ho (nob)	3150	33200	8800	18800	2100	12900	9050	1100	4150	14900	2850	Hq (ppp)	310	06	180	190	190	830	760	780	1/30	8200	3050	410	100	120	140	
1 60	1 50	1 88	1 22	0.91	0.81	-	1.07	1.28	1.01	0.95	-	1.04	0.77	0.68	0.93	0.84	Fe %	0.91	0.54	0.9	4.28	1.05 at 4	1 08	0.81	0.51	0.55	0.44	0.14	$\overline{}$		1.	4 60	1.58	98.0	1.07	0.77	98 0	0 80	0.48	0.58	0.74	0.58	0.67	0.58	
70	81	44	30	15	11	21	25	87	22	13	11	18	12	8	18	21	Cu (nom)	18	13		84				-	_	9	1	(mdd)	82	\$	44	54	15	22	13	-	0 7	12	80	8	8	00 0	8	
	185	187	88	52	33	32	38	87	48	+	48	52	217	21	4	8	(mod)	70	415	48	180	175	21	18	17	14	59	77	(ppm) C	123	66	278	172	73	62	87	28	150	503	15	9	0	0	6	
7	1	8	-	-	0.5	-	2	3	2	-	2	-	2	0.5	4	4	(ppm) Cr	4	2	7	300	330	80	2	2	2	2	+	Ca (ppm) Cr	9	4 0	0 1	8	3	3	2	6	¥ -	-	-	2	2		2	
-	-	0.5	2.5	0.25	0.5	0.5	0.5	12.5	2	-	0.5	2	-	0.25	0.5	1.5	(ppm) Co (100	0.25	0.25	0.25	0.25	0.25	0.25	0.25	0.25	0.25	0.43	pm) Ca	2.5	0.5	1 5 0	-	0.5	-	-	5.5	2 -	0.25	0.25	0.25	0.25	0.25	0.25	
90	2.48	207	9.75	7.87	10.55	10.75	10.25	11.1	8.83	8.8	8.61	8.83	1.96	15	15	15	Ca % Cd(0.25	15	10.1	00 2	15	15	15	15	12	0	ŏ		5.54		5.1	8.98	8.08	7.92	108	7 44						13	
210	┖						_1	130						190	190	280	Ba (ppm) Ca			_		210	80	130	100	210	320	0	O		820					- 1					180	80	018	2	
18	18	48	30	34	14	4	12	44	24	26	58	178	330	232	520	2140			198	462	540	1880	1420	972	138	138	334	30		12	4 4	18	8	38	99	104	200	78	78	99	28	14	77	44	
0.5	0.57	0.74	0.49	0.45	0.45		0.49			0.54				0.14		0.3	Al % As (ppm)									0.15	1		A % As (ppm)	0.65	0.96	54	62	48	7.2	24	0.48	32	70	.21	0.5	0.28	0.33	0.48	
0.4	-	_		0.1						0.1			-	0.1	_	0.1	Ag (ppm) A	0.8			0				\rightarrow		0.8		-	-					0 8 0			0 8 0		$\overline{}$		-1.	0 0	+	
18	5	5	3	9	2	2	-	2	12	13	51	189	715	3190	3520	6160	Au (ppb) Ag		3580	884	0 =	30	5180	4270	578	52	31	1	Au (ppb) Ag	10	0 0	11	7	37	103	138	184	378	580	283	0	0 0	17	-	
				-	+	1	1	1	1	-					20.	+	Au		1	+	+	-			1	+	+	H	Au (+	+	-			+	+	+	-		Н	+	+	+	+	
S.C. 50'		. 80.	5. 40	\$. 50'		. 20,	s. 50°	s. 50°	\$ 50'	Rb(s, 50'	s. 50°		1 20.	H.C.2,3 50°	Fault in H.C.3	200	Lithology	, HO,	HC.1 20	H.C.Z.3 80	2 10	Dike 55	3 36.	HC3 50'	3 50.	HC3 51	H C 4 47				20,			. 70.	Rb(s. 50°	00	50.	. 25.	1 20.	H.C.2,3 50"	3 80.	740 M C 3 50	3 50	3	
50 S.C.	100 S.C.	160 S.C.	200 Rbts.	250 Rbts.	300 Rbts.	350 Rbts.	N Rtxts	N RDIS	500 Rb4s	O REA	00 Rbds.	30 Rbts.	5 H.C. 1			DI HICI		0	ZHC	7	D H C 3	7 DIM	2 H.C	2 H.C	2 H C 3	3 HC	200		LIB	20 8 0	0 0 0	OSC	0 S.C.	O Rtxs.	O REG	450 RMS	500 Rbts	5 Rbts	545 H.C.1 20"	HC	OI	2 2	HO	200	
0			- 1			_	_	-1	_1			0 680	0 715			800	-		_	7 047	_	1					1000	-	2	1					350	-									
	5(100	160	200	250	300	350	404	451	200	55(800	680	715	750	800	F	_	4	790	1	862	Ц	Ц	1	852	1	1	From	9	100	150	200	230	300	330	450	500	525	545	900	7007	750	2000	
\$P-362	362	\$P-382	\$P-362	-362	SP-362	362	\$P-362	382	362	\$P-362	\$P-362	362	\$P-362	362	SP-382	SP-302	SP-366C	\$P-366C	SP-388C	2000C-100	SP-388C	3880	\$P-386C	\$P-366C	888C	SP-388C	SP-388C		\$P-157	57	SP-157	57	191	22	12	67	SP-157	57	57	57	\$P-157	167	SP-157	123	

Zn (ppm)	184	142	220	310	342	080	000	304	162	50	00	40	42	40	36	38	000	07	70 (000)	100	472	108	9 9 9	114	82	498	1150	496	34	30	20	24	20	16	20	18		Zn (ppm)	142	190	081	424	124	408	90	98	58	70	18	90	24	18	10	77	4
W (ppm)	2	2	5	2	2	0 4	0 40	2	2	a ur	3	0	0	10	0	0	מע	0	W (norm)	2	2	2	2	2	2	2	5	2	2	5	5	5	5	5	2	0 0			2	0	0 4	2 4	0 4	2	2	2	140	5	5	5	2	2	0	0 4	ō
	87	40	80	121	223	200	45	152	78	121	3 6	36	7	-	9	0 0	7		J (moon) V	140	44	54	8.5	48	20	249	880	274	13	+	3	3	2	2	7	3 6	1	\rightarrow	90	113	27	78	000	50	23	48	32	45	3	27	8	4	7	r a	R
	10	0	20	0	0 4	0 4	2 10	vo.	20	2	10	4	300	30	2 4	0 4	o ur	,	V (mon) II	-	2	45	2	20	90	20	2	5	5	2	5	S	2	2	5	מינו	Н	V (mod) II	0	0 4	2 4	2 40	2 4	2 4	NET.	100	10	2	5	50	2	2	0 4	2 4	c
Sr (ppm) T	88	A C	104	203	747	83	75	66	83	163	815	0 0	848	345	237	582	225		Sr (norm) T		70	81	159	108	88	104	121	80	210	572	1050	1185	978	1140	281	758			80	7 0 0	104	108	132	115	89	113	121	78	1445	15	388	1315	1330	128	1071
Sc (pom)	7 0	2 6	2 6	2	20	8	2	4	2	-	-	-	8	200	40		0		Sc (non) S		3	3	3	3	2	2	3	2	-	1	-	-	-	-	0	0		Sc (ppm) S	2 0	26	26	3	100	3	2	3	2	2	-	0	-	-	1	+	
(mod) oc	2.4	2 0	50	7.0	0.0	1.8	2.2	4	2.8	9.4	1.8	2.4	13.6	40	30	7.4	120		Sb (pom) S		2.8	2.4	3.2	2.8	1.4	8.4	17	7.8	7.4	0.8	0	0 1	0	0.2	21	0.2		1	2.0	3.4	-	2.4	8.4	2	1.6	1.8	1.4	1.8	0.1	2.8	8.8	0	4.4	CA	174
(mod) ou	0 0	9	e a	0 9	0 8	4	9	8	2	2	-	-	-	-	Ī	-	-		Pb (ppm)		80	80	9	9	4	4	8	8	2	2	-	-	-	1	1	4			0 4	9	10	8	80	8	8	8	4	4	2	-	7	7	2	2	. 7
T (Dom)	1110	2180	2180	1830	880	850	770	1980	1810	580	230	240	550	1350	2210	1520	830		P (mod) P	_	1110	1830	2270	1210	810	2540	890	1340	330	220	180	220	210	077	1380	1430	-	(ppm) Pt	3070	2810	410	950	2280	1390	390	800	930	1370	170	680	330	230	270	2980	2000
	37	44	80	B.S.	100	38	18	48	24	12	8	8	38	10	7	2	3		M (ppm) P	38	38	48	99	32	24	88	154	84	7	8	7	2	0	1	• •	4	- 0	N (mom) P (82	55	18	22	58	31	18	21	14	18	7	10	-	0 40	1	28	-
12	1	12	28	41	2	-	-	2	-	3	o	0	0	2	0	0	-		Mo (ppm) M	7	7	Ġ	22	6	3	22	83	22	7	ò	0	0	0 0	5 .	2	0	_	Mo (ppm) M (28	22	3	7	20	5	3	3	-	2	0	0,	- 0	00	0	2	
	22	19	18	101	21	28	24	21	23	4	20	21	29	18	8	10	18				32	25	15	15	17	36	18	35	24	21	18	22	47	1	3	2		Min (ppm) Mo (25	77	23	21	23	18	21	21	13	Z (21	* 00	23	22	18	14	
0.78	1.53	2.07	4.05	5 25	4.31	4.14	4.75	5.45	90.5	2.14	1.23	1.33	1.04	2.03	2.01	1.74	1.86			96.0	1.51	1.38	5.14	4.6	4.31	5.94	699	4.87	6.54	0.92	1.25	1.38	1.38	4 1 0	13	1.37	AAcr 04. NAm In		573	5.21	5.37	5.48	5.7	5.21	4.39	5.54	3.85	4.33	1.47	20.7	34	35	1.3	2.33	
0.38	-	0.38				0.27			_	_			0.28	0.17	90.0	0.07	0.03		×	0.27	_	0.27	_	-	4	-	4	_	_1	0.12	-1	0.10	200	000	0.07	0 11	K % M		0.24				0.33			_	_	_	1200	_	0.28	0.22	0.21	0.14	J
380	180	200	300	400	280	280	840	4110	2850	16900	37000	18800	89000	12800	9180	2050	2520		Hg (ppb)	300	230	370	330	180	120	310	890	830	2200	280	002	330	210	3850	150	120	Ho (ent)		390	390	00	160	400	280	300	480	820	1020	2400	3310	40	40	1330	8480	
3 1.89	-	-	-	-	_	_	4 0.78	_	-	-	0.68	0.68	-	0.75	9.0	0.41	0.3	\rightarrow	Fe %	-	-	-	-	0.99	-	1.03	_	_	0.78	-	_	0.00	_	1	-	0.33	1; w 9%	154		0							0.68	6/0	000	0.71	0.65	0.61	0.63	0.87	
63							14			13	=	11	45	1	8	4	3		Cu (ppm)	85	62	75	47	28	20	40	84	44	10	20 1	0	0	7	8	7	80	Cu (nom)	56	50	45	18	25	45	30	18	22	18	200	7 11	12	-	12	14	18	
181	122	118	53	48	39	42	38	31	32	78	3	3	24	43	22	15	37		Cr (ppm)	88	73	88	35	32	28	31	4	3/5	0	0 0	26	0 4	3	12	8	10	Cr (porn)	-	90	51	46	48	40	53	44	44	28	7.	230	25	2	3	3	58	
7	7	9	S	2	3	10	2	-	9	7	3	3	13	3	2	-	0.5			0	В	0	*	2	3	2	200	1	1	200	40	200	2	0.5	0.5	0.5	Co (pom) C		4	3	3	3	3	3	3	6	70	40	0.6	2	2	2	2	9	
	0.5	1.5	3.5	4.5	0.5	-	-	4 (700	6.0	0.75	0.25	0.5	0.75	0.25	0.25	0.25		(bbu)	0.25	-	6.0	1.5	- 0	0.5	4.5	011	0 00	0.70	0.75	0.36	0.25	0.25	0.25	0.25	0.25	(total)	1.5	2.5	2	0.5	2	*	-	1	- 0	60	30.0	0.50	0.25	0.25	0.25	0.25	0.25	
1.55	2.98	4.57	8.99	12.4	7.86	7.54	8 27	200	9.07	17.8	12	15	15	13.55	15	15	15	_	Ca % Cd	1.86	274	7/4		8 52	/2/	0.04	10.00	80.00	13.6	4,5	14	15	15	15	15	15	9 % Cd	2.91	9.72	10.5	8 26	8 46	80.6	8.69	7.47	9770	8 31	15	1	14.6	15	15	15	14.3	
270	200	190	920	870	510	420	980	480	070	210	410	180	110	220	670	250	110			70	06	00	410	320	070	040	300	270	200	170	140	280	110	200	80	120	Be (ppm) C		900	640	720	- 1	- 1	- 1	760	-1	240	330	370	400	180	170	880	450	
14	12	8	8	18	52	52	38	60	900	DAZ	480	584	5530	388	288	48	90	_		36	040	900	200	87	2 5	000	000	240	138	42	14	38	24	2540	18	20	(ppm) Be	30	18	14	8	20	28	24	42	000	44	14	88	214	2	8	258	588	
0.79			0.62	0.53	0.49	90	0.43	20.0	24.0	0 0	0.17	0.2	0.42	0.25	0.1	60.0	0.04	7		0.52	0.50	10.0	0.30	0.37	0.30	0.35	0000	0 10	0 18	0 0	0.28	0.18	0.13	0.04	0.13	61.0	1 % A	0.93	0.52	0.58	0.61	0.63	0.91	0.8	0.49	86.0	0.83	0 27	0.18	0.29	0.36	0.29	0.3	0.28	
9.0	0.4	0.0	0.8	-	0.1	0.2	0.2				0		0.1			0.1		$\overline{}$	Ag (ppm)	0.4	000	0		0		0.4		4 +		0		0.1			_		(mdd)	0.0	0.4	0 0	0	0	0.8	0.4	0 0	200	0 0	0 1	0.1	0.1	0.1	0.1		0.2	
	89	7	3	3	25	7	32	74	1350	0000	2400	2800	10000	1800	1450	69	105	-	-	200	0 0	2 0	200	7 8	0 00	707	26	338	4	0	0	12	18	86	9	3	Au (pob) Ag	0	-	0	0		*	2	28	40	55	0	150	610	0	0		760	-
						1	1	T	, AO.	200	1		. 20.		1	1	1	ľ	I	1	1	1	1	1	1	t			-			-				1	×.		1	1	1	t		00.3	+	t	t		.99			+		.96	
C 50	.C. 50°	.C. 20	bis. 30.	Ms. 50'	Dis 50.	Ebts 40	Dike 5	10c 20'	Fault in H C	C 2 60	2000	0 H C 3 50	auf in H.	HC34 50	C 4 50	H.C.4 50°	C.4 50	1	Limology	200	200	200	Ma. 60.	145 50°	200	Ma 50'	Me 25.	623 24	HC3 50	C 3 50'	C3 50	C3 50.	C.3 50.	C.4 36	H C.4 50"	C.4 25	Lithology	SC 45	78. 45.	50.	00 60	00 80	00.50	SOO PAUR IN KOS	30 SO	de 50'	Rbfs 40°	H.C. 1bxa 5	ult In H.C.	03 80.	C.3 50°	HC3 50	23.30	H.C.3.4Dxa	
50	100	120	150	200	200	540	570	600	ASO		7	750	800	850	900	950 H	1000	-	2	n S	1 4	300	380	305	38.	And	40.8	500 H	550 H	800 H	650 H	700 H	740 H	850 H	H 008	925 H	10	55	100 R	150 R	K 00 A	DOC 2	300 80	3000	450 64	500	540 Rt	545 H	600 Fa	650 H	700 H.	750 H	900	820 11	
0	20	100	120	150	450	200	546	670	500	000	000	700	750	800	850	006	DC#	1000		40	100	144	2000	250	300	380	400	475	2009	550	900	650	700	815	850	006	From	10		-1		1	007	340	400	450	9009	540	545	000	650	700	000	800	
SP-265	-285	-285	-285	-285	-285	-285	SP-405	2.285	2965	2002	202	-265	-285	SP-285	-265	-285	507-	Т	SP-243	OD 243	243	SP. 243	242	SP.743	243	SP.743	.243	243	SP-243	243	SP-243	243	243	243	SP-243	243	3390	330C	339C	3380	2000	2000	2000	3300	3300	3390	339C	339C	339C	33BC	339C	SP-339C	CORC	2285	-