UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

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DISTRIBUTION OF GOLD AND OTHER ORE-RELATED

ELEMENTS NEAR ORE BODIES IN THE

OXIDIZED ZONE AT GOLDFIELD, NEVADA

Ву

R. P. Ashley and J. P. Albers

Open-file report

This report is preliminary and has not been edited or reviewed for conformity with Geological Survey standards

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Distribution of gold and other ore-related elements
near ore bodies in the oxidized zone at Goldfield, Nevada

By R. P. Ashley and J. P. Albers

Abstract

The heart of the Goldfield mining district occupies 0.6 square mile within a 15-square-mile area of hydrothermally altered Tertiary volcanic rocks. Most of the ore shoots were irregular bodies of epithermal bonanza ore within a few contiguous silicified zones enclosed within clay-bearing altered rocks. In 1966, 278 samples of argillized and silicified dacite were collected from excavations at the Combination and January mines, which once yielded gold in commercial quantities. Semiquantitative analyses show that gold, silver, lead, bismuth, mercury, and arsenic are notably enriched in rocks of the cuts. All these elements except lead and mercury formed conspicuous ore minerals. Geochemical maps and one geochemical profile across strike show that relatively high concentrations of all these elements are restricted to silicified zones. This low-tenor metallization dispersed through silicified zones does not extend into adjacent clay-bearing rocks. During oxidation arsenic, copper, molybdenum, and zinc were more or less strongly leached from the silicified zones and the ore bodies within them, but these metals did not form distinct

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supergene halos in the surrounding argillized rocks. From the semiquantitative data available, the average amount of gold in 2 silicified vein material is between 2.4 and 3.8 ppm. 5 --8 10-11 12 13 14 15--16 18 19 20-21 22 23 24 25--

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We recommend that geochemical sampling to detect relict hypogene dispersion patterns in the Goldfield altered area, using oxidized rock samples, be restricted to the silicified zones. If only a few samples are collected from each silicified zone, analysis for lead, of all the elements tested besides gold, is most likely to detect significant gold metallization, even though the lead was only a minor constituent of the ores. Abundance of iron oxides is not a reliable guide to anomalous amounts of gold.

Introduction

General

This report discusses distributions of ore-related elements, particularly those closely associated with gold, in the oxidized zone in the vicinity of a mined gold-bearing vein at Goldfield, Nev. The work described herein is part of a broader study treating the geology and geochemistry of hydrothermally altered rocks in the vicinity of Goldfield.

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Figures 1 and 2 near here

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it is a small part of a 15-square-mile area of hydrothermally altered
Tertiary volcanic rocks extending to the east and north, termed the
"Goldfield altered area." The main district lies at the western
margin of the Goldfield Hills, a group of peaks with maximum relief
of about 1,200 feet, nearly surrounded by desert basins. Maximum
relief in the main district is only 180 feet, with elevations ranging
from 5,640 to 5,820 feet. The climate is arid; vegetation, sparse.
U.S. Highway 95 passes through the town of Goldfield. An all-weather
gravel road skirts the western and northern sides of the main district,
and several dirt roads traverse it.

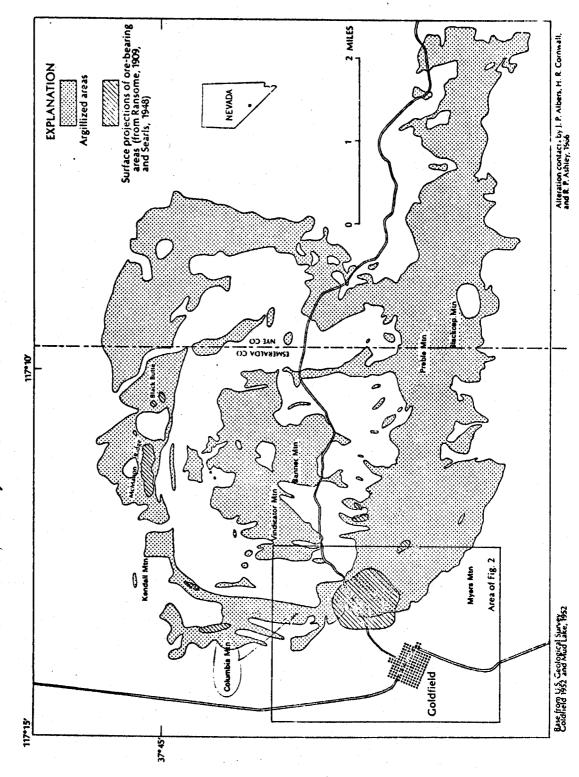
Most of the gold ore produced near Goldfield came from a 0.6-

(figs. 1 and 2). This area will be referred to as the "main district";

square-mile area immediately northeast of the town of Goldfield

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Fig. & - Map of the Goldfield distorct, showing areas of hydrothermal alteration and ore deposits.



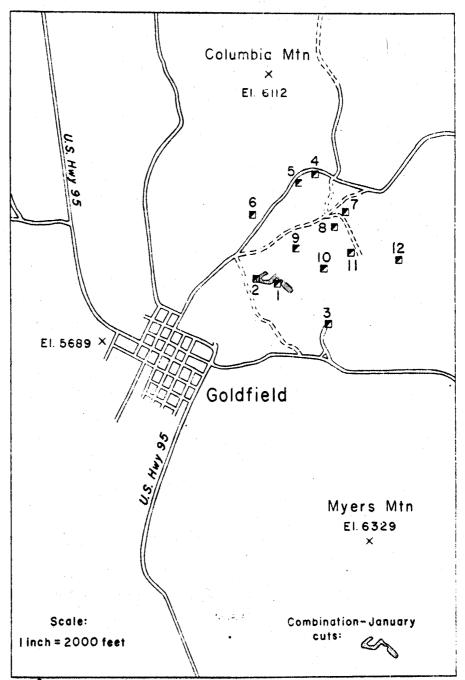


Figure — Map of Goldfield, Nevada, and vicinity, showing the locations of major mines and the Combination—January cuts. 1. Combination mine. 2. January mine. 3. Florence mine. 4. Laguna mine. 5. Red Top mine. 6. Silver Pick mine. 7. Jumbo Extension mine. 8. Clermont mine. 9. Mohawk mine. 10. Jumbo mine. 11. Grizzly Bear mine. 12. Merger mine.

In March 1966 the Davis-Goldfield Mining Corporation completed two 2 deep exploratory cuts in the main district which provided most of the samples for this study. These open cuts lie just east of the Combination and January shafts, along the vein system common to these two mines (pls. 1 and 2). The open cut walls, 15 to 50 feet high, expose many tunnels and stopes developed upward from the 80-foot level of the Combination mine, but the mine workings are almost completely inaccessible. Three months after excavation ceased, 278 samples for geochemical analysis were collected from the cut walls at 5-foot 10intervals. The cuts were mapped and sample locations were determined i 1 by planetable methods. The planimetric maps of the cuts (pls. 2-13) 12 do not show elevations and contours, because the land surface around the cuts and the cut floors both have relatively little relief, 14 whereas the intervening cut walls at the time of sampling and mapping were very steep, representing elevation changes between 15 and 50 16 feet along any given profile across the cut wall.

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The main objective of this study is to identify indicator elements suitable for geochemical exploration for gold; each element must be evaluated with the following requirements in mind. Ideally, amounts of an indicator element should correlate well with amounts of gold, and the range of values should be detectable by a reasonably inexpensive analytical procedure, with few samples falling below the detection threshold. Data also should be subject to less sampling error than gold data, and the element should form a dispersion aureole or halo larger than its associated economic gold deposit. These requirements imply that both the indicator element and gold were concentrated by means of the same processes. The 278 samples from the cuts and much smaller numbers of average-grade ore, high-grade ore, and unoxidized altered rock samples were studied to find indicator elements for gold.

The assumed exploration target is another heavy metals deposit

The essumed emploration target is another heavy metals deposit
in the vicinity of Goldfield like that mined in the main district.

Since a genetic relationship exists between hydrothermal alteration
and ore deposition, the entire 15-square-mile altered area, which
includes the main district, has potential for new deposits. The

Combination-January cuts expose oxidized altered and low-tenor metallized
rocks, so data from them should show geochemical relationships that

in bedrock emples collected from the surface
will be found throughout the altered area freewided that subsequent
geochemical exploration is limited to bedrock samples.

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Wilson (1944) evaluated several elements as indicators for gold in the Goldfield district. He showed that silver, bismuth, and tin are positively correlated with gold in the Goldfield Consolidated main vein on the 830-foot level of the Jumbo Extension mine (2,400 feet northeast of the Combination shaft, see fig. 2 and Searls, 1948, pl. 2). Although he did not find a clear relationship between gold, silver, bismuth, and tin in and near the Clermont vein on the 225-foot level of the Clermont mine (2,000 feet northeast of the Combination shaft), or in two surface localities, including one on the Jumbo vein (Jumbo mine, fig. 2, 1,200 feet east-northeast of the Combination shaft), Wilson concluded that bismuth and silver are promising indicator elements for gold. Since his samples yielding recognizable element correlations were entirely from unoxidized rocks, his results apply most directly to underground exploration. Except for preliminary results of this study (Ashley and Albers, 1969), no other reports concerning indicator elements for gold at Goldfield have been published

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In the first section of this report we identify potential indicator elements for gold. These include the following metals associated with gold in the Goldfield ores: copper, zinc, arsenic, antimony, bismuth, tellurium, mercury, lead, molybaenum, tin, and probably selenium. In the second section we examine geochemical maps of the Combination-January cuts for gold and for the seven of the above elements for which we have adequate data. In the third section, we attempt to determine which elements owe their spatial distribution in the cuts primarily to hypogene dispersion, and which owe their distributions largely to supergene processes. We also attempt to explain various interesting features seen on the geochemical maps, and various associations between the ore-related elements. The final section draws together information from the preceding sections pertinent to geochemical prospecting in the Goldfield altered area.

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The cooperation and assistance of Davis-Goldfield Mining
Corporation, owner of the investigated ground, made this project
possible. All underground information on the Combination and January
mines is from level maps compiled by Goldfield Consolidated Mines
Corporation (unpub. data), now held by Davis-Goldfield Mining
Corporation. Mr. M. G. Martin was particularly helpful in providing
these maps. D. H. Whitebread and L. D. Schultz assisted in geologic
mapping and sampling of the cuts. The late Martin C. Duffy, owner of
the Florence mine, conducted us through that mine, allowed us to map
and sample the limited workings still open, and discussed with us the
history of the mine. Mary E. Ashley coded the geochemical data for
computer input.

Geologic setting

The geology of the Goldfield area has been described by Ransome (1909, 1910a, b), Locke (1912a, b), Searls (1948), Albers and Cornwall (1968), and Albers and Kleinhampl (1970). The main district is at the western margin of a Tertiary volcanic center composed of silicic and intermediate tuffs and volcanic breccias, and rhyolite, quartz latite, trachyandesite, and rhyodacite flows. These volcanic rocks cover Ordovician metasedimentary rocks and Mesozoic granitic rocks that crop out in many small inliers to the north and northeast of the main district.

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The altered area and the position of the main district within it are shown in figure 2. At most localities the edge of the altered area shown in the figure represents the boundary between fresh rock and rock strongly enough argillized to be visibly bleached and locally stained with limonite. At some localities such as Blackcap Mountain, however, altered rocks are covered by younger unaltered volcanic rocks or alluvium, in which case the edge of the altered area is actually a contact with overlying materials rather than an alteration contact. Argillized rocks represent the bulk of the material within the altered area, but many silicified zones (veins) also appear, always surrounded by argillized rocks. The silicified zones are localized along, and delineate, the faults and fractures that served as conduits for the hydrothermal fluids that produced the alteration. Along the south side of the area, from the main district through Preble Mountain and continuing to the east edge of the map, these faults and fractures are very numerous, and trend northwest to nearly east-west, with steep dips both to the north and south. Alteration in the central part of the area, in the vicinity of Vindicator and Banner Mountains, was controlled by northeast-trending, east-dipping shingle faults. The fault blocks dip west, and are successively downdropped to the east. The west and north sides of the altered area are defined by a belt of altered rock which extends from the main district northward through Columbia Mountain to Kendall Mountain, then eastward through Black Butte. This belt reveals an arcuate structural pattern, with most of the faults and fractures alined approximately parallel to the inner margin of the belt. Although

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1 the faults here do not dip consistently inward toward the Vindicator Mountain-Banner Mountain area, the arcuate pattern suggests that ring fracturing occurred during the Tertiary volcanism, possibly accompanied by collapse, thereby forming at least a partial caldera (see Albers and Kleinhampl, 1970). 5 — 10-11 12 13 14 15-16 17 18 19 20-21 22 23 24 25-

9a (p. 10 follows)

The ore deposits of the main district were irregular pipes and sheets within seven or eight vein systems composed of silicified rocks, striking north and dipping at moderate to low angles to the east.

Dacite, andesite, and latite were altered to form these silicified zones and the argillized rocks which surround them.—

The

The terms "dacite," "andesite," and "latite" are the names Ransome gave to the three volcanic units that dominate the Tertiary section in the main district. We retain Ransome's nomenclature for this report.

By current volcanic rock classification systems (Rittmann, 1952; O'Connor, 1965), the dacite is a rhyodacite; the andesite includes both trachyandesite and rhyodacite flows; and the latite is a quartz latite.

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Combination-January, the westernmost of the major vein systems, is shown on level maps of the Combination and January mines (pl. 1).

The first level of the Combination mine was 80 feet below the shaft collar, at an elevation of 5,650 feet. Since the open cuts intersect workings that were reached from this level, a generalized geologic map of the cuts and adjacent surface has been substituted for the map of the first level. Early in the development of the properties, the workings of the Combination and January mines were joined; the Combination shaft serviced workings of the January throughout most of the history of production. Consequently, many workings of the January are accordant with those of the Combination, and different maps are not required for the two mines.

Dacite is the most widespread rock type in the two mines, occurring on all levels, but latite increases at the expense of dacite with increasing depth (pl. 2). Dacite rests directly on latite, and abuts andesite just east of the mine workings. Latite, andesite, and dacite form a stratigraphic sequence from base to top 3 miles east of the town of Goldfield; in the Combination-January area the dacite either intrudes the andesite, as suggested by Ransome (1909, p. 79-81) or is interlayered with the andesite, as suggested by Searls (1948, p. 11, 12). Geologic mapping elsewhere in the Goldfield mining district indicates that Ransome's interpretation of the dacite as locally intrusive into the andesite and latite is more likely correct. Locally on the west side of the open cuts, as much as 15 feet of sedimentary breccia of the Siebert Formation unconformably covers the hydrothermally altered volcanic rocks of the cuts (pl. 2). Mine dumps and 2 to 5 feet of alluvium cover much of the surface around the cuts (alluvium not shown on pl. 2). The positions of contacts shown on plates 1 and 2 are inferred where they are covered by these postalteration materials.

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The most conspicuous feature of the Combination-January vein 1 system is the abrupt change of strike at the January shaft from N. 60°-2 70° W. for the southern part of the system to N. 50° - 65° E. for the northern part. The acute angle thus formed persists to the second (130-foot) level, but opens progressively on the third (180-foot), 5 – fourth (230-foot), and fifth (280-foot) levels, and is not a notable feature on the sixth (380-foot) level (see pl. 1). The arcuate vein that passes through the January shaft at the surface and on the second level is nearly vertical at the shaft, but dips northwest north of the shaft and northeast south of the shaft. In the vicinity of the 10shaft, the dip of the vein decreases below the second level, reaching 11 65° E. on the fourth level. The January shaft and the vein both 12 continue to the fifth level, but these workings are not included on 13 the Goldfield Consolidated Mines Corporation maps, so we do not know their full extent and have omitted them from plate 1. This western 15part of the vein system pinches out below the fifth (280-foot) level. 16 The eastern part of the vein system dips steeply in the area east and 17 south of the Combination shaft, but northwest of the shaft, on the inside or eastern side of the sharp bend, dips are 50°-30° E. 19 decreasing with increasing depth. -/ The eastern and western parts of 20-21 -Footnote near here 22 23 the vein system are close together near the surface, whereas separation 24 is maximum at the fourth level. On the fifth and sixth levels all vein

material shown belongs to the eastern part of the system, which 1 continues to a maximum depth of 440 feet below the Combination shaft collar, corresponding to a minimum elevation of 5,290 feet. Most of the stoping was done from locations at the surface at elevations as high as 5,710 feet, to a point 330 feet below the Combination shaft collar, at an elevation of 5,400 feet. Ransome includes the western part of the vein system in his description of the January mine (1909, p. 216-220, pl. XVI), and includes the eastern part of the system in his description of the Combination mine (1909, p. 209-216, pls. XVII, XVIII). 10-11 12 13 15--16 17 18 19 20-21 23 24 25

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(p. 12a-1 follows)

1 5 -The ore bodies lying in the eastern part of the sharp bend had 10not yet been discovered at the time Ransome examined these mines. 11 Consequently we have no detailed information for this part of the 12 area on the distribution of silicified rock. It is likely, however, 13 part of the that the eastern vein extends northwestward into this ground and that 14 it includes these ore bodies. 15-16 17 18 19 20-21 22 23 24 25-

The shapes and orientations of veins belonging to the Combination-January vein system are probably controlled mainly by prealteration fractures. Also, on the fourth, fifth, and sixth levels, some veins conform to the shape of the dacite-latite contact. (See discussions by Ransome, 1909, p. 211-212, 217-218; and Locke, 1912b, p. 844, on form of the Combination-January vein system.) At many localities the veins have been fractured and brecciated due to post alteration movement, but in most cases these displacements are too small to significantly change the shapes of the veins.

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The Combination-January open cuts closely follow the Combinationbut they expect strengthers other than The weeks Themselves. January vein system (pls. 1 and 2), A conspicuous structure seen in the cuts is the set of northeast-trending faults exposed 160 feet northwest of the Combination shaft. Neither this set of faults nor the parallel-trending but southeast-dipping fault located 320 feet west-northwest of the Combination shaft, can definitely be identified on the second level or deeper levels. Possibly the dip of the northwest-dipping fault system progressively decreases with depth; if so, it could connect with the northeast-trending, 30°-northwest-dipping fault seen 250 feet northwest of the Combination shaft on the second level, before dying out at greater depth. The 140-foot-wide block between the two northeast-trending faults with opposing dips may be down-dropped, offsetting the silicified zone segment between the fault\$ to the west. Comparing the map of the cuts with the maps for the second, third, and fourth levels, however, the silicified zone lying east of the conspicuous set of northeast-trending faults certainly must be part of the zone that passes near the Combination shaft, and the silicified zone lying to the west of this set of faults certainly must be part of the zone that passes through or near the January shaft Since these two silicified zones appear to be separate at depth, they are likely also separate at the level of exposure represented by the cuts, implying that only small displacements are associated with the northeast-trending faults.

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(p. 16 follows

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Analytical methods

Gold values for samples from the Combination-January cuts were

Ine terms "value," "amount," and "concentration," as used in this report, mean quantity of an element expressed as weight per unit weight of rock. The specific units used in this report are percent and parts per million (ppm).

determined by atomic absorption spectrophotometry using hot hydrobromic acid extraction from 2-gram samples (Huffman and others, 1967). Three other groups of samples were analyzed by the cold hydrobromic acid-bromine method of Thompson, Nakagawa, and VanSickle (1968). Tellurium and zinc concentrations were also determined by atomic absorption spectrophotometry (Nakagawa and Thompson, 1968; Ward and others, 1969, p. 20-22). Antimony concentrations were determined by a solution-colorimetric method, and arsenic concentrations by the Gutzeit-apparatus confined-spot colorimetric method (Ward and others, 1963, p. 38-44). Mercury concentrations were determined by the atomic absorption technique described by Vaughn and McCarthy (1964), and Vaughn (1967). All other elements, including silver, barium, beryllium, bismuth, cobalt, chromium, copper, lanthanum, manganese, molybdenum, niobium, nickel, lead, tin, strontium, vanadium, yttrium, iron, magnesium, calcium, and titanium were determined by 6-step semiquantitative spectrographic analysis (Ward and others, 1963, p. 91-94; Grimes and Marranzino, 1968).

R. L. Miller, E. E. Martinez, F. Michaels, T. A. Roemer, J. A. 1 Thomas, J. D. Mensik, W. D. Goss, G. T. Burrow, G. D. Shipley, and 2 C. Huffman carried out the gold analyses. The analysts for tellurium were H. D. King and E. E. Martinez; for zinc, G. W. Dounay; for antimony, H. D. King; for arsenic, A. L. Meier, Z. Stephenson, and 5 -W. Campbell; for mercury, W. W. Janes, J. James, S. Noble, J. G. Frisken, and W. Campbell. A. W. Helz, W. B. Crandell, J. L. Harris, 7 H. W. Worthing, C. Heropoulos, H. Bastron, E. L. Mosier, J. M. Nishi, and J. L. Finley made the spectrographic analyses. 10-11 12 13 14 15-16 17 18 19 20-21 22 23 24 25

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Statistical methods

Element abundances were calculated using techniques described

-/"Abundance," as used in this report, means average concentration or weight proportion of an element in a given specimen or body of rock. Here it is calculated essentially by averaging analyses for some number of specimens of a given rock type.

by Miesch (1967). Frequency distributions for gold, mercury, arsenic, and zinc show moderate to strong positive skewness, so the data were transformed to common logarithms of the concentration values; a better statistical estimate of abundance is possible if the frequency distribution is relatively symmetrical. All other elements were spectrographically determined, with geometric reporting intervals, making it mandatory to convert the data for these elements to common logarithms to provide statistics valid for intercomparison (see Miesch, 1967).

The class intervals used for the gold, mercury, arsenic, and zinc histograms given on plates 4, 8, 9, and 12 were determined using a formula based on Sturges' rule (Sturges, 1926). For all other elements,

 $-/c_{i} = \frac{\text{MAX}_{i} - \text{MIN}_{i}}{\text{NM}_{i}}, \text{ where } c_{i} \text{ is the class interval, or class size,}$ $\text{MAX}_{i} \text{ is the largest data value, MIN}_{i} \text{ is the smallest data value, and }$ $\text{NM}_{i} \text{ is the number of classes.} \text{ NM}_{i} = 2.5 + 1.442726/\ln(N_{i}), \text{ where N}_{i} \text{ is the number of data values.}$

class intervals are equivalent to the geometric reporting interval used in 6-step spectrographic analysis: the sixth root of 10. Results are reported to the nearest number in the series 1, 0.7, 0.5, 0.3, 0.2,

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0.15, 0.1, etc. Analytical error (one standard deviation) is approxi-
    mately plus or minus one reporting interval.
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In order to compare abundances of various elements in various data sets, relative abundances were calculated as follows: First, arithmetic means and standard deviations were computed for each element using all data (transformed to common logarithms) within the analytical detection limits. The antilogarithms of the means thus calculated are geometric means. Some elements have "censored" frequency distributions: here some fraction of the observations fall below the lower analytical detection limit or above the upper analytical detection limit. For these elements, geometric means are too high if the censored data falls below a minimum detection limit, or too low if the censored data falls above a maximum detection limit. Using Cohen's method (see Miesch, 1967; Cohen, 1959, 1961), these means and their associated standard deviations were revised. Cohen's method assumes that the data outside the censor point would, if known, complete a normal (in this study, lognormal) distribution when combined with the known portion of the distribution. The revised geometric mean and standard deviation calculated by the method are those of this ideal complete distribution. Use of Cohen's method for log-transformed data, therefore, assumes that the frequency distribution is lognormal. Many of the elements have frequency distributions that depart substantially from lognormal, but the method gives a satisfactory estimate of the geometric mean as long as the total distribution is unimodal. Where more than 50 percent of the data for an element are censored, lognormality is a tenuous assumption. Here abundance estimates are not calculated, and the true abundance is assumed to be

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less than the detection threshold value (all such cases involve the
    lower detection limit). Where geometric means or deviations for two
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    data sets are claimed to be significantly different statistically,
    the difference between them was tested for significance at the
    95-percent-confidence level (Moroney, 1956)
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The abundance estimates derived by the above method are suitable for intercomparison, and they are used in figures accompanying the text. These estimates, however, are not the best estimates of true abundance possible, because the geometric mean generally gives values somewhat smaller than the true abundance. To obtain the best possible estimate of true abundance, one must calculate Sichel's <u>t</u> estimator, a statistical measure of central tendency for lognormal frequency distributions, designed to eliminate the negative bias inherent in the geometric mean. The <u>t</u> estimator gives a value close to the

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very high values, as is the arithmetic mean. To obtain <u>t</u>, the antilog of the geometric mean is multiplied by a correction factor approximately proportional in size to the antilog of the geometric standard deviation but also partly determined by the number of samples (see discussion by Miesch, 1967, p. B7-B8). Sichel's <u>t</u> estimator was calculated for each element from the geometric mean and standard deviation previously calculated with the aid of Cohen's method. These values are included in table of for the interested reader, even though they are not used for the data comparisons upon which this study relies. Approximate confidence intervals for Sichel's <u>t</u> were calculated for gold using an equation given by Aitchison and Brown (1963, p. 50).

5 --10-11 The \underline{t} estimator, developed by H. S. Sichel (1952, 1966), is not 12 to be confused with Student's \underline{t} , a frequency distribution function 13 commonly used to calculate confidence intervals for various statistical 14 measures. Sichel developed the t estimator specifically for evaluating ore blocks in South African gold mines. 16 17 18 19 20-21 22 23 24

Correlation coefficients were calculated for each pair of elements by computer using the log-transformed data, although using the untransformed data would have produced the same results because we calculated Spearman's rank correlation coefficient rather than the more widely used product-moment correlation coefficient. Flanagan (1957) showed that the rank correlation coefficient is particularly suitable for semifquantitative spectrographic data and is the only valid method of computing a correlation coefficient between an element determined chemically and an element determined spectrographically. The Spearman rank correlation coefficient is nonparametric, and therefore does not require the assumptions that must be made when using the product-moment correlation coefficient. These assumptions include a fundamentally normal distribution for each element, independence of successive data pairs, and homogeneity of variances. Since most of the elements investigated here were determined spectrographically, but several important elements were determined chemically, we have adopted Flanagan's method. Each correlation coefficient was calculated using only those observations having data within the detection limits for both elements involved. Since a different number of observations was used for nearly every correlation coefficient, the reliability of each coefficient is different. Each coefficient was subjected to a significance test, using Student's t. The correlation matrices (fig. 7) show which correlation coefficients are significant at the 99- and 95-percent confidence levels. Flanagan (1957) presents and explains the formulas for computing both the rank correlation coefficient and the

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significance test / and gives references to the statistics literature that carries the derivations of these formulas. Additional explanation of rank correlation coefficients and the way in which we use results of the significance test are given later in this report (p. kb the). 5 --10-11 12 13 14 15-16 17 18 19 20-21 22 23 24 25-

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Hydrothermal alteration and metallization

General

The Tertiary volcanic flows that occupied the Combination-January area were hydrothermally altered, metallized, and later oxidized to the depths now accessible. Hydrothermal alteration and metallization were related as follows. Strong fracture zones which cut the flows provided channelways for the hydrothermal solutions; the rocks in and adjacent to these fracture zones were silicified, and rocks farther from the fracture zones were argillized. According to Ransome (1909, p. 158), Locke (1912a, p. 800-801,) and Collins (1907a, p. 398), ore was associated with silicified zones (veins), and little ore extended into surrounding argillized rocks. Changes in one grade were often abrupt, but boundaries between ore and low-grade or barren rock were always gradational over at least a few feet (Ransome, 1909, p. 213, 218; Collins, 1907b, p. 435). Hydrothermal wallrock alteration was well advanced when metallization began, but some hydrothermal quartz and alunite formed contemporaneously with metal sulfides and gold (Ransome, 1909, p. 167, 169-170). Fracturing of the silicified ledges during the later stages of alteration produced local concentrations of gold and sulfides and provided relatively large open cavities in which rich ore formed. The bulk of the precious metal recovered, however, was apparently disseminated through volumes of rock within the silicified zones; most ore is therefore structurally controlled by prealteration fracturing (Ransome, 1909, p. 160-162). In the extensive barren parts of the Goldfield altered area the silicified zones show structural

relationships and alteration mineral assemblages identical to those 2 in metallized areas, so metals were deposited during and after wallrock alteration only where they were available to the hydrothermal system, presumably entering the system at deeper levels. 5 ---10-11 12 13 14 15 --16 17 18 19 20~ 21 23 24 25-

The fault zones and shear zones shown on plate 2, although they probably do not represent large displacements, record movement that occurred after silicification developed along northwest-trending fractures. The shear zones consist of many closely spaced fractures, whereas fault zones consist of one or several large breaks with few subsidiary fractures. Much of the rock exposed in the cuts, both silicified and argillized, is moderately to intensely fractured; the fractures have diverse orientations and at many localities show slicken sides representing movement in diverse directions. They are too small and too numerous to show on plate 2. Ore minerals, particularly in high grade ores, filled open spaces in shattered portions of the silicified The shattering resulted from movement that occurred after hydrothermal alteration was well advanced but before metallization was complete; metallization accompanied the later stages of hydrothermal alteration (Ransome, 1909, p. 215-216). In some of the shear zones and fault zones, clay minerals differ in abundance and proportions from adjacent argillized rocks, suggesting that these breaks formed before hydrothermal activity ceased. Supergene alteration could also be partly or wholly responsible for these differences in clay content, but we cannot rule out the possibility that at least some, possibly all of the postsilicification fault zones and shear zones shown on plate 2 originated before hydrothermal alteration and metallization and many small grantures ceased, even though some of these breaks, show slickensides that must postdate all alteration.

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All rock exposed in the cuts is within the upper part of the oxidized zone. The depth of oxidation at the Combination mine is 130 to 140 feet (Ransome, 1909, p. 177, 216), and water was encountered at 210 feet when the shaft was sunk (Co;lins, 1907a, p. 398). The depth of oxidation at the January mine is 180 feet and the original water level was 160 feet (Ransome, 1909, pl. XVI, p. 187, 219).

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To completely describe the petrographic changes produced by hydrothermal alteration, metallization, and oxidation, we should ideally have suites of unaltered rocks, unoxidized argillized rocks, unoxidized silicified rocks, unoxidized average-grade ores, unoxidized high-grade ores, oxidized angillized rocks, oxidized silicified rocks, and oxidized average-grade and high-grade ores. Of these nine groups, we were able to obtain satisfactory numbers of unaltered rocks, unoxidized silicified rocks, unoxidized average-grade ores, unoxidized high-grade ores, oxidized angillized rocks, and oxidized silicified rocks. The Combination-January cuts provided particularly large numbers of oxidized argillized and silicified rocks. Although some silicified rock samples from the cuts actually constitute average-grade ores, all high-grade loxidized ore has been mined out. The following sections describe the petrographic characteristics of unaltered dacite, then the locations and petrographic characteristics of the unoxidized rocks, and then the petrographic characteristics of the oxidized rocks. The petrographic descriptions in the latter section are more detailed because the

number of oxidized samples available is much greater than the number of unexidized samples. Oxidation produces few mineralogic changes but does produce some notable geochemical changes, discussed later in the report.

Petrography of unaltered dacite samples

The Combination-January cuts mainly expose altered dacite. Since no samples were collected from the few small exposures of altered andesite, this section describes only the petrography of unaltered dacite, for comparison with the unoxidized and oxidized altered dacite samples to be described in following sections. Seventeen samples collected from scattered outcrops east of the main productive area provided geochemical data used later in the report. Thin sections for three of these 17 samples provided the following data.

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Petrography of unaltered and ovidized

altered dacite samples

The Combination-January cuts expose soft clay-bearing altered rock and hard silicified rock equally well; both are the products of hydrothermal alteration of dacite and andesite. Since no samples were collected from the few small exposures of altered andesite, this section describes only the petrography of altered dacite and, for comparison, unaltered dacite. Three thin sections from rocks outside the compination-January area provide the data on unaltered dacite. Dacite is characterized by porphyritic texture, with 20 to 25 percent plagioclase phenocrysts 0.3 to 10 mm in diameter, about 5 to 8 percent each of biotite, hornblende, and augite phenocrysts as much as 2 mm long, 0.5 to 1 percent corroded quartz phenocrysts as much as 2 mm long, and 1 percent opaque grains 0.1 to 0.3 mm in diameter. The plagicclase phenocrysts show normal oscillatory zoning, and have a bulk composition of about An 50. The groundmass is one-half to two-thirds

microlites of sodic labradorite, a few percent minute opaque and mafic

grains, and the remainder glass. The groundmass shows good pilotaxitic

texture.

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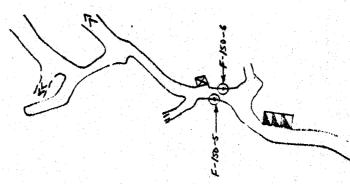
1 . Petrography of unoxidized altered rocks and ores Sampling of unoxidized materials from the Combination-January area would have been desirable, but was not possible due to lack of access and lack of ore samples. Thirteen unoxidized silicified dacite samples for petrographic and minor-element comparisons were collected from the Florence mine (fig. 3), 1,600 feet southeast of the Combination mine, 7 Figure 3 near here because the Florence is the only mine currently accessible. The Florence 10- and the Combination are both located on the same vein system. The 16 11 unoxidized average-grade ore samples are from mine dumps throughout the 12 main district (fig. 4) and the 15 high-grade ores are from mines 13 Figure 4 near here in several parts of the main district (fig. 5). Only two of the high-grade 16 Figure 5 near here 17 ore samples are unequivocally known to have come from the area under 19 Three other high-grade samples most likely came from the Combination mine but could also have come from any one of the other 21 major mines except the Florence. 22 23 24 25-

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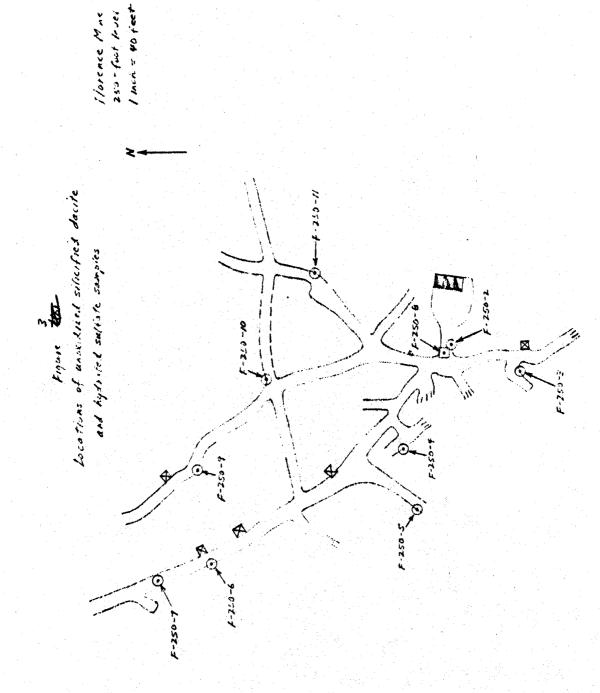
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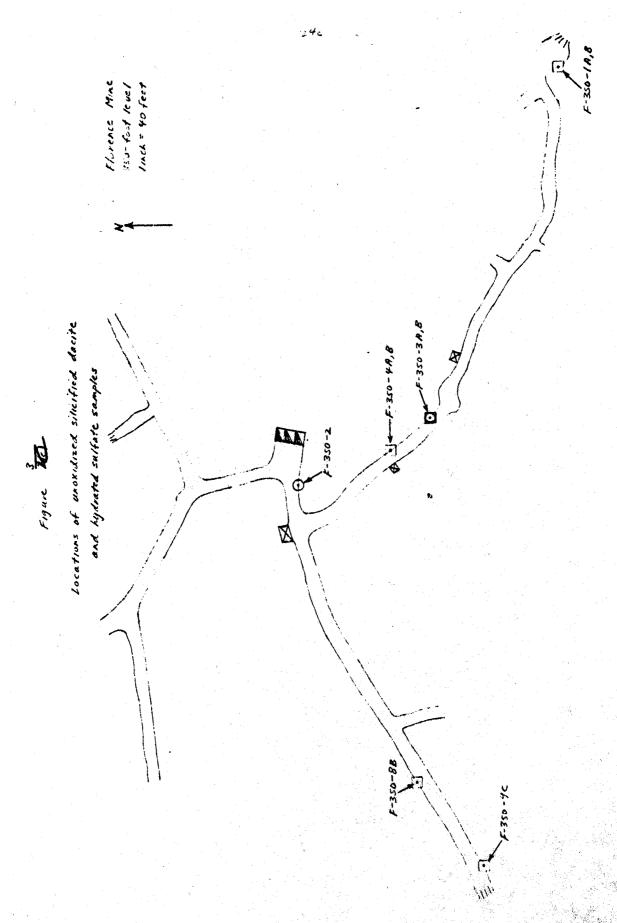


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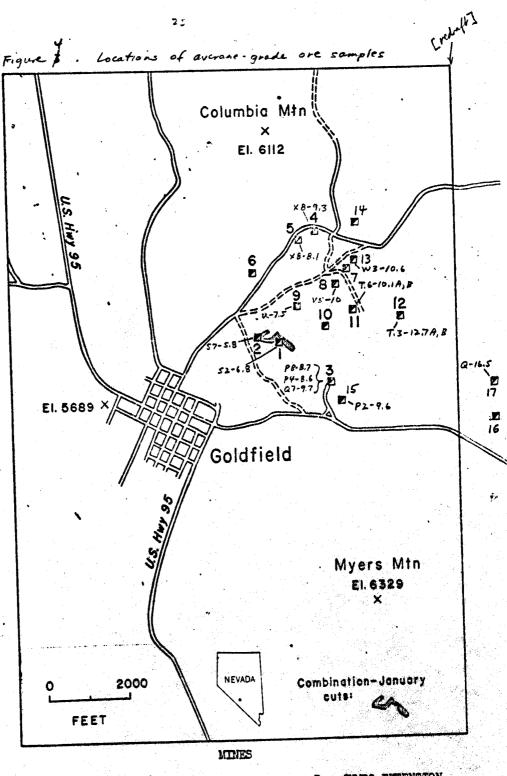
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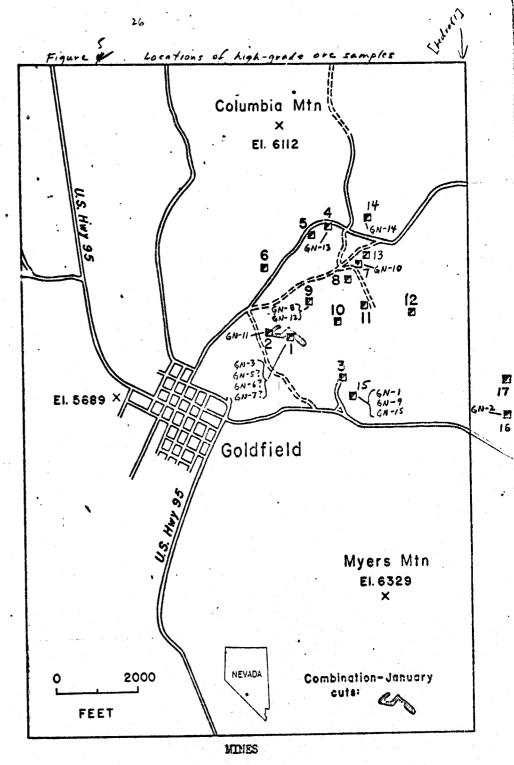








JUMBO EXPENSION 1. COLIBINATION (52-6.8) CLERMONT (V5-10) 8. 2. JANUARY (57-5.8) 9. MOHAUTK (4-7-5) 3. FLORENCE (Pa-8.7, P4-8.6, Q7-9.7) JULIBO 10. 4. LAGUNA (x8-9.3) 11. GRIZYLY BEAR (76-10.18,76-10.18) 5. RED TOP (x8-8.1) 14. KEWANAS 12. MERGER (T3-/2-7A, T8-12-78) SILVER PICK 15, VELVET (W3-10.6) Note: the Deep Mines has no recorded 15. LITTLE FLORENCE (P2-9.4)
16. GOLD KAP
17. DECP MINES (Q-16.5) production (searls, 1748, p. 18).



COMBINATION (6N-3, 6N-5?, 6N-6?, 6N-7?) JUMBO EXTENSION (611-10) 8. JAMUARY (6N-11) CLERMONT 9. MOHAWK (6N-B, 6N-12) 3. FLORENCE 10. JUL BO LAGINA (611-13) 11. ORIZZLY BEAR RED TOP MERGER VELVET 12. SILVER PICK 14 KEWANAS 13. (6N-14) -LITTLE FLORENCE (GN-1, GN-9, GOLD BAR (CN-2) GN-15) Notes: 1. GN-13 1s from the Mushett 16: lease, adjacent to the lanana. 2. The Geep Mines has no recorded production. DEEP MINES

3. Locaton (mine) anknown for GN-4-

27e (28 follows

The unaltered dacite is a porphyritic volcanic rock with plagioclase, biotite, hornblende, augite, and a few quarts phonocrysts in an aphanitic, partly glassy groundmass. A more complete description is given on p. 33. In unoxidized silicified dacite, the groundmass and phenocrysts are both replaced by quartz, alunite, and kaolinite, but the phenocrysts are richer in alunite and kaolinite than the groundmass. Pyrite partly replaces former mafic minerals and is also scattered throughout the altered groundmass. The oxidized silicified rocks described in the next section are petrographically the same as these rocks, except that hematite or goethite replaces the pyrite. Unoxidized average-grade ores contain 10 to 25 percent sulfide-bearing quartz aggregates which form veinlets that cut the silicified wall rock, or surround wall rock breccia fragments. Pyrite, stibioluzonite, and other sulfides (see p. in these veinlets and fillings form crusts or are intergrown with the quartz. Considerable alunite accompanies the quartz in many veinlets, and kaolinite fills scattered vugs remaining at the centers of the veinlets. Much of the material in the veinlets filled open fractures or open brecciæ, but some specimens show metasomatic effects extending several millimeters from the vein margins into the silicified wall rocks. Otherwise, wall rocks in these specimens are identical to unmetallized unoxidized silicified rocks; relict textures are generally well preserved. In a few specimens stibioluzonite is disseminated through the silicified wall rock. Here relict textures are obliterated, indicating that metasomatism locally extends at least several centimeters from veins into wall rocks. The high-grade samples could

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only be examined visually; sulfide minerals are conspicuous in all ı samples and abundant in some. Native gold is visible in several. 2 Details of ere mineralogy will be presented in a later section. 5 -10-11 12 13 14 15-16 17 18 19 20-21 22 23 24 25.

Petrography of oxidized altered dacite samples

The altered rocks exposed at the surface at Goldfield have been described in detail by Harvey and Vitaliano (1964). The criteria we from the Combination—January use for classifying individual oxidized altered—rock samples/are generally the same as the criteria they used for distinguishing several alteration zones and subzones. Thin sections from 30 of the samples from the cuts and X-ray diffractograms for 112 of the samples from the cuts provided the petrographic data which follows.

In this report we use the term "illite" to refer to a group of clay minerals having d_(OOl) approximately equal to 9.9Å, that do not expand when treated with ethylene glycal. We do not distinguish between 1MD and 2M₁ polymorphic forms, both of which occur in these rocks. The term "kaolinite" refers to any member of the kaolinite group except halloysite (kaolinite, nacrite, dickite). We have not attempted precise identification of kaolinite-group minerals. The term "montmorillonite" refers to a group of expandable clay minerals having d_(OOl) approximately equal to 14.7 to 15.5Å (samples air dried).

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Soft clay-bearing altered rocks, although considered a single group for geochemical comparisons, are described here in two subgroups those that contain montmorillonite, and those that do not. Montmorillonite-bearing rocks, restricted to a small area northwest of the Combination shaft (pl.2), have plagioclase phenocrysts that are partly to almost completely altered to aggregates of 1 to 10 $\mu\text{-long}$ illite and montmorillonite grains. Illite flakes and leucoxene granules replace biotite. Illite, leucoxene(?) or opaques, and in some cases minor quartz replace hornblende and augite. The groundmass is an aggregate of 1 to $5^{\circ}\mu$ -diameter quartz grains and 1 to $10^{\circ}\mu$ -long montmorillonite, illite, and in some rocks kaolinite flakes. Crystals of jarosite 2 to $10^{\circ}\mu$ in diameter are scattered through the groundmass of some specimens. These rocks have experienced the weakest hydrothermal alteration of any rocks exposed in the cuts; they were located farther from local sources of hydrothermal solutions than any other rocks described here. They belong to the montmorillonite subzone of the argillic zone described by Harvey and Vitaliano (1964, p. 568).

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Clay-bearing rocks with no montmorillonite contain abundant illite 1 and various amounts of kaolinite. In rocks with appreciable kaolinite, 2 this mineral replaces the plagioclase phenocrysts with aggregates of crystals less than 3 μ to 50 μ in diameter. In most rocks illite also occupies the plagioclase phenocrysts; the amount varies from a few flakes parallel to former crystallographic planes to 80 percent of the crystal, intergrown with the kaolinite. Biotite is replaced by illite and leucoxene; hornblende and augite are replaced by leucoxene, hematite, and minor quartz. The groundmass is an aggregate of 10-4 quartz grains with as much as 20 percent 1 to 5/p illite and as much as 30 percent diffuse patches of very fine-grained kaolinite. Scattered 1 to 5-u 11 granules of hematite, leucoxene, and in some rocks 10-u jarosite or 12 barite form about 10 percent of the groundmass. These clay -bearing rocks, 13 which belong to the illite-kaolinite subzone of Harvey and Vitaliano 14 (1964, p. 568-571), are the product of more intense hydrothermal alteration than the montmorillonite-bearing rocks. 16 17 18 19 20-21 22 23 24

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The rocks mapped as silicified rocks constitute the second group of samples used for geochemical comparisons (pl. 2). microcrystalline quartz with 15 to 65 percent alunite and kaolinite. Most rocks contain alunite and kaolinite in 10:1 to 1:3 proportions, but some have no kaolinite. Typically, alunite and kaolinite form about one-third of the rock. It is dense and hard, resembling chert, because most of the alunite and kaolinite is localized at the sites of former plagioclase phenocrysts, and sometimes former biotite phenocrysts, whereas the groundmass is mostly fine-grained intergrown quartz. Samples with more than 50 percent alunite and kaolinite become relatively soft, 10-(significant constituents because these minerals become important in the groundmass; these rocks are indistinguishable in the field from non+silicified clay-bearing rocks. Twelve such samples, found locally at the margins of visibly silicified zones, are included with silicified rocks on plate 2. These alunite- and kaolinite-rich zones are always substantially thinner than the harder silicified zones adjacent to them. Former plagioclase phenocrysts in silicified rocks are represented by randomly oriented aggregates of $20-100-\mu$ -wide alunite plates with 25 to 80 percent quartz (10 to 50μ) and kaolinite (2 to 10μ). groundmass is now a 2 to 20µ aggregate of quartz with 10 to 50 percent 20alunite or kaolinite or both, alunite generally the more abundant, and scattered granules of hematite, leucoxene or rutile, and in some rocks diaspore or jarosite. Subhedral areas containing 20 to 50 percent hematite, leucoxene or rutile, and rarely jarosite, but otherwise similar to the groundmass, represent biotite, hornblende, and augite

32 (p. 32 follows)

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phenocrysts. In some rocks, however, biotite is the site of coarse
1
    platy alunite with 20 to 30 percent fine-grained leucoxene or rutile,
2
    hematite, and minor quartz. The silicified rocks belong to the
    alunite-quartz zone of Harvey and Vitaliano (1964, p. 571). They
    form crudely tabular bodies which represent the rocks immediately
    adjacent to the fissures that conducted hydrothermal solutions.
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32a (p. 33 follows) S. GOVERNMENT PRINTING OFFICE: 1959 0 - 511171

365 (p. 37 follows)

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All the altered rocks show moderately to-well-preserved relict textures regardless of intensity of alteration. Relict quartz phenocrysts remain in all samples. 15-20-

Potential indicator elements:

elements associated with gold in Goldfield ores

The Tortiory volcanic flows that eccupied the Combination January

eres were hydrethermally altered, metallized, and later oxidized to

the depths now accessible. This section of the report focuses on

geochemical features produced by the metallization process; our objective is to identify all the elements that are associated with gold in the

Goldfield ores. Published descriptions of the ores and comparisons

between silicified rocks (13 samples), average-grade ores (16 samples),

and high-grade ores (15 samples), all unoxidized, served to delineate

the potential indicator element.

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Unoxidized ore typically contained pyrite (FeS2), bismuthinite ($\mathrm{Bi}_2\mathrm{S}_3$), stibioluzonite (famatinite)($\mathrm{Cu}_3(\mathrm{Sb},\mathrm{As})\mathrm{S}_4$), and native gold 2 _/Ransome (1909, p. 118-119) identified this mineral as famatinite, and presented an analysis showing the As:Sb ratio of the mineral to be nearly 1:1. Levy (1968) has shown that famatinite belongs to the tetragonal series luzonite-stibioluzonite, and has suggested that the name famatinite be dropped. Our X-ray data, when compared with Levy's data, indicate that the Goldfield famatinite is stibioluzonite. Ransome, 1909, p. 110-119, 165-166, 170). Collins (1907a, p. 398) 10reported tetrahedrite((Cu,Fe)₁₂(Sb,As)₄S₁₃) and small quantities of 11 chalcopyrite (CuFeS₂) and sphalerite (ZnS) from the unoxidized ores of 12 the Combination mine. Ransome (1909, p. 216) reported telluride ore 13 from the 280-foot level of the Combination (unoxidized ore). Tolman 14 and Ambrose (1934, p. 264-278) reported marcasite (FeS2), tennantite 15- $((Cu,Fe)_{12}(As,Sb)_4S_{13})$, goldfieldite $(Cu_3(Te,Sb,As)S_4)_7$ sylvanite 16 _/Palache, Berman, and Frondel (1944) give the formula 17 $Cu_{12}Sb_4Te_3S_{16}$. Thompson (1946), and more recently Levy (1968), consider goldfieldite a member of the tetrahedrite group. Levy gives the 19 formula Cu3(Te,Sb,As)S4. 21 (AgAuTe $_4$), hessite (Ag $_2$ Te), and petzite ((Au,Ag) $_2$ Te) in ores from other 22 mines in the vicinity. Searls (1948, p. 20) reports calaverite (AuTe2) 23 from a small vein developed by Newmont Mining Corporation about 0.3 24 mile west of the Florence mine. Ransome (1909, p. 112) described a few

occurrences of galena (PbS). Ransome's analyses of ore from the 1 Mohawk mine show copper, tellurium, bismuth, antimony, arsenic, gold, silver, zinc, and traces of lead (1909, p. 167, 169). X-ray diffraction and optical examination of the 16 average-grade ore samples confirm 5 ~ The ores examined contain approximately 0.1 to 2 ounces gold per ton and 0.1 to 6 ounces of silver per ton except for one sample bearing 30 ounces silver per ton. The average values for Goldfield, calculated using annual production figures for the most active period of the district, 1906 through 1918, are 1.56 ounces gold per ton and 10-0.35 ounces silver per ton (U.S. Geol. Survey, 1906-1918). 11 stibioluzonite and subordinate tetrahedrite-tennantite, which is often 12 enclosed in the stibioluzonite. Polished sections show that small 13 amounts of bismuthinite are commonly associated with tetrahedrite-tennantite. 14 15- Cursory examination of the 15 high-grade ores spectrographically 16 - Samples loaned by National Museum of Natural History, 17 Smithsonian Institution. 18 analysed for this study revealed no new major ore minerals. Searls (1948, 19 20- p. 18) reported minor but notable amounts of tin in ore from some of 21 the deeper ore bodies of the district. Preliminary microprobe data obtained by G. K. Czamanske show that stibioluzonite contains tin in 22 variable amounts: Concentrations of 0.5 percent tin are common and a 23 maximum of 2.7 percent tin was detected. No separate tin-bearing phase 24 25- was recognized. A polished section of one of the average-grade ores

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24a (p. 24b follows)

35a 35b

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provided this microprobe data and some additional microprobe data
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    referred to later in the report. The ores were thus characterized
2
    by copper, antimony, arsenic, bismuth, tellurium, gold, silver, zinc,
    lead, and tin, in approximate order of decreasing absolute abundance.
    The economically important elements were gold, silver, copper, and
    lead, in order of decreasing total value of production.
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24b (p. 25 follows)
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Elements suitable for geochemical exploration must occur in notable amounts in ore-bearing silicified zones, but not in barren silicified zones. To identify elements that were introduced mainly during metallization, rather than during silicification alone, we compare spectrographic data for unoxidized silicified rocks with similar data for unoxidized average-grade and high-grade ores. Analytical results and statistical data are given in tables 1 and 2. Data for 17

TABLES 1 and 2 NEAR HERE

samples of unaltered dacite are included for comparison with the altered and metallized rocks. The average minor-element content of these samples should be similar to that of the dacite in the Combination-13 January area before hydrothermal alteration. The samples comprising the three groups of unoxidized rocks and ores are few in number and from scattered localities, but they can be used for qualitative 16 comparisons because the same alteration mineral assemblages occur 17 throughout the district, and ore mineral assemblages from various 18 parts of the district have important features in common (Ransome, 19 1909, p. 165-169, 172-172). Unaltered dacite samples, unoxidized silicified dacite samples, unoxidized average-grade ores, and 21 unoxidized high-grade ores form a sequence of four groups whose 22 compositions show progressively stronger effects of the one-forming 23 process.

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Appendix A

Analytical dela, 17 analytical daute samples

Fe, 119, G, Ti 11 percent, all others in

parts per million

											Budan na
	Sample	An	Ag	Pb	Bi'	Hg	As	a	2n	Mo	Ba
	AAG-662	2(0.02)	2 (0.5)	20	2(10)	0.04	L(10)	4(2)	N(200)	(2)	700
		2(0.02)		10	4(10)	0.12	4 (10)	30	N(200)	L(2)	200
		4(0.02)		30	4(10)	0.18	4(10)	10	N(200)	4(2)	1500
		2(0.021		10	4(10)	0.24	(10)	2	N(200)	4(2)	700
•		2(0.02)		20	4(10)	0.10	4(10)	. 5	N(200)	4(2)	1000
1	(;	0.02	' 1	15	2(10)	n.d.	4(10)	5	NOOO	L(2)	700
	-1	4 (0.02)		20	2(10)	:	L(10)	2	N(200)	4(2)	200
		L(0,02)		20	L(10)	0.16	L(10)	5	N(200)	4(2)	300
		4(0.02)		20	2(10)		2(10)	50	N(200)	2(2)	1500
	抗	L(0.1)	N(0.5)	15	N(10)		4(10)	10_	N(200)	N(3)	150
	ADQ-869		N(0.5)	10		0.05	2(10)	15	N(200)	N(3)	1500
	1	4 4 (0.1)		15		0.04	4(10)	اسی ر	N(200)	N(3)	1500
	±	4(0.1)	N(0.5)	15	1	0.03	4(10)	20	N(200)	3	300
	EADQ- 942		N(0,5)	15		0.04	•	20	N(200)		1500
	2) 10	4(0.1)	N(0.5)			0.03	1	30	N (2.00)	3	1500
	<u> 0</u>	2(0.1)	N(0.5)			1(0.01)		30	N(200)	N(3)	150
	*/	4(0.1)	N(0.5)	:		6.07	L(10)	15	N(200)	N(3)	200
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		Note:	Lineans !	less th	an" the	defection	thresh	coffee y	own in	بالعرق والمائية	
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Sample	Be	Co	Co	La	Mn	Nb	Ni	52	V	· >
AAG-662	/	, 5	10	30	300	n.d.	5	700	100	20
AAG-682	. /	50	20	30	1500	n.d.	20	1500	200	20
AA6-688	4(1)	15	30	70	500	n.d.	15	1000	200	20
AA6-722	/	10	15	50	500	n.d.	15	700	700	20
AH6-724	/	4(5)	20	50	1000	n.d.	10	1000	100	20
AA6-773	/	2 7	20	50	500	n.d.	15	700	150	20
AAH-019	4(1)	10	30	50	700	n.d.	10	700	70	15
AA6-987	. /		50	50	700	n. L.	15	700	200	15
AAH-495	4(1)	7	30	50	1000	n.d.	7	1500	150	/5
ADR-865	NCIS	10	15	70	700	10	7	500	100	20
ADQ-869	N(I)	10	20	70	700	10	15	700	150	30
A0874	N(I)	15	15	70	700	10	10	700	150	30
ADQ-895	NCI)	15	30	70	700	10	15	1500	200	30
ADR-942	N(I)		15	70	2000	10	15	700	150	30
ADQ-943		100	30	70	1000	10	50	1000	150	30
ADQ-945	N(1)	20	30	70	2000	10	15	1500	200	30
ADR-068	N(I)	· /o	20	100	2000	15	10	1500	150	20
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Sample	Fe	Mg	Ca	77						
AA6-662	3.0	1.0	1.0	0.3						
AA6-682	7.0	2.0	3.0	0.5						
AA6-688	7.0	1.5	2.0	0.5		·				
AA6-722	3,0	1.0	1,0	0.2		·				
AA6-724	5,0	1.0	2.0	0.3				e facilities against against 1 January	1. Noted to 1.	
AA6-773	2.0	1.0	2.0	0.3				1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		
AAH-019	5.0	1.0	5.0	0.2						
PAG-987	7.0	1.5	5.0	0.7			Marie Confess (Marie Confess C	e state has all the harden a common		
AAH-495	10.0	2,0	3,0	0.7				an and a state of the state of		
ADQ-865	2.0	2,0	3.0	0.15	4		· · · · · · · · · · · · · · · · · · ·	···		
ADQ - 869		1.5	1.5	0.2					- Maria Maria	
0 ADQ-874	2.0	1.5	2.0	0.2						
ADR-895	3.0	1.5	3.0	0.5		*				
ADR-942	5,0	5.0	3.0	0.7			and the second second second second		e data standa e e e succes e de e e de un superingo y que e e	
ADQ -943	7.0	5,0	3.0	0.7						MARINE IN COMMERCIAL DE
ADR-945	7.0	5.0	7.0	0.7						
ADR- 068	5.0	2.0	3.0	0.7	· · · · · · · · · · · · · · · · · · ·	e commentation a subsequence some in se				
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Appointix B Analytical data, 13 unoxidized silicified docite samples from the Florence mine Fe, Mg, Ca, Ti in percent, all others in parts per mellion

San Contraction	And the distribution of the second				and the second of the second of the second of			page a sense a comp	<u> </u>	to made trans
Sample	Au	Ag	Pb	Bi.	Hg	As	Ca	2n	Mo	56
F-150-5	0.33	2	340	10	0.67	15	200	4(200)	لاد)	4(100)
F-150-6	0.39	1(0.5)	150	4(10)	0.30	30	100	2(200)	4(2)	4(100)
F-250-2	2(0.02)	0.5	70	4(10)	0.30	, 10	150	4(200)	2	4(100)
o F-250-3	0.75	4(0,5)	10	15	1,25	10	150	4(200)	2	4(100)
F-250-4	4, 32	1(0.5)	30	4(10)	0.42	<u> </u>	70	4(200)	4(2)	2(100)
F-250-5	L(0.02)	1(0.5)	100	4(10)	0.80		100	4(200)	15	2 (100)
F-250-6	1.79	5	300	200	1.05	10	300	200	4(2)	2 (100)
F-250-7	1.11	20	78	1000	10.4	320	\$ (5000)	4(200)	15	700
F-250-9	0.44	2(0.5)	300	4(10)	1.10	30	1000	L(200)	2	4(100)
F-250-10	0.96	0.7	100	4(10)	0.33		100	2(200)	4(2)	2(100)
F-250-11	L(0.02)	1(0.5)	150	L(10)	0.24	20	500	4(200)	7	4 (100)
£-350-2	0.05	4(0.5)	100	10	0.47	15	150	4 (200)	2	4(100)
F-350-38	4 (0,02)	15	15	200	4.90	6(320)	2000	1 (200)	4(2)	700
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WEST										
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Shrippile Sh	8										
F-150-6 L(20) 3000 L(1) L(5) 15 30 L(10) n.d. 10 500 F-250-2 L(20) 700 L(1) 10 30 30 50 n.d. 7 5000 F-250-3 L(20) L(100) L(1) 10 5 30 150 n.d. 10 L(100) E-250-4 L(20) 700 L(1) 5 15 30 100 n.d. 5 300 F-250-5 L(20) 5000 L(1) 150 70 30 100 n.d. 30 5000 F-250-6 L(20) 2000 L(1) 500 10 20 50 n.d. 70 1500 F-250-1 L(20) 5000 L(1) 300 20 20 30 n.d. 50 700 F-250-1 L(20) 3000 L(1) 70 30 50 100 n.d. 30 1500 F-250-10 L(20) 3000 L(1) 7 20 30 30 n.d. 7 1000 F-250-11 L(20) 3000 L(1) 150 15 50 100 n.d. 10 1500 F-250-12 L(20) 2000 L(1) 10 30 50 70 n.d. 7 500 F-250-12 L(20) 2000 L(1) 10 30 50 70 n.d. 7 500 F-250-12 L(20) 2000 L(1) 10 30 50 70 n.d. 7 500 F-250-12 L(20) 2000 L(1) 10 30 50 70 n.d. 5 150	Sampl	c Sn	Ba	Be	Co	Cr	La	Ma	Nb	Ni.	52
F-150-6 L(10) 3000 L(1) L(5) 15 30 L(10) n.d. 10 500 F-250-2 L(20) 700 L(1) 10 30 30 50 n.d. 7 5000 F-250-3 L(20) L(100) L(1) 10 5 30 150 n.d. 10 L(100) F-250-4 L(20) 700 L(1) 5 15 30 100 n.d. 5 300 F-250-5 L(20) 5000 L(1) 150 70 30 100 n.d. 30 5000 F-250-6 L(20) 2000 L(1) 500 10 20 50 n.d. 70 1500 F-250-7 L(20) 5000 L(1) 300 20 20 30 n.d. 50 700 F-250-7 L(20) 5000 L(1) 70 30 50 100 n.d. 30 1500 F-250-1 L(20) 3000 L(1) 7 20 30 30 n.d. 7 1000 F-250-1 L(20) 3000 L(1) 150 55 50 100 n.d. 10 1500 F-250-1 L(20) 2000 L(1) 10 30 50 70 n.d. 7 500 F-250-2 L(20) 2000 L(1) 10 30 50 70 n.d. 7 500 F-250-38 30 100 L(1) 4(5) L(5) 20 30 n.d. 5 150	F-150-	5 4(20)	5000	4(1)	4(5)	50	30	70	A.d.	5	2000
F-250-2 L(20) 700 L(1) 10 30 30 50 A.d. 7 5000 F-250-3 L(20) L(100) L(1) 10 5 30 150 A.d. 10 L(100) F-250-4 L(20) 700 L(1) 5 15 30 100 A.d. 5 300 F-250-5 L(20) 5000 L(1) 150 70 30 100 A.d. 30 5000 F-250-6 L(20) 2000 L(1) 500 10 20 50 A.d. 70 1500 F-250-7 L(20) 5000 L(1) 70 30 50 10 A.d. 50 790 F-250-7 L(20) 300 L(1) 70 30 50 10 A.d. 30 1500 F-250-7 L(20) 3000 L(1) 750 15 50 100 A.d. 10 1500 F-250-1 L(20) 2000 L(1) 150 15 50 100 A.d. 10 1500 F-250-1 L(20) 2000 L(1) 10 30 50 70 A.d. 7 500 F-350-38 30 100 L(1) L(5) L(5) 20 30 A.d. 5 150	•	j	3000	4(1)	4(5)	15	30	4(10)	n.d.	10	500
F-250-3 L(20) L(100) L(1) 10 5 30 150 A.d. 10 L(100) E-250-4 L(20) 700 L(1) 5 15 30 100 A.d. 5 300 F-250-5 L(20) 5000 L(1) 150 70 30 100 A.d. 30 5000 F-250-6 L(20) 2000 L(1) 500 10 20 50 A.d. 70 1500 F-250-7 L(20) 5000 L(1) 70 30 50 100 A.d. 50 700 F-250-7 L(20) 300 L(1) 7 20 30 30 A.d. 7 1000 F-250-7 L(20) 3000 L(1) 150 15 50 100 A.d. 10 1500 F-350-1 L(20) 2000 L(1) 10 30 50 70 A.d. 7 500 F-350-3E 30 100 L(1) 4(5) L(5) 20 30 A.d. 5 150	-		700	4(1)	10	30	30	50	n.d.	_ 7	5000
E-250-4 L(20) 700 L(1) 5 15 30 100 N.d. 5 300 F-250-5 L(20) 5000 L(1) 150 70 30 100 M.d. 30 5006 F-250-6 L(20) 2000 L(1) 500 10 20 50 M.d. 70 1500 F-250-7 200 1000 L(1) 300 20 20 30 M.d. 50 700 E-250-7 L(20) 5000 L(1) 70 30 50 100 M.d. 30 1500 F-250-10 L(20) 3000 L(1) 7 20 30 30 M.d. 7 1000 F-250-11 L(20) 3000 L(1) 150 15 50 100 M.d. 10 1500 F-350-12 L(20) 2000 L(1) 10 30 50 70 M.d. 7 500 F-350-38 30 100 L(1) L(5) L(5) 20 30 M.d. 5 150	·	,	4 (100)	4(1)	10	5	30	150	n.d.	10	L(100)
F-250-5 L(20) S000 L(1) 150 70 30 100 M.d. 30 5000 F-250-6 L(20) 2000 L(1) 500 10 20 50 M.d. 70 1500 F-250-7 200 1000 L(1) 300 20 20 30 M.d. 50 700 F-250-7 L(20) 5000 L(1) 70 30 50 100 M.d. 30 1500 F-250-70 L(20) 300 L(1) 7 20 30 30 M.d. 7 1000 F-250-71 L(20) 3000 L(1) 150 55 50 100 M.d. 10 1500 F-350-2 L(20) 2000 L(1) 10 30 50 70 M.d. 7 500 F-350-38 30 100 L(1) L(5) L(5) 20 30 M.d. 5 150		3		4(1)	5	15	30	100	n.d.	5	300
F-250-6 L(20) 2000 L(1) 500 10 20 50 H.M 70 1500 F-250-7 200 1000 L(1) 300 20 20 30 H.M 50 700 F-250-7 L(20) 5000 L(1) 70 30 50 100 H.M 30 1500 F-250-70 L(20) 300 L(1) 7 20 30 30 H.M 7 1000 F-250-71 L(20) 3000 L(1) 150 15 50 100 H.M 10 1500 F-350-2 L(20) 2000 L(1) 10 30 50 70 H.M 7 500 F-350-36 30 100 L(1) L(5) L(5) 20 30 H.M 5 150			5000	4(1)	150	70	30	100	n.d.	30	5000
F-250-7 200 1000 L(1) 300 20 20 30 4.d. 50 700 F-250-7 L(20) 5000 L(1) 70 30 50 100 4.d. 30 1500 F-250-70 L(20) 300 L(1) 7 20 30 30 4.d. 7 1000 F-250-71 L(20) 3000 L(1) 150 55 50 100 4.d. 10 1500 F-350-36 30 100 L(1) L(5) L(5) 20 30 4.d. 5 150		i.	2000	4(1)	500	10	20	50	n.d	70_	1500
E-250-7 L(20) 5000 L(1) 70 30 50 100 n.d. 30 1500 F-250-70 L(20) 3000 L(1) 150 15 50 100 n.d. 10 1500 F-350-1 L(20) 2000 L(1) 10 30 50 70 n.d. 7 500 F-350-38 30 100 L(1) L(5) L(5) 20 30 n.d. 5 150			1000	L(1)	300	20	20	30	n.d.	50	700
F-250-10 L(20) 3000 L(1) 7 20 30 30 M.d. 7 1000 F-250-11 L(20) 3000 L(1) 150 15 50 100 M.d. 10 1500 F350-2 L(20) 2000 L(1) 10 30 50 70 M.d. 7 500 F-350-38 30 100 L(1) L(5) L(5) 20 30 M.d. 5 150	42		5000	2(1)	70	30	50	100	n.d.	30	1500
F-250-11 L(20) 3000 L(1) 150 15 50 100 M.d. 10 1500 F350-2 L(20) 2000 L(1) 10 30 50 70 M.d. 7 500 F-350-38 30 100 L(1) L(5) L(5) 20 30 M.d. 5 150	*		300	2(1)	7	20	30	30	n.d.	7	1000
F350-2 L(20) 2000 L(1) 10 30 50 70 md. 7 500 F-350-36 30 100 L(1) L(5) L(5) 20 30 m.d. 5 150	- <u>4</u>	1		4(1)		15	50	100	n.d.	10	1500
F-350-38 30 100 L(1) L(5) L(5) 20 30 A.d. 5 150	1			201)	10		50	70	nd.	7	500
	2	વં		4(1)	4(5)	4(5)	20	30	n.d.	<u> </u>	150
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Sample	ν	Y	Fe	Mg	4	T1.				
F-150-5	200	4(10)	2.0	0.07	0.07	1.0				
F-150-6	150	4(10)	2.0	0,02	0.2	0.5				
F-250-2	300	4(10)	5,0	0.15	0.2	0.7		_		
F-250-3	50	140)	5.0	0.15	0.05	0.7	•			
° F-250-4		1(10)	3.0		L(0.05)	0.7		A COLUMN TO THE PERSON NAMED TO THE PERSON NAM	processing the second s	
F-250-5		4(10)	10.0	0.07	0.15	0.5				
F-250-6	30	4 (10)	7.0	0.2	L (0.05)					
F-250-7	70	4 (10)	20,0	0./	0.1	0.2		Angelo dello III di contra aggiorna di II deposi i		
1 F-250-9	200	L(10)	7.0	0.07	0.15	0.7				
F-250-10	150	L(10)	2.0	0.03	2(0.05)					
F-250-11	100	2(10)	5.0	0.2	0.1	0.5				
ÿ <i>F-350-</i> 2	100	L (10)	7.0	0.03	0.15	0.5				
F-350-3B	30	4(10)	1.0	2(0.02)	0.1	0.3				
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Appendix C Analytical data, 16 unoxidized average-grade ore samples Fe, Mg, Ca, Ti in percent, all others in parts per million

		Larrich Marie . V. 1930ap			p. Programmer	and bearing to a second	The second second second		man area and a second	.m 11.200
Sample	An	49	Pb	Bi	HJ	As	Cu	2~	Mo	56
57-5.8	70	100	300	1500	6 (6.0)	\$ (5000)	30000	4(200)	حی	10000
52-6.8	32	70	50	500	6(6.0)	3000	10000	4(200)	5	2000
T3-12.7A	3,0	20	200	20	2.7	9(5000)	50000	2(200)	10	3000
73-12.78	4.0	20	100	200	0.26	200	10000	700	15	700
4-7.5	63	1000	300	2000	0.20	6 (5000)	20000	4 (200)	2(3)	2000
Q-16.5	2.0	3	1500	70	0.22	30	1500	4 (200)	3	4(100)
P8-8.7	4.0	70	300	100	6(6.0)	6(5000)	30000	L (200)	4(3)	3000
×8-8.1	5.0	15	300	70	6 (6.0)	1000	10000	3000	4(3)	2000
P2-9.6	3.0	30	100	20	5 (6.0)	5000	10000	4(200)	4(3)	700
P4-8.6	1.0	5	300	_30	0.38	100	300	4 (200)	4(3)	2 (100)
Q7-9.7	23	30	300	30	6 (6.0)	4000	15000	2 (200)	5	1500
. ×8-23	1.0	15_	2000	200	1.0	150	10000	4 (200)	5	200
T6-10.1A	45	30	200	100	6 (6.0)	150	10000	4(200)	15	1500
78-10.18	2.0	200	200	1500	2.0	6 (5000)	70000	2 (200)	3	15000
V5-10	16	15	1000		2.7	6 (5000)	10000	2 (200)	3	1500
- V	/ •		7000	30	1	9.(9.: 2.	/		-	!
W3-10.6		100	2000	500		\$(5000)		1000		10000
W3-10.6									5	/0000
W3-/0.6										/0000
W3-/0.6										/0000
M3-/0.6										/0000
M3-10.6			2000	500						/0000
M3-10.6			2000							/0000
W3-/0.6		/00	2000	.500						/0000
₩3-/0.6		/00	2000	.500	3.5					/0000
₩3-/0.6	16	/00	2000			\$(5000)	70000			/0000
W3-/0.6		/00	2000			\$(5000)	70000			/0000

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Sample	SA	<i>7</i> e	Ba	Be	6	Cr	la	Mn	Nb	Ni
57-5.8	1000	7000	200	L(1)	<u>چ</u>	2	4(20)	10	L(10)	حى
52-6.8	206	4(1000)	200	2(1)	10	7	4(20)	10	2(10)	مد
T3-12.7A	700	4 (1000)	200	4(1)	7	. 5	2(20)	10	2(10)	15
73-12.78	10000	2(1000)	70	4(1)	100	10	2(20)	. 5	4 (10)	200
U-7.5	50	30000	300	4(1)	5	15	4(20)	10	4(10)	10
Q-16.5	20	4(1000)	500	4(1)	10	7	4(20)	2	4(10)	50
P8-8.7	700	4 (1000)	300	4(1)	10	3	2(20)	10	2(10)	30
X8-8.1	700	2 (1000)	500	2(1)_	5	15	4 (20)	15	4(10)	.5
P2-9.6	150	2 (1000)	300	4(1)	<u> </u>	5	4(20)	7	10	10
P4-8.6	4(10)	4(1000)	500	4(1)	5	2	2(20)		10	10
Q7-9.7	2000	4(1000)	500	2617	7	15	ل د <u>2</u> 0	10	2(10)	10
x8-9.3	10	L (1000)	500	4(1)	Z	<u>-</u>	L(20)	15_	2(10)	50
T6-10.1A	300	2(1000)	150	40)	15	3	4 (20)	7	4(10)	30
76-10.18	L (10)	7000	200	4(1)	15	3	4(20)	5	4 (10)	10
V5-10	2(10)	2 (1000)	150	4(1)	15	3	4(20)	7	2(10)	30
W3-10.6	2(10)	2 (1000)	300	4(1)	10	15	30	/	2(10)	15
	1) 10		- Fairly 1 - Fairly Mark 11 - Physics 1							
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sample	52	ν	Y	Fe	149	a	77.			
57-5.8	150	20	4(10)	5,0	0.005	0.07	0.07			
52-6.8	70	20	4(10)	5(10)	0.002	0.03	0.07		· · · · ·	
T3-12.7A	100	30	10	7.0	1 (0,002)	0.02	0.07		e de codo de	
73-12.78	30	100	4(10)	\$ (10)	2 (0,002)	0.015	0.015	•		
K-7.5	150	70	4(10)	3.0	0.005	0.03	0.2		· · ·	
Q-16.5	100	30	4(10)	6(10)	2 (0.002)	0.03	0.03	. 4	and the second s	
P8-8.7	200	20	4(10)	6(10)	0.005	0.1	0.01		e e e e e e e e e e e e e e e e e e e	
X8-8.1	500	150	2(10)	3.0	0.007	0.05	0.15		Pro Arriano Constituto del Constitut	
P2-9.6	150	70	4(10)	3.0	0,005	0.05	0.15		manner andre man appropriate and the	
2 P4-8.6	150	15	4(10)	3,0	0.015	0.07	0,07			
47-9.7	200	70	4(10)	2:0	2 (0.00Z)	0.1	0.15			
y 18-9.3	300	30	2010)	5(10)	2 (0.002)	0./	0.03			
76-10.1A	30	15	4 (10)	6 (10)	2 (0.002)	0.015	0.03		ada - na na sansananga anggapan danga ada na na	
T6-10.18	100	30	2(10)	1.5	2(0.002)	0,02	0.1		And the second section is a second	
× V5-10	700	70	2(10)	7.0	4 (0,002)	0.05	0.03		and the second s	
W3-10.6	200	100	4 (10)	5.0	2(0,002)	0.03	0.07			
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		+	4	1					L	+,

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Appendix D Analytical data, 15 unoxidized high-grade ore samples Bi, As, Cu, Sb, Te, and Fe in percent, all others in ppm

					the contract of the same	or an exercise recognization of				
	Bu	19	16	8;	Hg	As	Cu .	2n	Mo	<i>5</i> 4
6N-1	300	1000	150	1.0	1500	10	3.0	1500	15	7.0
6N-2	2000	5000	200	10	200	3	20	700	.200	5.0
6N-3	10000	500	300	0.5	30	0,3	1.0	300	10	0.15
GN-4	150	500	70	0.05	200	10	50	1000	20	7.0
6N-5	100	700	30	0.07	700	15	50	1500	7	7.0
6N-6	150	700	150	0.02	150	10	30	1000	20	7,0
6N-7	200	700	76	0.03	1000	15	50	1500	7	7.0
GN-8	30000	3000	2000	15	200	0.7	5,0	N(100)	N(2-)	7.0
5N-9	20000	100	70	0.07	20	N(0.02)	0.05	N(100)	N(z)	0.007
GN-10	500	2000	700	2.0	150	7	20	3000	50	5.0
6N-11	1000	700	100	1,0	3000	10	20	1000	N(2)	7.0
6N-12	7000	20000	1500	10	150	N(0.02)	0,5	N(100)	N(Z)	0.1
6N-13	200	300	2000	/0	500	2	5,0	20000	5	1.5
BN-14	50	1000	50	0.5	N(20)	7	20	3000	20	5.0
6N-15	20	10	30	0.01	50	N(0.02)	0.03	N (100)	N(2)	0.005
	1	<u> </u>		1	!			•		i
	Note: 4	all data	obtained	· boy 6-	step sea	neguantit	a five sp	ectrogo	apric .	
	Note: a analys to bet	is. Spector evolution	obtained tro of a afe thos	by 6- 10-fold . se clame	step ser	riquantit	afive sp chions we head co	rectrogo ore also	recorde.	
	Note: a analys to beth Special	is. Specifier evaluation	obtained tra of are those c param	by 6- 10-fold se cleme refers w	step ser and 100 at accuracy are user	fold dille	ative spections we high co	nectrogo nectro nectra nectra vaiues f	recorde.	
	Note: 4 Analys to bet Special	is. Spector evolute exposur	obtained tro of are the	by 6- 10-fold se cleme neters w	step sei and 100 at acco	miguantit. fold dille	ative spections we have the	nectro go ore also necetra Values f	recorde.	
	Note: a analys to beth Special	te. Spec ter evalue l'exposur	obtained tro of are those or param	10=fold (sc eleme mefers w	and 100-	fold dillering at a foota	Legin Co	values f	epnic recorde. Floris. or Mg.	
	Analys to beth Special	te. Spec ter evalue l'exposur	tro of a sec those param	10=fold (sc eleme mefers w	and 100-	fold dillering at a foota	Legin Co	ectroge ore also nectra Values f	epnic recorde. Flore. or Mg.	
	Analys to beth Special	ter evolute	tro of a a te those of param	10 fold of the second s	and 100-	fold dillering at a foota	Legin Co	ve also	recorde.	
	Analys to beth Special	ter evolue	tro of a a tro	10 fold of the second s	and 100-	fold dela correg at a to obta	Legin Co	ve also	Floris.	
	Analys to beth Special	ter evolue	tro of a a tro	10 fold of the second s	and 100-	fold dela	Legin Co	ve also	Floris.	
	analys to beth Special	ter evolue	tro of a param	10 fold of the second s	and 100-	fold dela	hega co	ve also	Floris.	
	analys to beth Special	ter evolue, l'exposur	tro of a param	10 fold of the second s	and 100-	fold dela	hega co	ve also	Floris.	

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	Sa	Te	cd	Ba	Be	6	Cr	la	Ma	NB
6N-1	30000	2,0	700	50	N(I)	100	3.	N(30)	50	NLT
6N-2	5000	10	100	7	N(I)	3	7	N(30)	. /	N(7)
6N-3	100	3,0	N(50)	700	N(I)	10	10	N(30)	7	N(7)
§ 6N-4	7000	0.7	300	70	N(I)	15	. 5	N(30)	10	N(7)
6N-5	7000	0.7	500	15	N(I)	150	N(I)	N(30)	3	N(7)
6N-6	3000	0.5	700	150	N(I)	15	3	N(30)	2	N(7)
₹ 6N-7	7000	1.0	700	30	NUS	20	2	N(30)	201)	N(7)
5N-B	30	5.0	300	30	N(I)	N(2)	. /	N(30)	/	N(7)
5N-9	N(7)	0.7	N(50)	100	N(I)	7	5	N(30)		10
6N-10	10000	5.0	200	20	NU	30	10	N(30)	NCIS	N(7)
6N-11	7000	5.0	150	<u> </u>	N(I)	N(2)	3	N(30)	10	N(7)
8 6N-12	N(7)	3.0	150		N(I)	N(2)	1.5	N(30)	N(I)	N(7)
6N-13	5000	0.3	700	300	N(I)	15	10	30	1.5	N(7)
6N-14	1000	N(0.1)	N(50)	20	N(1)	100	2	N(30)	ے	N(7)
6N-15	N(7)	N(0.1)	N(50)	500	N(1)	15	7	N(30)	7	N(7)
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	Ni.	52	ν	Y	FC	Mg	Ca	Ti		
GN-1	50	20	50	N(10)	0.2	15	50	100		
6N-2	1.5	N(5)	70	N(10)	0,2	2.	4(7)	. <i>30</i>		
GN-3	10	500	50	N(10)	1.0	20	500	1000		
6N-4	15	70	100	N(10)	1.0	15	70	150	Apper Man	
6N-5	100	10	50	N(10)	0.7	5	4(7)	. 70	The second secon	
6N-6	10	100	76	N(10)	1.0	5	50	300		
GN-7	15	15	100	N(10)	0.7	3	>	150	TO STORE THE MENT OF THE STORE OF	
GN-8	N(I)	20	N(3)	N(10)	0.7	3	20	150		
GN-9	1.5	150	36	N(10)	0.7	30	200	2000	Management (1) a con-	
6N-10	7	100	200	N(10)	1.5	10	20	30		ļ
6N-11	2	N(S)	70	N(10)	0.15	2	7	10		
GN-12	N(I)	N(5)	N(3)	NEID	0.02	2(2)	4(7)	10		
6N-13	20	700	70	N(IO)	2.0	20	500	1500		
6N-14	70	70	50	N(10)	7.0	50	30	20		
SN-15	3	700	70	10	1.5	50	2000	2000		
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Appendix & Statistical data, 17 unaffered davite samples

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Verex	Control of the series	Restration of	Jan Jan Harry	10 1 10 m	10 10 10 10 10 10 10 10 10 10 10 10 10 1)				
vic*	" Son My	Prop As Con	Co. St. Parcos.	10 15 132	9 34 31, 10g					or ordered again. I take these restrates
An	94									
49	100		_	_						
Pb	0	-2,79	0.13	_	-					
81	100			_						
Hg 41	1	-5.17	0.31	-5.23	0.38					
As	100	-								
Cu	6	-2.94	0.42	-3.00	0.47					
Zn	100						2-10-185 - 1-10-10-18-18-18-18-18-18-18-18-18-18-18-18-18-	gent and		
Mo	88		_		-		N - 10 - 10 - 100 - 140 -			
Ba	0	-0.84	0.19							
Be	59						on the second	als reposits /2 to where a p		
6	6	-2.87	0.33	-2.92	0.36					ļ
. Cr	0	-2.66	0.17			a de la companya de l			op a se se selection	-
<u> </u>	0	-2.24	0.14							
Ma	0	-1.08	0.24							ļ
NA Z	0	-2.98	0.06			.				1
Ni	0	-2.89	0.22							-
عرک او الم	0	-1.04	0.16	-						
ļ v	0	-/.83	0.14							 -
5 <u>~ </u>	0	-2.66	0.11	1						1
FL	0	0.65	0.24	t			Da tamm	ned for	16 . e. e. a. 6	
Mg	0	0.25	0.24		-		Determi	ed for	8 sampl	•
Ca	0	0.41	0.23					ļ		† . † · ·
Ti		-0.41	0.25	-						
			to the same of the same							
e e e e e e e e e e e e e e e e e e e		and the second of the second						A STATE OF THE STA	· [
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38 (p. 38ac follow).

statistical data, 13 unoxidized silicified dacite samples from the Florence mine Au -4.20 0.51 - 4. B6 3/ 1,12 Ag 0.61 1.06 -3,50 -4.41 Pb 0 -2,07 0.46 -2,20 Bi 0.78 -3./8 Hg -4.10 0.47 As 8 (6) -2.79 0.46 -2.66 _ Cu 8 (6) -/.52 0.60 -1.65 0.43 92 24 Mo -3.37 0.39 -3.7/ 0.51 46 56 85 Sa 85 -0.85 -0.96 0.51 Ba 8 0.63 4 (6) indicates 100 Be. -2.4/ 0.72 -2.75 percentage given is the 23 0.9/ Co number of samples below 0.36 Cr -2,68 0.30 -2.75 a lower defection limit. La -2.51 0.14 Mn 8 0.23 0.34 -2.19 NB no data Ni -2.91 0.38 -0.99 0.44 -1.09 0.54 5> V -1.97 0 0.31 100 Fe 0 0.67 0.32 Mg 15 -1.11 0.33 -1.23 0.42 -0.93 Can 0.19 -1.05 23 Ti 0 -0.33 0.21

Appendix H

Appendix I Statistical data, 16 unoxidized average-grade one samples

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	<i></i>	J. O. 1.	Post Star 10 de	R. Q. A. C.	Contraction	O'Sale ()					
	Au	0	-3./4	0.60							
	Ag	0	-2,45	0.60							
	PB	0	-1.49	0.45							
										a source consequence of a con-	
Š	Bi	0	-1.83	0.64	- 11	. 					
5	Hg	44(6)	-/.34	0.76	-0.49	1.21					and the second
Data	As	44(6)	-4.07	0.49	-3.41	0.90					· · · · · · · · · · · · · · · · · · ·
<u> </u>	a	0	0.09	0.58						un om i mengunki k	anner
MASS.	Z1	8/	<u> </u>					-			
DING.	M.	3/	-3.25	0.25	- 3.40	0,32					
. PE/	56	/3	-0.68	0.48	-0.89	0.72					
ANY. IN	5 4.	25	-1.52	0.82	-3.07	1.22					
300	<i>7</i> e	81									
2		·	·						-		
ē.	P -	0	-1.58	1.24							
Y PUBLISH	Ba	0	-1.58	0.24							
WESLEY PUBLISH	_Be	100									
DISON WESLEY PUBLISH	Be Co	100	-3.03	0.32							
A ADDISON WESLEY PUBLISH	Be Co Cr)00 0									
IN U.S.A ADDISON WESLEY PUBLISH	Be Co	100	-3.03	0.32							
FHED IN U.S.A ADDISON WESLEY PUBLISH	Be Co Cr)00 0	-3.03	0.32							
INJURAFHED IN U.S.A ADDISON-WESLEY PUBLISH	Se Co Cr La)00 0 0 94	-3.03 -3.25	0.31							
LITTICGRAPHED IN U.S.A ADDISON-WESLEY PUBLISH	Be Co Cr La Mn	100 0 0 94	-3.03 -3.25	0.32							
LITHUGRAFHED IN U.S.A ADDISON-WESLEY PUBLISH	Be Co Cr La Mn Nb	0 0 94 0 88	-3.03 -3.25 - -3.74 -2.74	0.32 0.31 0.31							
LITHUGRAFHED IN U.S.A ADDISON-WESLEY PUBLISH	Be Co Cr La Mn Nb Ni Sr	0 0 94 0 88 0	-3.03 -3.25 -3.74 -3.74 -2.74 -7.85	0.31 0.31 0.40 0.36							
LITHURAPHED IN U.S.A ADDISON WESLEY PUBLISH	Be Co Cr La Mn Nb Ni Sr	0 0 94 0 88 0	-3.03 -3.25 -3.74 -2.74 -7.85 -2.39	0.32 0.31 0.31 0.40 0.36							
LITTINGRAPHED IN U.S.A ADDISON-WESLEY PUBLISH	Be Co Cr La Mn Nb Ni Sr V	0 0 9# 0 88 0 0	-3.03 -3.25 -3.14 -2.74 -1.85 -2.39	0.31 0.31 0.40 0.36		0.44					
LITHURRAPHED IN U.S.A ADDISON WESLEY PUBLISH	Be Co Cr La Mn Nb Ni Sr V Y Fe	100 0 94 0 88 0 0 0 94 38(6)	-3.03 -3.25 -3.14 -2.74 -1.85 -2.39 -0.55	0.31 0.31 0.31 6.40 0.36	0.83	0,41					
LITHUGRAFHED IN U.S.A ADDISON WESLEY PUBLISH	Be Co Cr La Mn Nb Ni Sr V Y Fe Mg	100 0 94 0 88 0 0 0 94 38(6)	-3.03 -3.25 -3.14 -2.74 -1.85 -2.39 -0.55 -2.27	0.32 0.31 0.31 6.40 0.36 0.31	0.83	0,41					
LITHURBAFHED IN U.S.A ADDISON-WESLEY PUBLISH	Be Co Cr La Mn Nb Ni Sr V Y Fe	100 0 94 0 88 0 0 0 94 38(6)	-3.03 -3.25 -3.14 -2.74 -1.85 -2.39 -0.55 -2.27 -1.40	0.31 0.31 0.31 0.40 0.36 0.31 0.21 0.24	0.83						

Appendix J Statistical data, 15 unoxidized high-grade ove samples

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		, Age	, io	(v +	F.Y _)				*.	· · ·
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_	W.	de ontra line	Co. 20 (10)	Portal Paral	0 40, Co.	Print Con			control of the second s		
	Au	0	-/.20	0.94	_		1	, * * · · ·			
		0	-1.06	0.81					- milan ender 1 - milan - 10 Fair 1		
	<i>Pb</i>	0	-1.60	0.55					2		
Ş V	Bi	0	-0.33	1.05							-
i. ⊑ E	Hg		-1.61	0.61	-1.71	0.69					
Dafe	As	20	0.68	0,52	0.08	1.30				aparamana a salah apamana an ara salah salah sa	
¥ .	Cu	0	0.77	1.03							
MASS.	24_	27	-0.81	0.44	-/.26	0.86	1 A 1 A 1 A 1 A 1 A 1 A 1 A 1 A 1 A 1 A				
DING.	Mo	33	-2.75	0.45	-3.23	0.81					
, R	56	0	0.06	1.08							10.0 1
ANY. IN	_5a_	20	-0.57	0.83	-/.24	1.56					
COMP		/3	0.24	0.46	0,04	0.67					
Ş	cd	27	-/. 48	0.30	-1.79	0.58					
ã.	· · · · · · · · · · · · · · · · · · ·	0			-/.//	Ø, S B		The state of the s			
ĒSE	Ba		-2.41	0.75							
V NOSK	Be_	100	Water-								
- ADC	<u> </u>	20	-2.67	0,48	-2.95	0.72					
75 7 2	Cr	7	-3.41	0.3/	-3.46	0.36					
PHED 1	La	93									
400HA	Ma	20	-3.36	0.47	-3.55	0.59					
5	Nb	93	<u> </u>								
	Ni	13	-2.99	0.59	-3.17	0.73				and a special content of the second of	
	5-	20	-2.08	0.62	-2.42	0.89	1				
	v	/3	-2.2/	0.24	-2.44	0.56					
			-								
	. <u>У</u>	- 				· · · · · · · · · · · · · · · · · · ·		1		-	
	Fe	0	-0.25	0.62		A CONTRACTOR OF THE CONTRACTOR			1		
	Mg	7.	-3,02	0.47	-3.08	0.52					
	Ca	20	-2./7	0.74	-2.47	0.92	4	ļ			
	Ti	0	-1.87	0.77	-	-			•		

38c (p. 39 follows)

The comparisons provided by Figure 6 confirm that copper

Figure 6 near here

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dominates the ores, and antimony, arsenic, bismuth, tellurium, gold, silver, zinc, lead, and tin are relatively abundant, as one would 5expect from available information on the ores. Data for tellurium are incomplete, but there is no reason to believe that amounts greater umaxidista than 0.1 percent (1.000 ppm) reside in unaltered or silicified dacite. In addition to the above elements, mercury and cadmium are relatively abundant. Molybdenum shows modest enrichment in some ore samples, 10although the averages for molybdenum in the two groups of ores are not 11 unovidered [notably] larger than the average for silicified dacite. These elements 12 13 form a group which will subsequently be referred to as "ore-related elements." These are the elements to be investigated as geochemical 14 indicators for the Goldfield deposits. Barely detectable amounts of 15palladium appeared in seven of the high-grade ore samples, and small 16 but variable amounts of indium appeared in five of the high-grade ore 17 samples. Detectable tungsten (200-300 ppm) appeared in two high-grade 18 ore samples. Platinum, tantalum, and thallium were sought by spectro-19 graphic analysis, but none of the samples from any data set contained 20-detectable amounts of these elements. The only element not tested 21 22 that might be important as an indicator is selenium.

graphic analysis, but none of the samples from any data set contained detectable amounts of these elements. The only element not tested that might be important as an indicator is selenium. Ransome(1909, p. 134, 166) notes that selenium was found in tellurium-bearing oxidation products from two mines several miles northeast of the main district. Levy (1968, p. 131-132) notes an unidentified selenium-bearing mineral found as inclusions in goldfieldite.

1 ... o (geometre +16 (geometric deviation) 11 ----80. ar les Derrander dacite (1) or 1 samples

2 = unoxidized silicities decite (18 samples)

3 = unoxidized silicities decite (18 samples)

4 = unoxidized silicities decite or 18 (18 samples)

4 = unoxidized silicities decite or 18 (18 samples)

6 = unoxidized silicities decite or 18 (18 samples)

6 = unoxidized silicities decite or 18 (18 samples)

6 = unoxidized silicities decite or 18 (18 samples)

7 = unoxidized silicities decite or 18 (18 samples)

7 = unoxidized silicities decite or 18 samples

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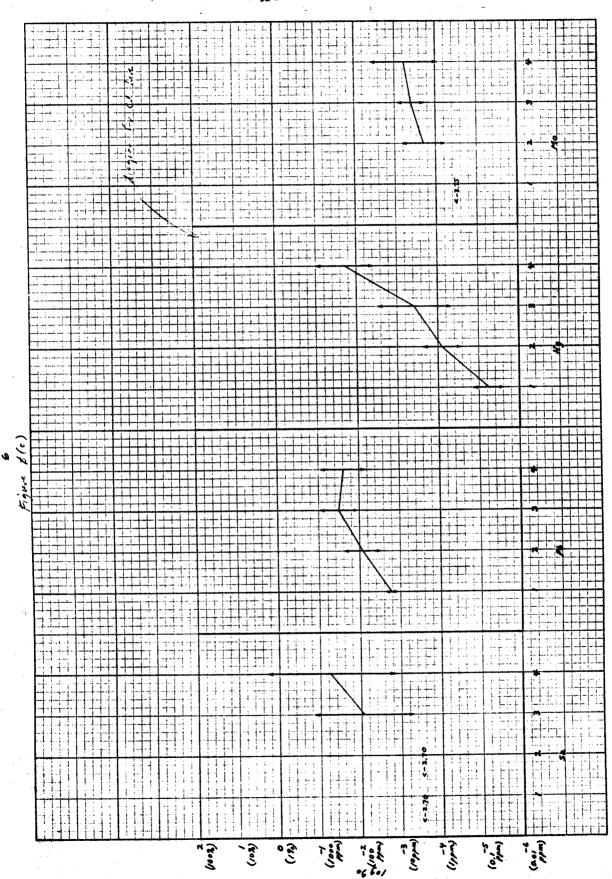
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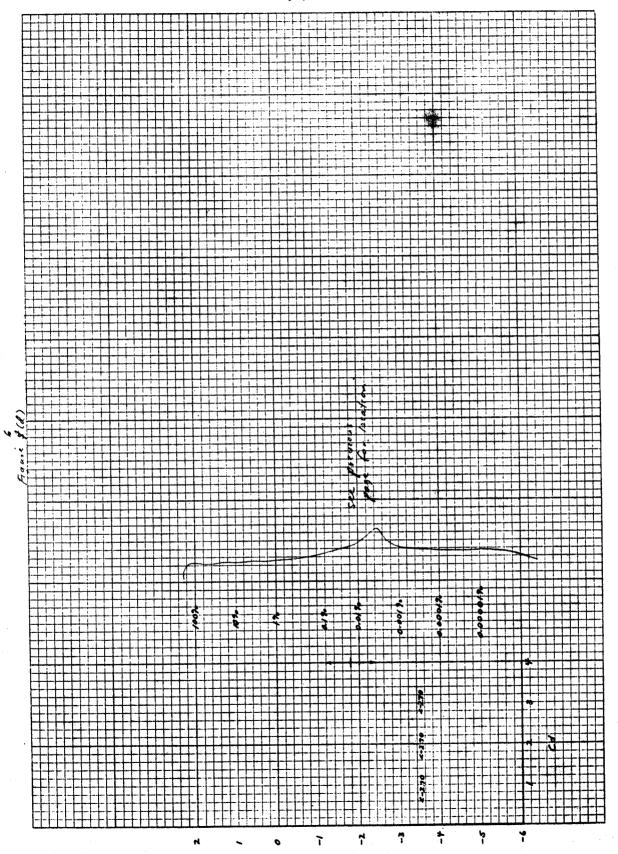
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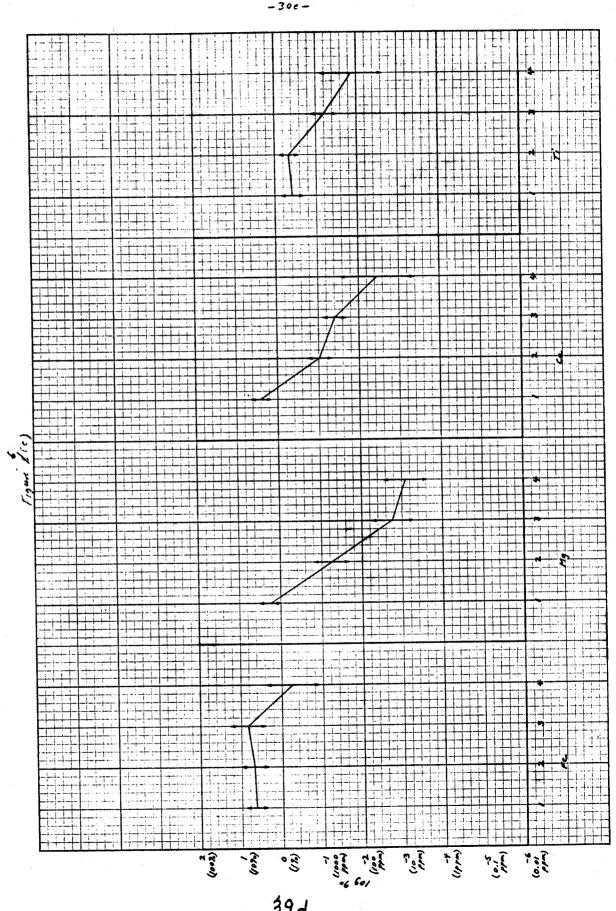


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The diagrams for iron, cobalt, chromium, nickel, and vanadium (fig. 6) are notably different from those for the elements described above. Regarding iron, unoxidized silicified dacite samples bear pyrite. much of which replaces former mafic minerals. Since the iron of this pyrite probably came from the preexisting mafic minerals, it is reasonable that unaltered and unoxidized silicified dacite have comparable amounts of iron, even though their mineral assemblages are vastly different. Some average-grade ore samples have pyrite associated with quartz veins and open-cavity fillings, in addition to that which replaces former mafic minerals in the silicified wall rock, producing iron contents notably larger than those seen in unoxidized silicified dacite samples. The average iron content, however, is not significantly larger than that calculated for unoxidized silicified rocks. High-grade ore samples, on the other hand, have much stibioluzonite and other ore minerals but relatively little pyrite, so these samples contain significantly smaller amounts of iron. Cobalt, chromium, nickel, and vanadium change little or not at all through the groups of samples. Cobalt and nickel (Ramdohr, 1969, p. 779) probably are retained in pyrite along with iron. These two elements are not, however, significantly less abundant in the high-grade ores, as is the case with iron; small amounts may have been introduced into the high-grade ores along with the abundant base and precious metals found therein. Cobalt and nickel data for the high-grade ores show significantly larger standard deviations than do the data for average-grade ores, suggesting that more complex processes involving addition, as well as removal influenced

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cobalt and nickel concentrations in the high-grade ores. It is not
    clear why chromium and vanadium are not more strongly depleted; we
    do not know which mineral phase(s) they reside in. Vanadium in the
    ores could reside in tetrahedrite-tennantite or stibioluzonite;
    substantial amounts of vanadium occur in colusite (Cu_{3}(Sn,V,As)S_{4}), a
    mineral closely related to both the luzonite and tetrahedrite-tennantite
    mineral series (Levy, 1967, p. 129).
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Magnesium, calcium, and manganese each show a pronounced and progressive decrease, indicating that depletion is directly related to intensity of hydrothermal activity. The same is true of titanium, barium, and strontium, except that each of these shows no significant difference in abundance between unaltered and unoxidized silicified dacite samples. Barium and strontium data, however, show significantly larger standard deviations in unoxidized silicified rocks than they do unaltered rocks, suggesting that hydrothermal alteration produced a net increase in barium and strontium concentrations in some samples and a net loss in other samples. Thin sections of the unoxidized silicified dacite samples show that titanium is retained in leucoxene (fine-grained anatase, TiO2) or in some cases rutile (TiO2), and barium and strontium are retained in barite (BaSO₄) or celestite (SrSO₄) or both. Barite and celestite form a continuous solid solution series, 15- but natural minerals are generally nearly pure BaSO4 or SrSO4 (Deer and others, 1962, p. 187-188, 197-198). The barite-celestite series mineral(s) recognized in thin section are too fine grained to separate easily for a more exact composition determination. It is not surprising to find barite (or celestite) along with the abundant alunite (KAl3(OH) (SO4)2) in the silicified rocks, since all these minerals are sulfates. Alunite abundance generally decreases with increasing amounts of ore minerals, and barite (celestite) probably decreases also.

Titanium, barium, and strontium could potentially form negative geochemical anomalies in areas with gold metallization. These elements could also conceivably be depleted, however, by particularly strong or long-enduring hydrothermal activity without metallization.

Use of negative anomalies, furthermore, requires particularly careful chemical analysis and involves greater difficulties in interpretation than use of positive anomalies; therefore these elements will not be considered further.

Lanthanum, niobium, and yttrium probably are depleted relative to amounts in fresh dacite, but little else can be said with the data available.

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Dispersion patterns of ore-related elements in the oxidized zone

General

The previous section of this report established that copper, antimony, arsenic, bismuth, tellurium, gold, silver, zinc, cadmium, tin, lead, mercury, and molybdenum characterize metallization at Goldfield. The following discussion examines the relationships between most of these elements in oxidized rocks of the Combination and January mines. It is particularly important to see whether these elements form primary aureoles or secondary halos around the ore bodies, and how closely related such relements are in the aureoles or halos.

-/Our use of the terms "primary," "secondary," "aureole," and "halo" is the same as that of Hawkes and Webb (1962). Since the Goldfield deposits are epigenetic (and hypogenetic), our usage is also concordant with revised definitions of the terms "primary," and "secondary" proposed by James (1967).

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Plate 2 is a geologic map of the open cuts at the Combination and January mines, showing rock types, structural features, and hydrothermal alteration zones. Sample localities are shown on plate 3. Specific localities mentioned in the text are numbered separately, and these numbers are given on plate 2 and the geochemical maps, plates

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The 278 samples that provide the data for this part of the report were collected about 4 feet above the floor of the large central cut. the smaller northwestern most cut, and the northwestern part of the cut east of the Combination shaft (see pl.2). Much of the latter cut is occupied by now-caved glory hele and stopes; samples here were taken about 4 feet above the top of the cave-in debris accumulated against the upper walls. The present walls are partly, perhaps largely, scarps left by blocks that broke off and slid or fell into the large caved area. Except for the caved area, the walls at the time of sampling and mapping were freshly excavated and very steep, representing elevation changes between 15 and 50 feet along any given profile across the cut wall. All rock exposed in the cuts is within the upper part of the oxidized zone. The depth of oxidation at the combination mine is 130 to 140 feet (Ransome, 1909, p. 177, 216), and water was encountered at 210 feet when the shaft was sunk (Collins, 1907, P. 398). of oxidation at the January mine is 180 feet and the original water level was 160 feet (Ransome, 1909, pl. XVI, p. 187, 219).

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Open workings, some forming precipitous holes, disrupt the sampling 2 scheme at most of the 14 locations where they intersect the cuts. Six of these are shown on plate 2 along the west wall of the largest cut between localities 2 and 7. These all lie within silicified rocks and 5 -field examination indicates that they are probably old stopes. seventh location (loc. 20), at the southeast end of the large caved area, in argillized rocks, is probably an access drift. Three more openings are located on the wall of the largest cut, between localities 6 and 7, and connect with the large opening shown at 7. last four appear in the vicinity of locality 8, and were probably originally 10-11 interconnected. All the workings represented by the latter two groups of openings are tunnels and possibly stopes that intersect the nearly 13 vertical cut wall at various elevations; the several near locality 8 14 almost overlap when projected onto a horizontal plane. For simplicity, 15therefore, these workings are not included on the plates. 16 the resulting problem of sampling bias in perspective, 12 of the existing 17 samples (samples 88, 89, 90, 91, 92, 161, 162, 185, 186, 187, 194, 195; 18 about 9 percent of the silicified rock samples) came from the walls or 19 floors of stopes, and probably contain less gold and other gold-related elements than the mined-out rock would have. Furthermore, several 21 additional samples would have been taken in the vicinity of locality 8 22 if it had not been previously disrupted by mining activity. 23

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Element distribution maps

Rocks affected differently by hydrothermal activity often show pronounced differences in amounts of ore-related elements, so plate 2 is used as a base for the element-distribution maps (pls. 413) which follow. The analytical and statistical data used to make the maps and accompanying histograms are given in tables 3 and 4. The cuts generally

TABLE 3 AND 4 NEAR HERE

parallel the vein system, but they expose one section approximately across strike immediately northwest of the Combination shaft (see AA', pls. 3-25). A geochemical profile for this section is shown for each element along with the geochemical map to help bring out relationships between the minor element distributions and various alteration and structural features, and to show the degree of contrast between oxidized metallized silicified and nonmetallized argillized rocks. This contrast may also be termed "anomaly height-to-background ratio," or "anomaly contrast." The histograms accompanying each geochemical map show the relative amounts of each element in silicified versus argillized rocks. The histograms also show the relative numbers of samples included in each range of values represented by a different map symbol. Distribution maps are included for gold, lead, silver, bismuth, mercury, arsenic, copper, molybdenum, and zinc.

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Table 3. Analytical data for oxidized silicified and argillized rock samples from the Combination-January cuts

SILICIFIED ROCKS, PERCENT

SAMPLE	AU. S	AG• &	PB • %	81,8	HG.8	AS, &	*• UO	Z.N. Z	% OW	BA, %
	0.000.0	0-001000	0.015000	0.0	0.000050	000900	0.002000	0.0	N 0.0	0.070000
7 7					0.000100	0.004000	0.000000	0.0	N 0 0 0	0.070000
8	0.000150	0.000200	0.050000		0.000000	0.0	0.00200	0.0	0.0	0.07000
4	0.000100	0.000100	0.007000	Z :	0.000010	0.012000	0.001500	0.0	2 2	0.07000
ις ·	0.000069	•	0.01000	200100	0.00010	0001000	0.000.00		2	0.07000
۰ ۵	0.000000	0.000300	0.05000	0.001000	0.000010	0.001000	0.010000	0.0	2000	0.070000
~ α			0.02000	0.001000	0.000020	0.001000	0.001500	0.0	0.0	0.070000
9	0.000340		0.02000	0.001000	0.000020	0.001000	0.003000	0.0	0.0	0.070000
10	0.000188	0.000300	0.020000	0.0	0.000020	0.012000	0.00500	0.0	0°0	0.070000
11	0.000279	0.000300	0.007000	0.003000	0.000050	0.016000	0.02000	0.0	0.003000	0.030000
30	0.000080	0.000150	0.0005000	0.001000	0.000011	0.001000	0.003000	0.0	0.001000	0.200000
32	0.000145	0.000300	0.015000	0.000500	0.000008	0.006000	0.003000	0.0	0.000700	0.02000
34	0.000150	•	0.030000	2 2	0.000011	001000	0.001000	0.002500	000000	0.015000
35	0.0	070000.0	0.030000	2 2	000000	001000	0.003000	0.0	0.000300	0.02000
30 7£	0.000290	0.000150	0.00000	z	0.000015	0.0	0.001500	0.0	0.000300	0.070000
38	0.000639	0.000200	0.030000	0.0	0.000015	0.001000	0.000500	0.0	0°0	0.100000
39	0.000430	0.000200	0.010000	0°0	0.000024	0.001000	0.000700	0°0	00800	0.100000
45	0.000288	0.000500	0.005000	0.00500.0	0.000020	0.001000	0.003000	0.0	O.0	0.02000
63	0.000000	0.000500	.003000	0.0	0.000013	0.006000	0.002000	0.002500	0.001500	0.07000
9 9	0.000000	0	0.0	Z :	0.000010	0.08000	0.001500	0.00500	0000000	0.00001
9	0.000140	0	0.020000	Z :	0.000008	0.004000	0.001500	0.0	0000000	0.10000
99	0.000000		0.010000	2 2	0.00000	0.00000	0.001000			0.30000
29	0.000275	•	0.012000	2 2	0.00011	0.001000	0.002000		00500	0.050000
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11	0.000030		0.015000	0.0	0.000015	0.002000	0.001000	0.00500.0	0.000500	0.050000
78	0.000240	0.0	0.050000	0.0	0.000017	0.016000	0.001500	0.004000	0.000500	0.05000
62	0.000036	0.0	0.015000	0°0	0.000011	0.014000	0.002000	0.014000	0.000500	0.07000
80	0.000150	0.0	0.00700.0	0.0	0.000015	0.02000	0.007000	0.0	0.001000	0.05000
81	0.000246	0.000500	0.02000	0.003000	0.000011	0.001000	000100	- L	2 2	0.0000
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86	0.00010	0.000500	0.015000	0.003000	0.000011	0.002000	0.001500	0.0	0°0	0.070000
87	0.000140	0.000300	0.030000	0.003000	0.000004	0.001000	0.001500	0.0	O.0	0.070000
88	0.000000	0.000300	0.020000	0.003000	0.000004	0.002000	0.001500	0.004000	0.0	0.070000
88	0.000055	0.000300	0.010000	0.0	900000*0	0.001000	0.000500	0.0	0.0	0.030000
06	0.000965	0.000100	0.030000		0.00000	0.001000	0.002000	0.0	O.0	0.05000
91	0.000940	0.000150	0.015000		0.000011	0.0	0.000100	0.0	200000	0.0000
92	0.000010	0.000050	0.015000	0.0	0.00000	0.00000	0001000	0.0	0000000	0.00000
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. 86	0.000030	0	0.001500	0.005000	0.001500	0.0	0.000200	0.150000	0.007000	0.000500
56	0.000030		0.001500	0.003000	0.002000	0.0	0.000200	0.200000	0.007000	0.000500
45		0.000100	0.001000	0.007000	0008000	0.001000	0.000200	0.030000	0.003000	0.000100
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101	0.000010			0.005000	0.001000		0.000300	0.015000	0.005000	00010
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37	7	0.015000	0.150000		
38	1.000000	0.015000	0.100000	0.200000	
39	2.	0.010000	0.100000	0.150000	
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9	4.	0.150000	0.100000	0.300000	
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89	4.	0.0	0.070000	0.150000	
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77	2.	•	0.00000	m,	
78	7	0.	0.0000	0.150000	
62	2	•	0.020000	0.200000	
80	4.	O.0	0.030000	0.150000	
. 81		0.007000	0.100000	0.300000	
8	ċ	•	0.02000	0.200000	
83	ċ	0.007000	0.100000	•	
84		0.007000	0.100000	•	
82	.	٠	000001	•	
88	•	0.0010.0	0.00001-0	•	
84	o (٠	0.0000	0.20000	
88		•	0.10000	•	
68	ċ	•	0.0000	٠	
26	• •	٠	0.00000	٠	
91	oʻ.	•	0.07000	0.00000	
76	ė (•	0000000	•	
	.		•	•	
110	- 0		•	•	
7 -	00000	٠		•	
2118	⊣		•	•	
-	000000	000.00	١		

MO.2 BA,2	0.000500 0.100000 0.001500 0.200000 0.001500 0.200000 0.0 N 0.070000 0.0 N 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000 0.0 0.000300 0.050000	ZZ
Z N 2	0.00 0.00 0.00 0.00 0.00 0.00 0.00 0.0	
£400	0.000500 0.003000 0.003000 0.003000 0.000300 0.000300 0.000300 0.0003000	0.000700
AS &	0.0	0.014000
HG• %	0.000010 0.000010 0.000011 0.000011 0.000011 0.000013 0.000013 0.000013 0.000013 0.000010	0.000011
81,2	0.00 0.00 0.00 0.00 0.00 0.00 0.00 0.0	0.0 0.0
PB, %	0.005000 0.007000 0.030000 0.030000 0.010000 0.050000	0.007000
A6, %	0.000200 0.00 0.00 0.00 0.00 0.000100 0	00200
AU, Z	0.000010 0.000010 0.0000236 0.0000160 0.000160 0.000160 0.000270 0.000270 0.000270 0.000270 0.000270 0.000270 0.000028	0.000040
AMPLE	120 121 129 129 161 162 163 164 165 167 171 171 173 173 174 174 177 177 178 179 170 171 171 171 173 174 174 177 178 178 179 170 170 170 170 170 170 170 170 170 170	221

» * ≻	ZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZ	
>*	0.005000 0.005000	
SR. 3	0.070000 0.050000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.050000 0.070000 0.050000	; ; ;
NI.	Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z	
~		i + +
NB, &	0.000000000000000000000000000000000000))))
N. N.	0.001000 0.200000 0.030000 0.01500 0.001500))))
LA,%	0.015000 0.0010000 0.00000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.00000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.00000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.00000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.00000 0.00000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.000	
CR, %	0.001000 0.003000 0.001500	•
% •00		•
BE, %	Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z	E
æ		•
SAMPLE	120 121 121 121 122 123 124 125 127 127 127 127 127 127 127 127 127 127	777

11,8	0.300000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000	0.200000 0.200000 0.200000 0.200000 0.200000 0.200000 0.300000 0.300000 0.300000 0.300000 0.300000
	000000000000000000000000000000000000000	
CA,%	030000 0100000 020000 030000 030000 030000 0100000 010000 010000 010000 010000 010000 010000 010000 010000 010000	050000 070000 100000 100000 100000 070000 070000 070000 070000 070000 070000 070000
ٽ	0.030000 0.020000 0.020000 0.020000 0.100000 0.100000 0.100000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000 0.070000	0.05000 0.050000 0.050000 0.100000 0.100000 0.100000 0.050000 0.070000 0.070000 0.070000 0.070000 0.070000
6 4	Z 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	0000 0000 0000 0000 0000 0000 0000 0000 0000
MG, %		01000000000000000000000000000000000000
FE , &	999998 999998 999998 999998 999998 999999	999999 999999 999999 999999 999999 100000 999999 000000 999999 999999 999999
		000000000000000000000000000000000000000
AMPLE	120 120 120 130 130 162 163 163 173 173 173 173 173 173 173 173 173 17	185 186 187 190 190 191 195 195 196 197 200 220 222
SAM	Ē 4 4	

SILICIFIED ROCKS, PERCENT

BA, X	0.030000	000001.0	0.01000	0.050000	0.100000	0.100000	0.100000	0.070000	0.050000	0.070000	0.01000	0.050000	0.050000	0.07000	0.100000	0.070000	0.100000	0.070000	0.070000	0.070000	0.070000	0.02000	0.01000	0.02000	0.05000	0.0000	0.050000	0.300000	0.05000	0.05000	0.050000	0.050000
MO, %	Z	Z	200	Z	Z	300	200	200	100	100	z	000	300	300	000	000	500	Z	Z	000	200	300	300	z	300	z	200	Z	z	500	z	z
X.	0.0	0.0	0.000500	0.0	0.0	0.000300	0.000500	0.000500	0.000700	0.000700	0	0.001000	0.000300	0.000300	0.001000	0.001000	0.000500	0.0	0.0	0.001000	0.000700	0.000300	0.000300	0.0	0.000300	0.0	0.001500	0.0	0.0	0.000500	0.0	0.0
2 · N 2	. ب	ب	_		_	_	_	_	ب	000	000	000			000	_	_	000	ب	_		ب.	ب	000	000	ب.		ب	_	ب	_	٠.
7	0.0	0.0	0.0	0.0	0	0.0	0.0	0.0	0.0	0.004000	0.004000	0.005000	0.0	0.0	0.004000	0.0	0.0	0.003000	0.0	0.0	0.0	0.0	0.0	0.003000	0.004000	0.0	0.0	0.0	0.0	0.0	0.0	0.0
CU,%	0000	.000	0020	0020	200	000	000	500	000	0009	000	0000	000	000	000	200	000	0020	000	0009	000	000	000	200	000	500	0000	1500	200	000	500	000
O	0.003000	0.007000	0.000700	0.000700	0.000500	0.001000	0.007000	0.001500	0.00500	0.015000	0.003000	0.02000	0.007000	0.003000	0.007000	0.001500	0.005000	0.000700	0.00200	0.005000	0.007000	0.003000	0.001000	0.001500	0.002000	0.001500	0.015000	0.001500	0.000200	0.001000	0.001500	0.002000
AS, &	0009	0001	000	0001	0001	000	3000	2000	3000	2000	0001	000	0001	0009	0009	0008	0001	0001	000	3000	0000	_	0001	0001	0001	0001	2000	0001	0001	2000	000	000
_	00090000	0.001000	0.001000	0.001000	0.001000	0.004000	0.08000	0.012000	0.008000	0.012000	0.001000	0.014000	0.001000	0.00900	0.016000	0.008000	0.001000	0.001000	0.004000	0.008000	0.010000	0.0	0.001000	0.001000	0.001000	0.001000	0.012000	0.001000	0.001000	0.012000	0.002000	0.004000
HG, %	0100	0100	3008	0100	9000	0100	0100	2004	0100	0100	0700	0029	0023	0100	2012	9019	2013	0200	2024	3018	0010	7064	2017	8900	2012	2015	0024	9000	0100	0019	0030	0600
	0.000010	0.000010	0.00008	0.000010	0.00000	0.000010	0.000010	0.00000	0.000010	0.000010	0.000020	0.000029	0.000023	0.000010	0.000012	0.000019	0.000013	0.000020	0.000024	0.00	0.000100	00.0	00.0	0.000068	0.000012	0.000015	0.000024	0.00000	0.000010	0.000019	0.000030	0.000030
8 i s	2000	Z	z	z	1000	3000	1500	z	3000	3000	2000	0.02000	3000	3000	3000	3000	Z	2000	3000	2000	7000	2000	1000	3000	2000	2000	Z	Z	z	z	1000	1000
	0.002000	0.0	0.0	0.0	0.001000	0.003000	0.001500	0.0	0.003000	0.003000	0.002000	0.02	0.003000	0.003000	0.003000	0.003000	0.0	0.002000	0.003000	0.007000	0.007000	0.007000	0.001000	0.003000	0.007000	0.005000	0.0	0.0	0.0	0.0	0.001000	0.001000
PB,%	0.02000	0.005000	0.005000	0.003000	0.005000	0.020000	0.010000	0.010000	0.030000	0000	0.05000	0.00000	0.030000	0.030000	0.02000	0.010000	0.02000	0.05000	0.010000	0.100000	0.05000	0.02000	0.015000	0.010000	0.030000	0.02000	0.005000	0.02000	0.00000	00000000	0.007000	0.005000
	0.02	0.00	00.0	0.00	00.0	0.02	0.01	0.01	0.03	0.030000	0.05	0.07	0.03	0.03	0.02	0.01	0.02	0.05	0.01	0.10	0.05	0.02	0.01	0.01	0.03	0.02	00.0	0.02	0.01	0.03	00.0	00.0
AG, %		_	Z	_	0.000150	0.000300	0.000150		0.001000	0.003000	0.000100	0.005000	0.003000	Z	Z	Z	Z	0.000100	0.001500	0.000150	0.002000	0.003000	0.000500	0.003000	0.003000	0.000000	0.000200		ب (0-000200	0.000200	0.000200
	0.0	0.0	0.0	0.0	0.00	0.00	00.0	0.0	00.0	00.0	0.0	00.0	0.00	0.0	0.0	0.0	0.0	00.0	0.00	00.0	00.0	0.00	00.0	0.00	00.0	00.00	00.0	0.0	0.0	00.0	00.0	00.0
AU, E	0.000030	0.000010	0.000037	0.000140	0.000116	0.000230	0.001407	0.000058	0.000210	0.000340	0.000330	0.000427	0.000119	0.000118	0.000290	0.000105	0.000180	0210	0156	0.000536	0.000419	0.000527	0.000226	0.000425	0.001860	0.000400	0.000267	0.000015	0.000049	0.000000	0.000057	0.000081
	0.00	0.00	00.00	00.0	00.0	00.0	00.0	00.0	00.0	00.0	00.0	00.0	00.0	0000	00.0	00.0	00.0	00.0	00.0	00.0	0.00	00.0	00.0	00.0	00.0	00-0	0.0	00.0	00.0	0,0	00.0	0.00
SAMPLE	223	240	246	249	250	251	252	254	255	256	257	258	259	260	261	262	263	264	265	266	267	268	269	270	27.	27.2	273	274	275	276	277	278
SA																																

× ,×	z	2	909	*	Z	z	z	z	Z	Z	z	z	Z	Z	Z	¥	Z	Z	Z	Z	Z	Z	Z	Z	Z	z	Z	Z	Z	Z	z	Z	
. •	0.0	9	001200	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	
>	0.00500	0.00700	0.007000	0.001000	0.00700	0.007000	0.007000	0.007000	0.015000	0.007000	0.005000	0.007000	0.007000	0.005000	0.015000	0.007000	0.010000	0.01000	0.002000	0.001000	0.007000	0.007000	0.010000	0.005000	0.007000	0.007000	0.007000	0.005000	0.010.00	0.010000	0.007000	0.007000	
SR,%	0.050000	0.050000	0.150000	0.050000	0.100000	0.070000	0.100000	0.070000	0.050000	0.050000	0.100000	0.100000	0.010000	0.150000	0.100000	0.030000	0.150000	0.010000	0.030000	0.100000	0.100000	0.030000	0.050000	0.030000	0000000	0.01000	0.007000	0.010000	0.030000	0.070000	0.100000	0.070000	
NI ,	Z	Z	z	Z	Z	Z	Z	z	z	Z	z	z	z	z	z	z	z	z	z	z	z	z	z	Z	Z	Z	z	z	Z	z	Z	Z	
~	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	
NB.	N 0.0	0.000100	0.000100	0.000700	0.000500	0.000500	0.000500	0.000100	0.0	0.000500	0.000500	N 0:0	0.0	0.000500	0.000500	0.000500	0.000500	0.000500	0.000500	0.000300	0.0	0.000300	0.0	0.0	0.0	0.0	0.000500	0.000500	0.000500	0.000500	N 0.0	N 0.0	
N. W	0.000300	0.001000	0.002000	0.000500	0.000500	0.001000	0.001000	0.003000	0.02000	0.005000	0.000500	0.007000	0.000100	0.070000	0.100000	0.070000	0.100000	0.005000	0.001500	0.003000	0.010000	0.003000	0.000100	0.000500	0.000700	0.000500	0.030000	0.001000	0.000700	0.001500	0.000500	0.000000	
LA.%	0.0	0.005000	0.007000	0.0	5000	20.0	25000	0.002000	20.0		0.0		. O	0.005000	0.005000	0.005000	0.005000	0.005000	N 0.0	0.0	5000	0.005000	0.005000	0.0		N 0.0							•
CR, &	0.002000	0.003000	0.003000	0.003000	0.003000	0.003000	0.003000	000000	0.001000	0.001000	0.002000	0.003000	0.003000	0.003000	0.002000	0.005000	0.002000	0.002000	0.001500	0.002000	0.062000	0.002000			0-002000	0.08280	0-002000	0.002000	0.002000	0.002000	00000	00200	3
£ 000	N 0.0	2	2	2	2	2	2	000100		2	2	. z	2	2	2	2	2	2	2	2	2	2	2 2	2	2	2	0.001500	•	}	: Z		2 2	
BE, %	z	Z	: 2	: 2	: 2	2 2	2 2	2 2	: 2	: 2	: 2	: 2	: 2	: 2	: 2	: 2	2 2	: 2	: 2	: 2	: 2	: 2	: 2	: 2	: 2	2 2	2	2	2 2	. 2	: 7	2 2	2
a	0-0	0				•	•	•	• c					• c	•		•	•		•	•	•			•		•		•		•	•	•
SAMPLE	223	070	246	07¢	030	25.30	250	720	255	256	25.5	950	250	240	26.1	107	707	776	107	697	296	197	092	270	ב ל ב ל	27.2	272	37.6	117	250	9 1	777	917

• .															•																	
11,2	0.30000	0.500000	0.30000	0.300000	0.300000	0.300000	0.300000	•	0.20000		٠	0.150000	0.300000		•		0.300000	•	0.300000	0.20000	0.300000	000008.0	0.300000	•	0.300000	٠	0.300000	0000000	0.300000			0.300000
CA, 2	0.050000	0.01000	0.010000	0.01000	0.150000	0.010000	0.150000	0.05000	0.00000	0.02000	0.010000	0.100000	0.05000	0.070000	0.100000	0.050000	0.500000	0.150000	0.05000	0.100000	0.10000	0.100000	0.200000	0.100000	0.150000	0.150000	0.030000	0.070000	0.150000	0.05000	0.050000	0.150000
M.C.	0.000700	0.007000	0.100000	0.005000	0.005000	0.005000	0.007000	0.00700	0.02000	0.015000	0.00500.0	0000000	0.00700	0.01000	0.02000	0.015000	0.010000	0.015000	0.010000	0.015000	0.015000	0.015000	0.015000	0.010000	0.010000	0.010000	0.050000	0.020000	0.020000	0.030000	0.010000	0.010000
FE 9.8	0.100000	0.100000	2.999998	0.300000	0.500000	1.499999	1.499999	1.499999	4.999998	2.999998	0.300000	166666.9	1.499999	1.499999	1.999999	1.999999	1.000000	0.500000	1.000000	1.999999	2.999998	1.000000	1.000000	0.70000	0.500000	0.700000	6.99997	0.500000	0.200000	2,999998	1.999999	1.499999
SAMPLE	223	240	246	546	250	251	252	254	255	256	257	258	259	260	261	262	263	264	265	266	267	268	269	270	271	272	273	274	275	276	277	278

SILICIFIED ROCKS, PPM

ВА,РРМ	700.00	700.00	700.00	700.00	700.00	700.00	700.00	700.00	700.00	700.00	300.00	2000.00	200.00	700.00	150.00	200.00	100.00	1000-00	1000-00	200.00	00.00	000.001	00.0001	00.007	3000-00	200.00	200.00	700.00	700.00	200.00	200-00	200.00	500-00	100.00	00.00	200.00	700.00	200.002	100.00	000	00.00	300.00	200.00	2000-00	200.00	1500.00	200.00	1000.00	700.00	700.00		
MO, PPM	N 0.0	0.0	0.0	Z 0.0	200	Z 0.0	N 0.0	N 0.0	N 0.0	N 0.0	30.00	10.00	7.00	3.00	10.00	3.00	3.00	N 0	3.00	Z 0.0	15.00	5.00	2.00	Z :	N 0 0	2.00	Z :	N 0.0	7.00	2.00	2.00	2.00	10.00	2 3	2 0 0	z :	2 2	2 2	2 2	2 2	2 :	2 :	2 :	N 0.0	10.00	3.00	2.00	Z 0.0	10.00	₹ 0.0		
ZN, PPM	10.0	0.0	0.0	10.0	1000	- 0.0	0.0	0.0	0.0 L	0.0 L	0.0 L	0.0	0.0 L	0.0 L	25.00	0.0 L	0.0 L	0.0	0.0 L	0.0	25.00	30.00	J 0.0	0°0 L	0.0 L	0.0	0.0	100.00	0.0	20.00	40.00	140.00	0.0	0.0	0.0	J 0.0	J 0 0	0.0	J 0 0	1 0°0	40.00	0.0	0.0	0.0	0.0	30.00	50.00	0.0	0.0	0.0		
CU,PPM	20.00	7.00	20.00	15,00	7.00	10.00	100.00	15.00	30.00	50.00	200.00	30.00	30.00	10.00	20.00	30.00	15.00	2.00	7.00	30.00	20.00	15.00	15.00	10.00	20.00	2.00	7.00	10.00	70.00	10.00	15.00	20.00	70.00	10.00	10.00	10.00	15.00	10.00	15.00	15.00	15.00	2.00	20.00	7.00	10.00	10.00	20.00	20.00	2.00	2.00		
AS, PPM	60.00	00-07	10.0	120-00	10.00		10.00	10.00	10.00	120.00	160.00	10.00	00.09	10.00	10.00	0.0	0.0 L	10.00	10.00	10.00	00.09	80.00	40.00	100.00	10.00	100.00	80.00	10.00	140.00	20.00	160.00	140.00	200.00	10.00	10.00	10.00	20.00	20.00	20.00	10.00	20.00	10.00	10.00	0°0	00.09	10.00	100.00	20.00	40.00	10.00		
н 6 , РРМ	0	000					0.50	0.20	07.0	0.20	0.50	0.11	0.08	0.11	90.0	90.0	0.15	0.15	0.24	0.20	0.13	0.10	0.08	90.0	0.11	90.0	90.0	90•0	0.15	0.15	0.17	0.11	0.15	0.11	0.19	0.13	0.11	90.0	0.11	0.0	0.04	90.0	90.0	0.11	90.0	0.10	0.10	0.10	0.10	0.08		
BI , PPM	2	2 2	2 2				15.00	00.01	00.01	N 0.0	30.00	10.00	5.00	0.0	0.0	0.0	0.0	0.0	0.0	20.00	N 0.0	0.0	N 0.0		0.0	0.0	N 0.0	N 0.0	2000	0.0	N 0.0	N 0.0	. N 0.0	30.00	30.00	N 0.0	N 0.0	N 0.0	30.00	30-00	30.00	N 0.0	0.0	0.0	₹ 0.0	0.0	0.0	Z 0.0	0.0	0.0	!	
P8.PPM	000	100.00	00.001	2000	00.00	100.00	130.00	00000	200.00	200-002	20.00	500-00	150.00	300.00	300.00	500.00	700.00	300.00	700.00	50.00	30.00	N 0.0	200.00	100.00	150.00	700.00	100,00	30.00	200.00	150.00	500.00	150.00	70.00	200.00	200.00	150.00	200.00	150.00	150.00	300.00	200.00	100.00	300.00	150.00	150.00	300.00	7.00	7-00	15.00	70.00)	
AG, PPM	•	00-01	1.50	00.7	00.1	00.2	3.00		9.00	00.0	00.4	1.50	3,00	0.30	0.70	2.00	1.50	2,00	2.00	2.00	2.00	N 0.0	N 0.0			Z 0 0		200	200			1 0 0		5.00	3.00	3.00	1.00	1.00	2.00	3.00	3.00	3.00	1.00	1.50	0.50	1.00	0.0	0.0	- 0	2 0 0		
AU, PPM	,	08.1	88.0	1.50	00.1	0.69	2.00	00.	3.10	0	1 • 00 2 20	7. 7. 7. 7. 7. 7. 7. 7. 7. 7. 7. 7. 7. 7	1.45	1.50	- 0-0	2.60	7.08	66.4	4.30	2.88	0.10	0.10	1.40	06-0	2.75	2.07	0-10	0.18	1.80	0.30	2.40	96.0	1.50	2.46	2.90	1.20	1.90	0.16	1.70	1.40	06.0	0.55	9.65	9.40	0.10	2.26	0.16			0.79	•	
SAMPLE		_	5	m	- ≢⊤1	'n	91	_	30 (^	3:	11	200	76	5 6	7 7	2 7	- X	2 2	45	63	3 2	59		29	. 84	9	6 22	<u> </u>	12	- 82	2	. e	8	82	83	**	82	98	87	60	6	6	2 5	6	101	711	117	91.	110	C	

DATE 5/28/71

SILICIFIED ROCKS, PPM

Y, PPM	0.0	0	200	2 0 0	Z 0	N :	Z :	Z :	O .	2000	0	2.00	2.00	7.00	7.00	2.00	2.00	2.00	2.00	7.00	2.00	200	20.00	N 0.0	20.00	2 :	2	Z :	•	•	•		•		•	2 2	•	2000	2000	3.00	3.00	3.00	3.00	20.00	10.00	•	• c	2 2	2	
V,PPM	10.00	70.00	70.00	100.00	100.00	10.00	10.00	70.00	70.00	20.00	200.00	100.00	70.00	70.00	150.00	10.00	10.00	10.00	10.00	30.00	20.00	20.00	100.00	10.00	70.00	70.00	70.00	70.00	100.00	70.00	70.00	10.00	70.00	30.00	30.00	00.00	00.0	20.07	70.00	70.00	15.00	10.00	20.00	100.00	50.00	20.00	30.00	٠	20.00	
SR,PPM	700.00	200.00	700.00	500.00	300.00	300.00	200.00	200.00	500.00	500.00	700.00	200.00	500.00	1500.00	1500.00	200.00	200.00	1500.00	2000.00	300.00	1000-00	100.00	1000.00	200.00	300.00	700.00	1000.00	200.00	200.00	100.00	1000.00	200.00	700.00	500.00	500.00	300.00	00.007	500.00	500.00	500.00	100.00	1000.00	150.00	300.00	150.00	200.00	20.00	100.00	00.00	
NI,PPM	N 0.0	× 0.0	Z 0.0	0.0	0	N 0.0	z 0.0	Z 0.0	0.0 N	0.0	N 0.0	2.00	2.00	3.00	2.00	2.00	2.00	2.00	2.00	2.00	3.00	Z 0.0	0	N 0.0	N 0 0	Z 0.0	2 0	0.0 N	N 0.0	N 0.0	0.0 N					2 2				0.0	0.0	0.0	N 0.0	3.00	3.00	Z .	Z :	Z :	2000	
NB . PPM	N 0.0	N 0.0	N 0.0	N 0.0	Z 0.0	N 0.0	N 0.0	N 0.0	0.0	Z 0.0	N 0.0	N 0.0	Z 0.0	N 0.0	0.0	× 0.0	Z 0.0	2 0.0	Z 0.0	10.00	Z 0.0	Z 0.0	N 0.0	Z 0.0	0	Z 0.0	Z 0.0	Z 0.0	0.0	0.0	0.0	0.0 0	0.0	N 0.0	N :	N 000	2 2	2 2	2 2	N 0.0	7.00	0.0	N 0.0	0.0 0	N 0.0	20.00	20.00	20.00	10.00	
Mdd.NM	150.00	15.00	15.00	30.00	2.00	20.00	20.00	20.00	30.00	70.00	300.00	50.00	50.00	70.00	100.00	20.00	20.00	15.00	20.00	30.00	300.00	300.00	30.00	10.00	30.00	10.00	10.00	7.00	30.00	7.00	7.00	7.00	15.00	70.00	10.00	7.00	00.0	10.00	2.00	15.00	3.00	7.00	10.00	100.00	10.00	15.00	10.00	٠	10.00	
LA, PPM	N 0.0	N 0.0	N 0.0	N 0.0	Z 0.0	0.0	0.0 N	0.0	₩ 0°0	N 0.0	0.0	30.00	N 0.0	70.00	70.00	0°0	20.00	50.00	30.00	70.00	50.00	N 0.0	70.00	70.00	70.00	N 0.0	70.00	70.00	70.00	0.0	N 0.0	N 0.0	₩ 0°0	N 0.0	O.0	N :	20.01	00.00		30.00	Z 0.0	50.00	30.00	50.00	20.00	100.00	0.0	200	10.00	
CR, PPM	10.00	10.00	10.00	10.00	15.00	10.00	10.00	10.00	10.00	10.00	10.00	20.00	15.00	30.00	20.00	20.00	20.00	15.00	15.00	10.00	10.00	20.00	20.00	20.00	20.00	15.00	20.00	15.00	15.00	15.00	10.00	15.00	10.00	15.00	15.00	20.00	20.00	00.00	20.00	10.00	3.00	15.00	15.00	10.00	15.00	30.00	N 0 0	(L)	20.00	
CO, PPM	N 0.0	N 0.0	N 0.0	N 0.0	N 0.0	0.0	2 0°0	0.0	0.0	N 0.0	30.00	5.00	0.0	5.00	N 0.0	N 0.0	0°0	N 0.0	2 0°0	7.00	N 0.0	0°0	N 0.0	Z 0°0	N 0.0	N 0.0	N 0.0	0.0	0.0	N 0.0	N 0.0	N 0.0	N 0.0	0.0	N 0.0	N 0 0	2 0	2 2	2 2	2		N 0.0	0.0	0.0	N 0.0	N 0.0	•		Z 0 • 0	
8E, PPM	0°0	0.0	0.0	0.0	N 0.0	0.0	N 0.0	N 0.0	N 0.0	0.0	0.0	0.30	0.30	0.50	0.30	0.30	0.30	0.30	0.30	0.50	0.30	N 0.0	0.0	N 0.0	0.0	N 0.0	0.0	0.0	N 0.0	N 0.0	N 0.0	0.0	N 0.0	N 0.0	0.0	Z 0 0	2 :	Z 2			0.50	0.50	0.0	0.10	1.50	N 0.0	0.0		N 0.0	
SAMPLE		7	m	· 4	M.	•	1	∞	6	10	11	30	35	34	35	36	37	38	39	45	63	49	65	99	19	89	69	74	22	11	78	62	80	81	85	83	8 8	\$ 32	9 6	- a	8	6	16	92	101	116	117	118	119	

SILICIFIED ROCKS, PPM

																																								-									
	3000,000	0	3000.00	3000.00	3000.00	3000.00	•	3000.00	000	1500.00	•	2000.00	•	2000.00	3000.00	•	2000-00	•	•	2000-00	20002	2000	3000	•	1500.00	_	2	2000-000		1500.00	000		000	•	1500.00	2000-00	2000-00	2000	2000	•	00.000		00.000			000	0	000	
CA, PPM		700.00	0	200-00	1000.00		•		•	500.	•		•	•	•	500	500	000	1000.00	300.00	200.00	00.007	00.001	200.00	200.00	00.00	00.001	200.00	200,002	700-00	500.00	300.00	•	700.	000	•	200	00.0001		00.001	1600.00	•	•		200.00		9		•
10. FYE	00 001		100.00	70.00	100.00	100.00	100.00	70.00	100.00	70.00	700.00	100.00	150.00	2000.00	200.00	-	150.00	o	•		200	000	3	.	3 0	2 00 00	3 6	2 2			. Z	0	70.00	100.00	•	20	100.00	100.00	100.00	150.00	00.00	00.07	00.00	0000	30-00		30.00	•	2
71,77,2	0	5000.00	666	2000		000	966	14999.99	19999.99	19999.99	9 0.0	666669	49999.98	6.6666	6.	6-6	6666	0	6.6666	19999.99	0.0	6666	6666	•6666	4999.	9999.	99999.	0000	*****	9000		6666	3000.0	2000.00	3000.00	49999.98	000	9666		000	000	•	2000.	*6*6669	000001			5000 O	
SAMPLE	•	-	u #	n <	+ w	•	~	· 00	6	10		30	32	34	35	96	37	38	39	45	63	49	65	99	67	89	69	7.	c	- 02	0 0	80	81	82	83	84	82	98	87	88	68	06	16	76	101	110	116	011	KIT

SILICIFIED ROCKS, PPM

																																٠																					
BA, PPM	1000.00	300.00	2000		00.00	700.00	700.00	300.00	700.00	200.00	200.002	00.00	00.000	100.00	700.00	500.00	500.00	700.00	700.00	700-00	200	000	200-000	200.00	200.00	200.00	700.00	500.00	700.00	700.00	700.00	500.00	500.00	500.00	200.00	0000	0000	00.007	00.00	00.00	300.00	100.00	700.00	200.00	500.00	1000.00	500.00	500.00	500.00	200-00	000	00.000	
MO, PPM	2.00	7.00	00	000	12.00	Z 0	Z 0.0	Z 0.0	5.00	Z	2		20.	Z 0.0	2000	Z 0.0	N 0 0	Z 0.0	2	2			3.00	2 0	3.00	2 0.0	0.0	N 0.0	0.0	7.00	10.00	10.00	3.00	2	2				200	2 0 0	3.00	Z 0.0	3.00	3.00	10.00	5.00	N 0.0	Z 0.0	10.00	2	2		
ZN, PPM	0.0	65.00		00.00	90.00	50.00	00.09	160.00	25.00	30.00		00.00)))	0.0 L	40.00	0.0 L	0.0	0.0	1 0 0	00.04			0.0	25.00	0.0	0.0	0°0	0.0 L	0.0	0.0	0.0	0.0	0.0	-	100			200	00.00	0.0	30.00	0.0 L	0.0 L	20.00	0.0 L	20.00	0.0	0.0	60.00	20-02			
CU, PPM	2.00	30.00		00.00	20.00	3.00	10.00	3.00	15,00	3.00	3	00.00	20.00	15.00	15.00	15.00	15.00	100-00	20.00	20.00		20.00	20.00	2.00	30.00	5.00	10.00	10.00	15.00	15.00	20.00	30.00	50.00	200	00.00		00.00	000	20.00	2.00	15.00	2.00	20.00	7.00	30.00	7.00	5.00	7.00	200.00	20.7		00.00	
AS, PPM	0.0	80.00		20000	160.00	10.00	120.00	10.00	100.00	20.00		00.07	220.00	160.00	10.00	20.00	80.00	160.00	140.00	00.071	00.01	00.01	140.00	140.00	160.00	120.00	100.00	20.00	40.00	80-00	80.00	10.00	60-00	00.00	200	000	00.01	00.00	00.07	0.0	10.00	10.00	10.00	10.00	10.00	10.00	10.00	0.0	160.00	140-00	00000	100.001	
HG, PPM	0.10	0.10	,	90.0	61.0	0.52	0.11	0.35	0.35	22.0	77.0	0.21	87.0	0.24	0.13	0.24	0.28	21.0	64.0	2 4 6	•	0.32	0.27	09.0	0.0	0.23	0.14	0.04	0.10	0.10	0.10	01.0	01.0			2.0	0.10	0.10	C1.0	0.17	0.10	0.11	0.13	0.15	0.08	0.10	90.0	90-0	90-0	· ·	100	01.0	
81,PPM	× 0°0	2		2	Z 0.0	Z 0.0	N 0.0	N C C	30.02	200	2 2	2 :	200	2.00	0.0	0°0	N 0.0	30.00	900		2 2	Z .	15.00	30.00	30.00	0.0	N 0	N 0.0	200	2000	200	15,00	2			2 2	2 2	2 2	200	0	× 0.0	0°0	N 0.0	N 0.0	0.0	0	0.0 N	2000	15.00	2000	٠	2	
PB. PPM	50.00	70.00		20.00	300.00	300.00	700.00	20.00	2000		00.001	200.00	300.00	500.00	150.00	200.00	200,00	200 002	0000		200-000	200.00	500.00	500.00	500.00	700-00	300.00	500,00	200,00	500.00	1500.00	200.00	00.002	00.007		00.00	200.00	00.00	300.00	200.00	7.00	150.00	150.00	500.00	700.00	70.00	50.00	500,00	00000	20.00	00.01	00.00	•
AG, PPM	2.00	2		2.0	0.0 L	Z 0.0	0.0		ò		J .	0.0	2.00	1.00	70.0	0.0 L	7.00	200		00.01	10.00	2.00	5.00	7.00	20.00	3.00	2.00	10.0	100-1	3,00		20.0	00.0	00.5	2.00	00.6	7.00	0°0	1.00	2.00	3.00	0.0 L	0.0 L	0.0	0.0	2.00	3,00	200			00.5	2.00	
AU, PPM	0.10	71.0		07.0	2.36	1.60	1.60	0 0	07.01	20.01	10.00	7.97	29.70	2.16	0.00	3.96	2.70	00	0.00		06-2	2.20	2.70	1.66	5.10	1-60	0.68	0.50	200	100	07. 4	200	72 80	00.67	00.0	0.28	0.20	2.10	0.58	1. 00	0.80	1.36	0.20	7.50	2.70	0.40	0.25	000		200		0.98	
SAMPLE	120	121	171	129	130	142	191	142	707	100	001	167	168	169	170	171	172	1.5	C 1	* ! !	<u>C</u>	176	177	178	179	180	181	182	182	781	501	701	100	701	007	691	261	161	192	193	194	195	196	197	65	199	200	202	330	777	177	777	

SILICIFIED ROCKS, PPM

Y,PPM	Z C		5	2	0	0	0	0	٥	0	0	0	0	N 0.0	0	0	0	0	0	0	a	0	0	0	Z 0.0	N 0.0	× 0.0	Z 0.0	N 0.0	0.0 N	0.0	0.0 N	0.0	N 0.0	0.0	N 0.0	0.0 N	0°0	10.00	10.00	0°0	10.00	30.00	N 0.0	•	N 0.0	0.0	20.0	
V.PPM	50.00		00.00	30.00	20.00	50.00	70.00	70.00	100.00	50.00	70.00	50.00	50.00	70.00	100.00	20.00	100.00	70.00	20.00	70.00	70.00	20.00	20.00	100.00	70.00	70.00	20.00	70.00	30.00	70.00	70.00	30.00	20.00	10.00	10.00	70.00	20.00	15.00	70.00	70.00	70.00	70.00	70.00	70.00	70.00	70.00	20.00	30.00	
SR, PPM	700.00	•	00.00	200.00	200.00	300.00	700.00	100.00	1000.00	500.00	700.00	700.00	700.00	1000.00	500.00	700.00	700.00	700.00	200.00	700.00	500.00	500.00	200.00	700.00	700.00	700.00	200.00	700.00	700.00	500.00	700.00	500.00	70.00	1500.00	300.00	1500.00	700.00	10.00	150.00	500.00	500.00	700.00	100.00	700.00	-	700.00	700.00	200.00)
NI , PPM	_	٠,	5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	N 0.0	0	0	0	0	0	0	0	0	0	0	0	0	0	,
NB PPM	00 00	•	2	•	•	2.00	N 0.0	0.0	Z 0.0	N 0.0	0.0	0.0	0.0	10.00	7.00	7.00	N 0.0	7.00	N 0.0	0.0	N 0.0	N 0.0	0.0	0.0	N 0.0	N 0.0	2.00	N 0.0	N 0.0	0.0	N 0.0	0.0	2.00	3.00	3.00	3.00	3.00	2.00	N 0.0	0°0	0.0	0.0		0.0		N 0.0		3.00	•
MN. PPM	00	00.01	10.00	2000.00	300.00	2.00	150.00	7.00	10.00	2.00	15.00	15.00	2.00	1.00	7.00	7.00	15.00	10.00	20.00	10.00	30.00	2.00	5.00	7.00	2.00	2.00	5.00	10.00	100.00	_	15.00	5.00	2.00	15.00	10.00	1.00	2.00	2 0°0	2.00	100.00	50.00	7.00	20.00	3.00	150.00	70.00	2.00	2.00)))
LA, PPM		150.00	70.00		0	0	0	0	0	0	0	a	0	0.0	a	0	0	0	0	0	0	0	0	0	0	N 0.0	30.00	50.00	0.0	0	N 0.0	0	0.0	50.00	30.00	Z 0.0	N 0.0	N 0.0	30.00	0.0	20.0	N 0	50.00	30.00	ċ		0.0	N 0.0	•
CR + PPM	•	•	ં	Z 0.0	15.00	15.00	O		0	7		3.00		15.00	0	15.00	15.00	15.00	15.00	15.00	7.00	7.00	15.00	15.00	20.00	20.00	15.00	20.00	50.00	15.00	15,00	5.00	2,00	20.00	15.00	20.00	20.00	N 0.0	0	15.00	30.00			30.00	20.00	20.00			3
CO, PPM		2	0	10.00	0	0	0	_		_	_	. ~			. ~			6	_		_		_	. ~		. ~								0	2000	0	0	0	0						0	0	0		
BE, PPM	-	_		_		•						2	_	_	-		2	20.0		_			_	_	Z 0 0	200	2000	_		-					. N												20.0	2	!
SAMPLE	,	120	121	129	130	142	161	162	163	166	167	891	691	170	2.	172	173	174	175	176	177	178	179	180	181	187	183	761	- K	186	187	38	180	190	161	192	193	194	195	961	197	8	001	200	207	220	221	222	j

SILICIFIED ROCKS, PPM

		,
TI, PPM		3000.00 3000.00 2000.00 3000.00 2000.00
CA, PPM		700.00 1000.00 500.00 150.00
MG, PPM		100.00 69999.94 300.00 50.00 7.00
FE, PPM		3000.00 3000.00 99999.94 2000.00 2000.00
SAMPLE	120 120 121 123 124 125 126 127 127 127 127 127 127 127 127 127 127	200 201 220 221 222

SILICIFIED ROCKS, PPM

0.0 L 200.00 20.00 0.10 60.00 70.00 0.0 L 0.0 N 300.00 0.0 L 50.00 0.0 N 0.00 10.00 7.00 0.0 L 0.0 N 10.00 7.00 0.0 L 0.0 N 10.00 0.0 L 0.0 L 0.0 N 10.00 0.0 L 0.0 L 0.0 N 10.00	_	IU*PPM AG*PPM	PB, PPM	BI, PPM	HG, PPM	AS. PPM	CU, PPM	ZN, PPM	MO. PPM	ВА,РРМ
50.00 0.0 N 0.10 10.00 70.00 0.0 L 50.00 50.00 0.0 N 0.08 10.00 7.00 0.0 L 5.00 0.0 L 0.0 L 0.0 L <	0.30 0.0	_	200.00	20.00	0.10	00.09	30.00	0.0 L	N 0.0	300.00
50.00 0.0 N 0.08 10.00 7.00 0.0 L 5.00 50.00 10.00 0.00 10.00 5.00 0.0 L 0.0 N 50.00 10.00 0.00 10.00 0.0 L 0.0 L 0.0 N 200.00 10.00 0.01 40.00 10.00 0.0 L 5.00 100.00 0.00 0.01 120.00 15.00 0.0 L 5.00 100.00 0.00 0.01 120.00 15.00 0.0 L 5.00 100.00 0.00 0.01 120.00 15.00 0.0 L 5.00 100.00 0.00 0.10 120.00 15.00 10.0 10.0 10.0 200.00 20.00 0.10 10.00 10.00 10.0		ي.	50.00	N 0.0	0.10	10.00	10.00	0.0	Z 0 0	1000.00
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700.00 200.00 0.29 140.00 200.00 50.00 10.00 300.00 30.00 0.123 10.00 70.00 0.01 3.00 200.00 30.00 0.12 160.00 70.00 0.01 3.00 200.00 30.00 0.12 160.00 70.00 0.01 10.00 200.00 30.00 0.13 10.00 70.00 0.01 10.00 200.00 30.00 0.24 40.00 20.00 0.01 10.00 100.00 30.00 0.24 40.00 20.00 0.01 10.00 100.00 70.00 0.18 80.00 50.00 0.01 10.00 200.00 70.00 0.18 80.00 50.00 0.01 10.00 200.00 70.00 0.01 10.00 10.00 10.00 10.00 10.00 200.00 70.00 0.04 10.00 10.00 10.00 10.00 10.00 2		0	500.00	20.00	0.20	10.00	30.00	40.00	0	200.00
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1000.00 70.00 0.18 80.00 50.00 0.0 L 10.00 500.00 70.00 1.00 100.00 70.00 0.0 L 3.00 500.00 70.00 0.64 0.0 L 30.00 0.0 L 3.00 150.00 30.00 0.68 10.00 15.00 0.0 L 3.00 300.00 70.00 0.12 10.00 20.00 40.00 3.00 200.00 70.00 0.15 10.00 15.00 0.0 L 15.00 200.00 0.0 N 0.24 120.00 15.00 0.0 L 0.0 N 200.00 0.0 N 0.0 N 10.00 2.00 0.0 L 0.0 N 100.00 0.0 N 0.10 10.00 0.0 L 5.00 70.00 10.00 0.30 20.00 0.0 L 5.00 70.00 10.00 0.30 40.00 0.0 L 0.0 L 0.0 N		o	100.00	30.00	0.24	40.00	20.00	7 0.0	2 0 0	700.00
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200.00 70.00 0.04 0.01 30.00 0.01 3.00 150.00 10.00 0.17 10.00 10.00 0.01 3.00 100.00 30.00 0.68 10.00 15.00 40.00 0.0 0.0 300.00 70.00 0.12 10.00 15.00 40.00 3.00 200.00 50.00 0.15 10.00 15.00 0.0 1 15.00 200.00 0.0 10.00 10.00 15.00 0.0 1 15.00 200.00 0.0 10.00 10.00 2.00 0.0 1 0.0 N 70.00 10.00 0.30 20.00 10.00 0.0 1 0.0 N 70.00 10.00 0.30 40.00 20.00 0.0 1 0.0 N		2	500.00	70.00	1.00	100.00	70.00	0.0	7.00	700.00
150.00 10.00 0.17 10.00 10.00 3.00 100.00 30.00 0.68 10.00 15.00 3.00 0.0 N 300.00 70.00 0.12 10.00 20.00 40.00 3.00 200.00 50.00 0.15 10.00 15.00 0.0 L 15.00 200.00 0.0 0.0 10.00 15.00 0.0 L 15.00 200.00 0.0 0.0 10.00 2.00 0.0 L 0.0 N 300.00 0.0 0.0 10.00 0.0 L 0.0 N 70.00 10.00 0.30 20.00 15.00 0.0 L 5.00 70.00 10.00 0.30 40.00 20.00 0.0 L 0.0 N		9	200.00	70.00	9.0	0.0 L	30.00	0.0	3.00	00.00/
100.00 30.00 0.68 10.00 15.00 30.00 0.00		0	150.00	10.00	0.17	10.00	10.00	7 0°0	3.00	100-00
300.00, 70.00 0.12 10.00 20.00 40.00 3.00 20.00 0.00 0.00 0.00 0.00 0.00		9	100.00	30.00	0.68	10.00	15.00	30.00	2 0 0	00.007
200.00 50.00 0.15 10.00 15.00 0.0 L 0.0 N 50.00 0.0 N 0.24 120.00 150.00 0.0 L 15.00 200.00 0.0 N 0.06 10.00 10.00 0.0 L 0.0 N 100.00 0.0 N 0.19 120.00 10.00 0.0 L 5.00 70.00 10.00 0.30 20.00 15.00 0.0 L 0.0 N 50.00 10.00 0.30 40.00 20.00 0.0 L 0.0 N		0	300.00	70.00	0.12	10.00	20.00	40.00	3.00	200-00
50.00 0.0 N 0.24 120.00 150.00 0.0 L 15.00 200.00 0.0 N 0.06 10.00 15.00 0.0 L 0.0 N 100.00 0.0 N 0.10 10.00 2.00 0.0 L 0.0 N 300.00 0.0 N 0.19 120.00 15.00 0.0 L 5.00 70.00 10.00 0.30 20.00 15.00 0.0 L 0.0 N 50.00 10.00 0.30 40.00 20.00 0.0 L 0.0 N		9	200.00	50.00	0.15	10.00	15.00	0.0	2	700.00
200.00 0.0 N 0.06 10.00 15.00 0.0 L 0.0 N 100.00 0.0 N 0.10 10.00 2.00 0.0 L 0.0 N 300.00 0.0 N 0.19 120.00 10.00 0.0 L 5.00 70.00 10.00 0.30 20.00 15.00 0.0 L 0.0 N 50.00 10.00 0.30 40.00 20.00 0.0 L 0.0 N		20	50.00	N 0.0	0.24	120.00	150.00	0.0	15.00	500.00
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300.00 0.0 N 0.19 120.00 10.00 0.0 L 5.00 70.00 10.00 0.30 20.00 15.00 0.0 L 0.0 N 50.00 10.00 0.30 40.00 20.00 0.0 L 0.0 N		-	100.00	N 0.0	0.10	10.00	2.00	0.0 L	200	500.00
70.00 10.00 0.30 20.00 15.00 0.0 L 0.0 N 50.00 10.00 0.30 40.00 20.00 0.0 L 0.0 N		9	300.00	Z 0.0	0.19	120.00	10.00	0.0	2.00	200.00
50.00 10.00 0.30 40.00 20.00 0.0 L 0.0 N		00	70.00	10.00	0.30	20.00	15.00	0.0	2:	500.00
		00	20.00	10.00	0.30	40.00	20.00	0.0	20.0	200.00

SILICIFIED ROCKS, PPM

3.00 0.0 N 0
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00.5
7.00

SILICIFIED ROCKS, PPM

TI,PPM	3000.00	2000.00	3000.00	3000.00	3000,00	3000.00	3000.00	3000.00	•	•	3000.00	1500.00	3000.00	3000.00	3000.00	3000.00	3000.00	3000.00	3000.00	2000.00	3000.00	3000.00	3000.00	3000.00	3000.00	3000.00	3000.00	3000.00		000	ċ	3000-00
CA , PPM	500.00	700.00	700.00	700.00	•	700.00	1500.00	500.00	700.00	700.00	700.00	1000.00	200.00	700.00	1000.00	500.00	2000.00	1500.00	200.00	1000.00	1000.00	•		1000.00	1500.00	1500.00	300.00	. 700.00	1500.00	200.00	200.00	1500.00
MG + PPM	7.00	70.00	1000.00	20.00	20.00	20.00	70.00	70.00	200.00	150.00	20.00	300.00	70.00	100.00	200.00	150.00	100.00	150.00	100.00	150.00	150.00	150.00	150.00	100.00	100.00	100.00	500.00	200.00	200.00	300.00	100.00	100.00
FE, PPM	1000.00	7000.00	29999.98	3000.00	5000.00	14999.99	14999.99	14999.99	49999.98	29999.98	3000.00	96°66669	14999.99	14999.99	19999.99	19999.99	10000.00	5000.00	10000.00	19999.99	29999.98	10000.00	10000.00	7000-00	5000.00	7000.00	96°66669	5000.00	2000.00	29999.98	19999.99	14999.99
SAMPLE	223	240	246	549	250	251	252	254	255	256	257	258	259	260	261	262	263	264	265	566	267	268	569	270	271	272	273	274	275	276	277	278

PRINT REPEATED BY OPERATOR

38 BA, LOG	N -1.1549	N -1.1549	N -1*1549	10751			-1-1	N -1.1549	7	7	? ٔ	0669*1- 64	7 7		17	1 N	7	T ~	7	7	1 1540	N -0.5229	-1-	T	~ Z	÷.	<u>.</u>	0106-1- 01	1 1	1 2	N -1.1549	N -1.3010	-1.0000	N -1.1549	N -1.1549	N -1.1549	N -1.5229	N -1.3010	o Z		6528-0-628	֓֞֞֞֞֞֞֞֞֞֞֞֞֞֞֞֞֟֞֟֞֟֞֟ ב		7 7	FCT • 7
\$ MD. LOGS	0.0		0.0	0.0			0.0	0.0	0.0	L -2.5229	T -3.0000	'n,	677C*C	· ·	1 -3.522	0.0	L -3.5229		•	١	0106.6-		1 -3.3010	0.0		ا پ		9.5010		0.0	١ 0٠٥	L 0.0	0.0	-	0.0	, 6		0°0	0.0	1	•	0106-6- 0	7.0000		•
ZN.1064	0.0			6						0.0		•	.	170071	6	0	0		-2-	7	.		S	Ó	•	•	-2-	7-	֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓	ò	0	•		>	Ċ	-2-	0.0	•	o	o (-2-	•7-			5
5% CU,LOG		-3.	L -2	7 :	ָי י	֓֞֜֜֞֜֜֜֝֓֜֜֝֓֓֓֓֓֓֓֓֓֓֓֓֓֓֡֓֓֡֓֡֓֓֡֓֜֝֡֓֡֓֡֓֡֓	-2-	-2-	-2.	-1-	-5-	-2-	ř.	7 -	1276-7	-	-3•	-2.	-5	-2.	-2-	י נ) (f)	-3.	-5	-3•	-2-	7 -	100	-3.	-3	7	90 -3.0000	7 6	1 ?	, E.	-5	<u>۔</u>	£ .		7	? ?	֓֜֞֜֝֜֜֝֓֓֓֓֓֓֓֓֓֓֓֓֓֓֟	
G% AS,LOG%	3010 -2.2218	-2	0	1		2010 -2 0000	וייו וייו	י ה	1	7	£-	-2	1	1	8239 0.0	ŀ		-	-5	-2•	7	7.	7236 -3.0000	7 7	- 6	-1-	-2-	÷.	9586 -1-8559	Ī	Ť		1)	3070 -2.0000		1	•		1	1	-2-	! 1	• • •	
16% HG, LOG	N4.30	N -4.00	N -5.00	N -5.0000		•••	• •	4-	7 -	229 -4.	-4-	3010 -5.09	N -4.95	N -5.2218	Z	N -4-8239	N -4-61	3010 -4-6990	7	N -5.0000	N -5.0969	N -5.22	M -4.9380	N -5.2218	N -5.2218	N -4.8239	N -4.8239	N -4-76	0866.4- N	5229 -4.9586 5229 -4.9586	i	į		-5-	• u	•		1	¥-	z	N -5.0000	٠	.n u	n (N -5.096
.06% 81,L06%		000 000				•		֖֓֞֝֟֝֓֞֝֓֓֓֓֓֓֓֓֓֟֝֓֓֓֟֝֟֓֓֓֓֟֝֓֓֓֟֝֟ ֓֓֓֓֓֓֓		•	-3.	-3.	229 0.0	ċ	0.0	ċ		-2.	•	2	°	•	239 0.0	ċ				0	-1.8239 0.0	ָרְ ה <u></u>	-2-	Ö	ö	o e	• 7 °	-1.52c4 +2.5c4 -		•					Φ.	• ·	1549 0.0
AG, LOG\$ PB, LOG\$		-3.8239 -2.0000					2.6990 -1.3010	7 7	` ī		7	7	-4.5229 -1.5229		-3.6990 -1.501U				10	z	z	Z	z:		2 Z	. Z	- -	_							ž.		3.5229 -2.6				000		-3	-2-	N -2.
AU, LOG\$ AG, I	7447					F	-2-	-3.5000 -3.5 -3.6685 -3.6	1	'	١			_	ı	3.1500 -5.6	, ,	-3.5406 -3.5	,					-3.6840 0.0					-4.4437 0.0	-3.8239 0.0	1	3.9208				•	-4-0456 -4-	,							4.5376 0.0
SAMPLE AU,1	, , , , , , , , , , , , , , , , , , ,	2 4-(3 - 3	4 -4.1	4-	•		90	10	1		32 -3.	•			36							1	89		- 3						83			•	,	98						1		119 -4.

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		0.0	0.0	0	0.0))	•		Ċ		•			1.1549	•	.3010	•	•	٠	•	٠	٠ •	•	٠,	ò		•	•	•	• •			0.0	0.0	0.0	•		0.0	3.5229	3.5229	3.5229	8.5229	2.6990		•		0.0		
>	-	0	0	0	0	.	•	5 C	Ó	, 0	0	13.0	13	13	6	-3	Ë	F	ï		1	.	7 9	,	7 0	> C	o c	o c	> C	, ,	,			_	٠.		<i>-</i>		ï	ï	`i' `	i `	1	ï		. ~	,		
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	SK, LUG2	.154	٠	•	•	•	٠		•	3010	•	•	•	• •			•		0.669.0	•	•	•	•	9	1.5229	٠	•	•	•	1.1249	3010	1.1549			1.5229	1.15	1.3010				-	٠	1.5229	i.	÷.	2.6990			
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	NI, LOG\$	0.0	0.0	0.0	0.0	0.0		0.0	0.0	٠ •	•	5 4	2007 6	5,522	, 4	, 40	9	•	9	Φ	3.52	0.0	0.0	0.0	0.0	0.0	0.0	•	•	•		٠	• •		•	0.0	0 0		0			٠.		•	•	•	0 0	•	
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	NB, LOG\$	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0 0))	0.0	•				0.0	0.0	0.0	3.0000	0.0	0.0		0.0	٠	٠	•	0		0.0	0 0) c		0	0.0	0.0	0 0	•		3,15	0.0	0.0	•	0.0		ሪ ፡	ก่ห	5	
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	36	30	30	239	229	549	06	06	066	53	5 (53	2010	3010	1249	3010	066	8239	0669	5229	5229	5229	5229	0000	59	00	0000	1549	5229	549	549	4 6	6579) }	64	49	00	1549	8239	5229	64	0000	0000	0000	239	0000	0000	2	
	MN, LOGS	1.823	ο oc	•	יט י	3.15	2.6990	٠,	5.69	٠	å,	. ;	٠	٠	•	2000	3 4	2.8	;	2	-	ä	2.	6	2.5229	3.00	3.00	3	5	3.1	3.1	n c	;,		3,1549	3	'n,	ㅎ -		, "	6	n	2.	٥.	2	'n.		•	
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	LA, LOG\$	c	•			0.0	0.0	0.0	0.0	0.0	0.0	0	2.522	0.0	7.7		2 6	10	,	7.7	2.3		2.1	2.1	2.1	0.0	-	2.1	2.15	0.0	0.0	0.0	0 0	•	0	0.0	2.30	0.0		$^{\circ}$	-2.301	.2.52	2.30	-2.30	-5.00	0.0	0.0	-2-15	
	تـ												1		1	1	(1	•	1	1		١	1	1		'	1	'								•			•	٠	'	•	•	,			•	
	25		3 8			8239	0000	00	00	00	00	8	06	39	62	0669		0770	0 0		0	06	0669	0669	06	8239	0669	8239	8239	39	0000	239	000	8528	0669	0669	0669	0669	0669	0000	8239	8239	0000	8239	229	Z	5229	0669	
	CR, LOG\$		•		,		חו	3.0000	•	m	G,	m.	2	\sim	2.	'n.	7 6	• •	20.00	3.0000	00.4	2.6990	2.6990	2.69	•	2.82	•		2.82	-2.8239	0	2.85	٠	7.84			-2.69	•	•	֓֞֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓	• 1			•	•	•	6	-2.6	
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	25	2	2 2	2 2	2 2	: 2	: Z	Z	Z	z	z	53	3010	Z	10	Z:	Z	2 2	2 2	2 0 %	۲	2	: 2	: 2	z	Z	Z	Z	Z	Z	Z	Z	Z	zz	Z	Z	z	Z:	Z:	Z 2	2 2	: 2	2	Z	Z	z	Z	Z	
	2907.00		•	•			0	0.0	0.0	0.0	0.0	?	'n	0.0	•	0.0	•	•	•		10				0.0	0		0.0	0.0	0.0	0.0	0.0	٠	•		0	0.0	0.0	•	3		•		0	. •	0.0	0.0	0.0	
	ပ											1	1		,					,																													•
	29	2	z :	2 2	E 2	: 2	: Z	z	Z	Z	z	z	5229	5229	3010	5229	5229	5229	67	67	2 0	7	: 2	: 2	: 2	: Z	: Z	: 2	Z	Z	Z	Z	Z	Z	Z 2	: 2	Z	Z	Z	010	2 6	2	649	239	Z	Z	Z	Z	
	BE, LOG%))	9 6) 	•		0.0	0.0	0.0	0.0	0.0	4.52	4.52	4.30	4.52	-4.52	24.92	6776.4	-4.3222	0106.						0.0		0.0	0.0	0.0	0.0	0.0	0.0			0.0	0.0	0	-4.3010	-4.5010	ñ c	2 4	-3.8239	0.0	0.0	0.0	0.0	
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	m	•	⊶ (7 (n .	t u	٠ د	٠,	- α	0	10	1	30	32	34	35	36	37	33	5 1	,	0 4	t d	6	200	- «	3 6	2	75	11	18	4	80	81	28	9 4	85	98	87	88	5 6	3 5	6	101	116	117	118	119	
	SAMPLE											. %																																					
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TI,L06%		ů	. 522	275.	. 522	. 522	. 522	. 522	. 522	. 522	.823		ø	-0.6990	0669.0-	-0.5229	-0.6990	0669*0-	0669*0-	-0.8239	•	669•	-0.5229	-0.5229	0669.0-	-0.5229	-0.8239	-0.6990		669	522	•	•	823		0669-0-	-0.6629	0669	9		'n	301		669		-0.690	S	-0.5229	.522	69	
CA, LOG\$		1.154	.154	.154	٠	• 000	000	.154	.154	. 522	.301	•	-1.3010	-1.5229	-1.0000	-1.3010	-1.3010	•	-1.0000	-1.0000	•	-1.3010	-1.1549	-1.0000	•	•	-1.1549	•	٠	•	-1.1549	٠	.	<u>.</u> ,	٠	-1-1549	-1.0000	4 -	-1.0000	-1.1549	-1.0000	, (6	15	~	52	•	-1.6990	-2.1549	.301	
MN, LOG\$		7	•	0	~		Ö	•	-2.1549	•	•	-1-1549	9	00	0669.0-	ഹ	-1.8239	8	8	•	000	-1.1549	000	-0.8239	0	.8239	0	.5229	Z .	•	٠	0	0	0	•	7.000	7•124 2 154	100		200	2 9	154	15	154	30	000	3	69.		.15	
_ FE, LOG		0.1761	٠	٠	4	9	. 154	٠	0.1761	٠		0		ø	0.8451					0.4771	0.3010	0.0	0.4771	0.690	0.4771	•	•	0.3010		.845	•	٠		669	•	•	6776	0.6930	: 4			0.669.0-	0	•	٠	-0.1549	9	-0.3010	0.1761	-0.3010	
SAMPLE	•		5	.	4	S	•0	~	80	σ.	01	11	30	32	34	35	36	37	38	39	45	63		· 65	99	29	89	69	74	75	7.1	8 2	4	08	# C	28	83	+ 000	6 8	2 6	- «	8	06	6	92	101	116	1117	118	119	

SILICIFIED ROCKS, LOG PERCENT

-5.0000	-2.6990	-2-3010	2	-5-0000	0.0	-3,3010	1 0.0	-3.3010	-1.0000
	N 0.0	-2.1549	2 2	-5.0000	-2.0969		-2.1871	-3.1549	-1.5229
	Z -	-2.3010	0.0	-5.2218	-1.6990	-2.5229	-2.3010	-3.0000	-1.1549
	0 0	-1.5229	2 2		3	-3.5229		Z 0 0	
	0.0	-1.1549	N 0.0	-4.9586		-3.0000	-2.2218	Z:	٦,
	N 0.0	-2.1549	0.0 N	-4.4559	-3.0000	-3.5229	-1.7959	0.0 N	-1.3229
	0.00	-2.0000		-4.6576		• •	-2.5229	Z 0.0	
	0°0	-1.3010	N 0.0	-4.6778	•	-3.3010	-2.5229	0.0	-1.1549
	-3.6990	-1.5229	0.0	-4.5528	-1.6021	•	0.0	-3.1549	—
	-4.0000	-1.3010	-3.3010	-4.6198		-2.8239	0	Z :	
	0.0	-1.8239	Z :	-4.8861	-3.0000	-2.8239	-2,3979	2 2	
	0°0 F	1.6990	2 2	2619.4-	0560-7-	-2.6239		2 2	-1-3010
	-2-1249	-1 1540	-2.5220	-4.8230	-1,7959	-2.0000		200	-
	-3,0000	-1.1549	-2,0000	-4-3768	-1.8539	-2.6990	0.0	20.0	-1.1549
	-3.0000	-1.3010	2000	-4.3979	-1.8539	-2.8239	-2.3010	N 0.0	-1.1549
3.6576	-3.3010	-1.3010		6764.4-	-3.0000	-2.6990	0.0	o.0	-1.1549
3.5686	-3.3010	-1.3010	-2.8239	-4.5686	-1.8539	-2.6990	0.0	-3.5229	-1.3010
-3.7799	-3.1549	-1.3010	-2.5229	-4.2218	-1.8539	-3.3010	•		-1.3010
-3.2924	-2.6990	-1.3010	-2.5229	-4.0458	-1.7959	-2.5229	•	229	-1.3010
-3.7959	-3.5229	-1.1549	Z 0	-4.6383	-1.9208	-3.3010	•	Z :	-1-3010
-4.1643	-3.6990	-1.5229	Z :	-4.8539	-2.0000	-3.0000	0.0	2 2	
-4.3010	0°0	-1.3010	2 2	-5.3979	-2.6990	-3.0000	 		-1,1549
3.0586	-3.5229	-1-3010	2 2	-5-0000	-2.0969	-2.8239	0.0	-3.1549	-1.1549
-3-3565	-4.0000	-0.8239	2000	-5.0000		-2.6990	0.0	-3,0000	-1.1549
3.2848	-3.6990	-1.1549	-2.8239	-5.0000	-3.0000	-2.5229	0.0	-3.0000	-1.3010
-2.6234	-3.6990	-1.1549	0°0	-5.0000	-2.2218	-2.3010	0.0	-3.5229	٠
-3.1331	-3.6990	-1.1549	N 0.0	-5.0000	-2.6990	-3.1549	0	2 ·	•
-4.5528	-3.5229	-2.3010	0.0	-5.0000	-3.0000	-2.5229	•	Z 0.0	٠
-4.6990	-3.6990	-1.6990	× 0.0	-4.8861	-3.0000	-2.3010	0.0	Š	-1.0000
-3.6778	0°0 F	-1.1549	000	-5.0000	-3.0000	-3.1549	0.0 L	0.0 M	-1.1549
-4.2366	-4.0000	6776*1-	2 2	6670*	0669.7	7 , 000	֓֞֜֜֓֓֓֓֓֓֓֓֜֜֜֜֓֓֓֓֡֜֜֜֓֓֓֡֓֡֓֡֡֡֓֓֓֡֓֡֡֡֡֡֡		1540
-4.0000	-5.6990	-1-3010	2 2	14.1090	13.0000 L	-2.8230		20	-1.5229
7 0446	6776-6	1 0220		0000°7				N 0 0	•
6.6990		-1.8239	200	-4.8861	-3.0000	-2.3010	0.0	229	•
-3-1249	0°0	-1.3010	0.0	-4.8239	-3.0000	-3.1549	-2.3010	-3.5229	-1.3010
-3.5686	0.0	-1.1549	. O	-5.0969	-3.0000	-2.5229	0.0	-3.0000	-1.3010
-4.3979	-3.6990	-2.1549	0.0	-5.0000	-3.0000	-3.1549	-2.3010	200	•
-4.6021	-3.5229	-2.3010	0.0	-5.2218	-3.0000	-3.3010			٠
-4.0458	0.0	-1.3010	0.0	-5.2218	0.0	•	•	Z 0.0	٠
-4.5229	0.0	-1.6990	239	-5.3979	-1.7959	•	-2.2218	္က	•
-4.3979	-3.6990	-2.1549	N 0 0 0	-4.9586	-1.8539	-3.1549	-2.1549	0.0	-1.1549
0000	•								

SILICIFIED ROCKS, LOG PERCENT

Y, LOG%	0	0	0	0	0	O .0	0 (0		5 6		٥ د	.	.		z :	Z :	2 :	Z 2	2 :	0.0	2 2		2 2		2 2	· c	. 0	-3.0000	-3.0000	0°0	000	-2.5229	0.0	0	N 0.0	0	0	
V,L06%	-2,3010	301	•	-2,3010	-2.3010	-2.1549	-2.1549	-2.0000	•	-2.1549	•	٠	·,	•			-2.1549	•	-2.1549	٠	•	٠	ο,	.154	7	301	 (n,	-2.1549	٠,	-2.5229	ø,		-2.1549	100	-2.8239	~	.154	7	~	7	٦.	~	.154	69•	-2.5229	
SR, LOGI	-1.1549	1.154	669	-1.3010	_:	-	٠	000	.301	1.154	7	 ,	1.000	1.301	1.154	1.154	1.154	1.301	1.154	1.301	<u>.</u>	1.301	1.154	1.154	.	5	7:	1.154	1.301	1.154	1,301	2.154	.	-1.5229	1 154	2.154	.823	<u>.</u>	1.3	-1.1549	-2.0000	-1.1549	-1.1549	-1.1549	. 15	-1.6990	
84	z	z	z	z	z	Z	z	Z	Z	Z.	Z	Z	Z	Z:	Z	Z	Z :	z	Z	z	Z	Z	Z	Z	Z	Z	Z:	Z	Z	Z:	Z:	Z :	z:	Z 2	2	z	z	z	z	z	z	z	z	z	z	z	
NI,L06%	0.0	0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0) 	•	0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	•	
NB.L06%	-2.6990	Z 0.0	-2.8239	0	-3,3010	•		0.0	•	0.0 N	O.0	•	-3.0000	•154	1549	•	-3.1549	٠	Z 0.0	0.0	-	_		_	•	-3.3010	Z 0.0	Z 0.0		2 0.0	0	3.301	3.522	ກ	7.756	-3.3010	0.0	. Z			0.0	2 0	N 0	N 0.0	-3.0000	-3.5229	
MN.LOG%		-3-0000	0	S	669	8	-3.1549	0	•	-2.8239	-2.8239	-3.3010	0	٦,	•	•	•	٠	•		٠	<u></u>	~	. •	-3.6990	•			-2.8239	•	•	•	2	m,	0.000	0100.0	3010	, <	,	(1)	2	3.5	1.82	2.1	•	69	
LA, LOG\$		-2,1549	•			0.0				0		0.	0	•	•	0	0.	۰.	•	0.0 N	0.0	20.0	0°0		•	-2.5229		N 0.0	N 0.0	٠		9	-2.3010	സ	2 3	2 2	•	1,0		, (301	•		3010	0	Z 0.0	
CR, LOG\$	۲	-2,5229		8239		-3.0000	•	S	•	ω,		~	-2.8239	9	.823	-2,8239	.823	•	.823	-3.1549	-3.1549	-2.8239	-2.8239		-2.6990	•	-2.6990	•	.823		301	-3.3010	2.699	\sim	7.09	0669.2-	0200	2.82	2 52	, ,	2.522	2.522	2.699	1 ~	2.6	1 (4)	
2907,03	•	2 2	0000	2000		2000	_	0	0.0	0°0	~ •	0	N 0.0	0°0	0.0	0.0		0	N 0.0	0	0	N 0.0	0	0		0.0	0	0	0	0		0		Z :	٠,	٥ (. c	2	`. o c		0.0	
**	:	2 2	: 2	: 2	: z	z	z	Z	z	z	z	Z	z	z	Z	z	z	Z	Z	Z	z	z	Z	Z	z	Z	z	Z	z	z	z	z	z	z	z	z	Z, 7	zz	2 2	2 2	2 2	: 2	2 2	z	: z	: z	
8E,L0G%	(•	•		•	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0		0.0	0.0			0.0	0.0	0.0	0.0	0		0.0			0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0	٠	5 0	٠		•		•	•	0.0		
SAMPLE	4	120	171	120	142	191	162	163	166	167	168	691	170	171	172	173	174	175	176	177	178	179	180	181	182	183	184	185	186	187	188	189	190	161	192	193	46.	195	101	100	100	200	207	220	221	222	

11,106%	0 . 5 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2		-1.0000 -0.8239 -0.6930 -0.6990 -0.6990 -0.6990 -0.5229 -0.5229	500 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
CA, L06%	1.522 1.301 1.699 1.699 1.000 0.699 1.154		-1.0000 -1.1549 -1.1549 -1.0000 -1.3010 -1.3010 -1.549	
MN, LOG\$	2.154 2.301 2.522 2.522 2.154 1.699 0.0	2.154 2.000 2.000 2.000 2.154 2.154 2.154 2.154	-1.8239 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -1.8239 -1.8239 -2.0000 -2.000	
FE, LOG\$	0.000 0.477 0.477 0.477 1.301 0.000 0.699	0000-000000	v O w N w N A W O 4 W N A v	, , , , , , , , , , , , , , , , , , ,
SAMPLE	120 121 129 130 142 161 162	166 168 169 170 171 173 174 175	177 178 180 182 182 184 185 186 187	190 191 193 194 196 198 198 200 220

SAMPLE	AU, LOG\$	AG, LOG%	PB,106%	81,106%	HG, LOG\$	AS, LOG%	CU, LOG\$	ZN, LOG2	MO, LOG%	BA, LOG\$
333	6665 7	0-0	-1-6990	0669*7-	-5.0000	-2.2218	-2.5229	0.0	N 0.0	-1.5229
240	6776-5-		-2.3010	N 0.0	-5.0000	-3.0000	-2.1549	0.0	Z 0.0	-1.0000
246	4318	N C -0	-2,3010	N 0.0	-5.0969	-3.0000	-3.1549	0.0	-3.3010	-1,1549
076	0258 2	- 0.0	-2.5229	2000	-5.0000	-3.0000	-3.1549	0.0	N 0.0	-1.3010
250	-3.9355	-3,8239	-2,3010	-3.0000	-5.0969	-3.0000	-3.3010	0°0	2 0.0	-1.0000
251	-3.6383	-3-5229	-1-6990	-2.5229	-5.0000	-2.3979	-3.0000	0°0	-3.5229	-1.0000
25.2	7 8517	-3.8239	-2.0000	-2.8239	-5.0000	-2.0969	-2.1549	0.0	-3.3010	-1.0000
750	-4.7366	0.0	-2.0000	N 0.0	-5.3979	-1.9208	-2.8239	0.0	-3.3010	-1.1549
255	-3.6778	-3,0000	-1.5229	-2.5229	-5.0000	-2.0969	-2.3010	0.0	-3.1549	-1.3010
256	-3.4685	-2.5229	-1.5229	-2.5229	-5.0000	-1.9208	-1.8239	-2.3979	-3.1549	-1.1549
257	-3.4815	-3.1549	-1.3010	-2.6990	-4.6990	-3.0000	-2.5229	-2.3979	0	-1.1549
258	-3,3696	-2,3010	-1,1549	-1.6990	-4.5376	-1.8539	-1.6990	-2.3010	-3.0000	-1.3010
250	-3.9226	-2.5229	-1.5229	-2.5229	-4.6383	-3.0000	-2.1549	0.0	-3.5229	-1.3010
260	-3.9281	N 0.0	-1.5229	-2.5229	-5.0000	-2.2218	-2.5229	0.0	-3.5229	-1.1549
261	-3.5376	0.0	-1.6990	-2.5229	-4.9208	-1.7959	-2.1549	-2.3979	-3.0000	-1.0000
262	-3.9788	0.0	-2.0000	-2.5229	-4.7212	-2.0969	-2.8239	0.0	-3.0000	-1.1549
263	-3.7447	2000	-1.6990	N 0.0	-4.8861	-3.0000	-2.3010	0.0	-3.3010	-1.0000
264	-3.6778	-4.0000	-1.3010	-2.6990	-4.6990	-3.0000	-3.1549	-2.5229	O.0	-1.1549
265	-3.8069	-2.8239	-2.0000	-2.5229	-4.6198	-2.3979	-2.6990	0.0	O.0	-1.1549
266	-3.2708	-3.8239	-1.0000	-2.1549	-4.7447	-2.0969	-2.3010	0.0	-3.0000	6461-1-
267	-3.3778	-2.6990	-1.3010	-2.1549	-4.0000	-2.0000	-2.1549	0.0	-3.1549	-1.1549
268	-3.2782	-2.5229	-1.6990	-2.1549	-4.1938	0.0	-2.5229	0.0	-3.5229	-I - 1549
269	-3,6459	-3,3010	-1.8239	-3.0000	-4.7696	-3.0000	-3.0000	0.0	-3.5229	-1.1549
270	-3.3716	-2.5229	-2.0000	-2.5229	-4.1675	-3.0000	-2.8239	-2.5229	N 0.0	-1.1549
271.	-2.7305	-2.5229	-1.5229	-2.1549	-4.9208	-3.0000	-2.6990	-2.3979	-3.5229	-1.3010
27.2	-3,3979	-3.6990	-1.6990	-2.3010	-4.8239	-3.0000	-2.8239	0.0	200	-1.1549
273	-3.5735	-3.6990	-2.3010	0°0	-4.6198	-1.9208	-1.8239	0.0	-2.8239	-1.3010
274	-4-8739	0.0	-1.6990	0.0	-5.2218	-3.0000	-2.8239	0.0	O .	-0.5229
275	-4.3054	1 0°0	-2,0000	0.0	-5.0000	-3.0000	-3.6990	0.0	O.0	-1.3010
276	-4.1549	-3.6990	-1.5229	0.0	-4.7212	-1.9208	-3.0000	0.0	-3.3010	-1.5010
277	-4.7441	-3.6990	-2.1549	-3.0000	-4.5229	-2.6990	-2.8239	0.0	0	-1.3010
278	-4.0915	-3.6990	-2,3010	-3.0000	-4.5229	-2.3979	-2.6990	0.0	o.0	-1.3010
)										

SILICIFIED ROCKS, LOG PERCENT

TI,L06%	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.6990	0669*0-	-0.5229	-0.8239	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	0669*0-	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229	-0.5229
CA, LOG%	-1.3010	-1.1549	-1.1549	-0.8239	-1.1549	-0.8239	-1.3010	-1.1549	-1-1549	-1.1549	-1.0000	-1.3010	-1.1549	-1.0000	-1.3010	-0.3010	-0.8239	-1.3010	-1.0000	-1.0000	-1.0000	-0.6990	-1.0000	-0.8239	-0.8239	-1.5229	-1.1549	-0.8239	-1.3010	-1.3010	-0.8239
MN.L06%	-3.1549	-1.0000	-2.3010	-2.3010	-2.3010	-2.1549	-2.1549	-1.6990	-1.8239	-2.3010	-1.5229	-2.1549	-2.0000	-1.6990	-1.8239	-2.0000	-1.8239	-2.0000	-1.8239	-1.8239	-1.8239	-1.8239	-2.0000	-2.0000	-2.0000	-1.3010	-1.6990	-1.6990	-1.5229	-2.0000	-2.0000
FE, LOG2	-1.0000	0.4771	-0.5229	-0.3010	0.1761	0.1761	0.1761	0.690	0.4771	-0.5229	0.8451	0.1761	0.1761	0.3010	0.3010	-0.0000	-0.3010	-0.0000	0.3010	0.4771	-0.0000	-0.000	-0.1549	-0.3010	-0.1549	0.8451	-0.3010	-0.6990	0.4771	0.3010	0.1761
SAMPLE	223	246	546	250	251	252	254	255	526	257	258	529	260	261	292	263	564	265	797	267	268	569	270	271	272	273	274	275	276	277	278

PRINT REPEATED BY OPERATOR

BA, %	0.070000 0.070000 0.070000 0.050000 0.150000	
M0,%	0.000700 0.000700	
ZN, %	0.0022000 0.003000	
2.U)	0.005000 0.003000 0.003000 0.002000	
AS. &	0.001000 0.001000	
%⁴ 9H	0.000000000000000000000000000000000000	
81.8		
PB, %	0.003000 0.00 N 0.00 N 0.001500 0.001500 0.001500 0.001500 0.001500 0.002000 0.002000 0.002000 0.002000 0.002000 0.002000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001500 0.00100 0.00100 0.00100 0.00100 0.00100 0.00100 0.0010	
AG. 2	0.000030 0.000030	
AU, Z	0.000019 0.000018 0.000020 0.000020 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010 0.000010	
SAMPLE	3 8 8 3	

× ,×	0.001000 0.001000 0.0	0.0 0.001000	0.000700	0.001500	0.000700	0.003000	0.003000	0.005000	0.001000	0.000100	0.000700	0.001000	0.000000	0.001000	0.02000	0.002000	0.001500	0.001500	0.000700	0-001000	0.000000	0.000000	0.000500	0.000.00	0.000100	0.001000	0.000,00	0.000100	0.000000	0.003000	0.002000	0.002000	0.002000	N 0.0	0.001000	0.001000	0.000100		
» >	0.015000 0.015000 0.010000	0.010000	0.007000	0.015000	0.007000	0.007000	0.010000	0.015000	0.010000	0.007000	0.007000	0.01000	0.010000	0.007000	0.007000	0.010000	0.010000	0.010000	0.007000	0.0000000000000000000000000000000000000	0.010000	0.010000	0.005000	0000000	0.010000	0.007000	0.007000	0.007000	0.005000	0.010000	0.010000	0.01000	0.001000	0.00010.0	0.007000	0.015000	0.010000		·
SR. %	0.050000 0.030000 0.070000	0.050000	0.070000	0.150000	0.100000	0.030000	0.150000	0.200000	0.150000	0.150000	0.050000	0.300000	0.20000	0.20000	0.150000	0.050000	0.02000	0.100000	0.070000	0.150000	0.05000	0.00000	0.015000	0.0000	0.02020	0.010000	0.050000	0.100000	0.010000	0.050000	0.150000	0.030000	0.02000	030000	0.050000	0.150000	0.030000	2000	i
NI ,	N N N	00500	0.000500	0.000700	0.000700	0.000700	0.000300	0.000500	0.000200	0.001000	0.000300	0.000300	0.000000	N 0.0	0.000700	00700	0.000300	0.0	0.000300	0.000300	20000	0.000200	0.000200	0.000200	0.000500	0.0	0.000200	0.000500	0.000200	0.000300	00200	0.0	2 2			0.000300	0.000300	0000000	•
NB • %	000	00200	0.0	. z	.N 0.0	0-000500 0-0		N 0.0	0.000500	0.0	0.000000	0.0 0.00 0.00 0.00	0.000500	N 0.0	Z 2	2 2	. o	0.000500		0.0 0.00 0.00	0.001000	0.0	O.0	2 000	0.000500	0.000500	00200	2 2	00500	0.000500		Z :	2 2			0.0	0.000500	0.00000	٠.
N. N.	0.020000	0.030000	0.300000	0.030000	0.020000	0.030000	0.007000	0.005000	0.007000	0.010000	0.007000	0.007000	0.003000	0.005000	0.003000	0000000	0.010000	0.010000	0.003000	0.005000	0.003000	0.002000	0.00700	0.02000	0.015000	0.015000	0.010000	0.007000	0.010000	0.010000	0.02000	0.003000	0.002000	0.015000	0.001000	0.020000	0.02000	0.000000	
LA, \$	0.010000 0.0 0.0	03000	0.005000	0.007000	0.007000	0.005000	0.010000	0.010000	0.007000	0.007000	0.010000	0.005000	0.007000	0.003000	0.005000	0.005000	0.007000	0.010000	0.005000	0.007000	0-007000	0.010000	N 0.0	0.015000	0.005000	0.010000	0.005000	0.007000	0-010000	0.007000	0.00500.0	0.007000	0.007000	0.007000	0.007000	0.010000	0.005000	0.00700	
CR.	0.002000	0.001500	0.002000	0.002000	0.002000	0.003000	0.002000	0.003000	0.002000	0.001500	0.001000	0.002000	0.003000	0.002000	0.001500	0.002000	0.003000	0.002000	0.002000	0.002000	0.003000	0.003000	0.001000	0.003000	0.003000	0.003000	0.003000	0.002000	0.003000	0.002000	0.003000	0.002000	0.001500	0.002000	0.001500	0.003000	0.003000	0.002000	
% 00	0.001000 0.001000	0001	0.001000	0.000700	0.000000	0.000700	0.000100	0.000100	0.000500	0.000300	0.000300	0.000300	2 2	0.0	0.001000	0.000700	0.000,00	N 0.0	0.000500	0.000500	0.000500	0.000500	0.000500	0.000100	0.001000	0.000500	0.000100	0.000500	00,000.0	200					2 2	00300	0.0	Z 0.0	
BE, 3	0.001000 0.0 N		0.000050	0.000030	0.000000	0.000050	0.000150	0.000100	0.000070	0.000050	0.000030	0.000030	0.000000	0.000030	0.000050	0.000200	0.000100	0.000200	0.000150	0.000100	0.000070	0.000010	0.000030	0.000030	0.000030	0-000030	0.000000	0.000030	0.000050	0.000200	0.000100				2 2	00100	0.000000	0.000100	
SAMPLE	13	1 51 2	11	8 6 10	20	21	27	24	52	27	28	29	31	7 4	41	45	£ 43	4	47	4.	Q 4	20.5	52	53	47.	25.	57	28	50	9	79	10	. 71	72	: 24 24	8	96	95	

8.17	0.200000	0.300000	0.300000	0.50000	0.500000			0.500000			0.300000			•	0.300000		0.500000	•	0.30000		•	0.50000	0.200000	0.500000	0.500000	0.500000	•	0.500000	• •		•	٠	0.300000	, "			5	0.500000
CA, &	0.500000		0.70000	0.100000	0.100000	1.499999	0.200000	0.300000	0.100000	0.70000	0.050000	0.150000	0.070000	0.070000	0.070000	0.070000	0.070000	0.150000	000000	0.100000	0.100000	0.200000	0.010000	0.100000	0.100000	0.100000	0.100000	00007000	0.0000	0.010000	0.100000	0.100000	0.070000	0.00000	0.100000	0.150000		0.150000
MG•8	0.500000	0.500000	0.700000		0.30000		•	0.500000		0.200000	0.020000	0.500000	0.150000	٠	0.300000		0.100000	0.70000	00000	0.20000	0.200000	0.20000	0.100000	0.500000		0000000	0.300000	0.30000		0.700000	0.500000	0.500000	0.150000	0.500000	0.100000	0.70000	1.000000	0.700000
FE, 2	6.999997 2.999998 0.300000			2.999998	1.999999	1.999999		0.500000		0.0	4.999998 4.999998			•	1.999999	1.999999		•	3666667	1.999999	•	٠	٠	8666667	66666	4.999998	2.999998	4.999998	1.000000	1.999999	266666.9	1.999999	1.000000			966666.6	66666	266666.9
SAMPLE	12 13	15	11	81	20	22	23	24 25	56	7.2	28 20	31	33	04	4 4	7 4 4 4 3	44	94	. •	* 4 & 6	20	21	52	2 4	32	26	25	80 0	60	61	62	02	17	21.	24	93	56	95

ARGILLIZED ROCKS, PERCENT

8A.%	0.200000 0.200000 0.100000 0.200000	0.030000 0.050000 0.020000 0.100000 0.150000	0.100000 0.200000 0.100000 0.070000 0.070000 0.030000 0.020000	0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.150000 0.050000 0.050000 0.050000 0.050000 0.050000 0.050000
MO, %	0.001000 0.001500 0.000300 0.000700	0.000500 0.000500 0.000300 0.000300 0.000300	0.001000 0.000500 0.000500 0.000700 0.000500 0.000500	0.000000000000000000000000000000000000
ZN, Z	0.004000	0.004000	0.017500 0.00 0.007000 0.005000 0.005000 0.005000	0.002500 0.0
* •00	0.005000 0.000700 0.001000 0.003000	0.003000 0.003000 0.003000 0.0001000 0.002000 0.002000	0.003000 0.001000 0.001000 0.001000 0.003000 0.003000 0.003000	0.001500 0.001700 0.001000 0.001000 0.002000 0.002000 0.002000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000
AS.	0.001000 0.002000 0.001000 0.001000	0.001000 0.001000 0.002000 0.001000 0.001000	0.004000 0.001000 0.002000 0.002000 0.002000 0.001000	0.001000 0.001000 0.001000 0.002000 0.002000 0.002000 0.001000 0.002000 0.002000 0.002000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000 0.001000
HG,%	0.000010 0.000020 0.000020 0.000010	0.00010 0.000040 0.000040 0.000010 0.000010 0.000010	0.000010 0.000010 0.000010 0.000010 0.000010 0.000010	0.000000 0.000010 0.000010 0.000010 0.000004 0.000004 0.000012 0.000013 0.000013 0.000013 0.000013 0.000013 0.000013 0.000013 0.000013 0.000013
BI•*			000000000	
PB, 3	0.002000 0.015000 0.010000 0.003000	0.02000 0.02000 0.02000 0.00700 0.001500 0.001500	0.001000 0.000700 0.000700 0.000700 0.000700 0.00700	0.000700 0.001000 0.005000 0.005000 0.005000 0.007000 0.007000 0.00700 0.00700 0.000700 0.000700 0.000700 0.000700 0.000700 0.000700 0.000700 0.000700 0.000700
AG• %	0.000030 0.000030 0.000030 0.000030	0.000070 0.000100 0.000050 0.000050 0.000030 0.000050	N L L N N N N N N N N N N N N N N N N N	
AU, X	0.000010 0.000187 0.000020 0.000019	0.000010 0.000048 0.0 0.00010 0.000010 0.000010	0.000017 0.000010 0.000010 0.000010 0.000010 0.000011	0.000010 0.000016 0.000050 0.000050 0.000000 0.0000000 0.0000000 0.0000000
SAMPLE	96 98 99 100	102 103 104 106 107 108	113 113 114 115 123 123 124	125 127 127 128 133 134 135 140 140 141 143 144 145 145 146 147 148 148 148 148 148

ARGILLIZED ROCKS, PERCENT

X . Y	0.003000	0.002000	0.003000	0.002000	0.002000	0.002000	0.001500	0.002000	0.00200	0.000300	0.000500	0.000500	0.000500	0.000100	02000							O.0	0		0.0	O.0	0.001000	0.001500	0.001000	0.001000	0.001000	0000000	0.002000	0.001000	0.00100	000100	000000	0.002000	000000	0.00200	0007000	00100	20000	0001	20000	03000	0000000	
> >	0.0010.0	0.010000	0.00700	0.010000	0.010000	0.010000	0.010000	0.015000	0.01000	0.007000	0.005000	0.005000	0.01000	0.010000	0.010000	0.007000	0.007000	0.001000	0.007000	0.005000	0.007000	0.007000	0.00700	0.007000	0.00500.0	0.007000	0.00700	0.003000	0.005000	0.007000	0.007000	0.005000	0.007000	0.00/000	0.007000	0.00700	000000	000100	000000	000200	0001000	000000	000000	000000	0002000	0000000	00010000	
SR. T	0.020000	0.030000	0.010000	0.030000	0.070000	0.0000	0.100000	0.150000	0.015000	0.070000	0.030000	0.070000	0.100000	0.100000	0.010000	0.010000	0.02000	0.050000	0.030000	0.050000	0.100000	0.003000	0.002000	0.030000	0.015000	0.050000	0.030000	0.030000	0.010000	0.050000	0.100000	0.100000	0.010000	0.050000	0.03000	0.0000	0000000	0000000	0.00000	0000000	0000000	0.00000	000000	0.030000	000000	000001.0	0.015000	
NI.	0.000300	0.000300	0.000300	0.000300	0.000300	0.000300	0.000300	0.000300	0.000300	0.000300	0.000300	0.000500	0.001000	0.001500	0.000100	0.0						0.0		0.0										O.0			2 2			2 2						2 2	0.0	
NB, 2	0.000500	0.000500	0.000500	0.000500			0.0	0.000100	0.000100	0.000100	0.000500	0.000500	0.0			0.0	0.0	0.0	0.0		0.0	0.002000	0.002000	0.002000	0.0	0.001000	0.0	0.0	0.001000	0.001000	0.001000	0.000500	0.000500	0.000500	0.000500	0.000500	0.000.0	0.00000	0.00000	0.00000	0.000000	00/00	Z 0.0	0.000,00	0.000500	0.000500	0.000500	
W.	0.030000	0.015000	0.010000	0.020000	0.02000	0.010000	0.020000	0.010000	0.010000	0.010000	0.010000	0.010000	0.030000	0.100000	0.050000	0.005000	0.020000	0.020000	0.010000	0.010000	0.007000	0.007000	0.010000	0.007000	0.010000	0.020000	0.015000	0.02000	0.007000	0.007000	0.010000	0.000300	0.007000	0.010000	0.010000	0.007000	0.005000	0.005000	0.005000	0.005000	0.003000	0.007000	0.005000	000010.0	0.005000	0.000300	0.000500	
LA, %	0.005000	0.005000	0.005000	0.005000	0.003000	0.005000	0.005000	0.00700	0.005000	0.007000	0.005000	0.007000	0002000	00001000	0.007000	0.007000	20.0	0.007000	0.005000	0.007000	0.010000	0.005000	0.0		0.010000	0.007000	0.007000	0°0	0.015000	0.007000	0.010000	0.007000	0.007000	0.015000	0.007000	00010		N 0.0	0.005000	0.007000	0.007000	7000	0.0	0.007000	0.010000	0.007000	0.005000	
CR• %	0.002000	0.002000	0.002000	0.002000	0.002000	0.002000	0.000200	0.007000	0.002000	0.002000	0.00200	0.002000	0.002000	0.002000	0.002000	0.002000	0.001500	0.002000	0002000	0.002000	0.000000	0.002000	0.002000	0.002000	0.002000	0.002000	0.003000	0.001500	0.002000	0.002000	0.003000	0.001000	0.002000	0.001500	0.002000	0.002000	0.001000	0.002000	0.002000	0.002000	0.002000	0.001500	0.001000	0.001500	0.001000	0.002000	0.002000	
% •00		0.0		0.0			.000300	0.000300	0.000300	0000000	000000	000000	0.001000	0.001500	0.001000	0.001000	0.000000	0-000100	0.00000	N 0 0 0				-		01000	0-000700	0-000100	2				0.0	0.0	0.0	0.0	2000		_	_	_	•	0.0		_	0.0	0.0	
BE, %	001000	0.000100	0.000000	0.000000	0.000150	0.000100	0.000010	0.000010	0.000100	0.000100	0.000100	0.0000	0.0000	0.000150	0.000100	0.1000	0.100100	0-1000	001000	0.000100	0.000100	001000	001000	00100	20000	2	2	2	2	2 2	0.0	. O. O	0.0	0.000100	0.000100		0.0	0.000100	0.0	0.000100	o.0	0.0	Z 0.0	0.0	0.0	0.0	0.000100	
SAMPLE	ŏ	76	. œ	8 8	901	102	103	201	101	507	100	- C	100			1112	113	411	411	122	122	126	125	136	127	128	131	132	133	134	135	136	137	138	139	140	141	143	144	1.45	146	141	148	149	150	151	152	

TI+2	_ , ,		0.200000 0.200000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000
CA,%		0.150000 0.200000 0.070000 0.050000 0.050000 0.100000 0.100000 0.100000	0.100000 0.100000 0.070000 0.050000 0.100000 0.100000 0.100000 0.150000 0.150000 0.150000 0.200000 0.200000 0.200000
MG, %			0.500000 0.500000 0.010000 0.010000 0.200000 0.700000 0.700000 0.700000 0.700000 0.700000 0.700000 0.700000
FE.8	2000 2000 2000 2000 2000 2000 2000 200		1.499999 1.499999 1.499999 1.499999 1.999999 1.999999 1.499999 1.499999 1.499999 1.999999 0.70000 2.999998 1.999999
SAMPLE	96 97 98 99 100 102 103 104 105 106 108	ととてごとこことととととととと	133 134 135 136 137 140 141 145 146 146 150

DEPLEN	
POCKS	
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D0039 PUBLICATION LISTING - USGS STATPAC (04/22/71)

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X.Y	0.0 N	0.000100	0.001500			0.0	0.001000	0.002000	0.002000			٠.	0.001200	0.001500	0.001200	0.001000	0.00000	0.002000	0.001000	0.001000	0.001000	0.001000	0.002000	0.001500	0005000		0.01000	2000	0.001000	0.003000	0.001000	0.001500	000100	0.001000	0.002000	20000	1500	2		0.001500	0.001000	0.001000	
× ^ >	0.005000	0.005000	0.007000	0.005000	0.005000	0.010000	0.020000	0.007000	000,000	0.003000	0.005000	0.001000	0.007000	0.00700	0.007000	0.007000	0.007000	0.007000	0.007000	0.007000	0.007000	0.007000	0.001000	0.007000	0.015000	0.007.00	0.01000	0.005000	0.007000	0.007000	0.010000	0.010000	0.0000	0.0000	000200	0001000	000000	000200	0.010000	0.01000	0.007000	0.007000	
SR . 8	0.050000	0.05000	0.020000	0.050000	0.030000	0.030000	0.070000	0.070000	0.030000	000000	0.03000	0.07000	0.050000	0.07000	0.050000	0.070000	0.00000	000000000000000000000000000000000000000	0.00000	0.150000	0.050000	0.000000	0.050000	0.050000	0.050000	0.001000	0.00000	000070-0	00000000	0.015000	0.070000	0.00001.0	0000000	0.00000	0000000		0,00000	000000	0.03000	0.150000	0.03000	00000100	i ! ! !
NI , %	z	zz	z	z	z	00	Z:	Z	z	z	Z 2	Z	_	ب	Z	z .	: ب	Z	J 2	Z	z	z	Z	Z	z:	Z:	z a	: z	z	Z	z	Z:	2 2	z z	2	2 2	Z 2	: 2	2	: 2	: z	z	:
Z	0.0		0	0.0	0.0	0.001500	0.0	0.0	0.0	0 0	0	0	0.0	0.0	0.0	0.0	0.0) 0		0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0		0	0.0	0.0	0.0)))) •		•		•		0		0.0) }
N8 • 8	0.000500	0.000500	0.000500	0.000500	0.000500	0.0	N 0.0	0.000700	00100		2 2	00300	0.0	0.000100	0.000500	0.000500	0.000500	2000	2 2	00500	0.000500	0°0	0.0	300	N 0.0	00200	2 2		01000	0.001000		Z :	N 002000	0.000,00	000000	0.000100	2000.0	00200	0.000.0	2	0100	N 0.0	
× N	0.005000	0.007000	0.001000	0.005000	0.020000	0.005000	0.015000	0.015000	0.003000	0.002000	0.01000	0.005000	0.100000	0.050000	0.200000	0.010000	0.070000	0.010000	0.000000	0.010000	0.010000	0.015000	0.005000	0.00700	0.010000	0.015000	0.010000	0.007000	0.007000	0.007000	0.003000	0.005000	0.003000	0.012000	00010-0	0.00000	0.012000	0.012000	0.015000	000000	0.01000	0.01000))))
LA,%	0.010000	0.00,000	N 0.0	0.0				0.0	0.007000	000/0	2 2	25000	0.005000	0.005000	0.005000	0.005000	0.007000	0.007000	0.0000	0.007000	0.007000	0.005000	0.007000	0.005000	5000	0.0	0.030000	0.001000	0.005000	0.005000	0.00700.0	0.007000	25000	N 0000	0.007000	0.001000	0.002000	000000	2000000	0000	000200	0.00500.0	1
CR,	0.001500	0.5100	0.002000	0.001500	0.002000	0.002000	0.002000	0.002000	0.002000	0.003000	0.002000	0-002000	0.002000	0.002000	0.001500	0.003000	0.002000	0.002000	0.002000	0.003000	0.002000	0.002000	0.002000	0.003000	0.00700	0.003000	0.005000	0.003000	0.003000	0.003000	0.005000	0.003000	0.003000	0.003000	0.002000	0.003000	0.002000	0.003000	0.003000	0000000	0000000	0000000	•
2.00		2 2	00700	0.0	0.0	0	0°0				0.0	2 2	0.001000	0.001000	0.000100	0.001000	03000	N 0.0	0005000	00700	0.000100	0.001000	0.0	0.0	N 0.0	× 0.0	0.0	2 2			0°0	0.0	_	v :	- ·		-		2 2			2	:
H	Z	z	2 2		00	00	00	00	8	0	00	2	: 00	00	00	00	50	00	2	<u>د</u> د	20	20	20	00	00	z	Z	Z 2		200	z	z	Z	z (200	<u>ء</u>	z:	z	3	2 2	Z (2	<u> </u>
8E, %	0.0	0 0		0.000100	0.000100	0.000100	0.000100	0.000100	0.000100	0.000100	0.000100	1000	0-000100	0.000100	0.000100	0.000100	0.000150	0.000100	0.000150	0.000150	0.000150	0.000150	0.000150	0.000100	0.000100	•	0.0	.	0.001000	0.000100	0.0	0.0	0.0	0.0	0.000150	٠	0.0	9	0000	•	֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓֓	•	•
SAMPLE	153	154	156	157	158	159	160	164	165	202	203	204	20,2	207	208	500	210	211	212	217	215	216	217	218	219	524	225	977	228	229	230	231	232	233	234	535	236	167	330	667	241	24.2) - -

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2.11	0.300000 0.300000 0.300000 0.300000 0.200000 0.200000 0.300000 0.300000	0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000	0.30000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000 0.300000
CA,%	0.200000 0.150000 1.000000 1.000000 0.070000 0.150000 0.150000 0.150000	0.300000 0.200000 0.200000 0.500000 0.500000 0.500000 0.300000 0.300000 0.050000 0.070000	0.030000 0.030000 0.030000 0.070000 0.070000 0.300000 0.100000 0.030000 0.100000 0.100000 0.030000
MG*	0.700000 0.700000 0.700000 0.700000 0.700000 0.700000 0.200000 0.500000 0.700000 0.700000	0.500000 0.500000 0.700000 0.700000 0.700000 0.700000 0.500000 0.500000 0.700000 0.700000 0.700000 0.700000 0.700000	0.300000 0.700000 0.200000 0.300000 0.500000 0.500000 0.500000 0.700000 0.700000 0.700000
FE .	2.999998 1.999999 1.499999 0.700000 1.499999 0.0 6.2999998 1.499999 0.0 2.999998 0.500000 0.500000		0.300000 1.499999 2.999998 2.999998 2.999998 0.500000 0.500000 0.0 0.0 0.0 0.300000 1.999999 6.999997 1.999999
SAMPLE	153 154 156 157 158 159 166 165 165 202 203 203 204 204	206 207 208 209 211 212 213 214 215 216 217 228 228	228 229 230 231 232 234 235 235 236 238 241 242 243

D0039 PUBLICATION LISTING - USGS STATPAC (04/22/71)

RGILLIZED ROCKS, PERCENT	HG,2 AS,2 CU,2 ZN,2 MG,2 BA,2	0.000004 0.01000 0.000500 0.0 t 0.0 N 0.030000	0.006000 0.001500 0.0 L 0.001000	0.008000 0.003000 0.0 L 0.001000	0.001000 0.005000 0.0 L 0.000700	0.001000 0.001500 0.0 L 0.0 N
	AS. Z	0.001000	000900*0	0.0800.0	0.001000	0.001000
ROCKS, PERCE	HG•3	0.000004	90000000	0.000000	0.000000	0.000004
ARGILLIZED	81,4	0.0	0.0	0.0	0.0	0.0
	PB • \$	0.000300	0.005000	0.005000	0.010000	0.000100
	AG, %	0°0	0.0	0.0	0.000100	0.0
	AU. E	0.0	0.000117	0.000000	0.000010	0.000010
	SAMPLE	244	245	247	248	253

X * X	0.0 0.00 0.0 0.0 0.0 0.0 0.0
> 2	0.007000 0.007000 0.010000 0.007000
SR, %	0.030000 0.200000 0.150000 0.100000
8,17	zzzz
Z	00000
NB•8	0.001000 0.0 0.0 0.000700 0.000500
MN, 8	0.005000 0.007000 0.005000 0.010000
LA, %	0.005000 0.010000 0.010000 0.005000
CR,	0.003000 0.005000 0.003000 0.003000
2.00	ZZZZZ
	00000
8E•48	0.000150 0.0 0.0 0.000150 0.0
SAMPLE	244 245 247 248 253

71.	0.200000 0.200000 0.200000 0.300000
CA+	0.070000 0.070000 0.070000 0.300000
MG, %	0.300000 0.100000 0.500000 0.700000
FE,*	0.070000 0.0 6 0.0 6 1.499999 1.999999
SAMPLE	244 245 247 248 253

ARGILLIZED ROCKS, PPM

BA, PPM	700.00	700.00	700.00	1000.00	200.00	200.00	2000.00	1500.00	1000.00	1000.00	200.00	300.00	1500.00	2000-00	200.00	200.00	2000.00	200.00	1500.00	1000.00	1000-00	200.00	1500.00	1500.00	700.00	200.00	1000.00	1000.00	1500.00	1000.00	1500.00	100.00	1000.00	2000-00	00.0001	1500.00	1200.00	1000-00	00.0001	1000-00	300.00	1000.00	2000-00	700.00	1500.00	700.00	1000.00	300.00	3000.00	700.00	
MO, PPM	7.00	7.00	N 0.0	7.00	10.00	N 0.0	2.00	3.00	0	2.00	3.00	0.0	0.0	3.00	200	20.00	15.00	10.00	2.00	Z 0.0	2.00	Z :		0	•	0	0	0	0	2 0	Z 0.0	3.00	3.00	50.00	00.7	2.00	3.00	00.0	20.0	2 :	200	10.00	3.00	Z 0.0	2.00	2.00	2.00	20.00	10.00	10.00	
N99.N2	220.00	00.09	30.00	0.0	25.00	40.00	35.00	0.0	25.00	30.00	25.00	0.0	0.0 L	0.0	0.0 L	0.0 L	80.00	40.00	0.0	J 0.0	0.0	0.0	0.0 L	0.0 L	0.0	0.0	0.0 L	0.0	0.0 L	0.0	0.0 L	0.0	40.00	30.00	0.0	80.00	80.00	80.00	80.00	80.00	80.00	80.00	0.0	0°0	0.0	25.00	0.0	40.00	30.00	30.00	
CU, PPM	50.00	10.00	30.00	30.00	20.00	10.00	20.00	15.00	15.00	20.00	10.00	10.00	20.00	20.00	20.00	20.00	2.00	20.00	30.00	10.00	20.00	20.00	70.00	10.00	20.00	10.00	20.00	20.00	30.00	20.00	15.00	10.00	30.00	20.00	20.00	30.00	20.00	30.00	20.00	30.00	20.00	15.00	10.00	10.00	20.00	15.00	7.00	20.00	20.00	30.00	
AS, PPM	10.00	140-00	10.00	100.00	80.00	10.00	10.00	40.00	0.0 L	10.00	10.00	10.00	7 0°0	0.0 L	0.0 L	10.00	20.00	10.00	20.00	0°0	00.09	0.0 C	0.0 L	0°0	0°0	0.0	10.00	10.00	10.00	0.0 0	0.0	40.00	0.0	10.00	0.0	10.00	10.00	10.00	10.00	10.00	10.00	80.00	40.00	10.00	10.00	10.00	120.00	00.09	40.00	20.00	
HG, PPM	01.0	0-13	90-0	0.02	90.0	90•0	\$0°0	90.0	90.0	0.0	0.15	0.11	0.08	0.11	90.0	0.08	0.04	0.08	0.08	90.0	0.20	0.10	0.10	0.10	90.0	90.0	90.0	0.08	0.10	0.04	0.11	0.08	0.08	90.0	0.08	0.08	90.0	0.08	0.08	0.04	90.0	0.08	90.0	90.0	90.0	0.04	0.17	0.11	90.0	0.04	
81,PPM	0				0	N 0.0	0	0	0	0	0	0	N 0.0	0.0	N 0.0	₹ 0.0	N 0.0	•		0.0	N 0.0	•	•	₩ 0°0	Z 0.0	0.0	N 0.0	0.0	0.0	N 0.0		0.0	0.0	N 0.0	Z 0.0	N 0.0	2												N 0.0		
P8 pPM	30-00	20.0	2	200	20.00	15.00	15.00	15.00	15.00	10.00	20.00	50.00	30.00	15.00	20.00	20.00	20.00	30.00	70.00	10.00	150.00	15.00	10.00	7.00	2.00	20.00	7.00	2.00	7.00	7.00	10.00	2.00	15.00	10.00	10.00	15.00	10.00	2.00	7.00	2.00	5.00	70.00	2 0 0	N 0.0	20.00	20.0	7.00	200.00	20.00	30.00	
AG, PPM	0,0	2 0	2000	200	0.30	0.30	0.30	0.30	0.20	0.30	0.30	0.20	0.20	0.30	0.20	0.30	0.10	0.30	2.00	0.30	1.00	0.30	0.20	0.20	0.20	0.20	0.20	0.20	0.20	0.20	0.20	0.30	0.50	0.30	0.30	0.30	0.30	0.30	0.30	0.20	0.70	0.20	N 0.0	N 0.0	N 0.0	N 0.0	0.0	0.30	0.50	0.30	
AU, PPM	01	۵ د	0.10	- 0	0.20	0.0 L	0.0 L	0.10	0.0 L	0.10	0.0	0.0	0.0 L	0.0 L	0.0 L	0.20	0.0 L	0.10	0.28	0°0	0.50	0°0	0.10	0°0	0.10	0.0 L	0.0 L	0.10	0.0 L	0.10	0°0	0.36	0°0	0.0 L	0.0 L	01.0	0.0 L	0.10	0.0 L	0.0	0.20	01.0	0.10	0.10	0.10	1.48	0.18	4.07	0.10	0.10	
SAMPLE	1.2	77	1 1	- 1	1 2	17	18	19	20	21	22	23	24	25	56	27	28	53	31	33	40	41	45	43	44	46	47	84	64	20	51	52	53	54	55	99	2.2	28	29	09	19	62	20	11	72	73	76	93	36	95	

ARGILLIZED ROCKS, PPM

*PPM	10.00 10.00 0.0 N 10.00 7.00 15.00 10.00 15.00 15.00	10.00 10	10.00 10.00 7.00 50.00
>	្គគ គ គ គ គ គ គ គ គ គ គ គ គ	ନ	AA K
V.PPM	150.00 100.00 100.00 100.00 100.00 150.00 100.00 100.00	100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000 100.000	70.00 150.00 100.00
SR, PPM	500.00 500.00 500.00 700.00 500.00 700.00 700.00 500.00 500.00	500.00 500.00 500.00 500.00 500.00 500.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00 700.00	500.00 500.00 300.00
SR,	500.00 300.00 700.00 500.00 700.00 1500.00 1500.00 1500.00 1500.00	1500.00 3 000.00 3 000.00 2 000.00 2 000.00 1 000.00	1500.00 300.00 700.00
NI PPM	0000 0000 0000 0000 0000 0000 0000 0000 0000	10000000000000000000000000000000000000	N 0000 N
NB, PPM			00000 00000 00000
MN, PPM	200.00 100.00 300.00 300.00 700.00 300.00 70.00 70.00 70.00	70.00 70.00 70.00 70.00 30.00 30.00 30.00 30.00 20.00 20.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00	200-00 200-00 200-00 300-00
LA, PPM	100.00 0.0 N 0.0 N 0.0 D N 30.00 70.00 70.00 70.00 30.00 100.00	100.00 100.00 100.00 50.00 70.00 50.00 100.00 100.00 100.00 100.00 50.00 100.00 50.00 100.00 50.00 100.00 50.00	70.00 100.00 50.00 70.00
CR, PPM	20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00	20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00 20.00	15.00 30.00 30.00 20.00
CO.PPM	10.00 10.00		
BE, PPM	10.00 0.00 0.00 0.50 0.30 0.30 0.30 1.50 3.00	0.50 0.50 0.50 0.50 0.50 0.50 0.50 0.50	0000 11.00 10.00 10.00
SAMPLE	23 22 23 26 26 27 28 28 28	221102110282110000000000000000000000000	75 76 94 95

ARGILLIZED ROCKS, PPM

TI,PPM	2000-00	3000.00	3000.00	3000.00	2000-00	3000-00	2000-000	00.000	00000	3000	00.000	3000-00	5000-00	5000.00	2000-00	2000-00	3000.00	2000-00	2000.00	2000-00	3000.00	3000.00	2000-00	5000-00	2000-000	2000	3000-00	2000-00	2000-00	2000-000	3000.00	5000.00	2000-00	2000.00	2000-00	5000.00	2000-000	3000	2000-00	3000-00	3000-00	2000-00	3000.00	3000.00	5000.00		2000.00	
CA, PPM	5000.00	5000.00	2000.00	2000-00	1500.00	7000-00	3000.00	00.0001	1000.00	00.00	14999.99	3000	1000,00	1000.00	7000.00	200.00	1000.00	1500.00	100.00	700.00	700.00	700.00	700.00	700.00	1500.00	1000.00	2000-00	1000-00	1000.00	100.00	1000,000	1000.00	1000.00	1000.00	1000.00	700.00	100-00	100.00	1000,000	1000-00	700-00	2000.00	1000.00	1000.00	1500.00	2000.00	1500.00	
MG, PPM	5000.00	2000.00	7000.00	2000-00	7000.00	7000-00	5000.00	5000-00	3000.00	00.0007	10000	7000	5000-00	7000-00	2000.00	200.00	700.00	2000.00	1500.00	1000.00	3000.00	3000.00	3000.00	7000.00	7000-00	2000.00	2000-00	2000-00	2000-00	1000	3000-00	5000.00	5000.00	3000.00	3000.00	3000.00	3000-00	7000	2000	2000	1500.00	5000-00	5000.00	1000-00	7000.00	10000-00	7000.00	٠
FE,PPM	6666	29999.98	3000	6666	6666	4999	9999	٠	6666	6666	19999.99	000-000	2000.00	5000.00	9 0.0	86*6666	86.66664	96*66669	14999.99	*6*66669			96666	29999.98	2000-00	9999	19999.99	19999.99	14999.99	00.000	20000.08	96°66669	69999.94	49999.98	29999.98	49999.98	49999.98	10000-00	400000	10000 00	10000-00	29999,98	86-66667	29999.98	96.66666	6666	%6°66669	
SAMPLE	12	13	14	15	9 1	11	18	61	20	21	77	23	56	3 %	27	78	67	31	33	40	41	7,5	43	55	46	14	%	64	20	10	27	35	55	26	25	28	65	095	10	7 6	2 5	72	73	92	6		96	

ARGILLIZED ROCKS, PPM

Y,PPM	30.00 20.00	30.00 N 0.00 N 0.00
V, PPM	100.00 10	70.00
SR, PPM	200.00 300.00 100.00 1500.00 1500.00 1500.00 1000.0	300.00 200.00 1000.00 150.00
NI,PPM		2222
NB + PPM	20000000000000000000000000000000000000	2.00 2.00 2.00 2.00 2.00 3.00 3.00
Mdd.NM	300 00 150 00 200 00 100 00 10	100.00 100.00 3.00 5.00
LA, PPM	50.00 50.00 50.00 50.00 50.00 70.00	70.00 100.00 70.00 50.00
CR. PPM	20000000000000000000000000000000000000	10.00 10.00 20.00 20.00
CO.PPM		Z Z Z Z Z 0 0 0 0 0
BE, PPM		00000 00000 00000
SAMPLE	9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9	148 149 150 151 152

ARGILLIZED ROCKS, PPM

TI,PPM	5000.00 3000.00 5000.00 5000.00 3000.00 7000.00		2000-00 2000-00 2000-00 2000-00 2000-00 2000-00 2000-00 2000-00	2000 3000 3000 3000 3000 3000 3000 3000
CA, PPM	1500.00 10000.00 700.00 1500.00 1000.00 2000.00			700.00 1000.00 1000.00 1000.00 1000.00 1000.00 1500.00 19999.99 2000.00 2000.00
MG + PPM	7000.00 5000.00 7000.00 10000.00 10000.00 2000.00 2000.00			2000.00 2000.00 2000.00 2000.00 7000.00 7000.00 7000.00 7000.00 7000.00 7000.00 7000.00 7000.00 7000.00
FE.PPM	49999.98 69999.94 19999.99 69999.98 29999.98 69999.96	6.6666 6.	9999.99	29999.98 14999.99 14999.99 19999.99 19999.99 14999.99 14999.99 29999.98 29999.98 29999.98 29999.98
SAMPLE	96 97 98 99 100 102 103	106 107 108 110 111 113 113	115 122 122 124 126 127 131 133 134	135 135 136 137 140 140 140 150 150

ARGILLIZED ROCKS, PPM

																															•																•				
BA, PPM	700.00	200.00	700.00	1000.00	2000.00	700.00	700.00	70.00	2000.00	300.00	500.00	2000-00	1000-00	2000	00.007	1000	1000.00	1000.00	1000.00	1000.00	1000.00	1000.00	700.00	1000.00	1000-00	200	00.00	0000	00.00	200-000	200.00	1500.00	1500.00	200.00	7000.00	1500.00	200.00	200.00	100.00	2000.00	300.00	200.00	2000.00	300.00	1000.00	300.00	1500.00	300.00	1000.00)))))	
MO, PPM	N 0.0	3.00	3.00	3.00	3.00	3.00	2.00	30.00	7.00	3.00	15.00	5,00	N 0			00.0	3.00	3.00	2.00	2.00	0.0	7.00	5.00	5.00	7.00	2		00.0	200	2	2.00	7.00	2.00	2.00	N 0 0	N 0.0	3.00	3.00	3.00	3.00	3.00	3.00	2.00	10.00	3.00	3.00	5.00	5.00	2,00)))	
ZN, PPM	40.00	0.0 L	50.00	20.00	20.00	00.09	25.00	30.00	0.0	0-0	-0.0	- 0.0	-		00.00	72.00	20.00	20.00	100.00	40.00	50.00	110.00	30.00	190.00	20.00	00.00	00.021	110.00	00.07	00.09	0.0	0.0	0.0	25.00	20.00	00.09	20.00	0.0 L	0.0 L	۰۰0 د	50.00	0.0 L	0.0	0.0	0.0 L	0°0	0.0	0.0	10.0)))	
CU.PPM	5.00	10.00	10.00	10.00	7.00	7.00	10.00	200.00	15.00	7.00	10.00		20.00	00.0	00.00	30.00	30.00	10.00	20.00	20.00	50.00	30.00	20.00	30-00		00.01	00.01	10.00	10.00	2.00	10.00	20.00	150.00	150.00	15.00	20.00	20.00	15.00	10.00	30.00	10.00	30.00	100.00	5.00	20.00	7.00	20.07	10.00	50.00		
AS, PPM	20.00	10.00	10.00	10.00	120.00	10.00	10.00	120.00	20.00	40.00	00.04				7 0.0	40.00	10.00	10.00	10.00	0.0	00.09	10.00	10.00		00.01	20.01	00.01	10.00	20.00	10.00	40.00	80.00	100.00	120.00	10.00	10.00	100.00	00.09	80.00	10.00	10.00	10.00	10.00	10.00	10.00	20.00	60.00	10-00	10.00) •	
HG • PPM	0.18	0.12	0.15	0.20	0.16	0.40	0.44	0.52	0.28	0.24	17.0	000	5.0	5.0	0.0	90.0	0.0	0.10	90.0	90.0	0.06	0.0	90.0	900	• •	0.0	90.0	0.0	0.08	0.04	90.0	0.08	90.0	0.10	0.04	0.04	0.04	0.11		90.0		0.04	0.17	90.0	0.10	0.10	0,10	0.20	0.10	•••	
BI + PPM			0.0									3 6				0	0	_			. N	2	2			2:	2	2000	200	0		0	0	0	0	0.0	N 0.0			0.0							-	2000			
PB, PPM	10.00	0.0	2.00	2.00	2.00	5.00	20.0	2	20.03	00.41	00.01	20.00	2 2	2 0	20.00	3.00	3.00	7.00	5.00	7.00	200.7	20.	000	000	00.00	20.00	10.00	7.00	7.00	3.00	N 0.0	3.00	3.00	20.00	N 0.0	N 0.0	50.00	50.00	10.00	N 0.0	7.00	7.00	5.00	50.00	50.00	20.05	20.00	000		70.00	
AG, PPM	0.0	2000	Z 0.0	N 0.0	20.0	N 0.0	2 2	2 2				2 2	2 2	2	Z 0.0	2 0 0	0.0	N 0.0	Z C	2	2			2 2	2 2	Z :	o. 0	N 0.0	0.0 N	Z 0.0	0.0	0.0	N 0.0	N 0.0	N 0.0	0.0	0.0	1.00	Z 0.0	7.00	2000	200	0.0	2000	20.0			12		2	
AU, PPM	0.19	0-0	0.0	0.0	1-00	1.67	81.0		•	•	۰	7	٥,	0.10	0.46	0.0 L	0.10	0.0	-0-0	-	1 6	•	2.0	•		7 0.0	0.0	0.0 L	0.10	0.10	0.0 L	0.10	0.0	0.0	0.10	J 0.0	0.0 L	0.10	0.16	96.0	0-20	0-0	0.10		-	3,60	0.00	70.0	97.0	o . o	
SAMPLE	153	154	155	156	157		0 6 1	160	201	+01	165	707	203	204	502	506	207	208	200	310	210	117	217	513	517	215	216	217	218	219	224	225	226	227	228	229	230	231	232	233	234	225	236	247	220	050	623	147	247	243	

ARGILLIZED ROCKS, PPM

W d d & Å	100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00	15000 15000 10000 2000 10000 10000 10000
Wdd 6 A	50.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00 100.00	100.00 100.00 100.00 100.00 100.00 100.00 100.00
SR, PPM	500 000 500	150.00 1000.00 1500.00 500.00 700.00 2000.00 300.00 1500.00
Mdd.IN		
NB PPM		10.00 0.00 7.00 7.00 7.00 7.00 7.00 10.00
MN, PPM	50.00 70.00 70.00 50.00 50.00 150.00 150.00 100.00 700.00 700.00 700.00 100	70.00 30.00 30.00 30.00 150.00 170.00 150.00 150.00 100.00
LA, PPM	100.00 50.00 0.00 N 0.00 N	50.00 70.00 70.00 70.00 70.00 70.00 70.00 50.00 50.00
CR, PPM	15.00 15.00 20.00	30.00 30.00 30.00 30.00 30.00 30.00 30.00 30.00
CO,PPM	00.00 00	
BE, PPM	N N N N N N N N N N N N N N N N N N N	N N N N N N N N N N N N N N N N N N N
SAMPLE	153 154 155 155 156 157 158 158 157 158 157 158 157 158 157 158 157 157 158 158 158 158 158 158 158 158 158 158	229 230 231 233 233 234 234 235 236 237 241 242 243

ARGILLIZED ROCKS, PPM

II, PPM	3000,00 3000,00 3000,00 3000,00 3000,00 1000,00 2000,00 3000,00 3000,00	3000.00 3000.00 3000.00 3000.00 3000.00 3000.00 3000.00 3000.00 3000.00	3000.00 3000.00 3000.00 3000.00 3000.00 3000.00 3000.00 3000.00 3000.00 3000.00
CA, PPM	2000.00 1500.00 2000.00 10000.00 29999.98 2000.00 1000.00 1500.00 1500.00	8000.00 8000.00 8000.00 8000.00 8000.00 8000.00 8000.00 8000.00 8000.00	300.00 700.00 700.00 300.00 300.00 700.00 700.00 700.00 1000.00 1000.00 1000.00 1000.00 1000.00
MG.PPM	7000.00 7000.00 7000.00 7000.00 7000.00 7000.00 2000.00 2000.00 5000.00 69999.94 7000.00	7000.00 3000.00 5000.00 7000.00 7000.00 7000.00 7000.00 7000.00 7000.00 5000.00 5000.00	7000.00 30.00 7000.00 7000.00 7000.00 150.00 7000.00 5000.00 5000.00 7000.00 7000.00 7000.00 7000.00 7000.00
FE, PPM		3000.00 19999.99 29999.98 7000.00 19999.98 19999.98 29999.98 29999.98 19999.98 19999.98	19999.99 99999.94 49999.98 3000.00 14999.99 29999.98 29999.98 29999.98 5000.00 5000.00 49999.98 69999.99 69999.99
SAMPLE	153 154 155 156 158 160 164 202 203	205 206 207 207 208 211 212 213 214 215 216 217	224 225 226 230 231 233 234 234 234 241

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	AU, PPM	AG . PPM	PB, PPM	81,PPM	нс, ррм	AS, PPM	CU,PPM	ZN.PPM	MO, PPM	BA, PPM
	0.0 L	2 0	3.00	0.0	0.04	10.00	2.00	0.0 L	N 0.0	300.00
	.17	0.0	20.00	0°0	90.0	00.09	15.00	0.0	10.00	500.00
	10	Z 0.0	20.00	N 0.0	0.10	80.00	30.00	0.0 L	10.00	700.00
,	10	1.00	100.00	0.0	0.80	10.00	50.00	0.0 L	7.00	500.00
253 0.	.10	0.0	7.00	N 0.0	0.04	10.00	15.00	7 0.0	N 0.0	1000.00

0.0 N 0.0 N 10.00 N

Y,PPM

70.00 70.00 100.00 70.00 V, PPM 300.00 2000.00 1500.00 1000.00 700.00 SR, PPM NI,PPM 10.00 0.0 N 7.00 5.00 NB,PPM ARGILLIZED ROCKS, PPM MN, PPM 50.00 70.00 50.00 100.00 50.00 100.00 100.00 50.00 0.0 N LA, PPM 30.00 30.00 30.00 30.00 CR, PPM ZZZZZ CO, PPM 1.50 0.0 N 0.0 N 1.50 BE, PPM 244 245 247 247 248 253 SAMPLE

ARGILLIZED ROCKS, PPM

TI,PPM	2000.00 2000.00 2000.00 3000.00 3000.00	
CA, PPM	700.00 700.00 700.00 3000.00	
MG,PPM	3000.00 1000.00 5000.00 7000.00 3000.00	
FE, PPM	700.00 0.0 6 0.0 6 14999.99 19999.99	
SAMPLE	244 245 247 248 253	

PRINT REPEATED BY OPERATOR

86		
BA, LOG	1.1549 1.1549 1.1549 1.1549 1.1529 1.	-0.8239 -0.8239 -1.0000 -1.5000 -1.5000 -1.1549 -1.1549 -1.5229 -1.5229
MO, LOG\$	-3.1549 -3.1549 -3.1549 -3.0000 -3.3010 -3.5229 -2.6990 -2.6990 -2.6990 -3.1549 -3.5229 -3.3010 -3.5229 -3.3010 -3.5229 -3.3020 -3.5229 -3.3020 -3.	-3.3010 -3.3010 -3.3010 -3.3010 -3.0000 -3.3010 -3.3010 -3.3010 -3.3010 -3.3010 -3.3010 -3.3010 -3.3010
ZN, LOG\$	-1.6576 -2.2218 -2.5229 -2.6021 -2.3979 -2.6021 -2.5229 -2.6021 -2.3979 -2.3979 -2.3979 -2.3979	2.0969 -2.0969 -2.0969 -2.0969 -2.0969 -2.0969 -2.0969 -2.0969 -2.0969 -2.0021 -2.5229
CU, LOG%	2.3010 -2.5229 -2.5229 -2.6990	
AS.106%	-3.0000 -2.0969 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -2.6990 -2.6990 -2.6990 -2.6990 -2.6990 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -3.0000	-3.0000 -3.0000 -3.0000 -3.0000 -3.0000 -2.0969 -3.0000 -1.9208 -2.2218 -2.2218
HG, LOG\$	5.0000 5.0000	-5.0969 -5.2218 -5.2218 -5.2218 -5.2218 -5.2218 -5.3379 -4.7696 -4.9696 -5.3979
%90	Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z	Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z
BI, LOG%		
PB, LOG\$	2.5529 -2.5529 -2.65990 -2.8239 -2.8239 -2.8239 -2.8239 -2.65990 -2.65900 -2.65990 -2.6590 -2	-3.000 -3.000 -3.3010 -3.1549 -3.6990 -2.1549 0.0 N 0.0 N -2.6990 -3.1549 -1.6990 -2.3010 -2.5229
A6,L06%	00.00 00	14.5229 14.5229 14.5229 14.6990 14.1549 14.1549 14.1549 16.00 10.0
AU, LOG\$	-4.7212 -4.7447 -4.7212 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.	-5.0000 -5.0000 -5.0000 -5.0000 -5.0000 -5.0000 -3.3904 -5.0000
SAMPLE	282222222222222222222222222222222222222	55 56 56 56 56 56 56 56 56 56 56 56 56 5

																						*.																							
Y, LOG\$	-3.0000	O.0	Š	-3.1549	-3.0000	-2.8239	-3.1549	-3.0000	6775.7-	-2.302.9	-2,8239	-3.0000	-3.1549	-3.1549	-3.0000	-2.5229	0106-6-	-2.6990	-2.5229	-2.6990	-2.8239	-2.8239	-3.1549	-3.0000	-3.0000	-3.1549	13. 3010	-3.1549	-2.8239	-3.1549	-3.0000	-3.1549	-3.1549	-2.1540	-2.5229	-2.6990	-2.6990	-2.6990	-2.6990	o.0	-3.0000	-3.0000	-2.1549	2122.47	
V.L06%	-1.8239 -1.8239	-2.0000	-2.0000	-2.1549	-2.0000	-1.8239	-2.1549	-1.8239	6461-7-	-3 6230	-2.1549	-2.0000	-2.1549	-2.1549	-2.0000	-1.8239	2 1540	-2.1549	-2.0000	-2.0000	-2.0000	-2.0000	-2.1549	-2.1549	-2.0000	-2.0000	-2.0000	-2.1549	-2.0000	-2.0000	-2.1549	-2.1549	-2.1549	-2.3030	-2.0000	-2.0000	-2.0000	-2.1549	-2.0000	-2.0000	-2.1549	•	-2.0000	-2.0000	
SR, L06%	-1.3010	-1.1549	-1.3010	-1.1549	0669.0-	-0.8239	-1.0000	-1.1549	6775-1-	00070-	-0.8239	-0.8239	-0.8239	-1.3010	-0.5229	-1.1549	0669*0-	-0.6990	-0.8239	-1,3010	-1.6990	-1.0000	-1.1549	-0.8239	-1.1549	-1.3010	1 6230	-1.1549	• (-1,1549	-1.1549	-1.3010	-1.0000	7.0000	-1-3010	-0.8239	-1.5229	-1.3010		-1.5229	-1.3010	φ,	-1.5229	C461-1-	
NI, LOG\$	000		2	-3.3010	-3.5229	-3.1549	3.154		-3-1549	-3.36.69	-3.6990		-3.0000	-3.5229	-3.5229	-3.6990	-3.5229	12 1540 N	N 0.0	649	-3.5229	0.0	-3.5229	-3.5229	<u>ر</u>	N 0.0		-3 6990	-3, 5229	-3.3010	0°0	-3.6990	-3.3010	73,4000	-3.5229	-3,3010	0.0	0.0	0°0	0.0	v 0.0	-3.5229	-3.5229	-2.36.6-	
NB, LOG\$	000		0.0 N	N 0.0	0.0	0.0	0.0	-3.3010	2 0	2 2	-3-3010	-3.3010	0.0	-3,3010	0.0		-3.3010	2 2	2 2	0.0	0.0	-3,3010	0°0		-3.3010	-3.0000	2 2	2 2	-3.3010		-3.3010	10	000	12 2010 M	-3-3010	d	0°0	0°0	0°0	0.0	O.0	N 0.0	-3.5010	'n	
MN, LOG2	-1.6990	-2.0000	-1.5229	-0.5229	-1.1549	•	-1.6990	.	-2.1549	2 2010	-2.1549	-2.1549	-	_	-	-2.1549	-2.5229	-2.5010	-2,1549	-2,5229	-2.0000	-2.0000	•	-2.3010	-2.3010	-2.5229	-2.6990	6461-7-	-1.8239	-1.6990	-1.8239	-2.0000	-2.1549	20000	-2.0000	-1.6990	-2.5229	-2.6990	•	-2.3010	-3.0000	ω,	0669*1-	-1.5229	
LA, LOG%	-2.0000 0.0 N	0.0	0.0 N	-2,3010	-2.1549	-2.1549	-2.1549	-2.3010	-2.5229	2 0000	-2.1549	-2.0000	-2.1549	-2.0000	-2.3010	-2.3010	-2.1549	6776.7-	-2.3010	-2.3010	-2.1549	-2.0000	-2.3010	-2.1549	-2.1549	-2.1549	8	U.U N	-2.3010	-2,1549	-2.0000	-2.3010	-2.1549	2 2010	-2.1549	4 (1)	-2.1549	-2.1549	-	-2.1549	-2.1549	-2.0000	-2.3010	-2.1549	
CR, LOG\$	-2.6990	~	2 0	-2.6990	-2.5229	-2.6990	-2.6990	-2.5229	-2.6990	0669*7-	6776-7-	-2.6990	-2.8239	-3.0000	-2.6990	-2.5229	-2.5229	2.6990	-2.6990	-2.5229	-2.6990	-2.6990	-2.6990	-2.6990	-2.5229	-2.6990	-2.5229	-3.0000	-2.5229	-2.5229	-2.5229	-2.5229	-2.6990	6776*7-	-2.6259	-2.5229	-2.6990	-2.8239	-2.6990	-2.8239	-2.8239	۲,	-2.5229	0660-7-	
2907.00	-3.0000	O.0		-3.0000	-3.1549	-3.1549	-3.1549	-3.1549	-3.1549	-3.1549	-3-1549	-3,3010	-3.1549	-3.5229	-3.5229	0.0	Z :	N 0000	-3.1569	-3-1549	-3.3010	0.0	-3.3010	-3.3010	-3.3010	-3.3010	-3.3010	-3.3010	-3-1349	-3,0000	-3.3010	-3.1549	-3.3010	-5.1549	2 2	2 2	0.0	N 0.0	0°0			229	Z 2	o. 0	
BE, LOG\$	-3.0000 0.0 N		N 0.0	-4.3010	-4.5229	-4.5229	-4.5229	-4.3010	-4.3010	-3.8239	-4.0000	-3.5229	-4.3010	-4.5229	-4.5229	-4.3010	-3.5229	-4.5229	-4.5010	-4-0000	-4.1549	-3.6990	-3.8239	-4.0000	-4.1549	-4.1549	-4.3010	-4.5229	6776-4-	-4.5729	-4.5229	-4.5229	-4.5229	-4.3010	-4.1549	-4.0000	N 0.0	0.0			0°0	-4.0000	-4.0000	-4.0000	
SAMPLE	12	14	5	17	18	19	20	21	22	- 23	24 25	26	27	28	53	31	33	0+;	41	7 7	5 5	94	4.1	48	64	20	21.	25	20.0	55	26	57	80 10 10 10 10 10 10 10 10 10 10 10 10 10	66	2 5	62	22	11	72	. 73	16	93	96	c k	

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84		000000000000		
11,106%	-0.6990 -0.5229 -0.5229 -0.5229 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.5229 -0.5229	-0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.5229	-0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010	-0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.3010
-	000000000000000000000000000000000000000	9999999999		
20			 	***
CA, LOG\$	-0.3010 -0.3010 -0.3010 -0.3010 -0.8239 -0.5229 -1.0000 -1.0000 -1.1549 0.1761	-1.0000 -1.0000 -1.3010 -1.3010 -1.8239 -1.1549 -1.1549	1.1549 1.1549 1.1549 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000	-1.1549 -1.1549 -1.0000 -1.0000 -1.1549 -1.0000 -1.0000 -0.6239
CA,				77777777000
MN, LOG	0.3010 0.3010 0.1549 0.1549 0.1549 0.1549 0.1549 0.1549	-0.3010 -0.1549 -0.6990 -1.6990 -1.1549 -0.3010 -0.8239 -0.5229	0.5229	-0.5229 -0.1549 -0.1549 -0.3010 -0.3010 -0.3010 -1.0000 -0.0549
Z.	99999999999	000000000000000000000000000000000000000		000000000000
290	0.8451 0.4771 0.5229 0.3010 0.3010 0.4771 0.4771 0.3010 0.3010	0.4771 0.3010 0.0 6 0.6990 0.6990 0.8451 0.1761 0.3010	0.5010 0.4771 0.4771 0.3010 0.3010 0.1761 0.1549 0.6990 0.8451 0.8451 0.6990	0.6990 0.0000 0.3010 0.3010 0.0000 0.4771 1.0000 0.4771 0.8471
FE,L06%	0.8451 0.4771 0.3010 0.3010 0.3010 0.1761 0.4771 0.3010 0.3010	0.477 0.0 0.0 0.699 0.845 0.176 0.301 0.301	0.5010 0.6771 0.6771 0.3010 0.3010 0.1761 0.6990 0.6990 0.6990 0.6990 0.6990	0.6990 -0.0000 0.3410 0.3451 0.3010 0.4771 0.6771 1.0000 0.8451
m m	12 14 14 16 11 16 17 17 18 19 22 23 24	25 26 27 28 29 33 40 41	574 574 574 574 574 574 574 574 574 574	59 60 61 72 73 73 73 95
SAMPLE				
رس				

SAMPLE	AU, LDG%	A6, LOG%	PB,L06%	81,106%	HG, LOG%	AS, LOG#	c0, L06%	ZN, LOG2	MO, LOG&	84,106%
96	-5.0000	-4.5229	-2.6990	N 0.0	-5.0000	-3.0000	-2.3010	397	-3.0000	0669*0-
97	-3.7282	-4.5229		_	0669**-	-2.6990	-3.1549	.5229	-2.8239	0669-0-
86	-4.6990	-4.5229	-2.0000	N 0.0	0669**-	-3.0000	000	0	-3.5229	-1.0000
66	-4.7212	-4.5229	-2.5229	_	-5.0000	-3.0000	2	•	-3.1549	•
100	-5.0000	-4.5229	٠	_	-5.0000	-3.0000	-2.8239	•	-3.1549	0669.0-
102	-5.0000	-4.1549	-1.6990	_	-5.0000	-3.0000	-3.1549	٠	-3.3010	•
103	-4.3188	-4.0000	-1.6990	_	-5.0000	-3.0000	٠	•	-3.3010	•
104	0.0	-4.3010	-1.6990	0°0	-4.3979	-3.0000	-2.5229	2.39	-3.1549	•
105	-5.0000	-3.8239	-2.1549	_	•	-2.6990	٠	•	'n,	•
106	-5.0000	-4.3010	-2.8239	 O	•	-3.0000		· •	-3.5229	-1.1549
107	-4.6990	-4.5229	-2.6990	•	-5.0000	-3.0000	2	ż	-3.5229	•
108	-5.0000	-4.1549	-2.8239	٠	•	'n,	2		-3.5229	-0.8239
109	7 0°0	-4.3010	-3.0000	0	ń	-3.0000	\sim	9756-1-	20000	0660-0-
110	-4.7696	-4.5229	-3.0000	0,0	0000	6765-7-	6775*7-	-1-1310	-3.0000	0000-1-
I :	-5.0000	-4.5229	7-15-1549	- ·	•	-2,6990	-3.0000	0.0	-3,3010	-1-0000
711	0000		-2.1540	•	-5,000	10	-3-0000	-2.1549	-3,3010	-1.1549
611	15,000		-3,1549	• •		-2.0000	2	2.3	-3.1549	-1.1549
111	-4-8239	200	-3,1549			0.0	•	-2.0969	-3.3010	-1.1549
122	0.0	0.0	-2,3010	0		-2.6990	-2.5229	-2,3010	-3.0000	•
123	0.0	0.0	-2.1549	0	-5.3979	-3.0000	-3.3010	0.0	-3.3010	-1.5229
124	-4.4318	0	0°0	0	-5.3979	-3.0000	-3.5229	0.0		-1.6990
125	-5.0000	_	-3.1549	•	-5.3979	-3.0000	٠	ż	0	-1.6990
126	-4.7959	0.0	-3.0000	_	-5.0000	-3.0000	'n,	2.522		-0.8239
127	-4.3010	N 0.0	-2.3010	0	'n.	-3.0000	'n o	'n.	20.0	0106-1-
128	0.0		-2.6990	0	ŝ	•	,	,	-5.5010	10.8239
131	0.0	N .	-2,3010	•	•	S	2	-2.3979	0106.6-	-1.0000
132		Z 2	-2.3010	2 2	-5.2218	0669-7-	-2.6990	-2.3979	N 0000	-1.8237
153	7 000	2 2	-2.3010	2 2	, K		10	-2.0458	7	-0.8239
135	0.0.0	2	-3,1549			3		-2.3979	-3.3010	-0.8239
136	0.0	000	-2.1549	N 0.0	-4.9208	-2.6990	-3.3010	-2.3979	0.0	-1.1549
137	-4.5229	0.0	-2.5229	0	-4.4089	-2.6990	-3.3010	-2.5229		٠
138	0.0	0.0	-2.5229	0.	-5.2218	-2.6990	-2.5229	0.0	0 (-1.3010
139	0.0	0.0	-3.0000	•	-4.8861	0.0	-3.1549	0.0	-3.5229	6461-1-
140	-4.3979	N 0.0	0.0	•	-5.2218	÷.	ċ	-2.5229	m c	6461-1-
141	0.0	-3.6990	-3.1549	_	-4.1959	-3.0000	٠	6776*7-	6776.6-	•
143	-4.2218	0°0	-3.1549	0	-4.6990	٠	-3.1549	0.0	-3.3010	•
144	-3.9586	0°0	-3.1549	0	-4.7959	-2.0969	٠	0	-3.3010	0000-1-
145	0.0	0.0	-3.1549	•	-4.9208	-3.0000	•	229	-3.3010	-1.5229
146	-4.7959	0°0	-3.1549	0.	-4.8861	•	٠	0	-3.5229	-1.0000
141	-4.1580		-3.1549	•	-4.8239	•	-2.5229	0	-3.1549	-1.3010
148	-4.0088		-3.0000	0	-4.6778		-3.1549	•	523	0000-1-
651	-4.6990		_	_	-4.3665	•	٠	٠	2 0.0	105.
150	0.0	0°0	-3.1549	0.0 N	•	-2.3979	000	•	ກຸດ	154
151	-5.0000	0	522	-3.3010	-4.6198	-2.2218	S	0.0	•	-1.3010
152	0.0	0.0	-2.5229	-3.3010	-4.5229	-2.0969	-3.5229	0.0	-3.5229	-1.1549

ARGILLIZED ROCKS, LOG PERCENT

Y,L06%	12.5529 12.6990 12.6990 12.6990 12.6990 12.6990 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.3010 13.30100 13.3010 13.3010 13.3010	, , , , , , , , , , , , , , , , , , ,
V,L06%	-2.0000 -2.00000 -2.0000 -2.0000 -2.0000 -2.0000 -2.0000 -2.0000 -2.0000 -2.00	-2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.3010 -2.3010
SR, LOG%	-1.6990 -1.5529 -1.1549 -1.1549 -1.1549 -1.1549 -1.1549 -1.1549 -1.1549 -1.1549 -1.1549 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229	
NI, LOG2		
NB, LOGI	-3.3010 -3.3010 -3.3010 -3.3010 -3.3010 -3.1549 -3.1549 -3.1549 -3.3010 -2.6990 -2.6990 -2.6990 -2.6990 -2.6990 -3.0000	ကိုကိုကိုကိုကိုကိုကိုကိုလိုကိုကိုကိ
MN.LOG%	-1.5229 -1.6990 -1.6990 -1.6990 -2.0000 -2.0000 -2.0000 -2.0000 -2.0000 -2.0000 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549	-2.0000 -2.1549 -2.3010 -2.3010 -2.3010 -2.5229 -2.1549 -2.0000 -2.3010 -2.3010 -3.5229
LA, LOG%	-2.3010 -2.3010 -2.3010 -2.3010 -2.3010 -2.3010 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549 -2.1549	-1.8239 -2.1549 -2.1549 0.0 N 0.0 N -2.3010 -2.1549 -2.1549 0.0 N -2.1549 -2.1549
CR, LOG%	- * * * * 1 * * * * * * * * * * * * * *	2.8239 -2.8239 -2.6990 -2.6990 -2.6990 -2.8290 -2.8239 -3.0000 -2.8239 -3.0000
00,106%	00000000000000000000000000000000000000	
8E, LOG3	44.4.1.1549 4.4.1	0.00 0.00
SAMPLE	96 97 98 99 100 100 100 100 100 100 100 100 100	138 139 140 141 144 145 147 148 150 151

D0039 PUBLICATION LISTING - USGS STATPAC (04/22/71)

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		TOTAL TOTAL

TI, LOG\$	-0.3010 -0.5229 -0.5229 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010	-0.5229 -0.5930 -0.6990 -0.6990 -0.6990 -0.5229 -0.5229 -0.5229	0.6990 0.6990 0.6990 0.6990 0.6990 0.6529 0.6529 0.6529 0.6529 0.6529 0.6529 0.6529
CA, LOG\$	-0.8239 -1.0000 -0.0000 -1.1549 -0.8239 -0.6990 -1.0000 -0.8239 -0.8239 -0.8239	-0.8239 -0.6990 -1.1549 -1.1549 -1.1549 -1.3010 -1.1549 -1.0000	-1.0000 -1.1549 -1.1549 -1.1549 -1.1549 -1.1549 -1.0000 -1.3010 -1.0000 -1.0000 -1.0000 -1.0000 -1.0000
MN.LOG%	-0.1549 -0.3010 -0.1549 -0.0000 -0.3010 -0.5229 -0.3010 -0.5229 -0.3010 -0.5229	-0.3010 -0.529 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.1549	-0.6990 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.3010 -0.1549 -0.1549 -0.1549 -0.1549 -0.1549
FE, LOG\$	0.6990 0.8451 0.3010 0.6990 0.6990 0.4771 0.8451 0.1761 0.1761 0.8451 0.8451	0.6990 0.4771 0.4070 0.4771 0.3010 0.3010 -0.5229 -0.3010 0.1761	0.8451 0.1761 0.1761 0.1761 0.1761 0.3010 0.3010 0.4771 0.4771 0.4771 0.3010
SAMPLE	96 97 98 100 102 104 105 106 108	111 112 113 122 122 124 126 127	132 133 133 133 140 140 144 146 150 151

SAMPLE	AU, LOG\$	AG, LOG\$	PB, LOG2	BI, LOG%	HG,L06%	AS, LOG%	2907.00	ZN, LOG%	MD, LOG%	BA, LOG%
						0007	0100	0706 6	2	-1.1540
153	-4.7212	Z 2	-3.0000	2 2	14.0208	-2-6990	-3.0000	0.0	-3.5229	-1.3010
154		-	2010		-4-8239	-3-0000	-3.0000	-2,3010	-3.5229	-1.1549
156	0.0		-3,3010	. 0	0669**-	-3.0000	-3.0000	-2.3010	-3.5229	-1.0000
157	-4-0000		-3.3010	0	-4.7959	-1.9208	-3.1549	-2.3010	-3.5229	0669*0-
158	-3,7773	. 0	-3,3010	0	-4.3979	-3.0000	-3.1549	-2.2218	-3.5229	-1.1549
159	-4.7447	0	0	0	-4.3565	-3.0000	ě	-2.6021	-3.3010	-1.1549
160	-4.3979	0	•	0	-4.2840	-1.9208		-2.5229	-2.5229	-2.1549
164	-4.0458		-2.3010	0	-4.5528	-2.6990		0.0	-3.1549	-0.6990
165	-4.2218	0	-2.8239	0.	-4.6198	-2.3979	'n,	0.0	-3.5229	-1.5229
202	-3.2076	0	.3010	0	-5.3979	-2.2218	-3.0000	0.0	-2.8239	-1.5010
203	-4.4437	0		0	-5.3979	-2.2218	5	0.0	-3.3010	0669-0-
204	-5.0000	0	0	•	-5.3979	0.0	'n,	0.0	2 :	0000
205	-4.3372	0	-2.3010	•	0.0 L	J 0.0	'n,	-2.5229	0.0	-0.6990
506	0.0		-3.5229	0	-5.2218	-2,3979	,	7.00.7	0100-6-	1.0000
207	-5.0000	0	-3.5229	0	-5.3979	-3.0000		-2.3010	-2.5229	0000
508	0.0	0	-3.1549	0	٠,	-3.0000	-3.0000	2 2000	230.00	1.0000
509	0.0	0	-3.3010	٠	, N	-3.0000	0669.2-	0000-7-	0100.0	1.0000
210	0.0	0	-3.1549	0	-5.2218	0.0	06697-	6166-7-	0100.0	0000
211	-4.5229	0	-3.1549	0	-5.2218	-2.2218	ż	-2.3010	20.0	0000
212	-5.0000	0	-3.5229	0	-5.0969	-3.0000	-2.5229	9866-1-	-3.1249	1.0000
213	0.0	0	-3.5229	_	-5.2218	-3.0000	0669*7-	6776.7-	0100-6-	1 0000
214	0.0		-2.5229	-	-5.3979	-3.0000		-1.7212	0105.6-	-1-0000
215	0.0	0	-2.6990	_	-5.2218	-3.0000	-3.0000	-2.5010	*	0000-1-
216	0.0	0°0	-3.0000	 0	-5.2218	-3.0000		-1.9208	Z 0.0	-1.5229
217	0.0		-3.1549	o.0	-5.3979	-3.0000	-3.0000	-1.9586	-3.3010	-1.3010
218	-5.0000	0.0	-3.1549	•	-5.0969	-2.6990	-3.0000	-2.1549	-3.5229	-1.3010
219	-5.0000		-3.5229	0	-5.3979	-3.0000	-3.3010	-2.2218	N 0.00	-1.3010
224	0.0	0°0	0°0	0	-5.2218	-2.3979	-3.0000	0.0	-3.5010	0669
225	-5.0000	0°0	-3.5229		-5.0969	-2.0969	-2.6990	0.0	-5.1249	-0.66.39
226	0.0	0°0	-3.5229	0	-5.2218	-2.0000	-1.8239	0 (0105-6-	-0.8639
227	0.0		-2.3010	•	-5.0000	-1.9208	-1.8239	v	0100-6-	1.0990
228	-5.0000		0.0	0	ŝ	-3.0000	-2.8239	v		0.101
525	0.0	O .	0		٠, ۱	-3.0000	0.00-7-	-2 2010	-2 5720 R	-1-3010
230	0.0	0.0	-2.3010	2.0	6160-0-	2 2210	0706-7-	•	-3.5229	1-6990
231	-5.0000	2	-2.3010		-5.2218	0177.7	-2.1549	0.0	-3.5229	-1.1549
787	-4.1939	2000	0000		-5,2218	-3.0000	-2.5229	0.0	-3.5229	0669*0-
253	4.0177	V+C1•C1	-3.1549		-5.3979	-3,0000	-3.0000	-2,3010	-3.5229	-1.5229
407	06.00		-3.1540		, ,	-3.0000	-2.5229	0.0	-3.5229	-1.6990
236	-5.0000		-3,3010	۰	4.	-3.0000	-2.0000	0°0	•	0669*0-
23.7	0.0	0	-2.3010	0	-5.2218	-3.0000	•	0°0	-3.0000	-1.5229
238	0.0	2000	-2,3010		-5.0000	-3.0000	•	0.0	.522	-1.0000
239	-4.5850	-4.0000	-2.3010		-5.0000	-2.6990	-3.1549	0.0	.522	-1.5229
241	-3.5129	7 0°0	-2,3010	0.0	-5.0000	-2.2218	-2.1549	0.0	.301	-0.8239
242	-4.5528	0.0	-3.3010	0	•	-3.0000	-3.0000	0.0	-3.3010	-1.5229
243	-4-1079	0.0	-3.0000	0	-5.0000	-3.0000	-2.3010	0.0	•	-1.0000

Y, LOG%	13.1549 13.1549 12.8239 12.6990 12.6990 12.6990 12.6990 13.0000	-2.6990 -2.6990 -2.6990 -2.8239 -2.8239 -3.0000 -3.0000
V,L06%	-2.3010 -2.3010 -2.3010 -2.3010 -2.3010 -2.3010 -2.1549	-2.1549 -2.0000 -2.0000 -2.1549 -2.1549 -2.0000 -2.0000 -2.1549
SR, LOG\$	-1.3010 -1.3010 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.3010 -1.3010 -1.3010 -1.5229 -1.3010 -1.5229 -1.3010 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229	-0.8239 -1.3010 -1.3010 -2.3010 -2.1549 -1.5229 -1.5229 -1.5229
NI,L06%	SZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZZ	
Z .		0000000000
NB.LOG%	-3.3010 -3.3010 -3.3010 -3.3010 -3.3010 -3.0010 -3.1549 0.00 N 0.00 N	549 549 549 549 549 500
NB.		
MN.106%	-2.3010 -2.1549 -2.1549 -2.1549 -2.3010 -1.8239 -2.3010 -1.8239 -2.3010 -1.549 -2.3010 -2.0000 -2.0000 -2.1549 -2.3010 -2.1549 -2.1549 -2.1549 -2.1549	-2.5229 -2.0000 -2.1549 -1.8239 -1.8239 -1.8239 -2.1549
LA, LOG\$	-2.0000 -2.3010 0.0 N 0.0 N 0.0 O 0.0 O 0.	-2.3010 -2.1549 -2.1549 -2.3010 -2.3010 -2.3010 -2.3010 -2.3010
CR.106%		2.5229 -2.5229 -2.5229 -2.5229 -2.5229 -2.5229 -2.5229
% 90	4 00000	
CO, LOG\$	0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0	
BE, LOG%	0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0	0.0 N 0.0 N 0.0 N 0.0 N 0.0 N 0.0 N 0.0 N 0.0 N 0.0 N 0.0 N
SAMPLE	153 154 155 156 157 158 159 160 165 160 203 203 203 203 203 203 213 215 215 215 216 227 228 228 229 229 220 220 220 220 220 220 220 220	233 234 235 236 237 238 241 242 243

TI, LOG%	-0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.6990 -1.0000	-0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229	-0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229	-0.5229 -0.6990 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229 -0.5229
CA, LOG\$	-0.6990 -0.8239 -0.6990 -0.0000 -0.6990 -1.1549	-1.0000 -0.8239 -0.8239 -0.3010 -0.3010 -0.529 -0.5090 -0.5090	-0.6990 -0.3010 -0.5229 -0.5229 -0.5229 -1.5229 -1.1549 -1.5229 -1.5229	-1.1549 -1.1549 -0.3010 -0.5229 -1.5229 -1.5229 -1.5229 -1.5229 -1.5229
MN, LOG\$	-0.1549 -0.1549 -0.1549 -0.1549 -0.3010 -0.1549 -1.0000		-0.3010 -0.1549 -0.3010 -0.1549 -0.3010 -0.1549 -0.5229 -0.5229 -0.5229	-0.5229 -1.8239 -0.1549 -0.3010 -0.3010 -1.3010 -0.1549 -0.3010 -0.3010
FE, LOG\$	0.4771 0.3010 0.1761 0.1761 -0.1549 -0.0000 0.1761 0.0 G	0.4761 0.8451 0.4771 -0.3010 0.3010 0.4771 0.3010 0.64771 0.3010	0.4771 0.4771 0.6990 0.6990 0.4771 0.3010 1.0000 0.6990 0.6990 0.6990 0.6990	0.4771 0.4771 0.4771 -0.3010 -0.5990 0.0 6 -0.5229 0.3010 0.8451 0.3010
SAMPLE	153 156 156 157 158 160	202 203 204 204 205 206 208 208 210 211	213 214 215 216 218 224 225 226 227 230	231 232 233 234 234 235 237 241 242 243

84,106%	-1.5229	-1.3010	-1.1549	-1.3010	-1.0000
MO,LOG%	N 0.0	-3.0000	-3.0000	-3.1549	N 0.0
2907'NZ	0.0	0.0	0.0	0.0	0.0
CU,LOG%	-3.3010	-2.8239	-2.5229	-2.3010	-2.8239
AS, LOG%	-3.0000	-2.2218	-2.0969	-3.0000	-3.0000
HG. LOG\$	-5.3979	-5.2218	-5.0000	-4.0969	-5.3979
81,106%	0.0	2 0.0	0.0	0.0	N 0.0
PB, LOG%	-3,5229	-2.3010	-2.3010	-2.0000	-3.1549
AG, LOG%	0°0	0.0	0.0	-4.0000	0°0
AU, LOGE	0.0	-3.9318	-5.0000	-5.0000	-5.0000
SAMPLE	244	245	247	248	253

					ARGILLIZED	ROCKS, LOG	PERCENT					
AMPLE	8E, LOG\$	£901.00	5 0	CR, L06%	LA, LOGZ	MN, LDG%	NB, LOG%	NI, LOG\$	89	SR, LOG\$	V, LOG2	Y,L06%
244	-3.8239	0.0	z	-2.5229	-2.3010	-2.3010	-3.0000	0.0	Z	-1.5229	-2.1549	0.0
245	0.0	0.0	Z	-2.3010	-2.0000	-2.1549	O.0	0.0	Z	-0.6990	-2.1549	-3.0000
247	0°0	0.0	z	-2.5229	-2.0000	-2.3010	0.0	0.0	z	-0.8239	-2.0000	0.0
248	-3.8239	0.0	z	-2.5229	-2.3010	-2.0000	-3.1549	0.0	z	-1.0000	-2.1549	-3.0000
253	N 0.0	0.0	z	-2.5229	N 0.0	-2-3010	-3.3010	0.0	z	-1.1549	-2.1549	0.0

TI,L06%	-0.6990 -0.6990 -0.6990 -0.5229
CA, LOG\$	-1.1549 -1.1549 -1.1549 -0.5229 -1.1549
MN,LOG%	-0.5229 -1.0000 -0.3010 -0.1549 -0.5229
FE, LOGZ	-1.1549 0.0 6 0.0 6 0.1761 0.3010
SAMPLE	244 245 247 248 253

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1.	עען	57	atistical Combi	nation -	January	ified roo	e sampl	es from	TRE		,
۷,	بر مصروت بعار	gentle ker	Carlo Carlo	10 10 10 10 10 10 10 10 10 10 10 10 10 1	Control of the second	2 4 19 19 60 CR	chi plan	ular (12), utanin nagarakan ka	nto difference non conservation of the second	· • • • • • • • • • • • • • • • • • • •	1 - F
				4060							
	Au	0.8	-3.87	0.57	-3.88	0.58	3.2				
	Ag	32	-3.51	0.45	-4.32	1.21	20				
	Pb	0.8	-1.73	0.45	-1.75	0.49	330				
>	Ei	65	1 .								
÷	Hg	0	- 4.86	0.30			0.18				
ลุโรษท	_		-2.49			+		THE RESERVE THE PROPERTY AND ADDRESS OF THE PERSON NAMED IN CO.	er valenda ja teksterioonen oleen eliveneen oleen olee		
V Dat	As	7.0	•	0.49	-2.55	0.52	58	2. 10. 20.10. 9.0000	an and the second of the secon	The special enterprises of the second	
₹	Cu	0	-2.78	0.42		<u> </u>	27				
MASS	Z,	72				-					
ADING	Mo	54	-3.23								
	Ba	0	-/.20	0.2/	_		700				<u> </u>
KINY. II	Ľe	88			<u> </u>						
46 COMP	_ Co	95									
ыы	Cr	2	-2.79	0.22	-2.82	0.26	18				
F Pu	_ La	64					_				
ON WEST	MA	0.8	-2.77	0.64	-2.78	0.66	52				
ADDISA	Nb	67	<u> </u>		_						
U.S.A	Ni	91									
. Q		0	-/.28	0.31		_	670		 		
GISAP	<u>V</u>	0	1	0.18	-	_	68		-		
		82	1	1							
	· <u></u> У	ļ									
	**************************************	1.6 (6)	-0.01	0.52	0.01	0.54	22000				
	Mg	7	-1.94	0.46	-2./0	0.73	320				
	Ca	0	-1.16	0.3/	<u> </u>		880		La de la También I Militario		-
	70	0	-0.63	0.14			2500		ļ.,		
	e e	<u>.</u>			!						
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		4		Mariana	الطشريا عاملا عاطا		<u> </u>	L	1 1	<u> </u>	

Appendix L

Statistical data, 149 argillized rock samples from the

Combination - January cuts

,) y'	54x+1		hada, 14	9 argilli January	red rock cuts	. Sa miple	s from)	TX E		•
K	y X	ما يو را	υ,		e s.		w.	* * * * * * * * * * * * * * * * * * *			
	X	No. le se	1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1		a ser of	6 10 10 00 00 00 00 00 00 00 00 00 00 00	Sick plan			grazani sarri sa sarri s	ra
	Ŵ,	40, 20, Xy	60, 50, 160	602 x 603 100	Con Con	C. L. L. C.	5° 9°				
	Au	40	-4.64	0,43	-4.95	0.55	0.25				
	Ag	58	-				-				
	Pb	11.	-2.8/	0.48	-2.95	0.60	29				
¥0.		98.7	_		-			. <u> </u>		Ma	ļ
James	Hg	/.3	-5.05	0.28	-5.06	0.29	0.11				
N Date	As	/3	-2.72	0.37	-2.95	0.45	19			er en	
ž.	Ca	0	-2.78	0.35	<u> </u>		23	an government after it we			
MG. MA	Z ^	47	-2.30	0.24	-2,58	0.37	38				
READ	Mo	26	-3.28	0.24	-3.4/	0.3/	5.1	1 NAT OF STREET	A CONTRACTOR OF THE PROPERTY O		
NY.	Ba Le	0	-1.12 -4.04	0.28	-4:3/	0.47	0.88				
COMPA	Ca	32 60		-	71.3/						
LISHING	C _r	0	-2.66	0.17			23				
EY PUBI	ta	/5	-2./8	0.14	-2.26	0.23	63				
N WESL	Mn	0	-2.02	0.46			170				
ADDISO	Nb	47	-3.22	0.15	-3.52	0.36	4:3				
1 A & D	Ni	65		_				a star of the star of the star of			ļ
LED IN	حرک	0	-1,28	0.36			730		A 1000000 10000000000000000000000000000		
HÖGRAF	· · · · · · · · · · · · · · · · · · ·	0	-2.11	0.13			82				
5	У	2.2	-2.98	0.25	-3.09	0.46	_/4	and the second s	ALL STREET COMMUNICATION		
	Fe	3.4(6)	0.34	0.39	0.37	0.41	37000				
	Mg	0	-0.44	0.47			6600				
	Ca	0	-0.87	0.39			2000				
		0	-0.51	0.14			3300				
			•								
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Tellurium and antimony maps are not included because replicate analyses indicated that sample preparation and extraction problems affected the atomic absorption and colorimetric analyses for these elements. Replicate determinations resulted in values from 10 to as much as 100 times higher than the first determinations. Possibly tellurium and antimony in these rocks are very finely divided and were not quantitatively liberated from the quartz matrix with the original grinding and extraction procedure. Unfortunately, the original splits for many samples were depleted before this problem was satisfactorily solved, so the limited newer data are not worth examining in detail. The data available suggest, however, that both tellurium and antimony are enriched in metallized silicified rocks relative to surrounding argillized rocks.

Maps were not prepared for tin and cadmium, the two remaining potential indicator elements (ignoring selenium), because only 17 of the 278 samples from the cuts had 3 ppm or more of tin, and none of the samples had as much as 50 ppm of cadmium.

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Gold

All but three high gold values (greater than or equal to 3 ppm) and most intermediate gold values (1 to 2.9 ppm) are in silicified rocks (pl. 4). The three high values in argillized rocks are all from fault or shear zones adjacent to silicified zones (see locs. 5, 6, and 18, pl. 4). Of the six intermediate values in argillized rocks, two are from fault zones between argillized and silicified rocks (loc. 9 and fault at loc. 6), and two more are within 5 feet of silicified zones, but are not associated with prominent structures (locs. 4 and 19). The remaining two intermediate values, at locality 3, are at least 20 to 30 feet southwest of the nearest silicified rock contact exposed in the wall of the cut or projected along the floor of the cut toward locality 4, and are not associated with structures. All remaining samples from argllized rocks have less than 1 ppm gold; many have less than O.1 ppm. (The detection threshold for gold for samples from the cuts is 0.1 ppm, whereas it is 0.02 ppm for the unaltered dacite samples and unoxidized silicified rocks discussed in the previous section).

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Gold

All but three high gold values (greater than or equal to 3 ppm) and most intermediate gold values (1 to 2.9 ppm) are in silicified rocks (pl. 4). The three high values in argillized rocks are all from fault or shear zones adjacent to silicified zones (see locs. 5, 6, and 18, pl. 4). Of the six intermediate values in argillized rocks, two are from fault zones between argillized and silicified rocks (loc. 9 and fault at loc. 6), and two more are within 5 feet of silicified zones, but are not associated with prominent structures (locs. 4 and 19). The remaining two intermediate values, at locality 3, are at least 20 to 30 feet southwest of the nearest silicified rock contact exposed in the wall of the cut or projected along the floor of the cut toward locality 4, and are not associated with structures. remaining samples from argllized rocks have less than 1 ppm gold; many have less than O.1 ppm. (The detection threshold for gold for samples from the cuts is 0.1 ppm, whereas it is 0.02 ppm for the unaltered dacite samples and unoxidized silicified rocks discussed in the previous section).

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1 The histograms reflect the substantially different amounts of gold in silicified versus argillized rocks. The geometric mean for 2 silicified rocks, 1.3 ppm, is more than 10 times that determined for argillized rocks, 0.11 ppm. The frequency distribution for silicified rocks is relatively symmetrical, with a rather broad peak between 5 ----4.00 and -3.50 log percent (between approx) 1 and 3 ppm). If ore samples had been available from the 12 locations at which old stopes 7 8 intersect the walls of the cuts, the upper tail of the histogram, 9 representing values of 10 ppm or more, would presumably be somewhat larger. A large number of values appear at the left end of each gold 10-11 histogram, just above the detection threshold. Most of the readings 12 that contribute to these large frequencies in the class interval 13 immediately above the detection threshold are readings of "0.1 ppm." 14 In fact, 44 of the 278 samples were read as 0.1 ppm, whereas only 17 samples were read between 0.1 and 0.2 ppm. This suggests that 15-16 analytical discrimination was poor for samples near the detection 17 threshold. Consequently, both readings of "0.1 ppm" and "less than 0.1 ppm" are given the same symbol on the map. 18 19 20-21 22

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Although silicification is by far the most important feature associated with relatively high gold values, the silicified zones do not show uniformly high values. The west edge of the vein system, exposed between localities 7 and 10 and on the wall to the west and north of the January shaft, shows the lowest values. The small silicified bodies at localities 14 and 15 also show low values. Along the wall between localities 2 and 5, even though the overall gold tenor is relatively high, many adjacent samples have substantially different amounts of gold. Moving southeastward along the vein system, individual exposures are unpredictable: relatively high values appear in the vicinity of locality 12, low values at locality 16, and high values again east of locality 19. Localities 18 and 20 are each represented by only one sample; more information would be needed to categorize these outcrops. The geochemical profile, which includes locality 12, shows nearly the maximum anomaly contrast that one would expect to see between metallized silicified rocks and argillized rocks, unless the sampling encountered ore-grade material. The geochemical profile also shows a sharp break within the large silicified zone at locality 12, produced by an isolated very low reading. Variations in values for the six samples between localities 6 and 7 provide another example. Irregularities like this apparently must be expected; such irregularities are a potential source of error when determining the average gold content of the silicified zones.

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Geologic processes must account for most of the abrupt variations in gold values, producing bunches of gold at the scale of an outcrop and even within a hand specimen. Before the oxidized-rock samples were processed for analysis, we removed and saved a 0.5- to 1-kg hand specimen from each sample. These hand specimens may be considered replicate samples from each of the localities shown on plates 3-3.

We obtained gold analyses for the replicate samples from 88 of the 129 oxidized silicified rock localities. We compared the original sample and replicate sample data for 65 of these localities, which showed 1 ppm gold or more in both the original analyses and the replicate analyses; analytical error is substantially lower above 1 ppm than below 1 ppm.

Using Garret's (1969) method for comparing combined sampling and analytical error to overall data variability by attilizing an F-test,

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SA (combined sampling and analytical variance) is 0.058, C_D^2 (data variance) is 0.11, and the value of F calculated is 1.90. Tables of the F distribution give 3.92 for degrees of freedom 1,128 at the 95-percent-confidence level (1.34 at the 75-percent-confidence level). Since our calculated F does not exceed 3.92, we must conclude that

The terms "sampling variance" and "sampling error" refer to

variation introduced into the data by samples or analytical portions of

samples that are not truly representative of the localities from which

error," on the other hand, refer to variation introduced into the data

they were taken. The terms "analytical variance" and "analytical

by imprecision in the laboratory analytical procedure.

chances are greater than one in 20 (but less than one in four) that combined sampling and analytical variance is responsible for the 2 variability seen in samples containing more than 1 ppm gold. Sampling error related to removal of material from an outcrop or face is therefore an appreciable problem. 10-11 12 13 14 15-16 17 18 19 20-21 22 23 24

Sampling error related to removal of an analytical portion from a crushed, ground, and mixed sample may contribute to variations in gold values also. Gold for atomic absorption analysis was extracted from only 2 grams of sample, an analytical portion very much smaller than the field sample, which was 2 to 3 kg. Gold was separated from approximately one-half-kilogram portions of the three samples having the highest gold values: the largest gold particle recovered was a thick flake 0.07 mm in diameter (W. J. Keith, unpub. data). Applicable here is the diagram prepared by Clifton, Hunter, Swanson, and Phillips (1969, p. C8) relating gold particle mass (and particle dismeter, for both spheres and flakes) to size of analytical portion expected to contain 20 gold particles for samples of various true grades.

If a sample contains 20 particles of gold, it is 95 percent probable that the true gold value will be within a range of values from approximately 50 percent more to approximately 50 percent less than the value obtained by chemical analysis (Clifton and others, 1969).

diagram shows that sampling error is acceptable for 0.07 mm flakes only if the true grade is larger than about 10 ppm. Gold flakes not quite twice as large (0.125 mm) or gold spheres of about the same diameter (0.062 mm) would produce sampling error for rocks having a true grade less than about 30 ppm, which is nearly the same as the highest reading (29 ppm) obtained in this study. Most of the gold is probably smaller than 0.07 mm and most 2-gram analytical portions probably contain more than 20 particles of gold, but it is likely that a relatively few

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larger gold particles significantly affect the readings for some samples. The replicate samples described above and unaltered dacite, unoxidized silicified dacite, and unoxidized average-grade ore samples were analysed using a 10-gram analytical portion. With an analytical portion of this size 0.062-mm gold spheres (or 0.125-mm gold flakes) 5 -would produce sampling error at true grades below about 6 ppm, and 0.062-mm gold flakes would produce sampling error only at true grades below about 0.8 ppm. We can expect, then, that in silicified rocks with enough gold to be of economic interest, sampling error due to removal of an analytical portion from a crushed, ground, and mixed sample 10should be a minor problem when using 10-gram or larger analytical 11 portions. 12 13 14 15-16 17 18 19 20-21 22 23 24

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The results described above indicate that any program meant to determine grades and tonnages of the low-tenor silicified rocks in the Combination-January area must be designed with caution. Gold analyses by atomic absorption probably should be done using 10-gram analytical portions. Fire assay, utilizing 29.167 grams of sample, would also 5 -be appropriate. Since the main objective of this report is delineating indicator elements for gold, some error related to field sampling and removal of analytical portions is not objectionable, but for a more precise determination of ore grade, more careful field sampling and analysis would be necessary. The replicate sample results indicate, 10however, that in this study we cannot attach a great deal of significance to the exact gold value determined for any given sample.

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With the foregoing discussion in mind, it is of interest to estimate the grade of unmined silicified rock exposed in the cuts, using the data at hand. The geometric means calculated for various elements in the several data sets discussed in this report are suitable for comparison with one another, but Sichel's t provides a better estimate of true abundance than the geometric mean (see introduction, p. 2). For the silicified rocks of the cuts, t is 3.0 ppm. With 95-percent confidence, the true abundance should be between 2.4 and 3.8 ppm (confidence interval calculated by method of Aitchison and Brown, 1963, p. 50). One part per million gold is equivalent to 0.0291667 ounce troy \$1 per short ton; specifically, 3 ppm is equivalent to 0.088 ounce per ton and the corresponding confidence interval is 0.070 to 0.11 ounce per ton. These amounts of gold, although subeconomic, are large enough to commend further exploration and evaluation of the remaining vein material.

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Lead

more) and intermediate lead values (70 to 150 ppm), whereas argillized rock samples show only 5 high values and 10 intermediate values (pl. 5). Some of the intermediate values in argillized rocks are adjacent to silicified zones, as at locality 12 and near localities 10 and 18. An intermediate value appears in a shear zone adjacent to silicified rock at locality 15, and a high value appears in a fault zone adjacent to silicified rock at locality 16. Four of the five high values for argillized rocks are grouped together north of locality 16. This group of relatively high-lead samples has no distinctive structural or alteration feature associated with it.

The anomaly contrast shown by the profile is representative of what one might expect along any traverse across the strike of the silicified zones in the vicinity of the cuts. The anomaly contrast is as strong as that for gold itself, and in some possible profiles it might well be stronger.

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Consistently high lead values appear within the silicified zones 1 in the areas characterized by high gold values. Between localities X and 12, scattered high lead values accompany intermediate gold values. Other smaller silicified outcrops all show some degree of consistency 5- between gold and lead values, but it would be difficult to predict gold tenor from lead readings. The histogram for silicified rocks (pl. \$) shows large frequencies in four class intervals above the geometric mean: 200, 300, 500, and 700 ppm, so the frequency distribution looks skewed toward the higher values, with a very small upper tail above 700 The frequency distribution for gold, by comparison, shows a long, relatively smoothly, declining upper tail above the geometric mean. 11 12 Thus even though an area may show some very high gold values, along with 13 some high and intermediate values, almost all the lead readings are between 200 and 700 ppm. Lead values from silicified rocks therefore 14 give no idea about the highest gold values to be anticipated, but on the other hand, lead values of 200 ppm or larger indicate that gold values 16 17 of at least 1 ppm will be found nearby (locality 16 is an exception). A different method of chemical analysis offering greater analytical 18 19 precision for rocks with true lead values between 100 and 1,000 ppm might improve the correlation between gold and lead for individual 21 samples, but geologic factors, rather than analytical precision, may 22 account for the limited range of variation for lead in relatively goldrich samples (see interpretation section, p.7/-11/ 23 "ceiling" on lead values has an advantage for geochemical exploration, 24 in that lead might well show broader, less ragged anomalies over potential gold ore bodies than would gold itself.

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Silver

All high silver values (greater than or equal to 10 ppm) appear in silicified rocks (pl. &). All intermediate values (3 to 7 ppm) but one are also in silicified rocks; the single exception appears at locality 17 in argillized rocks associated with a fault zone. Of the (8) low values (between 1 and 2 ppm) in argillized rocks, (3) are adjacent to silicified zones (localities, 12, 13, and 19), and one other is in a fault zone (southeast of locality). Almost all the argillized samples and 48 of the silicified rock samples (37 percent) bear less than 1 ppm silver. Obviously, no notable silver halo exists in argillized rocks. The geochemical profile shows anomaly contrast typical of what might be expected many places in the cuts. One can anticipate that profiles run northeast of the January shaft, however, might produce anomalies twice as strong as those shown on [the] profile, gold and silver patterns in the silicified zones are grossly similar, even though larger areas within the zones show relatively low silver values.

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The semiquantitative spectrographic data available for silver s not precise enough to allow a good determination of silver tenor in the silicified rocks. Even an approximate estimate is not possible because the detection threshold for 26 of the silicified rock samples (20 percent) was 1 ppm, due to spectral interference, rather than the usual 0.1 ppm. In calculating the geometric mean for the silicified rocks by Cohen's method, the lower of these two detection limits, 0.1 ppm was used, possibly biasing the result toward low values. Furthermore, the histogram for silver in silicified rocks is very irregular; possibly more data with higher analytical precision would show a polymodal frequency distribution. The resulting geometric mean shown with the histogram on plate 6 has such a large standard deviation associated with it that it has little meaning. The data does suggest, however, that fire assays of samples taken northeast of the January shaft should show some silver values as high as 1 set troy per short ton. Delite

ounce part per million by weight is equal to 0.0291667 as troy per short ton.

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Bismuth

	by direct	
2	Only two, argillized rock samples yielded detectable amounts of	
3	bismuth (5 ppm). Forty-five silicified samples show readings ranging	موسف ال ال
4	from 5 to 200 ppm, but the majority of silicified samples (65 percent)	
5-	also have less than 5 ppm bismuth. Clearly, the 5-ppm detection	
6	threshold provided by the semiquantitative spectrographic technique	
7	employed is too high to provide enough data for many conclusions about	
8	geochemical relationships. Bismuth cannot be fully evaluated as an	*
9	indicator element for gold, but it obviously shows a strong preference	
10-	for silicified zones, and most of the intermediate and high bismuth	
11	values appear northeast of the January mine shaft, in an area characteri	zed
12	by intermediate and high gold values (pl. X).	
13	Only three readings above the detection threshold appear along	
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Only three readings above the detection threshold appear along
the geochemical profile line, so the profile for bismuth is not included
with plate v.

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Mercury

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builisel The silicified rocks of the cuts contain more mercury than do the argillized rocks. The averages for the two data sets are not greatly different (see histograms on pl. 8), but they are significant statistically, at the 99-percent-confidence level. Plate & bears this difference out for most of the area of the cuts: silicified rocks along the south and west sides of the largest cut, continuing into the area northeast of the January shaft, have many intermediate (0.11 to 0.30 ppm) and high (greater than 0.30 ppm) mercury values. Throughout the southeastern cut (locality 16 to locality 3), silicified and argillized rocks are similar. The silicified rock east of locality 10 has intermediate to high gold values, but it is expressed no more distinctly on the mercury map than is locality 16. On the other hand, the area northeast of the January shaft with high gold values certainly has the most mercury. The consistently intermediate to high mercury values south of locality & and between localities & and liffurthermore. suggest that mercury halos may extend 30 to 35 feet into argillized rocks from silicified zone contacts in gold-bearing areas. The width of this possible halo, however, is small relative to the width of the adjacent silicified zone: about 90 feet at locality \mathring{g} . Also, the geochemical profile reveals that the anomaly contrast is very low. The histograms jointly show that the total range of mercury values is barely two orders of magnitude, compared to at least three, and probably four orders of magnitude for gold, so one cannot expect mercury anomalies to be very strong. Under these circumstances, relatively narrow halos extending into argillized rocks are not particularly valuable for reconnaissance sampling. They might be of minor value for a detailed sampling program in a small area.

Arsenic

The Gutzeit colorimetric method used for arsenic determination has a reporting interval that becomes larger with increasing values, but is arithmetic within certain ranges of values, & Values reported for the oxidized rocks of plate \(\) are: less than 10, 10, 20, 40, 60, 80, 100, 120, 140, 160, 180, 200, and 250 ppm. Corresponding log percent figures are: less than -3.000, -3.000, -2.699, -2.398, -2.222 -2.097, -2.000, -1.921, -1.854, -1.796, -1.744, -1.699, and -1.602. Within this range of values the interval is arithmetic except for the upper and lower ends of the range. The histograms (pl. $\tilde{\chi}$) are based 10on the Logarithms of the reported values, however, to avoid overemphasizing values of 200 ppm or more, and to make the histograms more easily comparable with those on the other plates. The resulting histograms, with class intervals determined by Sturges' rule, have null class intervals below -2.500 log percent. Other class intervals, determined arbitrarily, will not eliminate the problem unless one reduces the number of class intervals to three or four, and this, in turn, eliminates variations shown in the upper parts of the histograms Such treatment also subdues the fact that a disproportionate number of values for both silicified and argillized rocks were read as 10 ppm; 10 ppm was reported for 49 of the 129 silicified rocks (38 percent), and 74 of the 149 argillized rocks (49.7 percent), whereas less than 10 ppm was reported for only 7.0 percent of the silicified rocks and 12.8 percent of the argillized rocks. Analytical discrimination was apparently poor near the detection threshold for

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61a (p. 616 follows)

the Gutzeit colorimetric test. Readings of less than 10 ppm and 10 ppm are therefore given the same symbol on plate. This same problem arose with gold readings near the gold detection threshold of 0.1 ppm and was treated similarly. 5 -10-11 12 13 14 15-16 17 18 19 20-21 22 23 24 25-

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6/16 (p. 6/2 follows)

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The geometric means indicate that oxidized silicified rocks have about twice as much arsenic, overall, as do oxidized argillized rocks. Since the problems with reporting interval and analytical discrimination produce discontinuities among the smaller data values, as shown by the histograms, the validity of these geometric means is questionable. The percentages of samples read as "less than 10," "10," "20," and "40" ppm, however, are generally similar for the two data sets, so improved analyses would likely have similar effects on both frequency distributions and both geometric means. The upper parts of the two frequency distributions, clearly different, probably would not be substantially changed. Both geometric means would likely be decreased with better analyses, but the difference between them would probably not be changed much.

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Although oxidized silicified rocks contain more arsenic than oxidized argillized rocks, as is the case for the preceding elements, the relationship between arsenic and gold is not consistent (compare pl. 5 with pl. 4). Intermediate (60 to 80 ppm) and high (100 ppm or more) arsenic values accompany moderate to high gold values for most, but by no means all, samples northeast of the January shaft. Arsenic values are relatively lower than gold values for the area between localities 6 and 7, relatively higher than gold values for the area between localities 8 and 10, relatively lower in the vicinity of locality 12, and relatively higher at locality 16. The other silicified rock outcrops show a few samples with concordant gold and arsenic values, but these are but a small percentage of the silicified rock samples.

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Intermediate and high arsenic values appear at many localities in argillized rocks. Some are associated with fault zones and shear zones, but many are not. Halos around silicified zones do not exist. The geochemical profile shown is typical for argillized rock areas, but gold-bearing silicified rocks could show greater arsenic anomaly contrast in profiles taken at some other locations. With the small range of variation in the data, however, even a profile taken in the northwestern part of the cuts would show only modest anomaly contrast. For geochemical exploration, arsenic is not reliable as an indicator for gold.

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Copper

except for a group of high-gold samples in the vicinity of locality that also show high copper values. The histograms for silicified rocks and argillized rocks are much the same, and the two data sets have nearly the same range of values. The geometric means are not significantly different statistically at the 95-percent-confidence level and the geometric standard deviations are not much different. Throughout the map area, fault zones and shear zones are not particularly favorable to high copper values relative to less broken ground.

Obviously copper cannot be used as an indicator element for gold in the silicified rocks. The geochemical profile (pl. 20) bears this out. The question arises whether some areas of argillized rock with intermediate and high copper values, such as those between localities and 15 and between localities 16 and 17 represent a halo. One or more longer profiles across strike would be helpful in answering this question, but the low values between localities 9 and 11 and the lack of a systematic decrease in values from locality 15 to locality 19 indicate that a halo is not consistently developed.

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Molybdenum

Most silicified outcrops show only a few scattered intermediate mol/bolenum and high values with no consistent relationship to gold values. Some intermediate (7 to 10 ppm) and high (15 ppm or more) molybdenum values are scattered throughout the high-gold area northeast of the January shaft (pl. 21), The overall tenor of molybdenum is higher in oxidized argillized rocks than it is in oxidized silicified rocks. The intermediate and high molybdenum values in argillized rocks are also scattered, but about half of these values are associated with faults and shear zones (west of loc. 2, at loc. 5, near loc. 6, north of loc. 13, near loc. 14, and south of loc. 16). No coherent halo is developed in argillized rocks. The geochemical profile demonstrates that molybdenum values cannot be predicted on the basis of alteration and structural criteria.

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Except for a few scattered intermediate (50 to 90 ppm) and high (100 ppm or more) values, zinc is relatively scarce in the silicified rocks (pl. 12). Intermediate zinc values appear near silicified zones between localities 12 and 17, 16 and 17, and at locality 14, suggesting that zinc may form a geochemical halo. As is the case with coper, however, other areas near gold-bearing silicified rocks, particularly between localities 15 and 19 and between localities 2 and 14 are not enriched in zinc, so again a halo is not consistently developed. group of intermediate and high values found along the cut wall east of locality 14 and 100 feet northeast of locality 16 show no systematic relationship to faults or shear zones.

Zinc

The geochemical profile for zinc is not particularly informative, Trongwill be considered in the next section of the report.

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Interpretation of the geochemical maps:

Separating effects of hypogene and supergene processes

General

The distribution of gold and other ore-related metals in the Combination-January cuts is the result of two processes: hydrothermal alteration culminating in metallization, and oxidation. Here metallization is a primary (and epigenetic) geochemical dispersion process, whereas oxidation and accompanying dissolution and redeposition of the epigenetic elements constitute a secondary geochemical dispersion process. The following section describes the evidence for and the degree to which the map pattern for each element is the result of secondary rather than primary dispersion processes.

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Mobility of elements in the zone of oxidation The Goldfield hydrothermally altered rocks, both silicified and 2 argillized, have ubiquitous pyrite (Ransome, 1909, p. 113-114; and our observations), which makes strong supergene alteration possible. the prevailing arid climate, surface water descends to the water table, reacting with pyrite above the water table to yield Fe2+, Fe3+, HSO,, and SO, . Usually most of the iron reacts further to form limonite minerals (goethite, FeO(OH), and hematite, Fe₂O₃). Where ore is exposed to meteoric waters, other sulfides and sulfosalts are also oxidized to $^{10-}$ yield metal cations and sulfate ion (SO $_{\mu}$). Studies on mine waters (Baas Becking and others, 1960; Sato, 1960) and on experimental oxidation of iron and manganese (Sato, 1960) indicate that Eh values in oxidizing 13 sulfide ore deposits may be as high as +0.6 to +0.86 volt with low 14 accompanying pH values of 2 to 3. Any metal whose compounds (hydroxide, carbonate, chlorize, and particularly sulfate) are soluble in these 16 -Garrels and Christ (1965) consider a species with an activity of 17 10^{-6} m or greater soluble in the framework of geologic processes. 18 19 relatively high-Eh and low-pH aqueous solutions will be more or less mobile in this environment. Highly mobile (soluble) elements, such as copper, may be thoroughly leached, destroying their primary dispersion 22 patterns. Metals having intermediate mobility, such as molybdenum, are 23 transported only short distances and precipitated, some along with limonite. Some metal cations are easily reduced to the native metal and precipitated

(mercury), or form stable compounds that remain essentially in place

1 (lead). The latter elements should reveal relict primary dispersion 2 patterns. Copper, and silver to a lesser extent, may be redeposited below the water table to form a supergene enrichment zone. The samples from the cuts at Goldfield, however, lie approximately in a plane 5 -parallel to the ground surface and within the upper part of the oxidized zone, so supergene enrichment effects can be ruled out. Ransome (1909, p. 170-174), in describing Goldfield oxidized ores and changes in the ores with depth, does not mention supergene enrichment at or near the water table, but we have found indications that supergene enrichment occurs at least locally (p. 92). Leaching effects should 10-11 be essentially the same in rocks of the cuts as at the ground surface. 12 13 14 15-16 17 18 19 20-21 22 23 24

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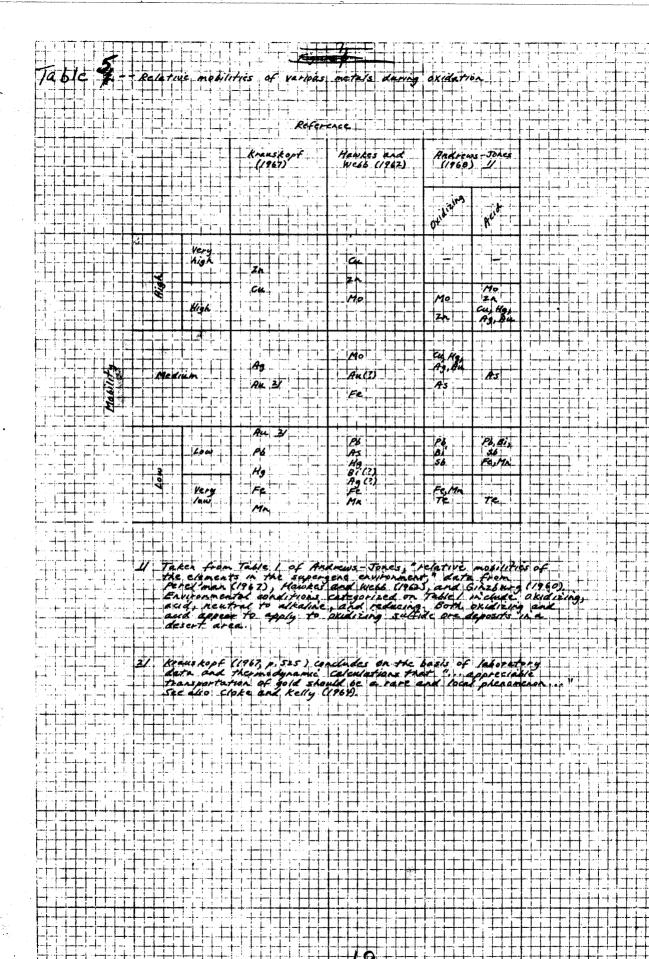
A number of authors give relative mobilities for many of the Table 5 metals being considered here. Figure 7 summarizes information given

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by three authors, who have in turn assembled data from other sources. The figure shows that we may expect zinc to be highly mobile, copper and molybdenum to be moderately to highly mobile, silver and gold to be moderately mobile, arsenic to be immobile to moderately mobile, and lead, bismuth, antimony, and tellurium to be immobile. Krauskopf (1967) and Hawkes and Webb (1962) agree that Hg should show low mobility in solution, but Hawkes and Webb indicate that Hg may be very mobile in the vapor phase.

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During oxidation, groundwater passing through rocks bearing pyrite and other sulfide minerals dissolves and carries away mobile elements, accomplishing leaching. Where such ground water encounters mine workings, it evaporates, leaving coatings and crusts of hydrated sulfates which incorporate the metals actively being leached (see for example Lovering, in Morris and Lovering, 1952). Such coatings are common on the walls of mine workings beneath oxidizing sulfide ore bodies, and hydrated iron sulfates are usually the most abundant constituents of the coatings; these hydrated iron sulfates, formed from iron released by the oxidation of pyrite, are intermediate products that are eventually converted to hematite and goethite (Blanchard, 1968, p. 51-55). Eight samples of secondary hydrated sulfates from walls of workings in the Florence mine (table 6 and fig. 2) contain relatively large amounts of copper and zinc, lesser but

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notable amounts of bismuth and molybdenum, and cobalt, nickel, and manganese as well. None of the seven hydrated iron sulfate samples contain as much as 1 ppm silver, but the gypsum sample contains 3 ppm silver. None of the samples has as much as 7 ppm lead. The detection thresholds for gold (15 ppm), arsenic (200 ppm), antimony (100 ppm), and tellurium (1,000 ppm) are too high to allow significant amounts to be recognized, except for antimony in gypsum (700 ppm). All the sulfates sampled except gypsum are readily soluble in water, so they were separated and cleaned by hand picking; consequently not enough material was evailable for extensive chemical testing.

	Sample	Mineralogy	Ag	Bi	Cu	Mo	2n	
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	F-250-8	Halotrickite (Fe, Mg) Alz (504) 4.22H20	/v_(/)	1 .	15000	30	7300	
- -	F-350-1A	Halotrickite	N (I)		5000	15	700	
	E-360-12	Mala a			A COLUMN CONTRACTOR CO		ļ	
	F-350-18	Melanterite Fesou.7H20	N (I)		700		, 556	
	F-350-4A		N(I)	20		10		
		(Fe, Mg) Fe4 (504) (04)2.			Conference of the Conference o			
	F-350-48		N(i)	50	2000	10	1000	
		coquimbite mixture Fez (504) 3 9H20						
	F-350-8B	Halotrichite-	N(I)	/0	50	N(2)	3000	
	1	rozenite mixture Fesoy 4H20						
	F-350-9C	ROZENITE FESO4.4H20	N(I)	N(7)	200	7	2000	
	1			+				
-	F-350-3A	Gypsum Casoy-2H20	3	70 .	100	N(2)	N((00)	
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	N= not All Other	detected at limit figures in ppm. D	ata for	other one	-related e	lements a	ere: Au,	NCIS)
	N= not All Other	les; Pb, NCT) for all	samples;	As, N(20	o) for all	samples;	Ed, N(50)	N(15) for a
	N= not All Other all samples; Te. N(1000	sigures in ppm. D les; Pb, N(1) for all Sb, N(100) for all so for all samples.	ata for samples; exples exples	other ore As, N(20 cept 700 f	-related e b) for all for F-350- resent in	lements a samples; 3A; Sn, N notable a	Cd, N(50) (17) for all	for all samp
	N= not All Other All samples; Te, N(1000 Mr. (3030 except for	signes in ppm. D les; Pb, N(7) for all Sb, N(100) for all so o) for all samples. o), Except for F-350 F-350-39 (N(1))	ata for samples; exples exples	other one As, N(20 cept 700 f ements p Ca (50-100	related e o) for all or F-350- resent in o), except	samples; 3A; Sn, N notable a for F-350	Cd, N(50) (C1) for all acquets in -3A, EN(2) >,	for all samp
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1 The workings in the Florence mine generally follow a north-trending 2 silicified zone which forms the southern extension of the Combination vein system, as previously mentioned. The silicified zone and ore shoots 3 within it dip steeply, and the Florence shaft passes close to the silicified zone from the surface all the way down to the 350-foot (7th) level. Silicified rock is oxidized to the 100-foot (2d) level, but little oxidation appears on the 150-foot (3d) level and none below it (Ransome, 1909, p. 227). A small flow of water has entered the 350-foot level since at least 1908 (Ransome, 1909, p. 230). Sample localities F-350-3 and F-350-4 are crusts from the only part of the mine that was 10--11 wet at the time of sampling. The sulfates at localities F-350-8 and 12 F-350-9 form incrustations also; this part of the mine was probably 13 wet at some time in the past. Localities F-250-8 and F-350-1 yielded 14 fluffy efflorescences typical of most of the mine; these apparently form in dry parts of the mine as small quantities of ground water reach 15the walls and evaporate. 16 17 18 19 20-21 22 23 24

The sulfate sample localities are related to known ore bodies as follows. Sample F-250-8 lies between 50 and 100 feet beneath and to the east of the stope developed in the Sweeney lease, one of the largest single ore bodies in the district (Ransome, 1909, p. 154, 225-228; Newmont Mining Corporation, unpublished maps). Localities F-350-3 and F-350-4 lie about 120 to 200 feet from the Sweeney stope, being a few tens of feet to the south and east of a vertical projection of the stope to the 350-foot level. Another smaller stope lies approximately 60 to 100 feet above and immediately to the south of these two sample localities. A stope approximately 250 feet above the localities was entirely of oxidized ore (Ransome, 1909, p. 228). Locality F-350-1 is from the wall of a stope along a fault that may continue to the northeast into the Engineers' lease (Ransome, 1909, pl. XXXV, p. 232). Localities F-350-8B and F-350-9C lie between 120 and 230 feet below and to the west of the Sweeney stope. No ore bodies are known to have been mined from the ground immediately above these samples.

Most of the mine workings from which the sulfates came were made between 1903 and 1905 (Ransome, 1909, p. 225-226 and plate XXXV; M. C. Duffy, oral commun., 1966). The tunnel from which samples F-350-8B and F-350-9C came (fig. 3) is a crosscut driven about 1919 (Searles, 1948, p. 20). Thus the sulfate coatings, which were sampled in September 1966, are the result of 47 years to over 60 years of supergene leaching of low-grade unmined silicified rocks, and possibly unknown small bodies of high-grade ore.

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Comparing amounts of elements shown in table 6 with amounts shown in figure 6, the potential indicator elements for gold can be qualitatively arranged in order of decreasing mobility as follows: zinc, molybdenum, copper, bismuth, silver, lead. Information is inadequate or lacking for gold, arsenic, and mercury. This arrangement must be approximate because geochemical information is lacking for average-grade ores taken from the Florence mine. Silver was unusually scarce in ores from the Florence (Ransome, 1909, p. 230), so silver

Secondary sulfates such as those analyzed here were used by
Lovering (in Morris and Lovering, 1952) to determine supergene
mobilities of gold, silver, lead, copper, and zinc in the Tintic
district, Utah. His results generally agree with ours, even though
his samples contained larger amounts of these metals, and he
considered differences in mobility related to several different wall
rock environments.

might be more mobile than indicated.

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Iron and managanese

Iron and manganese are included in table 5 because they are both easily oxidized and precipitated as hematitie, goethite, and various manganese oxides and hydroxides. Several other elements, among them arsenic, copper, molybdenum, and zinc, may be precipitated with these minerals or adsorbed by them (Hawkes and Webb, 1962, p. 162-177). Since these four elements are potential indicator elements for gold, it is appropriate here to discuss the distribution of iron and manganese in the rocks of the Combination-January cuts.

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Iron does not seem to show consistent preferences for any lithology 2 or structural environment. Many high and very high iron values are associated with fault zones, but the majority of such values are not The most important thing to note is that unaltered dacite, unoxidized silicified dacite, and unoxidized average-grade ores contain 4.3, 4.7, and 6.8 percent iron, respectively (geometric means, see fig. 6 and tables 1 and 2), but the oxidized silicified dacite and oxidized argillized dacite of the cuts average only 1.0 and 2.3 percent iron, respectively (see histograms, pl.) and a 10- column, tables 3 and 4). This means that even though hematite and 11 limonite stain many of the rocks of the cuts conspicuously, much 12 iron has been removed. Figure 6 shows that essentially all iron is 13 retained during hydrothermal alteration and ore deposition except 14 in the high-grade ores, which occurred in relatively small volumes; 15-the iron removed from the cuts, therefore, must have been leached 16 during oxidation. In the few unoxidized argillized rocks we have 17 examined from the Florence mine and elsewhere in the Goldfield altered 18 area, pyrite mostly replaces former mafic minerals, just as it does in unoxidized silicified rocks. Before oxidation, therefore, the 20-amounts of iron in silicified and argillized rocks must have been similar, so that the oxidized silicified rocks were more strongly 22 leached than the oxidized argillized rocks. The most likely explanation for this is that numerous small postalteration fractures remain open in the oxidized silicified rocks, whereas in oxidized argillized rocks they 25 are squeezed shut. These small fractures have diverse orientations and

are easily visible in the silicified rocks along the cut walls. 1 2 Locally they are so abundant that from a distance the rock looks brecciated. The same small fractures in argillized rocks, on the other hand, are much tighter and can only be seen readily on freshly excavated 5- surfaces. The histograms show that both oxidized silicified and argillized rocks have a wide range of iron values. Probably differences in porosity, permeability, and other properties affecting movement of ground water and reactivity of ground-water solutions would have to be considered to explain details of the iron distribution map. The present petrographic division of the data, in spite of the obvious differences 10in physical properties between silicified and argillized rocks, is 12 -An exception is the advanced argillic rocks (12 samples), included 13 14 with silicified rocks because of their alunite-bearing mineral assemblages. These rocks appear physically similar to argillized rocks (see p. 36). 17 inadequate for evaluating these factors. Clearly iron, even though it 18 is quite immobile once it forms limonite, has undergone considerable 19 redistribution during oxidation at Goldfield. It will be useful to compare the iron map with the maps for potential indicator elements 21 to assess the extent of supergene dispersion of each element. 22 23 24 25

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Manganese should behave in the same way iron does during
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    pxidation (see table 5). Manganese, like iron, is depleted in oxidized
                                           Sichel's <u>t</u> 52 ppm
    silicified rocks (geometric mean 17 ppm) relative to oxidized argillized
                         , Sichel's I 170 ppm
    rocks (geometric mean 96 ppm) (see table 4), but manganese, unlike iron,
 5- is strongly and progressively depleted during hydrothermal alteration
    and ore deposition (see fig. 6 and discussion, p. 28).
    likely that manganese was removed from both silicified and argillized
    rocks of the cuts during hydrothermal alteration, and probably more
    was removed at that time from silicified rocks. The secondary sulfates
 10-previously described show 100 to 300 ppm manganese, which indicates
    some movement during oxidation. Manganese therefore was probably
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    removed during both hypogene and supergene alteration so that it is of
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    little use for determining whether other elements have undergone
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    supergene redistribution; consequently a manganese distribution map
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 15-is not included, and manganese will not be considered further.
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Correlation matrices

Finally, before discussing individual potential indicator elements, it is convenient to introduce correlation matrices for these elements and iron (fig. //). A correlation matrix facilitates making comparisons

FIGURE 7 NEAR HERE

between the various elements and visualizing the results. statistical methods used to generate each matrix were discussed on further discussion of the meaning of numbers given in the matrices follows. The upper triangle of each matrix gives the correlation coefficients. The lower triangle gives the number of data pairs used for calculating each coefficient; where either or both date values of a data pair were outside the detection limits, that data pair was rejected. The maximum correlation coefficient is 1.0, corresponding to ferfect corresponding to perfect correlation, in which case ranked lists of samples for each of the two elements being compared would be identical. A correlation coefficient of zero indicates complete absence of correlation, and a coefficient of -1.0 indicates perfect negative correlation, in which case ranked lists would be exactly reversed. By no means are all the positive correlation coefficients large enough to constitute significant positive correlations, or are all the negative correlation coefficients close enough to -1 to constitute significant negative correlations. As the number of data pairs

increases, however, the smallest coefficient that may be considered

significant at a given statistical confidence level becomes progressively

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Figure 1 a participled and argulished rocks from the Conditation to January orea Speanman Correlation coefficients for silicitied rocks An Po As 01 11, Hg As Cu MO ZALL Aca -0.14 0.18 0.4 0.46 0.06 0.21 -0.34 0,02 -----0.11 -0.25 6.0 2:14 -0.16 0.03 0.27 0.21. 1 0.5 Rg 0.77 0.09 0.50 0.29 -0.16 -0.06 0.24 0.19 +0.02 0.0 -0.20 -0.06 0.15 +0.05 0.16 0.27 00.5 43 0.7.2 0.09 2,46. 722 128 040 ----37 2.4. 24 60 6.12 0.17 36 136 Correlation coefficient
Significant At 757a Confidence Rve 997a Confidence level

e :	sure 1/2 continued	
Jpearma Jpearma	an Cornelation Coefficients for angillized hocks	
	(149 samples)	
Au Ph	Rg Bi Hg Rs Cu Ma 2n Fe	
Au 0.22	0.27 - 0.24 3.17 +0.19 0.12 +0.25 0.02	
76	0.40 - 0.12 0.19 0.04 0.30 -0.23 0.12	
-Ag 36 61	- 0,23 0,17 +006 -030 0,05 0,22	
2	0	
Hg 87 /30	61 2 0.74 -0.20 -0.15 -0.17 0.03	
As 23 //4	48 2 112 070 013 -0.05 0.27	
Ga 89 /32	62 2 147 180 0.23 0.17 0.42	
Ma 69 199	43 2 769 764 710 +0:4 660	
7/ TA 49	32 0 77 74 70 6/ 0/4	
## B5 /25	147 145 125 145 145 105 78	
	Numbers of date pains	
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smaller. Coefficients statistically significant at the 99-percent-1 confidence level, and coefficients statistically significant at the 95percent-confidence level but not at the 99-percent-confidence level are each designated in the matrix. All other coefficients are taken to indicate no significant correlation. The numerical value of a given correlation coefficient is of little importance; the important thing is whether it is large enough, considering the number of data pairs used, to indicate a significant degree of positive or negative correlation. The statistical confidence levels provide arbitrary but objectively 10defined cut foffs for deciding which correlation coefficients are 11 large enough to consider geologically important. Coefficients 12 significant at the 95-percent-confidence level are taken to indicate a 13 meaningful association or correlation in the geologic context, and 14 coefficients significant at the 99-percent-confidence level are taken to indicate a strong correlation. 16 17 18 19 20-21 22 23

77e 66 (p. 67 TOLLOWS)

In the following discussion gold is presented first, and the remaining ore-related elements are presented in order of increasingly strong supergene leaching and redistribution. We used the same order in the previous section, so that it is easy to compare the sections. 5- The position of bismuth in this order is rather arbitrary because most of the analytical data for bismuth falls below the detection threshold. 10-15-20-

Interpretation of data for the potential indicator elements
Gold

The gold distribution pattern in the Combination-January cuts 3 (pl. 4) is mainly the result of hypogene dispersion processes little modified by supergene dispersion, and therefore is essentially a relict primary dispersion pattern. Plate 1 shows that some parts of productive silicified zones were not ore grade, but most of the ore, whether primary or oxidized (above 130 to 180 feet depth) was from silicified zones. The fact that moderate and high gold values in the cuts are 10- almost entirely within silicified zones indicates that the gold map battern is essentially a hypogene dispersion pattern. The low-tenor 11 oxidized silicified rocks of the cuts may be thought of as belonging 12 to a hypogene gold aureole connecting and extending outward from the 13 ore bodies but restricted to the silicified zones. The fact that the 14 15- ore bodies had assay walls (Ransome, 1909, p. 213, 218; Collins, 1907b, p. 435) is consistent with this conclusion.

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Some supergene movement of gold might be expected, since gold is thought to be capable of at least an intermediate degree of mobility (table 5), but supergene dispersion has not greatly modified the relict primary dispersion pattern, Both Ransome (1909, p. 170-174, 216) and Spurr (1905, p. 138) felt that the gold of rich oxidized ores, some of which occurred with limonite in fractures, was concentrated to some degree during oxidation, but even where notable enrichment occurred rich sulfide ore probably existed previously. All the gold-bearing samples found along faults in oxidized argillized rocks (three high and two intermediate gold values) have high iron contents due to abundant hematite or hematite-goethite mixtures (see loc. 6 (two samples), 5, 9, and 18, pl. 4 and 3). The abundant ferric oxide and hydroxide indicate that the high Eh's (0.9 volt or more) and low pH's (less than 2 to 5) necessary to dissolve gold may well have developed at these localities (see Cloke and Kelly, 1964). These faults may have formed, however, before hypogene activity ceased; if so, gold might have been transported to these sites by either hypogene or supergene processes, or both, since hypogene gold left along faults might be 18 particularly susceptible to supergene mobilization. The correlation .19 diagrams (fig. 8), furthermore, show no significant gold-iron correlation for either silicified or argillized rocks, indicating that supergene 21 dispersion of gold cannot be generally important, even if gold moves 22 short distances to produce enrichment along fractures. This result 23 agrees with Cloke and Kelly's (1964) data on gold solubility and 25 Krauskopf's (1967) calculations and conclusion that significant migration

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of gold occurs only locally and for short distances. Even if gold moves
    as much as a few feet, the effects would not be visible at the scale of
    sampling of a geochemical exploration program.
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Lead

The map patterns for lead is a relict primary dispersion pattern; the patterns for gold and lead are very much similar (plf. 4 and 9).

Lead is generally thought to be immobile during oxidation (table 5) because it forms a very stable sulfate (anglesite, PbSO₄) and carbonate (cerussite, PbCO₃). It is not found in supergene sulfates from the Florence mine (p. 59, 5). Anglesite is likely the predominant lead-bearing mineral in the oxidized zone. Lead values tend to show smaller changes between adjacent samples than do gold values (compare geochemical profiles, plf. 4 and 9), indicating that lead is more evenly distributed through the rocks than gold is. Since lead should be less mobile than gold during oxidation; we infer that this is mainly, if not entirely, a hypogene rather than supergene feature.

The correlation diagrams show that gold and lead are correlated in argillized rocks and strongly correlated in silicified rocks.

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Two features shown by the histograms for lead (pl. 3) are presumably the result of hypogene processes: the negatively skewed distribution of logarithms of lead values in oxidized silicified rocks, mentioned earlier (p. 美), and the bimodal distribution of lead in oxidized argillized rocks. The histogram for gold in silicified rocks has a distinct upper tail which would be even better developed if mined-out ore bodies intersected by the cuts had been sampled. With the strong correlation between lead and gold in oxidized silicified rocks and the similarity between the lead and gold maps, we would expect the histogram for lead to show a better developed upper tail. The histogram for lead in oxidized silicified rocks shows instead a sharp drop above 700 ppm, with very few values of 1,000 ppm or more. Figure 6 shows that few values greater than 1,000 ppm occur in either average-grade or high-grade ores. Amounts of lead in these two groups of samples are similar, whereas all other ore-related elements show at least some enrichment in the high-grade ores, and most show considerable enrichment. Another manifestation of this apparent ceiling on lead values was the paucity of lead-bearing minerals in the ores, although some lead production is recorded (U.S. Geol. Survey, 1912-1924, U.S. Bur. Mines, 1934-1946). Ransome (1909, p. 112) reported galena from only a few localities in the main district. No lead minerals have been reported in either oxidized or unoxidized ores of the Combination and January mines. Preliminary microprobe examination of one unoxidized average-grade ore sample by C. G. Czamanske shows that neither stibioluzonite nor tetrahedrite-tennantite contains as much as 0.1

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percent lead, but bismuthinite may contain amounts on the order of 0.5 1 percent. The relative scarcity of lead thus seems to be characteristic 2 of the Goldfield district and most likely relates to lack of available lead at the source of the ore-bearing solutions. Ores obtained in the deeper, eastern parts of the district contained considerable 5 copper and tin (Searls, 1948, p. 17-18; Wilson, 1944), suggesting some zoning of those metals, but lack of information on lead content of the ores makes it impossible to evaluate zoning in the district with respect to lead. 10-11 12 13 14 15-16 17 18 19 20-21 22 23

ı Regarding the bimodal frequency distribution of lead in oxidized 2 argillized rocks, the lower mode lies below the average amount of lead 3 in unaltered dacite (17 ppm or -2.8 log percent), and the upper mode lies above it, so some rocks have been depleted and other rocks have 5 been enriched in lead. The dacite represented by the lower mode was probably leached of lead during the hydrothermal activity that produced argillization. Most of the values comprising the upper mode, in the range 20 to 150 ppm, are found in fault zones or within 10 feet of 9 silicified zones or fault zones. The 20 to 150 ppm values near silicified 10zones likely represent a narrow hypogene aureole extending from the 11 silicified zones a short distance into the argillized zones. 12 of the faults originated during the later stages of hydrothermal 13 . alteration, the values associated with these structures could also 14 represent low-tenor hypogene metallization. We have no explanation for 15localities enriched in lead but devoid of structure, such as that 16 north of locality 6 where four high values appear. We infer that 17 igneous lead was leached from the dacite during the earlier stages of 18 hydrothermal activity, and ore lead was added locally to the resulting 19 argillized rocks during the later stages of hydrothermal activity. Other ore-related elements may have had a similar history in argillized 21 rocks, but if so, ranges of values, detection limits, and supergene 22 redistribution effects obscure the bimodal frequency distributions. 23

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Silver

Almost all intermediate and high silver values are found in the 2 silicified zones, with the highest values in the same general areas as those in which high gold values occur. This fact argues against much supergene movement of silver, even though silver often shows at least moderate supergene mobility (table 5). Supergene silver mobility is relatively low at the Florence mine, as deduced from the analyses of secondary sulfates (p. 50, 68). Ransome (1909, p. 119-120, 171) reported chlorargyrite (cerargyrite, AgCl) and probable minor embolite 10- (Ag(Cl, Br)) in the oxidized zone at Black Butte and McMahon Ridge, miles northeast of the main district. Much of the ore in that part of 11 the Goldfield altered area was oxidized, and yet it generally had higher 12 silver-to-gold ratios than did ores from the main district (Ransome, 1909, 13 p. 246-251). Schaller (1941) reported probable minor iodyrite in oxidized 15- ore that likely came from the Combination-January area. Burgess (1911) reported silver halides in the Tonopah district, 25 miles north of 16 Goldfield. In the upper part of the oxidized zone at Tonopah, silver 17 was not carried far from the original sulfide ore before it was precipitated as chlorargyrite, the most abundant silver halide mineral. Boyle (1968, 20-p. 188-207) indicates that much silver should remain in the oxidized zones of sulfide ore deposits in semiarid and arid areas, with silver halides the most abundant silver minerals in the upper parts of the oxidized zones. The foregoing observations point to a relatively low 23 supergene mobility for silver in the oxidized zone at Goldfield, and 24 25-indicate that the silver map shows essentially a relict hypogene pattern

Silver and iron are positively correlated in oxidized argillized rocks, indicating some movement of silver and redeposition with limonite. Silver must therefore be more mobile than lead. The supergene mobility of silver relative to gold is difficult to determine based on association with limonite, because gold is associated with abundant limonite minerals at several conspicuous localities, but overall, gold and iron are not correlated (see p. 70). 10-15--20-

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Although high gold and silver values occur in the same areas. 1 gold and silver show no significant correlation in oxidized silicified 2 rocks or argillized rocks, and this is probably mainly a hypogene 3 feature. Wilson (1944) analyzed silicified rock samples from the Goldfield Consolidated main vein (Jumbo Extension mine, 830-foot 5 level), Clermont vein (Clermont mine, 225-foot level), and from two surface localities, one of which is on the Jumbo vein (p. 6). He found a consistent relationship between gold and silver in the vein on the Jumbo Extension 830-foot level but found no consistent relationship at the other three localities. We can gain some information by 10comparing other ore-related elements with gold and silver in oxidized 11 silicified rocks. Bismuth, mercury, and copper all show residual 12 13 highs northeast of the January shaft in spite of supergene effects, 14 important particularly in the case of copper. These three elements correlate with both gold and silver, but gold-silver and lead-silver 15correlations are conspicuously missing from this tightly knit group 16 of undoubtedly hypogene element associations, suggesting that the 17 variations in gold-silver ratio have a significant hypogene component. 18 19 20-21 22 23

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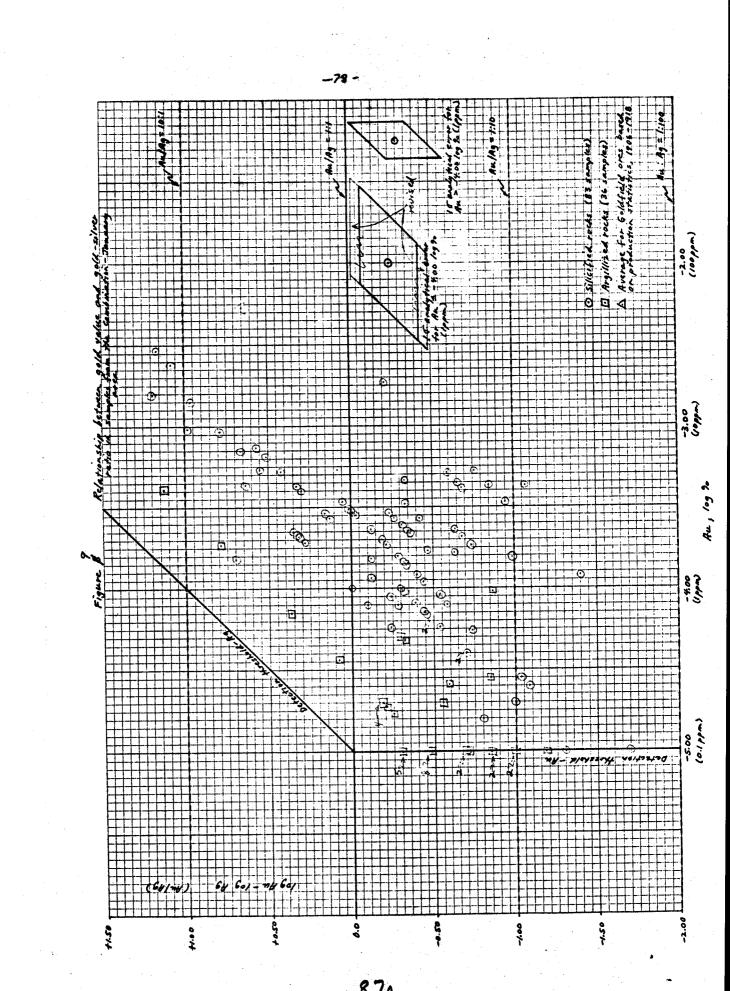
Figure 9, designed to explore the lack of gold-silver correlation further, is a plot of gold-silver ratio against gold value for all samples from the Combination-January cuts that have detectable amounts of both gold and silver. The gold-silver ratios have a wide range (15:1 to 1:50), and the plotted points show considerable scatter, which accords with the lack of statistical correlation. Analytical error for both gold and silver likely accounts for a good deal of the scatter, as the one-standard-deviation boxes accompanying figure 9 show. Sampling error related to removal of analytical portions may

Figure 9 near here

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be a problem for gold (see p. 12-12) in addition to usual analytical error. Supergene redistribution of both elements, even though we consider it insignificant in terms of the geochemical maps, probably also accounts for some of the scatter, since leaching may have changed the original gold-silver ratios more or less in some samples. The correlation data discussed above suggests that a significant part of the data scatter is also due to differences in the proportion of gold and silver deposited by the ore-bearing fluids, but sampling and analytical error aside, we cannot determine precisely how important supergene versus hypogene effects were in producing the wide range of gold-silver ratios without being able to compare gold-silver ratios for oxidized and unoxidized portions of individual ore bodies in the Combination-January vein system.



Although figure 9 shows a wide range of gold-silver ratios, the ratios generally increase with increasing gold values; if a significant part of the data scatter is due to hypogene processes, this trend is probably also the result of hypogene processes. The high ratios at high gold values cannot be satisfactorily explained by supergene 5 --leaching of silver relative to gold in richer ores, because production data from Ransome (1909) and from U.S. Geological Survey statistics 7 (1903-1918) indicate that high gold-silver ratios were characteristic 8 of most ore-grade material from the main district whether oxidized or unoxidized. as shown in table 7. If the increase in gold-silver 10-11 The reverse is true for some ores from mines in outlying areas 12 (Ransome, 1909, p. 171-172, 250). 13 14 15-TABLE 7 NEAR HERE 16 ratios with increasing gold values is indeed due to hypogene processes, 17 the ore-mineral paragenesis should reflect this increase. 18

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ome Goldfield, ores Reference	Kansome, 1909, p.171	18id, p.171	ibid, p. 230	1612, p. 167	ibiż,p. 167, 169	1846, 9.169	انارم ، انار		4.5.6cological Survey, 1906-1918, Miseral Resources Of the United States Garad Volumes].
Table of Gold-silver ratios in some Goldfield ores Location of ore, Reference Reference	Upper levels of the Combination Mine, best ove processed before 1905 (mostly oxidized).	Upper levels of the Combination mine, best of the above lot of Nob tons lentively oxidized)	Florence mare, 250-foot kuel, unoxikiried.	Composite Sample of 9 lots of our decision of the Holeant mine (caply 1907). These and Lollouning Holeant mine samples under the samples	gandyses of rich ore from the Holowal made on material recovered from ore thieves.	6 anchyses of rich ove from 1220-foot level, flages-Nonaette 1226, p. 169 workings, Horank Muré.	Carload of ore shiped from Hayes-Homeste lease, Mohamk mine, January 1907, the rickest as your of ever whoped out of the district	Combination mine, 230-foot level, unoxidized.	Based on total production for perror 1906-1918, the main productive period. Mostly unoxidized.
E E	1166 Skort tons	Short fons	Unknown	Unknown	Unknown but Small	Unknown	47.8 Skort tons	Hank Sample	3,776,609 Short tons
E. 12	7.55.7	30:1	28.8:/	8:1	5:0:/ to 1:9:/	##:/+o s:5:1	8.1.71	<i>#:1</i>	1:944
37.37 24	693	088/	3460	20,000	4500 -	12,200-	20, 900	23	53.5

Tolman and Ambrose (1934) determined the ore mineral paragenesis pyrite and marcasite, followed by famatinite (stibioluzonite), tennantite, and sphalerite, followed by bismuthinite, followed by goldfieldite, followed by gold-silver tellurides and native gold, with considerable overlap between adjacent minerals in the sequence. They found at least small amounts of tennantite associated with stibioluzonite in all the high-grade ore specimens they examined, but in a figure showing sequence of precipitation of the ore metals. they show both silver and gold being precipitated simultaneously at the end of the ore paragenesis, implying that virtually all of the silver accompanied gold late in the depositional sequence. According to Goldschmidt (1954, p. 190, 194), tennantite-tetrahedrite series minerals may bear even more silver than galena. In the average-grade ores we examined, which all contain more silver than gold, pyrite and stibioluzonite are generally the only abundant minerals that belong to the ore-mineral sequence, but in several samples, subordinate amounts of tetrahedrite-tennantite are associated with the stibioluzonite. Except for small amounts of bismuthinite and very small amounts of native gold, none of the other minerals characteristic of rich ores appear in the average-grade samples. Preliminary microprobe analyses suggest that tetrahedrite-tennantite is indeed an important host mineral for silver in the Goldfield ores, although the silver content of the tetrahedrite-tennantite is variable. Concentrations of 0.5 to 1.5 percent silver are common in tetrahedrite-tennantite, whereas stibioluzonite contains less than 0.1 percent.

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We suggest that silver, mainly in tetrahedrite-tennantite, was deposited early in the ore-metal sequence, accompanying copper. Copper and bismuth continued into the middle of the sequence, but overlapped gold (and, by inference, lead), which were deposited late along with lesser amounts of silver. With this suggested two-stage introduction of silver, most of the silver in many of the ores could have been introduced early in the paragenesis. Also, the amount of gold introduced late in the paragenesis at any given spot could have been largely independent of the amount of silver introduced earlier. Bismuth, deposited in the middle of the paragenesis, and copper, deposited both early and in the middle, overlap both early silver and late gold and middle-to-late lead, producing several positive correlations between silver, bismuth, and copper, and between gold, lead, bismuth, and copper, but no correlation between gold and silver or lead and silver associated with hypogene aureoles in ore-bearing silicified rocks.

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Figure 9 shows that in oxidized argillized rocks, the gold-silver ratio is generally higher than it is for oxidized silicified rocks with comparable amounts of gold. No sulfide mineral other than pyrite has been reported from unoxidized argillized rocks, nor have we observed any sulfide other than pyrite in our few observations of unoxidized argillized rocks in the Florence mine and elsewhere in the Goldfield altered area. Possibly the small amounts of lead, gold, and silver in these rocks were all introduced by ore-bearing fluids late in the paragenesis, explaining the high gold-silver ratios, but leaving unexplained the lack of correlation between gold and silver that persists in oxidized argillized rocks. In argillized rocks silver correlates with iron, indicating different behavior, and possibly greater supergene mobility, than it shows in silicified rocks, where it probably forms silver halides nearly in place. The strong lead-silver correlation in argillized rocks may also be due wholly or partly to supergene processes. Silver may be associated with lead-bearing oxidation products such as anglesite (PbSO4) or plumbojarosite (PbFe6(SO4)4(OH)12 instead of silver halides (see Boyle, 1968, p. 192-195). Since lead moved little during oxidation, presumably the silver associated with lead moved to the sites of oxidizing lead-bearing minerals. Distances the silver moved, however, would not necessarily have to have been very large.

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Bismuth

Our data show that bismuth correlates with gold, lead, and silver in the Combination-January area in spite of oxidation (fig. 8). Wilson (1944) noted that bismuth correlated with gold, silver, and tin in vein material from the 830-foot level of the Jumbo Extension mine, but these correlations did not exist at any of the other three localities he sampled (p. 6). Our bismuth and silver maps are similar 6 and D, although more samples have undetectably small amounts of bismuth than is the case with silver. With the large number of samples having amounts of bismuth below the detection threshold, effects of oxidation are difficult to evaluate. There are only two oxidized argillized samples with detectable amounts of bismuth, so correlation coefficients cannot be calculated for argillized rocks. Since bismuth in silicified zones, however, correlates with gold, lead, and silver, which all show essentially relict hypogene dispersion the bismuth my (pl.4) patterns, plate 7 must also show essentially a relict hypogene pattern, as one might predict from figure 8. On the other hand, Ransome (1909, p. 121-123, 213, 219) reported bismite (Bi203) in oxidized ore, in some cases partially filling prismatic cavities in quartz left by leaching of bismuthinite (Bi2S3). The supposed bismite was subsequently reidentified by Schaller (1941) as bismoclite (BiOC1), another secondary bismuth mineral. Ransome's observations thus indicate partial removal of bismuth during oxidation, as do bismuth contents of the secondary sulfate samples (p. Better data for bismuth, therefore, might well reveal some definite supergene effects. Analyses with a detection limit at least an order of magnitude lower than ours (at least 0.5 ppm instead of 5 ppm) would probably be necessary to provide adequate data for bismuth.

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Mercury

Mercury correlates strongly with gold, lead, and silver in suitived sold and silver in suitived rocks (fig. 7), indicating that hypogene processes were important in producing the features seen on the mercury map (pl. 1). These relationships are somewhat difficult to appreciate by visual comparison of the mercury map with the gold, lead, and silver maps, due to the rather small total range of mercury values and strongly overlapping ranges for silicified and argillized rocks (see histograms, pl. 1). Mercury also correlates with copper, as do silver and gold; much copper was leached during oxidation, but enough remains to give correlations with these three ore-related elements (see p. 1). Mercury thus forms a hypogene dispersion aureole extending outward from the ore bodies but restricted to silicified rocks, just as gold, lead, and silver do.

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In the case of mercury, we must not only consider whether secondary dispersion took place, but also how it took place. Mercury is unique in that secondary dispersion can occur by diffusion of mercury gas. Mercury probably has low mobility in the low pH and moderate to high Eh solutions that must have been involved in oxidation 5 at Goldfield (Krauskopf, 1967, p. 516). Diffusion of mercury gas, therefore, should account for most posthydrothermal movement of mercury away from the silicified zones at Goldfield. Migration of mercury gas conceivably could have become more important than hypogene dispersion as hydrothermal activity waned, and could have continued to 10the present, regardless of when oxidation took place. If various sulfide minerals, however, carried most of the mercury in the ores, then little mercury would have been free to disperse until the sulfides were destroyed by oxidation. 15--20-

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Although some of the features of the geochemical map and some of the element correlations may be explained by gaseous diffusion, we believe that hypogene dispersion alone provides an adequate explanation. Gaseous diffusion of mercury might explain the fact that differences in mercury values between adjacent samples in the cuts are mostly rather small, diffusion having evened out sharp variations. On the other hand, the lack of distinct gradients outward from silicified zones argues against extensive supergene migration of mercury gas. Intermediate and high mercury values between localities 3 and 4 may represent a narrow mercury halo around the silicified zone exposed in the cut wall northwest of locality 4 (see p. 4). This is the only area within the cuts showing a halo, but it is, after all, adjacent to a particularly wide segment of the Combination-January vein system (pls. 1 and 2) and the vein material to the northwest is definitely enriched in mercury. Explaining the narrow mercury halo between localities 3 and 4 as due solely to diffusion, however, raises another problem besides lack of evidence for a diffusion gradient: is hard to understand how several samples at localities 5 and 6 escaped being enriched in mercury.

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Regarding ways in which various element correlations could result from gaseous diffusion, mercury is correlated with gold in oxidized argillized, as well as oxidized silicified rocks; this could be due either to hypogene association, or to amalgamation of gaseous mercury with native gold, or both. But mercury is also strongly correlated with silver in both silicified argillized rocks, and free native silver has not been reported in any oxidized Goldfield ores (Ransome, 1909, p. 171). Silver in the oxidized zone is probably present mainly as silver halides, or, in argillized rocks, it may be associated with oxidized-zone lead-bearing minerals (p. 16). We do not know how silver 10and mercury might be associated mineralogically in the oxidized zone. Thus the mercury-gold and mercury-silver correlations may or may not be the result of gaseous diffusion of mercury, but the mercury-silver correlations move probably are not. Gold and silver, regardless of whether in they have experienced some supergene movement, 15probably still show dominantly hypogene patterns, so whether or not gaseous diffusion played an important role, the mercury-gold and mercury-silver associations in these rocks cannot be taken as evidence for supergene dispersion of mercury. Mercury shows negative correlations with copper and molybdenum in argillized rocks: both copper and 20molybdenum have undergone considerable supergene redistribution in argillized rocks, as will be explained later (p. 88-882 If gaseous diffusion effects were predominant, we would expect to see fewer and weaker correlations between mercury and other metals in

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oxidized silicified rocks, and no significant correlations, except

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possibly with gold and less likely with silver, in oxidized argillized rocks.
     We infer that hypogene dispersion features dominate the mercury map.
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With the strong association between mercury and gold, lead, and silver in silicified rocks, the lack of correlation between mercury and bismuth is surprising. Published descriptions of mineral paragenesis are of no help because mercury-bearing minerals have not been reported from any part of the main district, even from high-grade ores, although mercury definitely was enriched in the ores (fig. 1, and Ransome, 1909, p. 113). Relatively large amounts of mercury are known to occur in tetrahedrite-tennantite (Ramdohr, 1969, p. 554; Chan, 1969), which could explain the correlation with silver, which also occurs in tetrahedrite-tennantite, but does not explain the association with gold and lead. Without a detailed investigation of the amounts of mercury in all the various ore minerals, we cannot attempt further explanation.

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The analytical data for unoxidized silicified rocks, average-grade ores, and high-grade ores (fig. 3) show much more mercury in these groups of samples than in rocks from the cuts. One might conclude that much mercury has been lost from the silicified rocks of the cuts, presumably by secondary migration. We hesitate to compare the data of figure 3. however, with that for the cuts because the low-grade ore samples, Florence mine samples, and unaltered dacite samples were prepared by a different laboratory using different procedures than samples from the cuts (the high-grade samples were hand-ground). Different sample preparation procedures can produce greatly different mercury yield from the same sample, due to loss of mercury during grinding (Crosby, 1969, p. 189-191). Relative differences between samples done by a given method tend to be retained if the samples are treated differently and rerun, even though absolute amounts of mercury obtained may be considerably different. Thus there should be no problem making comparisons between rocks from the cuts fand between the groups of samples shown on figure 5, but we cannot compare the former body of data the latter one.

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Secondary dispersion of mercury to form anomalies in soil or other overburden above ore bodies is very common and has been documented but helds little promuse as a geochemical exploration technique at boilhold. in many studies, A To produce these anomalies, mercury is transported as gas, in solution, or by mechanical movement of mineral and rock fragments, the relative importance of these three mechanisms varying from place to place, depending on permeability and adsorptive capacity of the soil or overburden, climate, and topography (Koksoy and Bradshaw, 1969). Often primary dispersion aureoles exist along with the secondary halos, but not in every case even if secondary halos in soil are well developed (see particularly Friedrich and Hawkes, 1966). The Goldfield ore bodies are in the category of deposits with limited primary dispersion of mercury, because hypogene aureoles are essentially restricted to silicified rocks, and gas-phase dispersion has not substantially modified the primary dispersion pattern. relationships have been observed in several other districts: Pachuco Real del Monte, Mexico (Friedrich and Hawkes, 1966), Achisai, Kazakhstan (Furzov, 1958), Ivrindi, Turkey (Bradshaw and Koksoy, 1968), and probably Cripple Creek (Gott and others, 1967). We did not attempt a soil survey for mercury in the Combination-January area because much of the surrounding area is covered with mine dumps or otherwise Since mercury at Goldfield probably does not move in disturbed. solution during oxidation, and since relatively little has moved in the gas phase, mechanical dispersal of mercury-bearing silicified rock detritus is the only way that soil anomalies could form. Low to moderate topographic relief throughout the altered area and low

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oxidized anomaly contrast between metallized silicified rocks and surrounding argillized rocks thus become two factors that limit the possibilities for developing mercury soil anomalies at Goldfield. 5--10-15--20--

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Arsenic

Arsenic is the first element to be discussed that shows a strong correlation with iron, indicating significant supergene redistribution, The arsenic map, however, also shows at least one definite relict hypogene feature: a group of high and intermediate arsenic values in silicified rocks northeast of the January shaft, an area with particularly high gold, lead, silver, bismuth, and mercury. relict hypogene feature is less pronounced than it was for the forgoing elements, and a lack of correlation with gold, lead, silver, and bismuth results, although correlation with mercury persists. noted previously (p. 53), silicified rocks in other parts of the cuts do not show a consistent relationship between gold and arsenic. Silicified rocks have enough residual arsenic so that their average arsenic content is higher than the average arsenic content of argillized rocks, as is the case with the preceding dominantly hypogene elements. More severe leaching can reduce the average amount of an element in silicified rocks to a figure below that for argillized rocks, as is We conclude that arsenic has undergone some supergene redistribution throughout the area, although the hypogene pattern has not been completely erased northeast of the January shaft.

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The arsenic-copper and arsenic-molybdenum correlations in silicified rocks are of interest because both copper and molybdenum have also been partly leached during oxidation, as will be described later. The copper and arsenic in silicified rocks were both derived from stibioluzonite,—which was relatively abundant even in low-grade

—Ransome's analysis of famatinite (stibioluzonite) shows almost

-Ransome's analysis of famatinite (stibioluzonite) shows almost 50 percent Cu₃AsS₄, the remainder of the mineral being Cu₃SbS₄ (Ransome, 1909, p. 118-119).

ores and also from smaller amounts of tetrahedrite-tennantite.

Stibioluzonite and tetrahedrite-tennantite in the silicified rocks
were oxidized along with pyrite, releasing iron and antimony along
with the copper and arsenic. No molybdenum minerals have been reported
in the ores, so without analyses for individual ore minerals the paragenetic relationships of molybdenum are unknown. Even though considerable
iron was removed from silicified rocks during oxidation (proff),
enough remained to produce the strong iron-copper-arsenic-molybdenum
association, which must be due to sorption of all three elements on
limonite (note the strong copper-iron and molybdenum-iron correlations
in siglicified rocks), and probably also to coprecipitation of hydrated
iron arsenates (particularly scorodite, Fe(AsO4)·2H₂O), copper and
copper-iron arsenates and hydrated arsenates, and possibly hydrated
iron molybdate (ferrimolybdite, Fe (MoO₄)₃·8H₂O) with limonite.

It is puzzling that arsenic does not correlate with copper and molybdenum in argillized rocks, especially since copper and molybdenum are strongly correlated there, and all three metals are again strongly correlated with iron. As mentioned earlier, we have never seen any sulfide other than pyrite in argillized rocks. Thus, any small amounts of arsenic and copper introduced into argillized rocks during hypogene metallization may have existed in forms other than stibioluzonite and tetrahedrite-tennantite, and whatever mineral phases these were, perhaps they did not behave as stibioluzonite and tetrahedrite-tennantite did during supergene leaching. The mineralogic composition of the argillized rocks, particularly the greater abundance of clays relative to silicified rocks, may also have affected the mechanism of redeposition of arsenic and copper in the oxidized zone as the water table moved downward, although arsenic and copper are not associated with any particular clay -- 108 - 108d) mineral (p. \$8+89a, fig.

Conditions in argillized rocks during oxidation probably favored formation of secondary lead-arsenic minerals, which could account for the lead-arsenic correlation in argillized rocks.

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Copper

With the exception of the high values in silicified rocks northeast of the January shaft, copper values in oxidized silicified and argillized rocks look similar (pl. 10). Average copper values for oxidized silicified and argillized rocks are indeed essentially the same, and the histograms (pl. 10) confirm that the ranges of values for the two groups of rocks are similar.

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Copper was the most abundant metal in both average and high-grade ores, and since the ores and their associated hypogene gold, lead, and silver aureoles were restricted to silicified rocks, the silicified rocks surely contained more copper than did argillized rocks prior to oxidation. Copper is relatively abundant in the unoxidized silicified rocks of the Florence mine (300 ppm geometric mean, fig. 6). Before oxidation, silicified rocks of the Combination-January cuts likely had at least as much copper as the Florence mine silicified rocks, and locally more. The average copper contents of oxidized silicified 10rocks (geometric mean 16.4 ppm) and oxidized argillized rocks (geometric mean 16.7 ppm) are not much above the 10 ppm copper found in unaltered dacite (fig. 6). Obviously considerable copper has been leached from silicified rocks during oxidation: more, no doubt, than for any element yet discussed. Whether the oxidized argillized rocks have 15also suffered overall removal of copper cannot be determined without unoxidized argillized rocks for comparison.

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Copper correlates strongly with iron in both oxidized silicified and argillized rocks, another indication that copper has undergone considerable supergene redistribution. Arsenic and molybdenum also correlate strongly with iron in both oxidized silicified and argillized rocks, and copper correlates with both these elements in silicified rocks and with molybdenum in argillized rocks to form an arsenic-coppermolybdenum-iron supergene association. We can offer no detailed explanation for the strong negative correlation between copper and mercury in oxidized argillized rocks. The ultimate cause is presumably greater supergene movement of copper relative to mercury. Many geochemical field studies, as well as laboratory studies on the solution chemistry of copper, have shown that copper is moderately to highly mobile in oxidizing sulfide ores (table 5 and Garrels and Christ, 1965, p. 240). Behavior of copper at Goldfield is no exception, judging from rocks of the Combination-January cuts and from the Florence mine secondary sulfates (p. 59, 63).

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In spite of the relatively high mobility of copper, hypogene element associations in oxidized silicified rocks have not been completely destroyed, Three again; high values, appear northeast of the January shaft (pl. 10) associated with high gold, lead, silver, bismuth, 5- mercury, and arsenic values, and correlations between copper and gold, silver, and mercury persist. As explained above, the copper-arsenic association is due primarily to their mutual association with limonite and therefore is mainly a supergene phenomenon, but hypogene association may contribute to this correlation indirectly in that both copper and 10- arsenic were locally released from ores together in relatively large amounts, and consequently large amounts of both elements found their 11 way into the limonite formed at these localities. Hypogene association 12 13 could also contribute directly to the copper-arsenic correlation if the 14 rocks northeast of the January shaft contain relict (unoxidized) disseminated stibioluzonite and tetrahedrite-tennantite. If unoxidized 16 sulfides exist in the samples from the cuts, however, they must be 17 fine grained and their amounts must be small, because we have not seen 18 sulfide grains in the rocks, and we are convinced that substantial 19 proportions of the iron, copper, and arsenic present before oxidation have been removed by supergene leaching. The copper remaining with gold, 21 silver, and mercury must not be associated with limonite, because none 22 of these three elements correlate with iron. This copper could occur 23 as azurite (Cu₃(CO₃)₂(OH)₂), or malachite (Cu₂(CO₃)(OH)₂), both reported by Ransome (1909, p. 108-109, 216) as rarely staining some oxidized and partly oxidized ores. It is unlikely that much copper resides in

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these two minerals, however, because pH values of solutions in the oxidized zone at Goldfield probably were generally too low to form them (Garrels and Christ, 1965, p. 240). We have not seen chrysocolla ($\operatorname{CuSi0}_3 \cdot 2\operatorname{H}_2 0$) or turquoise-group minerals (particularly turquoise, ${\rm CuAl}_6({\rm PO}_4)_4({\rm OH})_8\cdot {\rm 4H}_2{\rm O}$, and chalcosiderite, ${\rm CuFe}_6({\rm PO}_4)_4({\rm OH})_8\cdot {\rm 4H}_2{\rm O}$, and as far as we know none of these minerals have been previously reported. Of the 39 oxidized silicified rock samples we X-rayed, 30 contained kaolinite, and (9) were free of kaolinite. Average amounts of copper for these kaolinite-bearing and kaolinite-free samples are essentially the same and both are very close to the average for all 129 silicified 10-11 rock samples. Copper thus is not notably associated with clay in the silicified rocks, and the form of copper in these rocks remains 12 13 unidentified. 14 15-16 17 18 19 20-21 22 23 24

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In the oxidized argillized rocks copper and other elements that were mobile during oxidation could be associated with clays as well as with limonite, because clay minerals are capable of ion exchange and sorption (Grim, 1968, p. 185-233; Carroll, 1959); however, we do not see this association. The argillized rocks of the Combination-January area all contain kaolinite and illite in various proportions, and 22 samples from the area north of locality 18 contain montmorillonite as well (pl. 2). Although pH of solution, clay-mineral particle size, and metal-ion concentration all effect the amount of metal sorbed by clays, capacity of the common clays to sorb copper, molybdenum, and zinc generally increases in the following order: kaolinite, illite, montmorillonite (Heydemann, 1959; Jones, 1957; Chu, 1969). Montmorillonite is particularly effective in sorbing many ore metals, mainly because of its large cation exchange capacity (Perel'man, 1967, p. 100). Cation exchange is important in the sorption of copper and zinc, whereas anion exchange is important for molybdenum and arsenic, which form molybdate (MoO_4^{-2}) and arsenate (AsO_4^{-3}) ions in solution. Molybdate and arsenate, however, readily form very insoluble compounds with ferric iron (ferrimolybdite, Fe₂(MoO₄)₃·8H₂O, and scorodite, Fe(AsO4).2H2O) so that we would expect the clays to be less important than iron in determining the supergene redistribution of these two In the oxidized argillized rocks kaolinite, illite, and montmorillonite occur in various proportions; total clay contents generally are between 20 and 60 percent. Ideally, we should compare copper, molybdenum, and zinc abundances in rocks with known amounts

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(p. 108a follows)

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of each clay mineral. Clay mineral percentages are difficult to estimate accurately, however, so we have merely divided the argillized rocks into three groups, as follows: montmorillonite-bearing, montmorillonite-free with illite dominant over kaolinite, and montmorillonite-free with kaolinite dominant over illite. The groups include only samples that we X-rayed in the course of our petrographic examinations. Figure 10 gives the results. Data for iron is included

Figure 10 near here

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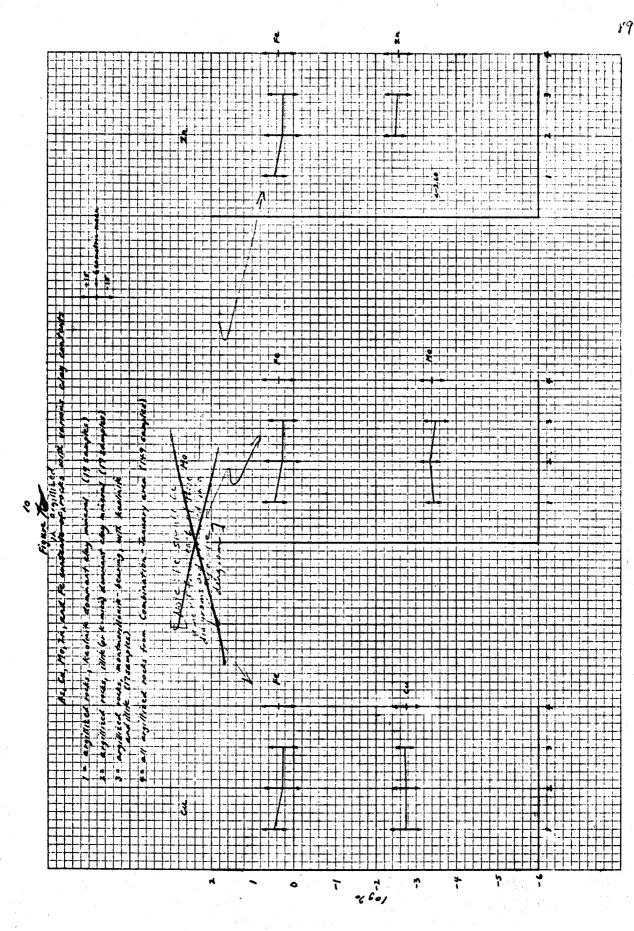
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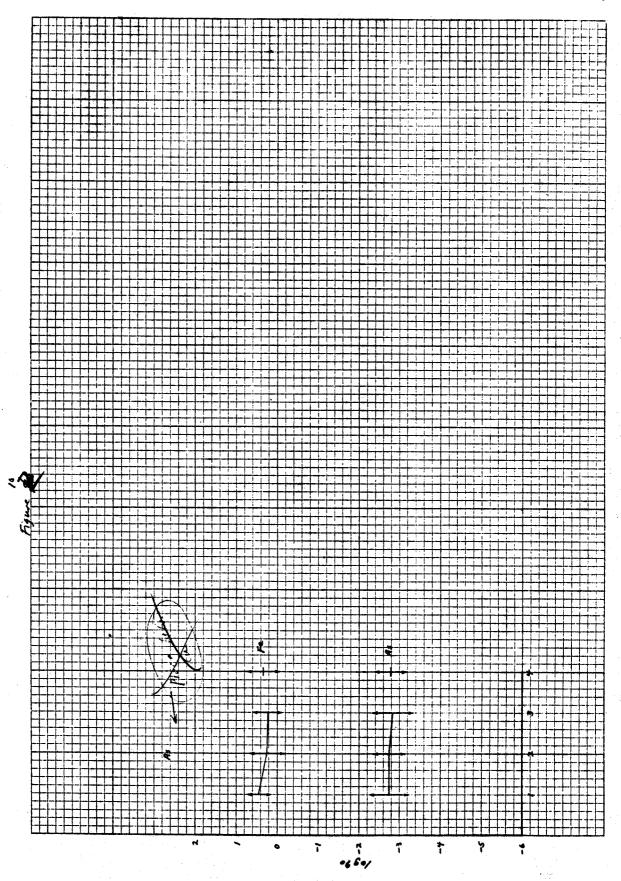
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for comparison; the three groups of samples have similar iron contents. Arsenic, copper, molybdenum, and zinc do not show significant differences between the three data sets, with the possible exception of zinc in kaolinite-dominant rocks. Since the average for zinc in these rocks may not be much below the detection threshold at -2.60 log percent (25 ppm), we cannot say definitely that clay content has an effect on zinc. Although the three data sets are not as precisely defined as we would like them to be with respect to relative amounts of the various clay minerals, we conclude that the clay content of argillized rock has little effect on supergene redistribution of the above four elements. At least, abundance of iron must be much more important than clay content.

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Questions worth considering are what happened to the copper that was leached from the oxidized silicified rocks, and why it did not 2 form a supergene halo. Copper commonly forms zones of supergene enrichment immediately below the water table. No such zone has been reported in the mines at Goldfield, but some enrichment probably 5 escaped notice because copper content was irrelevant in determining the worth of near-surface ores. Covellite (CuS) largely replaces tetrahedrite-tennantite and partly replaces stibioluzonite in two of our unoxidized average-grade ore samples, showing that some supergene enrichment of copper took place at least locally. Perhaps not enough 10ore has been oxidized to produce a pronounced supergene enrichment 11 zone; the pre-Siebert Formation and present-day erosion surfaces in the 12 Combination-January area probably truncate the upper parts of the lodes. 13 Even if part of the copper from metallized silicified rocks moved 14 laterally into the surrounding argillized rocks during oxidation, 15copper values are not consistently high in argillized rocks, being 16 related particularly to limonite distribution, so that copper does not 17 form a distinct and coherent supergene halo around ore-bearing oxidized 18 silicified rocks. We have no way of knowing whether a hypogene copper 19 aureole extended into the argillized rocks, but even if it did it was 20destroyed during oxidation. Perhaps if the lodes were more deeply 21 eroded, we would see greater development of a supergene enrichment 22

zone or a supergene halo, or both.

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Molybdenum

Some high molybdenum values occur northeast of the January shaft along with all the preceding elements, but features of the molybdenum map are largely the result of supergene redistribution (pl. 21). relict hypogene association shown by rocks northeast of the January shaft is so weak that molybdenum shows no correlations with gold, lead, silver, bismuth, or mercury in oxidized silicified rocks (fig. 8). This is probably due both to greater supergene mobility of molybdenum than any previously discussed element and to lack of pronounced enrichment of molybdenum in the ores. Although molybdenum is not as strongly enriched in the ores as are many other ore-related elements (fig. 6), during metallization molybdenum probably was somewhat enriched in silicified rocks relative to argillized rocks. histograms show that oxidized silicified rocks contain less molybdenum than oxidized argillized rocks, although an average was not calculated for oxidized silicified rocks because slightly less than half the silicified samples contained detectable molybdenum. The best evidence for substantial supergene mobility of molybdenum, however, is the similar behavior of molybdenum and iron.

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The association between molybdenum and iron is as strong or stronger than the association between copper and iron; it is strong enough to be readily visible by comparing the molybdenum and iron geochemical maps and profiles (pls. 11 and 18). Possible supergene minerals producing these associations have already been discussed in the section on arsenic, the most important molybdenum mineral probably being ferrimolybdite (Fe₂(MoO₁₄)₃·8H₂O). Jones (1957) showed that hydrous ferric oxide is highly effective in sorbing molybdate (MoOh=) anions from acid solutions. Significant amounts of molybdenum in the Florence mine supergene sulfates (p. moving downward through oxidizing rocks does indeed contain molybdenum. We conclude that molybdenum, like copper, is relatively mobile in the oxidized zone, at least until solutions carrying it come in contact with limonite. The correlation matrix for argillized rocks however) shows one important difference between the supergene behavior of copper and molybdenum: the strong lead-molybdenum correlation. correlation could be due to formation of wulfenite (PbMoO,), which, although not reported at Goldfield, is a common secondary mineral in oxidized ore deposits containing lead. Takahashi (1960, p. 1105-1108) and Williams (1963, p. 1121-1122) have shown, however, that wulfenite is only conditionally stable in an oxidizing environment with sulfate and carbonate present. Since we do not know the mineral forms of lead and molybdenum in either silicified or argillized rocks prior to oxidation, we cannot investigate the lead-molybdenum correlation further, nor can we determine why lead and molybdenum are correlated in



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oxidized oxidized argillized rocks but definitely are not correlated in silicified rocks. Molybdenum shows no correlation with calcium in either 2 silicified or argillized rocks (not shown on fig. 7); so powellite (CaMoO), must not be an important supergene molybdenum mineral at Goldfield. 5 ---10-11 12 13 14 15-16 17 18 19 20-21 22 23 24 25-

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The relatively high supergene mobility of molybdenum presumably produced the negative molybdenum-mercury and molybdenum-silver correlations, just as relatively high supergene mobility of copper presumably produced the strong negative copper-mercury correlation.

Zinc

No vestiges of the strong hypogene metal concentrations northeast of the January shaft remain (pl. 2). The hypogene behavior of zinc must have been similar to that of the other ore-related elements (fig. 6), and consequently it must have been enriched in the silicified rocks during metallization. Sphalerite (ZnS) appeared as a minor constituent in ores of the Combination mine (Collins, 1907a, p. 398). The histograms show that zinc, rather than being abundant in the oxidized silicified rocks, is depleted in oxidized silicified rocks relative to oxidized argillized rocks. Furthermore, zinc shows no positive correlations with any of the preceding elements but does show negative correlations with gold in silicified rocks and lead in argillized rocks (fig. 8). Both the latter elements have experienced only minor supergene redistribution. We conclude that zinc has been strongly leached from the rocks of the cuts and leached more strongly from the silicified rocks.

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The geochemical map (pl. 22) shows that the distribution of zinc in argillized rocks near silicified zones is too erratic to form a distinct halo. Zinc may be coprecipitated with limonite (Hawkes and Webb, 1962, p. 55, 164, 377), but here zinc shows no correlation with iron, so it has not been partially retained in the oxidized zone by copreciptation with limonite or sorption by limonite, as have arsenic, copper, and molybdenum. Sorption by clays probably was not effective in retaining zinc in the oxidized zone (p. 89-89-

Apparently zinc was more mobile during oxidation than any other element considered in this study. The relatively large amounts of zinc in the Florence mine supergene sulfates attest to this high mobility (p. 59, 65, and table 2). High supergene mobility for zinc is in accordance with results of previous work summarized in table 5.



Discussion of indicator elements for gold and methods of geochemical sampling

Geochemical exploration in the vicinity of Goldfield should concentrate on the numerous silicified zones well exposed throughout the hydrothermally altered area. High values for gold, lead, silver, bismuth, and mercury, the elements showing dominantly relict hypogene dispersion patterns, are found mostly in silicified rocks. Even though all these elements except bismuth show at least a few intermediate and high values in argillized rocks, none of them consistently form hypogene aureoles extending a significant distance into argillized rocks. They do, however, form aureoles surrounding ore bodies within silicified rocks. The geochemical profiles constructed for gold, lead, and silver show that these elements in particular form aureoles that contrast sharply with much lower values found in adjacent argillized rocks. None of the elements that were notably leached and redistributed during oxidation, including arsenic, copper, molybdenum, and zinc, have moved outward from oxidized ore-bearing silicified zones into the surrounding oxidized argillized rocks to form distinct supergene halos. Hence bedrock sampling should be restricted to silicified rocks.

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Three of the nine ore-related metals considered in this report are potentially useful guides to ore for geochemical prospecting; usefulness of the remaining six metals is more or less limited. Gold analyses are indispensable as a guide to ore. Lead is reliable as an ore guide and should be particularly useful for reconnaissance sampling programs, because lead is even less mobile than gold during oxidation, and its aureoles are characterized by less sample -to-sample variation than accompanying gold. Silver analyses are also worthwhile because gold-silver ratios vary considerably, and amounts of silver cannot be predicted from gold or lead values. Amounts of silver exceeded associated amounts of gold in some ores, particularly in mines away from the main district (Ransome, 1909, p. 171-172), so potential ore bodies might well be missed if silver were not determined. We consider mercury optional in a geochemical survey because its aureoles are not distinguished by outstandingly high values. Information on mercury, might be more interesting in other parts of the Goldfield altered area than it is in the Combination-January area, because Ransome (1909, p. 113) reported mercury showings at an otherwise unmetallized locality about 4 miles northeast of the main district. Our data is inadequate to fully evaluate bismuth, so we cannot determine how useful bismuth might be in a geochemical survey. Arsenic, copper, and molybdenum are too strongly leached to detect anything but a fairly near-surface, extensive, and relatively high-grade ore occurrence such as that northeast of the January shaft; that is, an occurrence that likely would have been discovered already.

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0 so strongly leached during oxidation that it seems useless. 1 cannot rule out the possibility that arsenic, copper, molybdenum, 2 or zinc could form well-developed supergene halos extending into 0 oxidized argillized rocks around ore todies in other parts of the Goldfield altered area, but from the results of this study, we 5 -cannot commend a sampling program for these elements utilizing bedrock samples from argillized rocks. 7 10-11 12 13 14 15-16 17 18 19 20-21 22 23 24 25 -

Realizing that our data are derived from only a small part of the Goldfield altered area, we tentatively suggest the following minimum gold, lead, and silver values to be considered anomalous for bedrock samples from silicified zones. A geochemical survey should not miss any significant anomalies if gold values equal to or greater than 0.3 ppm, lead values equal to or greater than 70 ppm, and silver values equal to or greater than 1 ppm are considered anomalous. Values in the ranges of "high" values shown on plates 4, 5, and 6 (3 ppm or more for gold, 200 ppm or more for lead, and 10 ppm or more for silver) should certainly be worthy of further investigation. Treating a large number of samples from the entire Goldfield altered area to determine geochemical background and anomalous values is beyond the scope of this report and is the subject of additional work.

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0 In many areas visible concentrations of limonite minerals are a 1 good guide to anomalous amounts of metals, but at Goldfield usefulness 2 of this guide is limited. Qualitative comparison between iron content 3 0 and relatively intense color of either red (hematite dominant) or yellow-red (goethite dominant) hue (Goddard and others, 1948) indicates 5 that intense color and iron content are definitely well correlated. Unfortunately iron content is not correlated with gold, lead, silver, 7 bismuth, or mercury content in silicified rocks (see fig. 8). Thus, although limonite-rich samples are likely to have larger amounts of 10arsenic, copper, or molybdenum than limonite-poor samples nearby, 11 they will not necessarily have large amounts of gold, lead, silver, 12 bismuth, or mercury. We expect that a sampling program utilizing 13 limonite-rich altered rocks or limonite scrapings from fractures, and analyzing for arsenic, copper, or molybdenum, would produce 14 results difficult to interpret. At many localities silicified 15zones are so numerous that the source of anomalous arsenic, copper, or 16 molybdenum could be hard to find, particularly if the highest values 17 showed up in argillized rocks; detailed sampling for gold would then 18 19 be required. A better approach would be to selectively sample silicified rocks, collecting limonite-rich samples wherever they are available, but taking care not to ignore silicified zones showing little limonite. 21 The limonite-rich samples could be tested for unusual amounts of 22 arsenic, copper, and molybdenum in addition to gold, lead, and silver. 23 Such samples should be just as likely to show anomalous gold, lead, 24 and silver as limonite-poor silicified rocks. Whether the information gained over doing gold, lead, and silver alone would be worth the cost of the additional arsenic, copper, and molybdenum analyses is open to question.

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We doubt that soil sampling surveys would be of much use in the Goldfield altered area. Should such a survey be undertaken in one of the more poorly exposed parts of the area, the anomaly contrast between oxidized silicified and oxidized argillized rocks shown by geochemical profiles give an indication of the maximum anomaly contrast that might be encountered, assuming mainly mechanical dispersal of the ore-related elements in colluvium. Lead would be the best indicator in a colluvial-soil survey. Soil-covered parts of the Goldfield altered area appear to be dominantly of colluvial origin, although residual soils that are actually soft argillized rocks appear locally.

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The consulting field remains in the deldrums. Some months ago, I tought that the election would profide the investment confidence which has been lacking. N w. I am counting on the end of the Vet Nam problem. Should that fail, there is no answer except to await the unexpected.

Kitty and I would love to have you stop bye and see us in Reno, at any time; of that does not work, the chances are that I will be in Sunnyvale, one of these days, checking on three grandchildren, andqif so, how about a session and lunch.

My very best to you and your good wife who coded the geochemical data so efficiently.

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· August 1995 · Prince Common of All Street exploration was

Dr. Roger P. Ashley, U. S. Geological Survey, 345 Middlefield Road, Menlo Park, California 34025.

Dear Roger:

Your consideration and thoughtfullness, as reflected in the copy of your open-file report on ore-related elements at Goldfield, has indeed been appreciated.

Report and maps arrived on Friday and I am slowly and methodically going over it. Knowing me, as I hope you do, there are parts, under "Statistical Methods", which lose me. H wever, except for that Assetion, which is an indication of my age and the fact that my "continuing education" has not included that field. I am following the rest with great interest.

Believe me, it is a welcome addition to my personal files and will be of real help, insamuch as interest in Goldfield and its dormant, low grade but economic potential remains very much alive.

Somemonths ago I spent a lot of time with the material you so kindly provided, working out structural and alteration patterns (in an ex-company geologist manner) and hoping that Harold would proceed. N thing new has been heard and I can only conclude that negotiations failed.

Since last seeing you there have been other 'tangents' to keep me off the streets and away from the pool halls. None pays any bills but I believe that given time the desk 'experience' will be productive.

The absolute dating of plutons, starting from the California maps released by the Survey and adding a wealth of material available for Nevada. Oregon and Idaho, provides a different approach to Nevada lineaments and how one might prospect; and, of course, the new Prate Tectonies and ocean floor spreading with its subduction sone has caused sleepless nights.

The consulting field remains in the doldrums. Some months ago, I tought that the election would profide the investment confidence which has been lacking. N w, I am counting on the end of the V et Nam problem. Shauld that fail, there is no answer except to await the unexpected.

Kitty and I would love to have you stop bye and see us in Reno, at any time; of that does not work, the chances are that I will be in Sunnyvale, one of these days, checking on three grandchildren, and if so, how about a session and lunch.

My very best to you and your good wife who coded the geochemical data so efficiently.

Sincerely,

David LeCount Evans

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The state of the second second second second second Dr. Roger P. Ashley. U. S_ Geological Survey, 345 Middlefield Road, Menlo Park. California 94025.

Dear Roger: The total section of the control of the

Your consideration and thoughtallness, as reflected in the copy of your open-file report of ore-related elements at Goldfield, has indeed been appreciable.

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Report and maps arrived on Frigy and I am slowly and methodically going over it. Knowing me. as I hope you do, there are parts, under "Statistical Methods", which lose me. H wever, except for that assection, which is an indication of my age and the fact that my continuing education has not included that field. I am following the rest with great interest.

Believe me, it is a welcome addition to my personal files and will be of real help, instance as interest in Goldfield and its dormant. lew grade but economic potential remains very much alive.

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