

chlorite, sericite and quartz. Chlorite-quartz-biotite alteration is most often associated with relatively higher copper grades; chlorite-quartz-biotitized dacite porphyry between 820 feet and 907.7 feet, for instance, averages 0.35%.

Drill hole BSW-8 (TD 566 feet), was drilled to test a small, intense VIP anomaly for an hypothetical, near-surface, high-grade chalcocite enrichment blanket. The hole provided an explanation for the anomaly--but failed to find the anticipated ore body--in penetrating 92.8 feet of oxidized and quartz-sericitized quartz latite porphyry and quartz-sericite rock, 20 feet of enriched, quartz-sericitized intrusion breccia, averaging 0.3% Cu, and 456 feet of quartz-sericitized intrusion breccia and quartz-feldspar porphyry averaging about 10 volume percent pyrite, and 0.05% Cu as chalcopyrite and cupriferous pyrite.

Drilling, geologic mapping, geochemical sampling and induced polarization have shown Buckskin Southwest to be a tight porphyry copper sulfide system with abnormally steep alteration and mineralization gradients. The system has been penetrated by only one deep hole (BSW-7), which bored through a 510 foot zone averaging nearly 0.2% Cu. Alteration within this zone is highly complex, alternating between varieties of quartz-biotitization and quartz-sericitization. Such an alteration pattern suggests that BSW-7 penetrated the transition zone between the potassic core of the system and its quartz-sericite halo. The unusually steep mineralization gradients suggest that ore grade copper mineralization might occur in the potassic core within a few hundred feet laterally of BSW-7.

An hypothetical Buckskin Southwest ore body would be relatively small--100 to 200 million tons--but high grade, probably in the 0.8 to 1% range. The deposit would most likely be situated north, northwest, or northeast of BSW-7, and would be concealed beneath 750 feet to 1,000 feet of barren quartz-sericite rock and/or postmineral cover.

## RECOMMENDATIONS

The writer recommends deepening two shallow drill holes completed in 1970--BSW-1 (TD 292 feet) and BSW-2 (TD 200 feet)--to 2,000 feet. Completion of these two proposed drill holes will provide a definitive test of the target hypothesis outlined under "Summary and Conclusions".



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for CC blanket & IP

## 1972 DRILLING

Two diamond drill holes were completed at Buckskin Southwest during 1972. The deeper of the two, BSW-7, was drilled to test for the presence of a deep-primary porphyry copper deposit, vertically zoned beneath a quartz-sericite "lithocap". Buckskin Southwest 8 was drilled to test a small, but highly intense, VIP anomaly for an hypothetical, near-surface, high-grade chalcocite enrichment blanket.

Deeper

Buckskin Southwest 8 was completed, with discouraging results, at a depth of 566 feet (Figure 7). The hole penetrated 92.8 feet of oxidized and quartz-sericitized quartz-lathite porphyry and quartz-sericite rock, 20 feet of enriched, quartz-sericitized intrusion breccia, averaging 0.3% Cu as chalcocite, and 456 feet of quartz-sericitized intrusion breccia and quartz-feldspar porphyry averaging 0.5% Cu and cupriferous pyrite and chalcopyrite.

105  
Probably should read 0.05% Cu  
RWR

Buckskin Southwest 7 (TD 1,715 feet), completed in mid-1972, penetrated beneath 750 feet of quartz-sericitized intrusion breccia, a 510-foot zone of complexly altered dacite porphyry, intrusion breccia and quartz-feldspar porphyry averaging 0.16% Cu as chalcopyrite. All rock types within this zone are heavily quartz-veined and altered to various combinations of biotite, chlorite, sericite, and quartz. Chlorite-quartz-biotite alteration is most often associated with relatively higher copper grades; chlorite-quartz-biotitized dacite porphyry between 820 feet and 907.7 feet, for instance, averages 0.35% Cu.

## EXPENSES

Drilling	\$25,395.35
Assaying	1,562.19
Temporary salaries and wages	990.07
Travel and expenses	1,629.55
Materials and supplies	408.28
Freight and haulage	77.68
Licenses and fees	257.75
Roads and drill sites	798.00
Rent	35.00
Utilities	20.50
	<hr/>
	\$31,174.37



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Drill hole BSW-8 (TD 1,125 feet), was drilled to test an intense VIP anomaly for a hypothetical, near-surface, high-grade chalcocite enrichment blanket. The hole provided an explanation for the anomaly, but failed to find the anticipated ore body, in penetrating 92.8 feet of oxidized and quartz-sericitized quartz latite porphyry and quartz-sericite rock, 20 feet of enriched, quartz-sericitized volcanic breccia, averaging 0.3% Cu, and 1,012 feet of quartz-sericitized volcanic breccia, and quartz-feldspar porphyry averaging about 7 volume percent pyrite, and 0.02% Cu as chalcopyrite (Figure 6).

or 45.6' of .05% Cu

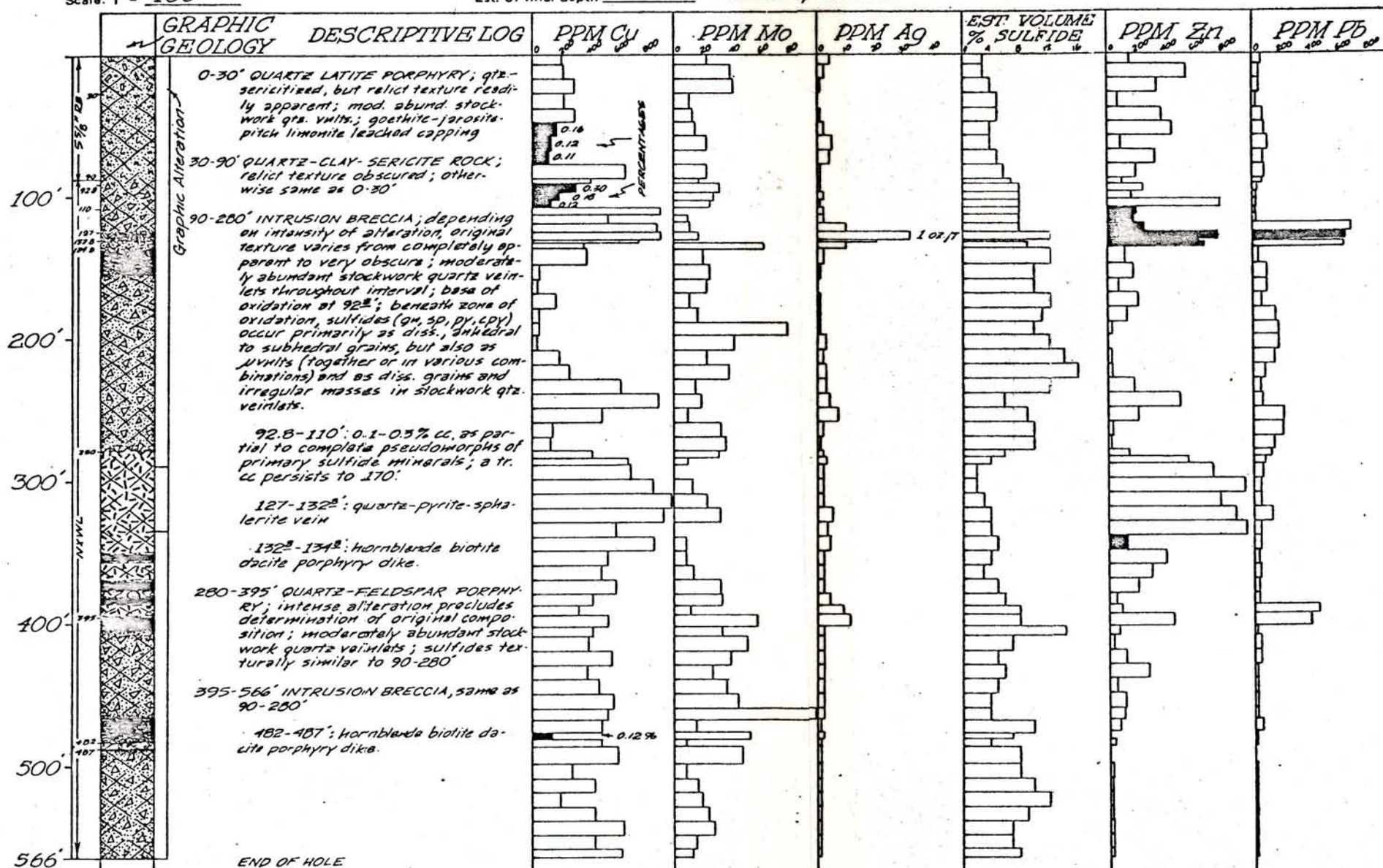


NORTHWEST DISTRICT Name BUCKSKIN SOUTHWESTDRILL HOLE NO. BSW-8 PAGE 1 OF 1

Summary Drill Hole Log

Code 05-01-0411BEARING VERTICAL DIP uPurpose of hole TEST OF HYPOTHETICAL CHALCOCITE  
ENRICHMENT BLANKET & SHALLOW VIP ANOMALYLOCATION: N. 10,670' E. 12,340' (Approximate)COLLAR ELEV. APP. 5290'TOTAL DEPTH 566' BY J.B. HulenSTART 5/11/72 COMPLETED 5/25/72drafted by J.B. HulenScale: 1" = 100'

Est. of final depth





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NORTHWEST DISTRICT  
Summary Drill Hole Log

Name BUCKSKIN SOUTHWEST  
Code 005-01-0411  
DOUGLAS CO., NEV.

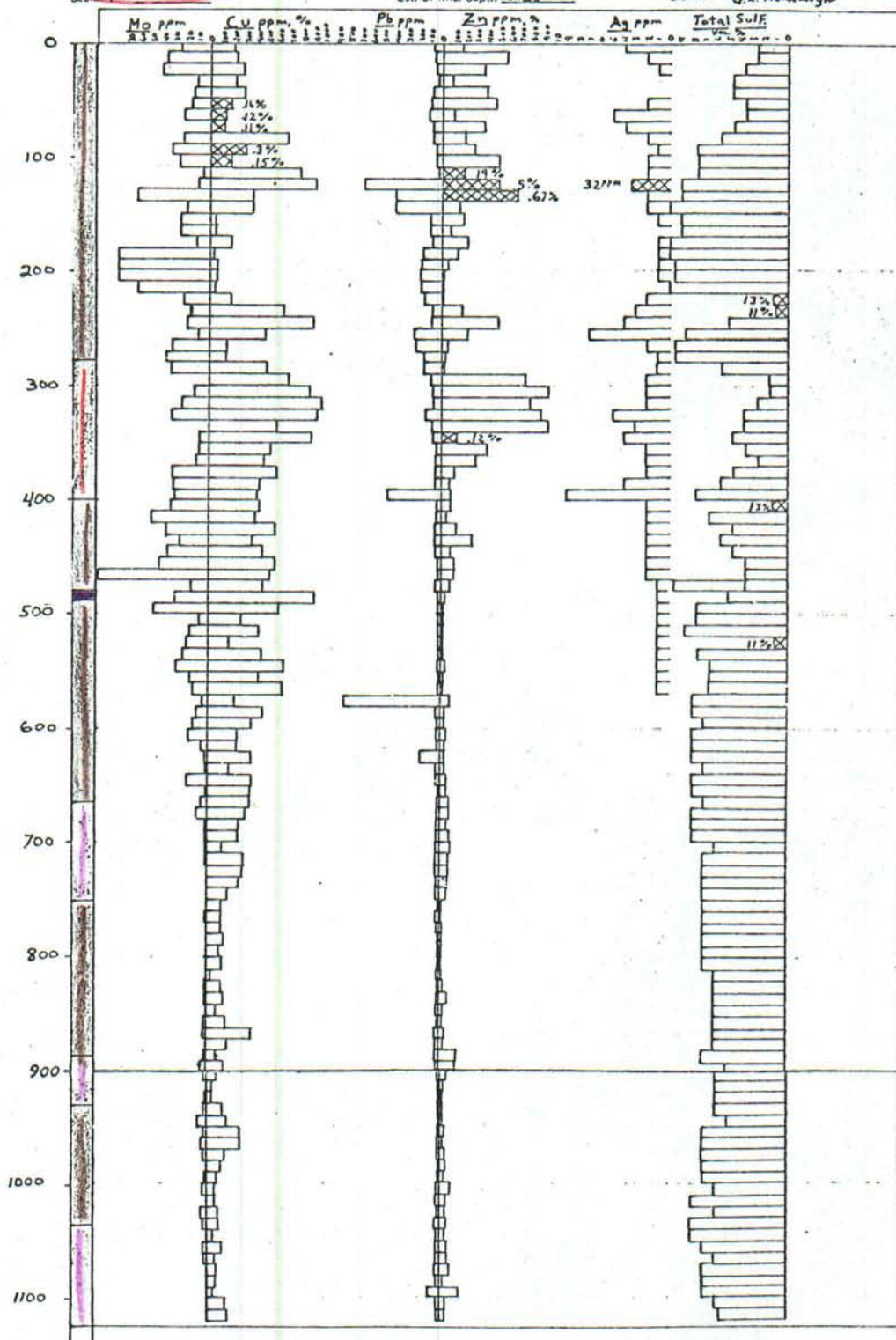
DRILL HOLE NO BSW-8 PAGE      OF       
BEARING      DIP 90°  
LOCATION: N.      E.       
COLLAR ELEV. 5290  
TOTAL DEPTH 1125' BY       
START      COMPLETED     

**BSW - 8**

Scale: 1" = 100'

Est. of final depth 1125'

DATA: G.R. Heinemeyer



SEE FIGURE (5) FOR EXPLANATION  
OF LITHOLOGY



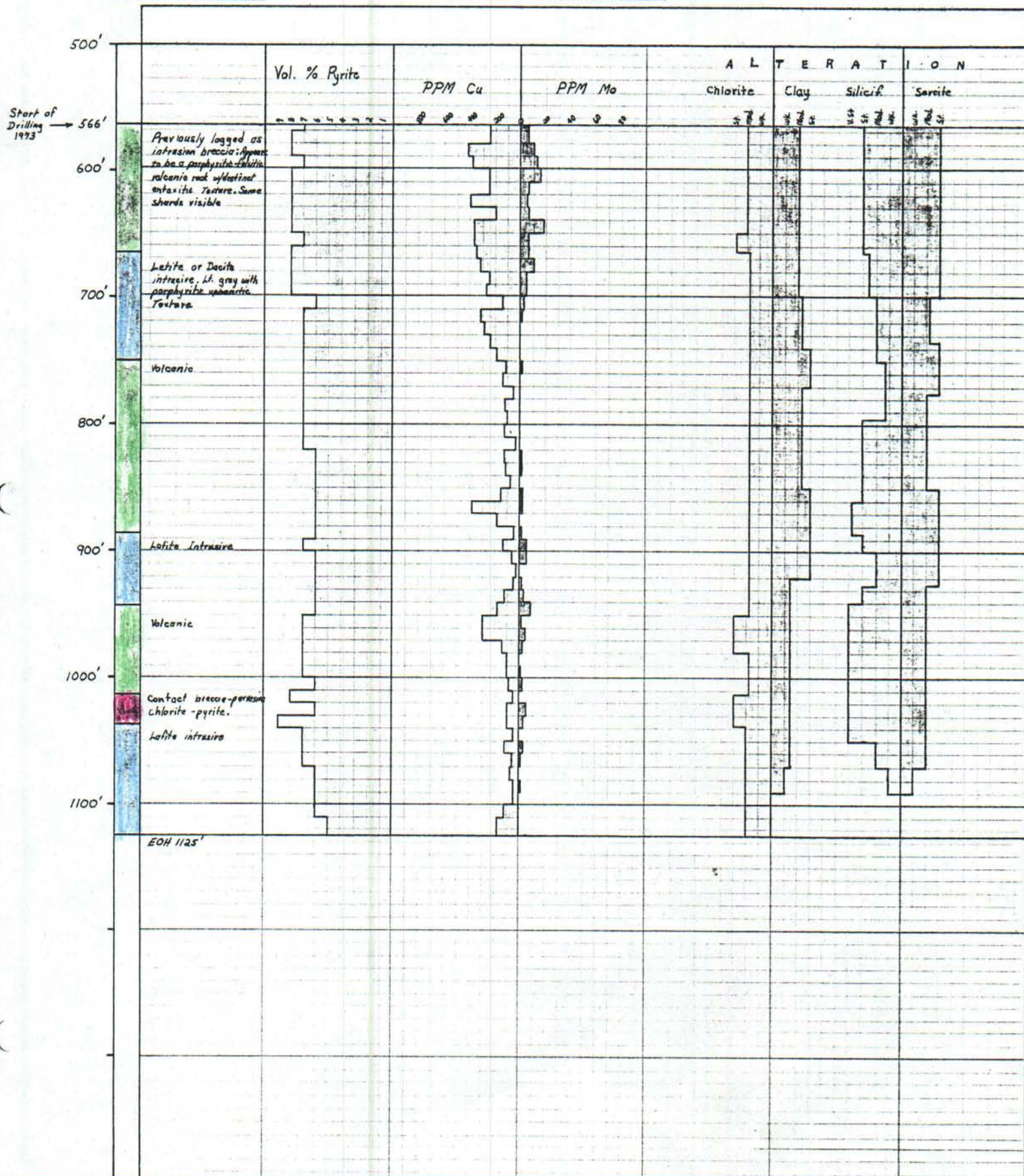
NORTHWEST DISTRICT

Name Buckskin SouthwestDRILL HOLE NO. 8 PAGE 8 OF 15

## Summary Drill Hole Log

Code 005-01-0411BEARING \_\_\_\_\_ DIP 90°Purpose of hole To complete assessment work for 1993

LOCATION: N. \_\_\_\_\_ E. \_\_\_\_\_

COLLAR ELEV. 5290'TOTAL DEPTH 1125' BY \_\_\_\_\_START 8/4/93 COMPLETED 8/20/93Scale: 1" = 100'Est. of final depth 1125'



\* SHATTERED: &gt; 20 PLS./FT

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SHATTERED: > 20 PLS./FT.

Core Size	From	To	Interval	CORE RECOVERY		Max. Lgth. Ft.	UNITS PER FOOT			No. Pcs.	Ft. Grvl.
				Feet	%		FRACTURES				
							0-30°	30-60°	60-90°		
NC	460'	464'	42'	22'	50	022'	SHATTERED *				
XXWL	464'	470'	62'	62'	100	036'		"			
↑	470'	480'	102'	102'	100	033'		"			
	480'	490'	102'	102'	100	049'		"			
	490'	495'	52'	52'	100	040'		"			
	495'	504'	92'	92'	100	032'		"			
	504'	514'	102'	102'	100	043'		"			
	514'	521'	72'	72'	100	042'		"			
	521'	531'	102'	102'	100	032'		"			
	531'	538'	72'	72'	100	036'		"			
	538'	545'	72'	72'	100	041'		"			
	545'	548'	32'	32'	100	042'		"			
	548'	550'	22'	22'	100	043'		"			
	550'	552'	22'	12'	50	020'		"			
	552'	555'	32'	32'	100	032'		"			
	555'	557'	22'	22'	100	025'		"			
	557'	567'	102'	102'	100	038'		"			
	567'	572'	52'	52'	100	033'		"			
	572'	577'	52'	52'	100	030'		"			
	577'	583'	62'	62'	100	036'		"			
	583'	592'	92'	92'	100	032'		"			
	592'	599'	72'	72'	100	045'		"			
	599'	605'	62'	62'	100	030'		"			
	605'	612'	72'	72'	100	046'		"			
	612'	618'	62'	42'	66	040'		"			
	618'	623'	52'	52'	100	036'		"			
	623'	628'	52'	52'	100	036'		"			
	628'	635'	72'	72'	100	056'		"			
	635'	642'	72'	72'	100	041'		"			
	642'	651'	92'	92'	100	032'		"			
	651'	653'	22'	22'	100	024'		"			
	653'	658'	52'	52'	100	040'		"			
	658'	662'	42'	42'	100	035'		"			
	662'	667'	52'	52'	100	032'		"			
	667'	671'	42'	42'	100	032'		"			
Y	671'	673'	22'	22'	100	024'		"			
	673'	682'	92'	92'	100	042'		"			



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Core Size	From	To	Interval	CORE RECOVERY		Max. Lgth. Ft.	UNITS PER FOOT				
				Feet	%		FRACTURES			No. Pcs.	Ft. Grvl.
							0-30'	30-60'	60-90'		
NXW	682	687'	52'	52'	100	046'	SHATTERED				
	687	691'	42'	42'	"	035'		"			
	691	696'	52'	52'	"	039'		"			
	696	698'	22'	22'	"	027'		"			
	698	704'	62'	62'	"	041'		"			
	704	707'	32'	32'	"	036'		"			
	707	709'	22'	22'	"	025'		"			
	709	713'	45'	45'	"	032'		"			
	713	721'	75'	75'	"	027'		"			
	721	726'	52'	52'	"	026'		"			
	726	730'	42'	42'	"	031'		"			
	730	740'	102'	102'	"	032'		"			
	740	747'	72'	72'	"	031'	EXC. 742-747 1	8	0	31	
	747	755'	82'	82'	"	051'	5	8	20	65	
	755	764'	92'	92'	"	041'	14	9	26	78	
	764	774'	102'	102'	"	045'	9	12	25	83	
	774	784'	102'	102'	"	045'	4	15	12	97	
	784	790'	62'	62'	"	055'	SHATTERED				
	790	795'	52'	52'	"	061'		"			
	795	804'	92'	92'	"	053'		"			
	804	809'	52'	52'	"	051'		"			
	809	815'	62'	62'	"	062'		"			
	815	824'	92'	92'	"	053'		"			
	824	831'	72'	72'	"	041'		"			
	831	840'	92'	92'	"	062'		"			
	840	845'	52'	52'	"	052'		"			
	845	850'	52'	52'	"	042'		"			
	850	860'	102'	102'	"	061'	(2	" except 855-858' 7 4 16 48)			
	860	866'	62'	62'	"	042'	0	8	9	43	
	866	869'	32'	32'	"	065'	SHATTERED				
	869	871'	22'	22'	"	031'		"			
	871	872'	12'	033'	33	011'		"			
	872	875'	32'	32'	100	032'		"			
	875	876'	12'	12'	"	017'		"			
	876	880'	42'	42'	"	032'		"			
	880	882'	22'	22'	"	025'		"			
	882	883'	12'	033'	33	011'		"			



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Core Size	From	To	Interval	CORE RECOVERY		Max. Lgth. Ft.	UNITS PER FOOT			No. Pcs.	Ft. Grvl.
				Feet	%		FRACTURES				
							0-30°	30-60°	60-90°		
NYWL	883	885'	22'	22'	100	032'	SHATTERED				
↑	885	890'	50'	50'	"	032'		"			
	890	893'	30'	30'	"	028'		"			
	893	898'	50'	50'	"	036'		"			
	898	902'	40'	40'	"	035'		"			
	902	907'	50'	50'	"	029'		"			
	907	912'	50'	50'	"	028'		"			
	912	917'	50'	10'	20	028'		"			
	917	919'	20'	20'	100	030'		"			
	919	924'	50'	50'	"	030'		"			
	924	929'	50'	50'	"	036'		"			
	929	932'	30'	30'	"	030'		"			
	932	937'	50'	50'	"	032'		"			
	937	942'	50'	50'	"	029'		"			
	942	947'	50'	50'	"	029'		"			
	947	953'	60'	60'	"	022'		"			
	953	958'	50'	50'	"	018'		"			
	958	963'	50'	50'	"	051'		"			
	963	967½'	45'	45'	"	048'		"			
	967½	972½'	50'	50'	"	061'		"			
	972½	977'	45'	45'	"	061'		"			
	977	982	50'	50'	"	041'		"			
	982	987	50'	50'	"	045'		"			
	987	992	50'	50'	"	038'		"			
	992	997	50'	50'	"	021'		"			
	997	1002	50'	50'	"	043'		"			
	1002	1007'	50'	50'	"	045'		"			
	1007	1012'	50'	50'	"	032'		"			
	1012	1017'	50'	50'	"	038'		"			
	1017	1020	30'	30'	"	029'		"			
	1020	1025'	50'	50'	"	021'		"			
	1025	1030'	50'	50'	"	044'		"			
	1030	1035'	50'	50'	"	026'		"			
	1035	1040'	50'	50'	"	013'		"			
	1040	1044'	40'	40'	"	016'		"			
↓	1044	1049'	50'	50'	"	030'		"			
	1049	1055'	60'	60'	"	040'		"			



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Core Size	From	To	Interval	CORE RECOVERY		Max. Lgth. Ft.	UNITS PER FOOT			No. Pcs.	Ft. Grvl.
				Feet	%		FRACTURES				
							0-30°	30-60°	60-90°		
NXWL	1055	1062	72'	72'	100	041'	SHATTERED				
↑	1062	1066	42'	40'	"	020'		"			
	1066	1071	52'	32'	60	034'		"			
	1071	1075	42'	42'	100	021'		"			
	1075	1080	52'	52'	"	019'		"			
	1080	1084	42'	42'	"	032'		"			
	1084	1091	72'	72'	"	032'	12	5	4	71	
	1091	1098	72'	72'	"	041'	13	10	8	62	
	1098	1100	22'	22'	"	040'	17	14	7	50	
	1100	1103	32'	32'	"	041'	8	3	2	26	
	1103	1107	42'	42'	"	052'	7	2	9	35	
	1107	1112	52'	52'	"	042'	6	4	11	46	
	1112	1117	52'	52'	"	025'	SHATTERED				
	1117	1122	52'	52'	"	036'		"			
	1122	1127	52'	52'	"	034'		"			
	1127	1130	32'	32'	"	033'		"			
	1130	1135	52'	52'	"	028'		"			
	1135	1138	32'	32'	"	045'		"			
	1138	1141	32'	32'	"	040'		"			
	1141	1146	52'	52'	"	028'		"			
	1146	1151	52'	52'	"	041'		"			
	1151	1156	52'	52'	"	030'	4	16	13	53	
	1156	1160	42'	42'	"	046'	SHATTERED				
	1160	1165	52'	52'	"	039'		"			
	1165	1169	42'	42'	"	035'		"			
	1169	1174	52'	52'	"	042'		"			
	1174	1179	52'	52'	"	033'		"			
	1179	1184	52'	52'	"	041'	23	9	6	62	
	1184	1189	52'	52'	"	041'	21	11	7	69	
	1189	1194	52'	52'	"	061'	SHATTERED				
	1194	1198	42'	42'	"	032'		"			
	1198	1203	52'	52'	"	033'		"			
	1203	1208	52'	52'	"	042'		"			
	1208	1210	22'	22'	"	034'		"			
	1210	1215	52'	52'	"	040'		"			
	1215	1220	52'	52'	"	041'	9	5	1	11	
Y	1220	1225	52'	52'	"	044'	8	6	4	22	



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Core Size	From	To	Interval	CORE RECOVERY		Max. Lgth. Ft.	UNITS PER FOOT				No. Pcs.	Ft. Grvl.
				Feet	%		FRACTURES					
							0-30°	30-60°	60-90°			
NXWL	1225	1230	52'	52'	100	043'	10	6	4	19		
↑	1230	1240	102'	102'	100	056'	9	13	8	36		
	1240	1250	102'	102'	"	054'	13	6	12	52		
	1250	1260	102'	102'	"	041'	8	9	10	46		
	1260	1266	62'	62'	"	028'	SHATTERED					
	1266	1275	92'	92'	"	042'		"				
	1275	1282	72'	62'	86	042'		"				
	1282	1292	102'	102'	100	045'		"				
	1292	1302	102'	102'	"	054'		"				
	1302	1311	92'	92'	"	038'		"				
	1311	1321	102'	102'	"	028'		"				
	1321	1331	102'	102'	"	029'		"				
	1331	1341	102'	102'	"	040'		"				
	1341	1351	102'	102'	"	036'		"				
	1351	1361	102'	102'	"	036'		"				
	1361	1371	102'	102'	"	045'		"				
	1371	1380	92'	92'	"	042'		"				
	1380	1390	102'	102'	"	043'		"				
	1390	1400	102'	102'	"	036'		"				
	1400	1410	102'	102'	"	032'		"				
	1410	1420	102'	102'	"	032'		"				
	1420	1425	52'	52'	"	029'		"				
	1425	1435	102'	102'	"	035'		"				
	1435	1445	102'	102'	"	040'		"				
	1445	1455	102'	102'	"	050'		"				
	1455	1465	102'	102'	"	042'		"				
	1465	1475	102'	102'	"	028'		"				
	1475	1485	102'	102'	"	037'		"				
	1485	1495	102'	102'	"	041'		"				
	1495	1505	102'	102'	"	045'		"				
	1505	1513	82'	82'	"	048'		"				
	1513	1523	102'	102'	"	051'		"				
	1523	1533	102'	102'	"	041'		"				
	1533	1539	62'	62'	"	040'		"				
	1539	1549	102'	102'	"	037'		"				
	1549	1559	102'	102'	"	032'		"				
Y	1559	1565	62'	62'	"	046'						







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SHATTERED : &gt; 20 PCS./FT.

Core Size	From	To	Interval	CORE RECOVERY		Max. Lgh. Ft.	UNITS PER FOOT				No. Pcs.	Ft. Grvl.
				Feet	%		FRACTURES					
							0-30	30-60	60-90			
NXWL	90'	92'	2'	2'	100	0.30	SHATTERED	*				
↑	92'	100'	7'	7'	100	0.35	"					
	100'	105'	5'	5'	100	0.20'	"					
	105'	110'	5'	5'	100	0.25	"					
	110'	115'	5'	5'	100	0.30	"					
	115'	118'	3'	3'	100	0.40	"					
	118'	121'	3'	3'	100	0.20	"					
	121'	124'	3'	3'	100	0.30	"					
	124'	127'	2'	2'	100	0.15	10	16	14	49		
	127'	132'	5'	3'	78	0.10	SHATTERED					
	132'	140'	8'	7'	94	0.20	2	9	21	51		
	140'	144'	4'	3'	75	0.22	SHATTERED					
	144'	148'	4'	4'	100	0.15	"					
	148'	155'	7'	5'	80	0.20	"					
	155'	160'	5'	4'	90	0.50	"					
	160'	168'	8'	8'	100	0.11	0	14	31	106		
	168'	172'	4'	4'	100	0.11	SHATTERED					
172'	179'	7'	7'	100	0.51	"						
179'	186'	7'	6'	86	0.16	"						
186'	193'	7'	7'	100	0.30	"						
193'	201'	8'	8'	100	0.60	"						
201'	211'	10'	10'	100	0.30	"						
211'	221	9'	9'	100	0.30	"						
221	231	10'	10'	100	0.41	"						
231	235	10'	10'	100	0.20	"						
235	241	6'	6'	100	0.51	"						
241	251	10'	10'	100	0.30	"						
251	259	8'	8'	100	0.12	"						
259	269	10'	10'	100	0.20	"						
269	275	6'	6'	100	0.20	"						
275	285	10'	10'	100	0.12	"						
285	295	10'	10'	100	0.15	"						
295	305	10'	10'	100	0.50	"						
305	312	7'	7'	100	0.11	"						
312	315	3'	3'	100	0.16	"						
↓	315	325	10'	10'	100	0.61	5	7	27	62		
	325	335	10'	10'	100	0.30	SHATTERED					



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Core Size	From	To	Interval	CORE RECOVERY		Max. Lgth. Ft.	UNITS PER FOOT			No. Pcs.	Ft. Grvl.
				Feet	%		FRACTURES				
							0-30	30-60	60-90		
NXWL	335'	345'	102'	102'	100	016'	SHATTERED				
▲	345'	355'	102'	102'	100	033'		"			
	355'	365'	102'	102'	100	021'		"			
	365'	373'	82'	82'	100	036'		"			
	373'	383'	102'	102'	100	023'		"			
	383'	388'	52'	52'	100	025'		"			
	388'	395'	72'	22'	28.6	022'		"			
	395'	397'	22'	22'	100	020'		"			
	397'	405'	82'	82	100	032'		"			
	405'	408'	32'	32'	100	031'		"			
	408'	415'	72'	72'	100	021'		"			
	415'	425'	102'	102'	100	025'		"			
	425'	435'	102'	102'	100	040'		"			
	435'	443'	82'	82'	100	031'		"			
	443'	452'	92'	92'	100	021'		"			
	452'	459'	72'	72'	100	021'		"			
	459'	466'	72'	72'	100	052'		"			
	466'	471'	52'	42'	80	031'		"			
	471'	476'	52'	35'	70	031'		"			
	476'	485'	92'	92'	100	022'		"			
	485'	493'	82'	42'	50	031'		"			
	493'	496'	32'	32'	100	032'		"			
	496'	505'	92'	92'	100	026'		"			
	505'	512'	72'	72'	100	023'		"			
	512'	521'	92'	92'	100	031'		"			
	521'	529'	82'	82'	100	022'		"			
	529'	536'	72'	72'	100	032'		"			
	536'	542'	62'	62'	100	031'		"			
	542'	551'	92'	92'	100	041'		"			
	551'	560'	92'	92'	100	040'		"			
Y	560'	566'	62'	62'	100	022'		"			



BLP = fbbhp on map

Drill Hole No. 133W 7 Elevation \_\_\_\_\_ Inclination \_\_\_\_\_ Page No. 1 of \_\_\_\_\_  
 Collar Coords. \_\_\_\_\_ Bearing \_\_\_\_\_ Logged by: Quack (see memo) Date \_\_\_\_\_

Mineralization					Core Rec.	Ground Cond.	Geology	Analyses		
					464		464 - 542 bxa - volcanic? textures difficult to see because of very strong Qtz-sericite			
					500		alteration various rock types present but all look volcanic to me.			
					20		strong pyrite: 5-10% ore ~ 6-7% (dissom 90%)			
					45		542-562 <sup>MVB silicified</sup> extreme silicification of a Vfg rock - prob a felsite <sup>breccia</sup> TS 7-8% py			
					60		562-626 MVB bxa frags very distinct 577-579 - still look like			
					80		577-579 (very ground up - possibly faulted) a volcanic matrix to me.			
					600		high sulfide - Androm log seems good as Est. M. strong sil-ser.			
					20		Note Selective purification of frags.			
					40		626-653 MVB still bxa but in a foliated volcanic matrix - looks like <u>dry bone</u> .			
					60		<u>bone like</u> extreme sericitization here			
					80		650-689 ground up - poss. fault. - intermittent unbroken zone			
					700		(694-707) definite fault gouge - 707-740 + highly broken - usually only small pieces remain			
					20		653-732 MVB - as before			
					40		740-784 MV bxa tapering off until no more frags readily visible at 741			
					60		740-784 - greenish - otherwise same			
					80		732-784 - Dark brown - grainy texture - almost silty or sandy volcanic - w/ <u>thre</u> tuffs			
					100		a few remnants but 90% of py is diss.			
					130		784-803 - bxa on upper margin similar to the BP of BEMC in 133W 1 - but brit matly gone - also like B1P in Blue 1+2			
					800		HA-MV 784-803 HA hornblende andesite (BEMC = HBDP) still a volcanic to me.			
					130		803-809 fault Qtz - py (+black mineral) veins increasing - mostly vertical upy seen.			
					80		809-813 MVB chl alt			
							813-816 B1P dike - act bxa - chloritized			
							816-826 MVB			

(33) Item 15  
 Strong ser  
 Mod ser  
 chl



Drill Hole No. 17

Elevation \_\_\_\_\_

Inclination \_\_\_\_\_

Page No. 2 of \_\_\_\_\_

Collar Coords. \_\_\_\_\_

Bearing \_\_\_\_\_

Logged by: \_\_\_\_\_

Date \_\_\_\_\_

Mineralization	Core Rec.	Ground Cond.	Geology	Analyses
	80		826-842 - dike BLP	
	80		same as BSW 1 DP at 756	
	60		842-872 MVB(AH) mixed	
	80		872-890 Fault BXA - some competent pieces	
	80		890-904 MVB strongly broken - v. small frags - no arg or gouge so prob not a fault piece.	
	80		904-908 Sericitized rock - prob. in fault zone - crushed. white - w/ a volc.??	
	20		908-917 MV - as before (re some chl some ser veins & weak stockwork of very	
	20		917-936 faulted MV or MVB - clay matrix indicating some fault gouge.	
	40		936-940 MV-MVB	
	40		940-945 Fault gouge	
	40		945- MV dark to light green - chl some of very apparent foliation ~ 45% ka	
	80		952-953 Small qtz-ser alt dike - a porphyry - similar to BLP but may not be same.	
	1000		958-1000 - fault zone - high silica low py	
	20		1002-1039 faulted - intermittently Sericitized, small islands of unbroken material - slides etc seen	
	40		1039-1096 fault gouge - clay + ser - white	
	60		gouge	
	80		1096-1100 MVB sericitized & silicified. some large qtz veins - veins often broken therefore	
	100		1100-1107 - Fm porphyritic andesite 1 mm phenocrysts of plagioclase in brown aphan. gneiss.	
	20		1107-1149 MVB - intermittently re-brecciated - some gouge in places	
	40			
	60		1149-1166 MVB - foliated ~ 55° - also broken but less than preceding zone.	
	80		possibly HA foliated?	
	80		1166-1175 stockwork veins & altered rock - HA? - HA & FPA maybe same early w/ different chl	
	1200		1175-1191 Good HA grading into good FPA at 1191 but HA has lots of brot that stand out	
			Some qtz veins - s	
			1191-1200 - FPA more veins in IPA than Ha - perhaps they are composite & transitional.	



Drill Hole No. 1

Elevation \_\_\_\_\_

Inclination \_\_\_\_\_

Page No. 3 of       

Collar Coords. \_\_\_\_\_

Bearing\_\_\_\_\_

Logged by: \_\_\_\_\_

Date \_\_\_\_\_

[illegible]



SCALE \_\_\_\_\_  
START \_\_\_\_\_  
REMARKS \_\_\_\_\_

LOGGED BY \_\_\_\_\_  
COMPLETION \_\_\_\_\_

COORDINATES \_\_\_\_\_  
DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
BEARING \_\_\_\_\_

PROJECT \_\_\_\_\_  
ELEVATION \_\_\_\_\_

Page \_\_\_\_\_ of \_\_\_\_\_  
DRILL HOLE # BSW 7  
TOTAL DEPTH 1715  
INTERVAL 1240 to 1373

ASSAYS  
BCL -

LITH.

GEOLOGY

Rock Type, Alteration, Mineralization, Structure

ALTERATION

PERVASIVE VEINS

Script clay Q Calc Gyp

1240-1261 FPA Fine porphyritic Andesite - normally dark brown w/ fsp xtal 1/2 mm ~10-15%  
4-10% biotite clumps or bands - up to 3mm - but w/ strong alt.  
Alt - strong new ser in & very strong Qtz veining - strong stockwork 8"/inch  
Minz veins - strong - py & Qtz py + Qtz hematopy  
TS 3-4%

1261-1266 Fault Gouge? - strong ser - small pieces - no solids core -

1266-1289 BIP Broken fsp - clay + mafic - highly altered - ser & clay - (Some high pressure)  
whitish - greenish when wet  
some Qtz veining - 2-3/foot 1/2 to 1"

1289-1301 BIP? Brownish - tan - also a little finer grained + more competent -  
Alt - ser - perv. no clay  
TS ~ 1% some 1/2" veins - ie 1287  
this rock cuts the stockwork in the Fine grained volcanic - (or HAZ) (or FPA)  
it is also cut by later Qtz veins  
NO - see 1362

1301-1310 MV? Strong Qtz vein stockwork in fine grained volcanic? cut by BIP above  
many barren Qtz veins cut & offset by later - steep Qtz - py veins  
TS - 2-3% stockwork is early - may be answer w/ the emplacement of the MV or the where it is.

1310-1315 BIP(?) - prob BIP 1312 but some strange silicification that is extremely selective  
& it occurs bound on a side by a Qtz vein is very confusing - almost looks  
like another rock for the color is so different - but it basically is a 1/2" calc vein  
stockwork Qtz from 1312 to fault at 1315 - earlier (1301) it clearly post dates the stockwork.  
(is this an older look alike?)

1323-1361 BIP - as in BSW1  
Alt - Strong ser to chl + wk ser - (remnant chl?)  
strong mag in BIP in weaker alt zones.  
TS - 1-2% - veins - maybe 2-3% in - stronger serite zones. w/ oncompy also.

1361-1368 MV - stockwork Qtz Fine grained stockworked volcanic? or mafic - stockwork has been refaulted - offset veins

1368-1373 BIP - as before - strong ser

Strong Mod  
stockwork

Weak Veining

Strong  
stockwork

Weak  
stockwork

Strong  
stockwork



SCALE \_\_\_\_\_  
START \_\_\_\_\_  
REMARKS \_\_\_\_\_

LOGGED BY \_\_\_\_\_  
COMPLETION \_\_\_\_\_

COORDINATES \_\_\_\_\_  
DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
BEARING \_\_\_\_\_

DIP \_\_\_\_\_

PROJECT \_\_\_\_\_  
ELEVATION \_\_\_\_\_

Page    of     
DRILL HOLE # BSW 7  
TOTAL DEPTH \_\_\_\_\_  
INTERVAL 1373 to 1510

ASSAYS

LITH.

GEOLOGY

Rock Type, Alteration, Mineralization, Structure

ALTERATION

PERVASIVE VEIN

Ser chd clay S-1 S-2 S-3 S-4 S-5 S-6 S-7 S-8 S-9 S-10

1373-1380 MU - stockwork qtz in ~~ophanitic~~ volcanics cut by FPA at 1380 strong ser  
1380-1382 FPA - FPA dikelet clearly cuts volcanics & stockwork - but is offset itself  
by many small flat fractures (w/ hem).  
1382-1410? MU - stockwork - some silicification as well as seradhesion  
TS - 1-2%

1410 - ~1457' FPA brown turny bleached at 1444 losing some texture - argl?  
some veining - late effects by small faults - a fraction to several  
inches displacement TS  $\leq 1\%$  v.  
very broken w/ slicks from 47 to 57

FPA brown some vein some chl

1450 broken  
FPA  
1450-1459? bleached  
1459-1480 Fault? zone & BKA - strong ser TS  $\leq 1\%$

1480-1510 PBL Porphyritic biotite latite - could be phase of BLP  
biotite phenos - weak foliation 20-30° to CA, mag present  
strong ser to 1496 - then weaker ser + arg. of plag. then weaker ser to 1508  
TS ~ 1-2% stronger cp4 visible up at 1511



Page      of       
 DRILL HOLE # B5W 7  
 TOTAL DEPTH       
 INTERVAL 1510 to 1650

ASSAYS					LITH.	GEOLOGY Rock Type, Alteration, Mineralization, Structure	ALTERATION						
							PERVASIVE				VEIN		
							Ser	chl	clay		SI	Q	Chl
					10	1510-1553 PBL as before - grey - brot essentially intact - some wk chl TS - 2 1/2 lf 2mm zone V. small py fractures 1556 1/2" Qtz py-cpy vein (flat - 20° bit) + strong qyp	W	W				V	S
					520	PBL							S
					30	chil' zone at 1553 2" but cant see contact w/ andesite (MV?)							S
1-2					40		S						S
					1550	1533-1570 Mva: Andesite - vfg w/ 4-7% 2-3mm py + cpy	W					V	S
					60		W						S
4-7					70	1590-1572 Bxa - Mvb or fault? no good core to see							S
8					80	1572-1592 PBL as before grey - a few qb veins + strong qyp						W	S
					90	can't see contact w/ BLP - Robbed?							S
2-4					1000	1592-1650 - BLP same as before in this hole - but note a difference between PLP of BSW 7 + BLP of BSW 1 at 456' - there is a slight external difference and there are the rare qb phenos present in BSW 7. Py - 20-25% 2-7mm black brot 1-2mm - 4-6% (weakly chloritized) mag present Alt - weakly chloritized - mostly black brot + shiny plag - rare qb phenos							M
1-3					10	Mm2 = 1-3% py - druse + on hairline fractures 1631 - 1/4" qb-py veinlet w/ bleached ser halo.							M
					20								M
1-2					30								M
					40								M
					50								M
					60								M
					70								M
					80								M
					90								M
					1000								M



SCALE \_\_\_\_\_  
START \_\_\_\_\_  
REMARKS \_\_\_\_\_

LOGGED BY \_\_\_\_\_  
COMPLETION \_\_\_\_\_

COORDINATES \_\_\_\_\_  
DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
BEARING \_\_\_\_\_  
DIP \_\_\_\_\_

PROJECT \_\_\_\_\_  
ELEVATION \_\_\_\_\_

Page \_\_\_\_\_ of \_\_\_\_\_  
DRILL HOLE # BSW 7  
TOTAL DEPTH 1715  
INTERVAL 1650 to 1715

ASSAYS

LITH.

GEOLOGY

Rock Type, Alteration, Mineralization, Structure

ALTERATION

PERVASIVE VEIN

1-2

4-7

1

50

60

70

80

90

1700

steep & uneven contact

1650-1698

Blp as before - partly

1698-1712

Mva & MVB

1710-1712

dark brown mostly w/ 4-7% dolomite

one or two remnants of Blp (-2")

1712-1715

TD-Bottom

Blp as before.



chlorite, sericite and quartz. Chlorite-quartz-biotite alteration is most often associated with relatively higher copper grades; chlorite-quartz-biotitized dacite porphyry between 820 feet and 907.7 feet, for instance, averages 0.35%.

Drill hole BSW-8 (TD 566 feet), was drilled to test a small, intense VIP anomaly for an hypothetical, near-surface, high-grade chalcocite enrichment blanket. The hole provided an explanation for the anomaly--but failed to find the anticipated ore body--in penetrating 92.8 feet of oxidized and quartz-sericitized quartz latite porphyry and quartz-sericite rock, 20 feet of enriched, quartz-sericitized intrusion breccia, averaging 0.3% Cu, and 456 feet of quartz-sericitized intrusion breccia and quartz-feldspar porphyry averaging about 10 volume percent pyrite, and 0.05% Cu as chalcopyrite and cupriferous pyrite.

BSW 7

alt — Drilling, geologic mapping, geochemical sampling and induced polarization have shown Buckskin Southwest to be a tight porphyry copper sulfide system with abnormally steep alteration and mineralization gradients. The system has been penetrated by only one deep hole (BSW-7), which bored through a 510 foot zone averaging nearly 0.2% Cu. Alteration within this zone is highly complex, alternating between varieties of quartz-biotitization and quartz-sericitization. Such an alteration pattern suggests that BSW-7 penetrated the transition zone between the potassic core of the system and its quartz-sericite halo. The unusually steep mineralization gradients suggest that ore grade copper mineralization might occur in the potassic core within a few hundred feet laterally of BSW-7.

An hypothetical Buckskin Southwest ore body would be relatively small--100 to 200 million tons--but high grade, probably in the 0.8 to 1% range. The deposit would most likely be situated north, northwest, or northeast of BSW-7, and would be concealed beneath 750 feet to 1,000 feet of barren quartz-sericite rock and/or postmineral cover.

## RECOMMENDATIONS

The writer recommends deepening two shallow drill holes completed in 1970--BSW-1 (TD 292 feet) and BSW-2 (TD 200 feet)--to 2,000 feet. Completion of these two proposed drill holes will provide a definitive test of the target hypothesis outlined under "Summary and Conclusions".














(35)  
Item 15

J. B. HULEN 6/13/72



# EXPLANATION BSW-7 & 8

(35)  
Item 15

LITHOLOGY	STRUCTURE
 HORNBLENDE BIOTITE	 FAULT ZONE
 DACITE PORPHYRY (Post-mineral Leach 1592' in BSW-7)	 MINOR FAULT
 QUARTZ-FELDSPAR PORPHYRY (ALTERATION PRECLUDES DETER- MINATION OF ORIGINAL COMP.)	 STOCKWORK FRACTURING
 QUARTZ-SERICITE ROCK	 QUARTZ-CALCITE VEIN
 INTRUSION BRECCIA	 QUARTZ VEIN
 BIOTITE HORNBLENDE QUARTZ LATITE PORPHYRY	

## EXPLANATION OF GRAPHIC ALTERATION LOG

- ☐ CHLORITE-QUARTZ-BIOTITE Complete alteration of original rock matrix to fine crystalline aggregate of dark brown chlorite and dark grayish-green biotite, moderate to intense stockwork quartz veining. Vein widths range from 1 mm to 10 cm, and 2 mm, in which moderate sericite is seen to the quartz minor sericite calcite, epidote hematite and pyrite pyrrhotite may be locally present.
- ☐ QUARTZ-BIOTITE Same as chlorite-quartz-biotite except minor to moderate secondary chlorite.
- ☐ SERICITE-QUARTZ-BIOTITE Same as chlorite-quartz-biotite, except no chlorite; moderately abundant, fine crystalline light gray sericite texturally similar to the biotite and chlorite of chlorite-quartz-biotite alteration.
- ☐ QUARTZ-SERICITE All original minerals replaced with a fine crystalline quartz-sericite aggregate, weak to intense stockwork quartz veining, epidote hematite and pyrite pyrrhotite locally present.
- ☐ CHLORITE-QUARTZ-SERICITE Same as quartz-sericite, except moderately abundant, light-dark grayish-green chlorite.
- ☐ CLAY-SERICITE Original rock feldspar partially-completely altered to very fine crystalline, light gray to white clay-sericite aggregates; original matrix minerals partially completely altered to light-dark grayish-green chlorite occasionally, with minor clay and/or sericite, minor epidote hematite and pyrite pyrrhotite locally present.
- ☐ PROPYLITIC Original rock matrix minerals altered to fine crystalline, light-dark grayish-green chlorite; original rock feldspar partially altered to calcite and light, mixed with clay and sericite along cleavage planes; minor epidote and/or quartz locally present.

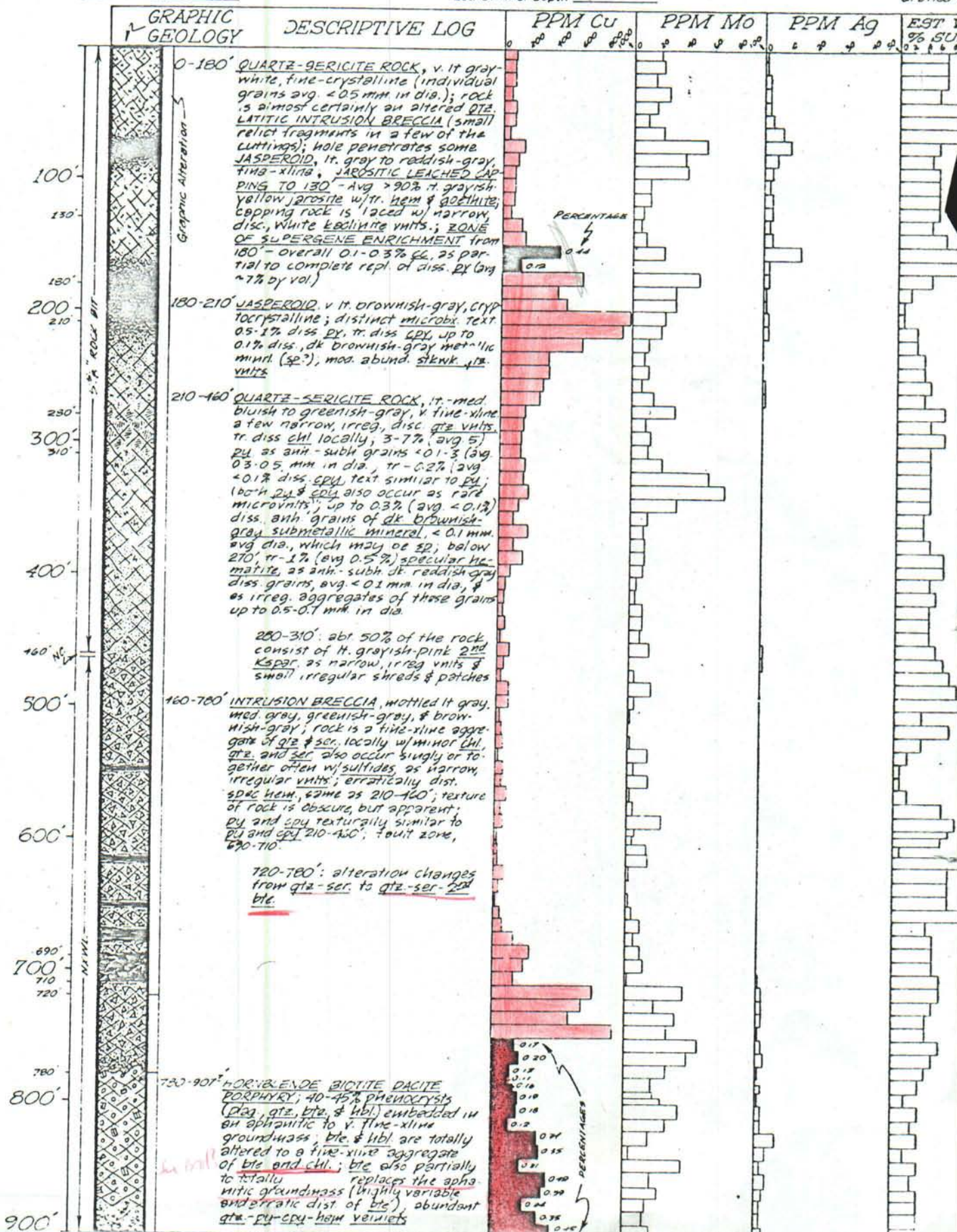


## Summary Drill Hole Log

Code 05-01-0411BEARING VERTICAL DIP 0Purpose of hole TEST OF DEEP, PRIMARY, PORPHYRY COPPER  
ORE TARGETLOCATION: N. 10,000' E. 10,000'COLLAR ELEV. APP 5440'TOTAL DEPTH 1715' BY J.P.START 4/25/72 COMPLETED 6Scale: 1" = 100'

Est. of final depth \_\_\_\_\_

drafted \_\_\_\_\_





DISTRICT Name BUCKSKIN SOUTHWESTDRILL HOLE NO. BSW-7 PAGE 1 OF 2Hole Log Code 05-01-0411BEARING VERTICAL DIP W

\* NOTE: BLACKED-IN INTER

EST OF DEEP, PRIMARY, PORPHYRY COPPER

LOCATION: N. 10,000' E. 10,000' (Approx)

VALS INDICATE 10X

COLLAR ELEV. APP 5440'

INCREASE IN GRAPHED

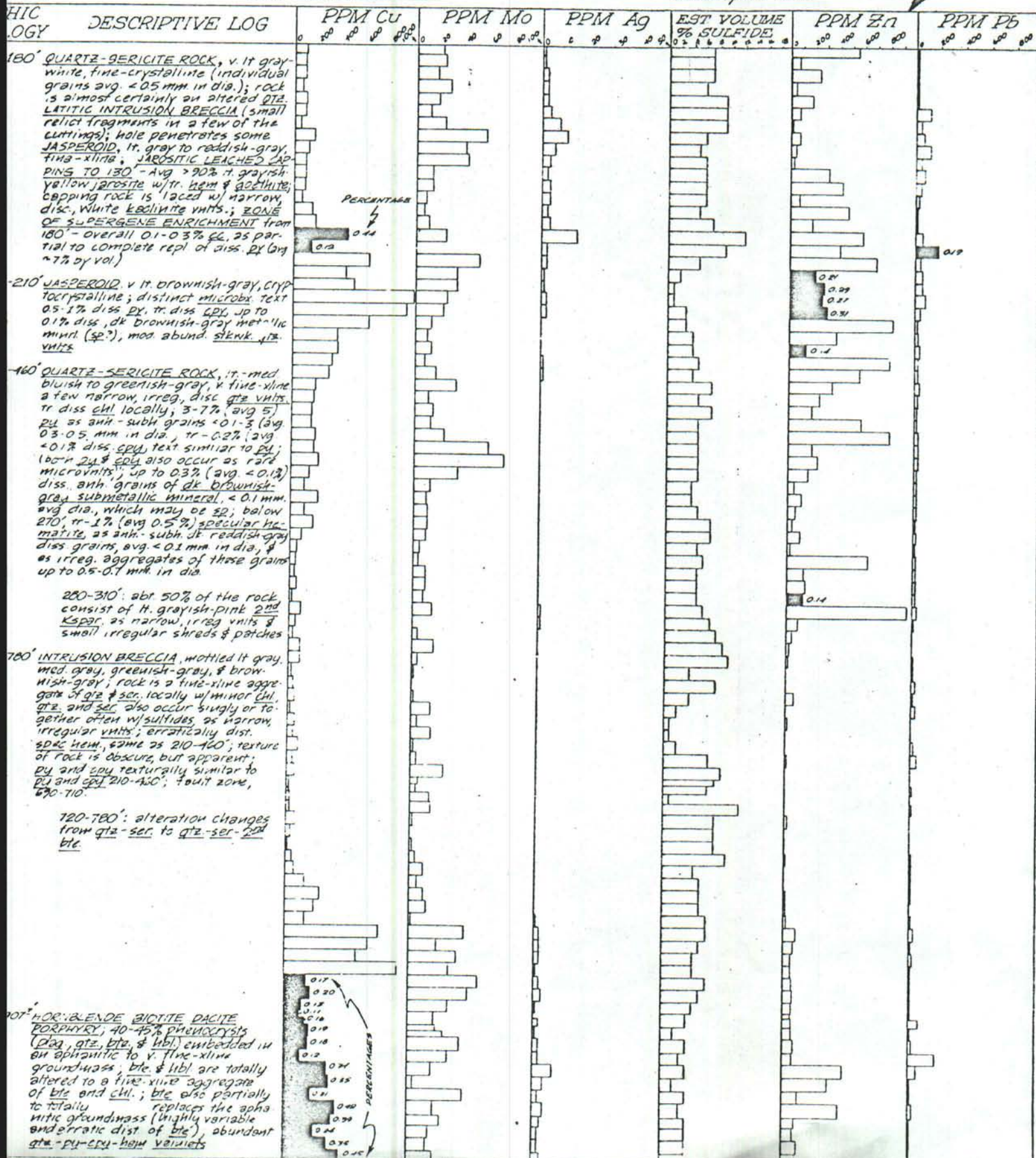
TOTAL DEPTH 1715' BY J.B. Hulen

VALUES

START 4/25/72 COMPLETED 6/25/72

Est. of final depth

drafted by J.B. Hulen





## Summary Drill Hole Log

Code 05-01-0-111

Purpose of hole SEE PAGE 1 \*BEARING            DIP           

LOCATION: N.            E.           

COLLAR ELEV. \_\_\_\_\_

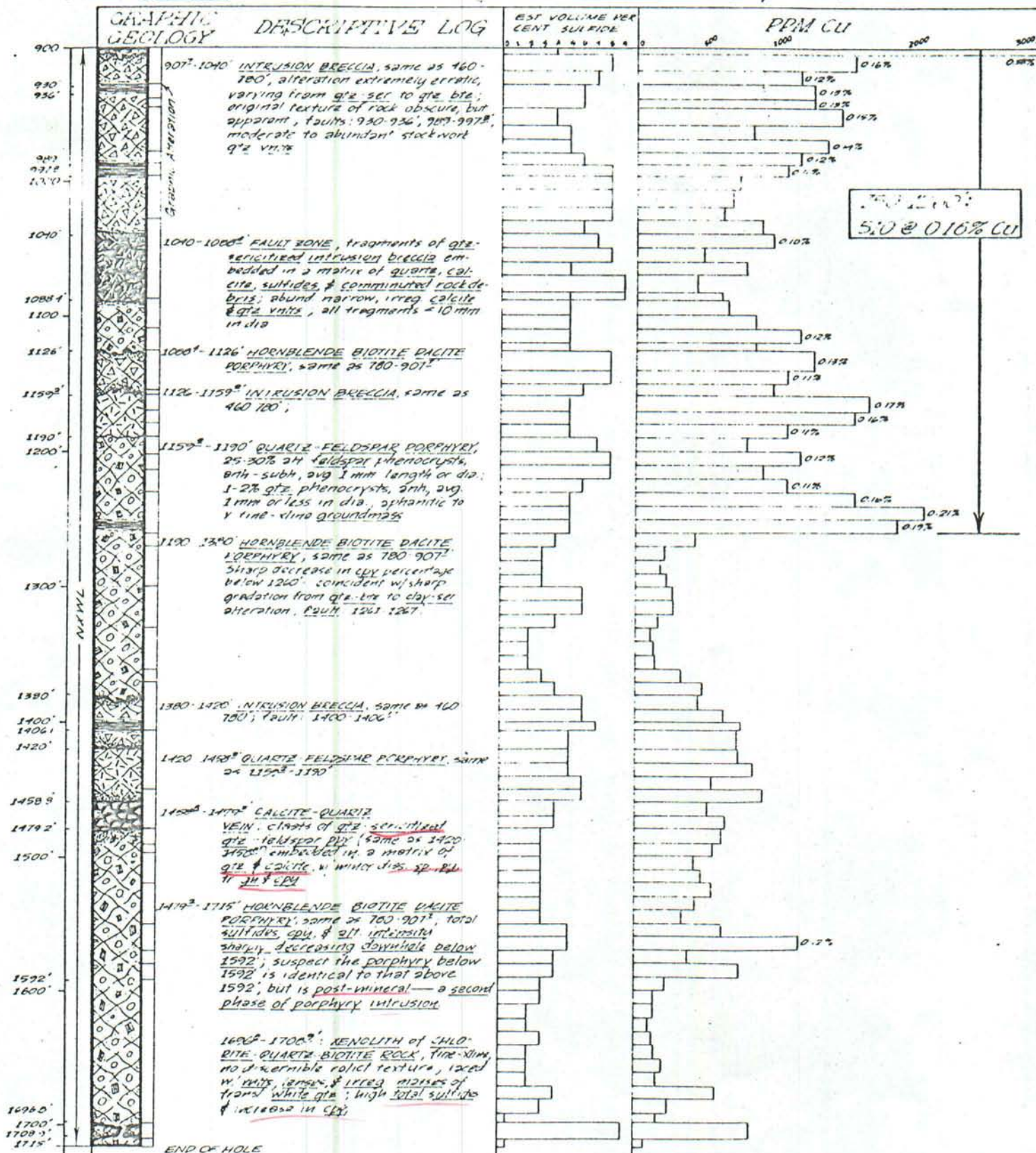
TOTAL DEPTH 1715' BY J.B. Huler

START  COMPLETED 

drafted by J.B. Hulien

Scale: 1" = 100'

Est. of final depth





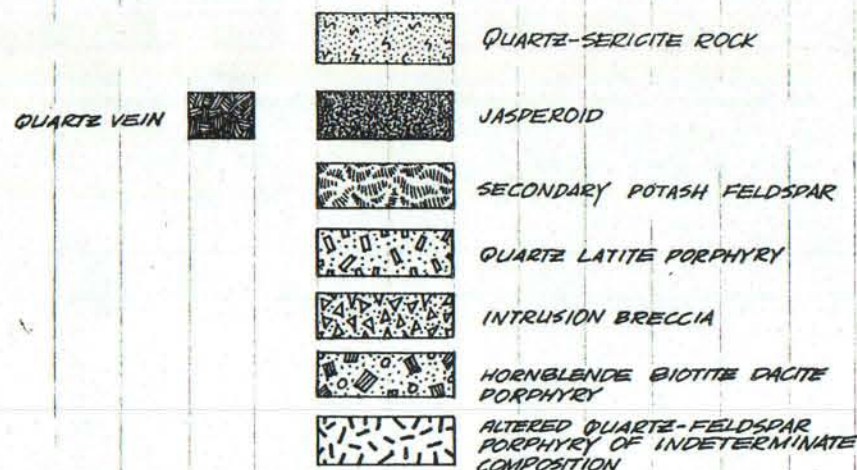
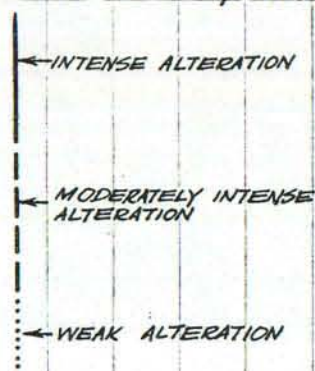
Logged by  
JEFFREY B. HULENBUCKSKIN SOUTHWEST EXAMINATION  
DOUGLAS COUNTY, NEVADA

DRILL HOLE BSW-7 &amp; BSW-8

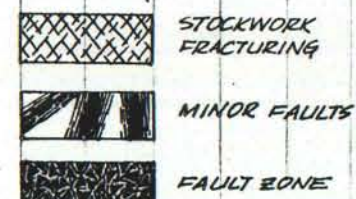
## -EXPLANATION-

## GRAPHIC GEOLOGY

## Rock Types

Alteration Mineralogy:  
Alteration Intensity Graphs

## Structure



## BSW-7: DESCRIPTION OF LITHOLOGY, STRUCTURE, MINERALIZATION, AND ALTERATION IN IMMEDIATE VICINITY OF DRILL SITE

Quartz-sericitized INTRUSION BRECCIA, oxidized, formerly sulfide-bearing; color varies from lt. yellowish gray to lt.-med. reddish gray; original texture is largely obscured through alteration, but essentially 25-30% clasts, angular to subrounded, embedded in a finely porphyritic (or fragmental) matrix; clasts range in size from 2-3 mm. to at least 50 mm. in dia. and range in texture (relict) from ophanitic to porphyritic; rock is now a fine-xline aggregate of qtz, ser., and clay (prob. kaolinite) — all three minerals also occur singly or together as narrow, irregular, discontinuous veinlets; small, irregular pods of dense, lt. yellowish to reddish-gray jasperoid are locally present; small (up to 25 mm. in dia), irregular vugs in the jasperoid are commonly lined with minute, water-clear, euhedral quartz crystals; rock has oxidized to form a jarositic leached capping; jarosite occurs as 1) diss. resinous, lt. yellowish to orange brown grains & grain aggregates partially to totally filling former sulfide sites — which range in size from < 0.1 mm. to 5 mm. (avg. 0.5 mm.) in dia. & 2) irreg. discontinuous veinlets, lt. yellowish-brown, up to 15 mm. (avg. 1 mm.) in dia, filling stockwork fractures; distance between these fractures/veinlets averages about 50 mm.; a little lt. reddish to purplish hematite is locally present w/the jarosite; a little diss. fine-xline pyrite, commonly w/chalcocite microcrins, is locally present in the jasperoid pods;

Est former sulfide content in vicinity of drill site avg. about 1%; a little lt. grayish-blue turquoise, powdery, is locally present as narrow, discontinuous veinlets, commonly with jarosite; one small (~ 6" wide) dike of hornblende biotite dacite porphyry at the drill site; this porphyry is lt. brownish-gray, clay-sericitized, w/about 4% diss. goethite.

CONT'D



## BEAR CREEK MINING COMPANY - NORTHWEST DISTRICT

CODE NO.: 05-01-0411

Started April 25, 1972 Completed June 6, 1972 Project Buckskin Southwest Hole No. BSW-7 Page 1 of 13  
 Coords: N 1 E 1 Collar Elev. App 5440' Bearing 1 Vertical Angle 90° Total Depth  
 Descriptive Location: Approximately 1900' N29°E of the SE corner of section 11, T.13 N., R.23 E., southwestern Buckskin Range, Douglas Co., Nevada  
 Objective: Deep test of hypothetical primary porphyry copper ore target, allegedly concealed beneath 1000-1500' of cupriferous quartz-sericite rock  
 Stopped Because \_\_\_\_\_ Scale: 1" = 10' (0-60') Logged By: J.B. Hulen  
1" = 20' (60-70')

f \* former sulfide content

Core Recovery		Graphic Alteration	Graphic Mineral (%)	Vol. % Sulf. Est.	GEOLOGY		Thin Section Location & Description	Analyses							Sample Nos.
Size Interval	% Rec				Interval	Description		Interval	PPM Cu	PPM Mo	PPM Pb	PPM Zn	PPM Ag	PPM Au	
0-10'	5% RBZ	not applicable		1-6 *	0-10'	Oxidized, formerly pyritic QUARTZ-SERICITE ROCK, varies from lt. gray to yellowish to brownish-gray; v. fn. x-line qtz-ser. aggr.; mod. abund. jarositic limonite, varies from BJ, 25, to all J; these limonites occ. as partial-compl. pseud. after S <sup>2+</sup> s, avg. gr. size ~ 0.1-0.5 mm, and as narrow (< 1 mm.) vults. comm. w/ qtz.; a few narrow (< 1 mm.) qtz. vults. some w/ a little pitch limonite; tr. cc in more siliceous chips, as diss. gr. < 0.5 mm. in dia; tr. diss. hem.; tr. diss. & vult. clay		0-10'	85	20	10	540	1	-0.1	A2126
10-20'				5-7 *	10-20'	SAME AS 0-10' % of orange-brownish gray chips (flooded w/ limonite decreases from ~ 10% (0-10') to about 5% in this interval; rock is almost certainly altered INTENSIVE BRECCIA; a few chips contain ~ 0.5% diss. bright greenish-gray micaceous material, prob. either chlorite or biotite; cc may be sl. more abund., but still only tr.		10-20'	80	22	10	65	1	-0.1	A2127
20-30'				5-7 *	20-30'	SAME AS 0-10' Except appearance of pitch limonite in ~ 15% of the chips; this limonite occurs as diss. irreg. grns. < 1 mm. in dia, and as open irreg. vults. in which the limonite forms dk. reddish-brown, resinous, botryoidal crusts on the vein walls; some of the more siliceous chips (prob. jasp.) contain up to 0.5% cc, but cc overall is prob. ~ 0.1% (maybe more acc. ~ 0.05%)		20-30'	80	16	10	240	2	-0.1	A2128
30-40'				1-6 *	30-40'	SAME AS 20-30' Except cc (or a dark gray metallic mineral which greatly resembles cc) has increased to ~ 0.2% overall; cc is generally assoc. w/ those chips that are more siliceous and usually faced w/ v. narrow qtz. vults.		30-40'	55	26	-10	40	2	-0.1	A2129
40-50'				7-9 *	40-50'	SAME AS 0-10' cc has dramatically decreased to tr. overall - prob. contamination from above (20-40'); 20% (app.) of the chips contain 3-5% extr. fine-x-line brown micaceous minrl. (bte?) (disseminated)		40-50'	85	17	120	175	2	-0.1	A2130
50-60'				7-9 *	50-60'	SAME AS 40-50'		50-60'	80	11	50	70	3	-0.1	A2131
60-70'				7-9 *	60-70'	SAME AS 40-50'		60-70'	40	22	70	15	7	-0.1	A2132



Logged By J.B. Hulen  
F\* former sulfide contentProject Buckskin Southwest Hole No. BSW-7Page 2 of 13  
From 70' To 210'

Core Recovery		Graphic Geology	Graphic Alteration					Graphic Mineral (%)					Vol % Sulf Est	GEOLOGY		Thin Section	Analyses										
Size & Interval	% Rec.		silice.	ser.	clay	qtz. vults.	chl.	ble.	py	cpy	cc	bn		MoS	gn	sp	Interval	Description	Location & Description	Interval	PPM Cu	PPM Mo	PPM Pb	PPM Zn	PPM Ag	PPM Au	Sample Nos.
70-80'	5 1/8" RB 24	NOT applicable														* 4-6	70-80'	QTE-SER. ROCK, same as 0-10', exc. appearance of a little py, coated w/dk. gray steely cc, in abt. 5% of the chips; abt. 10% of the chips are actually siliceous enough as to be sericitic JASPEROID. These chips laced w/narrow, irreg. qtz. vults; a few waxy, lt. gray, irreg. clay vults. present in some chips.		70-80'	155	55	120	70	10 (0.29 oz/T)	0.003 oz/T	A2133
80-90'																* 3-5	80-90'	This interval ~ 60% QTE-SER. ROCK (same as 70-80') and ~ 40% JASPEROID, lt. gray, extr. dense, w/ ~ 3% diss., v. fine-xline py, partially totally replaced w/dk. gray steely cc		80-90'	105	39	50	40	5 (0.15 oz/T)	-0.1	A2134
90-100'																* 4-6	90-100'	Same as 70-80', exc. no py or cc, plus addition of ~ 5% pitch limonite, texturally similar to that occurring in the interval 20-30'; 10-20% of the QTE-SER chips contain ~ 1% unid. lt. brown transl-opaque MnO2, as diss. gr. aggregates up to 0.5 mm. in dia. (ind. grns. < 0.1 mm. dia.) - aggregates are irregular		90-100'	85	41	80	25	3 (0.09 oz/T)	-0.1	A2135
100-110'																* 4-6	100-110'	Same as 90-100'		100-110'	85	13	50	315	2 (0.06 oz/T)	-0.1	A2136
110-120'																* 4-6	110-120'	Same as 90-100'; The unid. lt. brown, transl-opaque GRAN. BACTERIATES may be ANDALUSITE, as found a mile or two south in qtz-ser. INTUSION BRECCIA at the BLUE DANUBE mine; may also be rutile? doubt it.		110-120'	80	13	20	425	3 (0.09 oz/T)	-0.1	A2137
120-130'																* 4-6	120-130'	Same as 90-100'. A number of the QTE-SER. chips show a def. porphyritic (relict) texture - phenos. are anh. subh., < 1 mm. in dia, commonly lath-shaped; these fragments are probably cut from altered clasts in INTUSION BRECCIA.	** Base of oxidation at ~ 130' **	120-130'	55	8	10	70	3 (0.09 oz/T)	-0.1	A2138
130-140'																* 3-5	130-140'	QTE-SER. ROCK, unpx, lt.-med. gray, v. fine-xline (< 0.1 mm); qtz. makes up the total rock; for alt. description, refer to 0-10'; 3-5% py, as diss. anh. grains & grain aggregates from < 0.1 mm. to 1.5 mm. (avg. ~ 0.5 mm); a few of these microrimmed w/dk. gray steely cc; unid. MnO2 of 90-100' still present in part - at any rate, this MnO2 may be repl. of orig. MnO2 in the BRECCIA.	about 85% of py grains are 2nd bte	130-140'	140	6	10	475	2 (0.06 oz/T)	-0.1	A2139
140-150'																* 5-7	140-150'	Same as 130-140', exc. py, w/cc microrims, has increased to 5-7%		140-150'	170	10	60	180	3 (0.09 oz/T)	-0.1	A2140
150-160'																* 10-12	150-160'	Same as 130-140', exc. py/w/cc, increasing to 10-12%; cc much more abund. - some py grains totally replaced w/dk. gray steely cc; many py grains are furnished reddish to purplish metallic color (bn microrims?); ~ 15% of sample is mottled, pyritic - appearing JSEP, laced w/narrow, clear qtz. vults, and ~ 1% py		150-160'	* 0.44% *	23	* 0.19% *	590	14 ** (0.41 oz/T)	-0.1	A2141
160-170'																* 7-9	160-170'	Same as 130-140', exc. total S2- decreasing to 7-9%; the rock is totally devoid of relict texture - just an extr. fine-xline aggr. of qtz. and ser. (and S2-); one chunk of solid cc in this interval 1x2x2 mm. in dimension; one py grain 3 mm. in dia.		160-170'	* 0.12% *	-1	110	410	2 (0.06 oz/T)	-0.1	A2142
170-180'																* 3-5	170-180'	Appearance of 25-30% jasperoid, v. lt. gray-white, cryptocrystalline, laced w/lt. irreg. qtz. vults, diss. w/ < 0.5% total sulfide (py, tr. gpx); drilling extr. slow in this rock; sample otherwise same as 130-140 w/ ~ 15% contamination from the oxide zone; hard to distinguish color. cupriferous py from valid cpy w/extr. small grains - both gel-brass.		170-180'	610	50	50	700	2 (0.06 oz/T)	-0.1	A2143
180-190'																* 1.5-2	180-190'	JASPEROID, same description as 170-180'; tr. cc as microrims on py and cpy; some (5-10%) contamination from oxide & overlying sulfide zones; drilling continues extremely difficult; estimated total sulfides in jasp. have increased to ~ 15-2%		180-190'	425	33	60	* 0.21% *	1 (0.03 oz/T)	-0.1	A2144
190-200'																* 0.5-1	190-200'	Same as 180-190'; JASPEROID has distinct microbrecciated texture; contains a tr. ser. & chl.; a few small (< 0.5 mm. dia.) vugs are present in the jasperoid - these are lined w/minute, water-clear qtz. xk.; ∴ suspect jasperoid is epithermal.	related to intrusion of T. dacite ppy?	190-200'	490	33	60	* 0.29% *	1 (0.03 oz/T)	-0.1	A2145
200-210'																* 0.5-1	200-210'	Same as 180-190'		200-210'	970	25	90	* 0.27% *	2 (0.06 oz/T)	-0.1	A2146



## BEAR CREEK MINING COMPANY - NORTHWEST DISTRICT

Logged By J.B. Hulen

Project Buckskin Southwest

Hole No. BSW-7

SCALE: 1" = 20'

Page 3

Of 13

From 210'

To 350'

Core Recovery		Graphic Geology	Graphic Alteration										Graphic Mineral (%)										Vol % Sulf Est	GEOLOGY		Thin Section		Analyses																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																	
Size & Interval	% Rec.		silic	ser	clay	qtz	chl	bit	k-sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py		sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py	sp	py



DESCRIPTION OF SPOT CORE SAMPLE, 460-464

QUARTZ-SERICITE ROCK, mottled lt. gray, med. gray, & lt. brownish-gray; rock is a fine-crystalline aggregate of quartz and sericite w/ disseminated py, tr. cpy, and ~ 0.5% specular hematite; qtz greatly predominates over ser — but relative % of qtz & ser. varies throughout, as irregular patches; qtz and ser. occur singly or together, often with sulfides, as narrow-irregular discontinuous veinlets; a few of the qtz veinlets are locally open, and lined with tiny (< 1 mm. in length), water-clear, euhedral qtz crystals; 5-7% pyrite, anomalously yellowish (cupriferous?), as diss., anh.-subh. grains, avg. ~ 0.3 mm. in dia. (< 0.1-2 mm. in dia.), and as narrow, irreg. disc. veinlets, with or without qtz and/or ser.; chalcopyrite, text. similar to pyrite; ~ 0.5% dark metallic gray specular hematite, as diss. grains, subh.-euh., < 0.1 mm. in dia., and as aggregates of these grains up to 1 mm. in dia.; poss. tr. 2nd qtz. (could be ultrafine-gr. spec. hem.); rock appears to be an intensely altered INTRUSION BRECCIA; fragmental texture, though, badly obscured.



## BEAR CREEK MINING COMPANY - NORTHWEST DISTRICT

Logged By J.B. HulienProject Buckskin SouthwestHole No. BSW-7Page 4 Of 13  
From 350' To 490'

Core Recovery		Graphic Geology	Graphic Alteration							Graphic Mineral (%)							Vol % Sulf Est	GEOLOGY		Thin Section Location & Description	Analyses							Sample Nos.
Size & Interval	% Rec.		silicif.	ser.	qtz. v. m.	chl.	btz.	Ksp. m.	py	cpy	MoS <sub>2</sub>	sp						Interval	Description		Interval	ppm Cu	ppm Mo	ppm Pb	ppm Zn	ppm Ag	ppm Au	
350- 360' 5% RB	not applicable																	350- 360'	ITE-SER. ROCK, lt. greenish-gray to lt. gray, v. fine-xline; 3-4% diss. fine-xline py & tr. cpy; <1% diss. spec. hem.; <1% diss. v. fine-xline, greenish-gray chl.; possible tr. diss. 200 Ksp. m.; cpy tends to be associated w/both chl. & hem.; both diss. & v. m.; both py and rare cpy occur as anh-subh. grains, <0.1-2 (avg ~0.5) mm dia., and as rare microvnlts.		350- 360'	100	13	20	50	-1	-0.1	A2160
360- 370'																		360- 370'	same as 350-360'		360- 370'	205	10	20	105	-1	-0.1	A2161
370- 380'																		370- 380'	same as 350-360', exc. py increasing & chl decreasing; qtz may also be increasing; spec. hem. occurs as diss. aggregates, up to 1 mm. (avg. ~0.2 mm) of almost cryptocrystalline scales, just barely visible under high power		370- 380'	100	11	20	65	-1	-0.1	A2162
380- 390'																		380- 390'	same as 370-380, exc. dramatic increase in py; more cpy; perhaps a little more spec. hem.		380- 390'	175	11	10	35	-1	-0.1	A2163
390- 400'																		390- 400'	same as above, exc. up to 1% spec. hem. in some chips; a few chips contain up to 20% total sulfide		390- 400'	80	10	20	100	-1	-0.1	A2164
400- 410'																		400- 410'	same as above, exc. py decreasing; chl. virtually nonexistent at this level		400- 410'	50	5	10	25	-1	-0.1	A2165
410- 420'																		410- 420'	same as above, exc. py found more commonly as microvnlts - perhaps 10-15% microvnlts; 85-90% diss.; also, spec. hem. has decreased to <0.1%		410- 420'	45	4	20	650	-1	-0.1	A2166
420- 430'																		420- 430'	ITE-SER. ROCK, v. lt. gray, v. fine-xline; a few v. narrow (<0.5 mm. wide) qtz. vnlts, generally w/ little py; tr. diss. dk. greenish-gray chl.; v. fine-xline; ~0.5% diss. spec. hem. (same as 370-380); % of qtz. in qtz-ser. rock has increased over higher intervals; diss. & microvnlts py & tr. cpy; rock becoming incr. harder to drill; ~15% contamination		420- 430'	40	13	20	130	-1	-0.1	A2167
430- 440'																		430- 440'	same as 420-430', exc. no chl.; qtz. continues to increase.		430- 440'	55	10	20	155	-1	-0.1	A2168
440- 450'																		440- 450'	same as 420-430'; several chips of qtz.-cpy-MoS <sub>2</sub> veinlet present; quartz still seems to have increased even more.		440- 450'	20	8	20	* 0.14% *	-1	-0.1	A2169
450- 460'																		450- 460'	same as 420-430'		450- 460'	85	15	30	960	-1	-0.1	A2170
460- 461' NC	50																	460-	Refer to description of spot core sample, 460-461' Entire interval is similar		460- 470'	45	-1	10	90	1	-	A2161
461- 470'	100																		-INTRUSION BRECCIA-									
470- 480'	100																	470- 480'	same as 460-470', exc. ~1% diss. spec. hem.; a few irregular bands in which hem. occurs as lensoid to irregular-elongate masses, up to 1.5 x 0.5 mm. in x-section, consisting of extremely tiny xls. (<0.1 mm. in dia.); also tr. diss. fine-xline chl.		470- 480'	35	4	-10	45	-1	-	A2262
480- 490'	100																	480- 490'	same as 460-470'; py v. erratic in distribution; varying, in individual areas, from ~1% to 25% - est. overall ~9%.		480- 490'	75	17	-10	25	-1	-	A2263



BEAR CREEK MINING COMPANY - *NORTHWEST* DISTRICT

Project Buckskin SouthwestHole No. BSW-7

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From 490' To 630'

Core Recovery		Graphic Geology	Graphic Alteration								Graphic Mineral (%)						Vol % Sulf Est	GEOLOGY		Thin Section	Analyses						
Size & Interval	% Rec.		silicified	ser.	clay	pyrite	chl.	pyrrho.	spec. hem.	py	cpv	gh	sp	MOS	Interval	Description		Location & Description	Interval	PPM CU	PPM MO	PPM PB	PPM ZN	PPM AG	PPM ALI	Sample Nos.	
490-495'	100														9-10	490-500'	Quartz-sericitized INTRODUCTION BRECCIA; for full description see description of spot core sample, 460-464', preceding page 4; fault: 491-492E; stress-py is erratically distributed, varies fr. 2% to 30%, but avg. 9-10%.		490-500'	75	5	10	10	-1	—	A2264	
495-504'	100															9-10	500-510'	Same as 490-500', except ultrafine-granular gte. has greatly increased relative to ser.; spec. hem. has dropped out; total sulfides have decreased; poss. tr. 2nd bte.; py increasing somewhat toward base of interval.		500-510'	20	1	10	5	-1	—	A2265
504'-514'	100															3-4	510-520'	Same as 490-500', exc. no spec. hem. between 510' & 515'; a few fib. = 1cm. wide, are locally present; again: py very erratic in distribution—locally up to 40-45%.		510-520'	30	-1	30	5	-1	—	A2266
514-521'	100															7-8	520-530'	Same as 500-510', exc. a little spec. hem. appears at 528'.		520-530'	35	1	10	65	-1	—	A2267
521-531'	100															3-4	530-540'	Same as 500-510', exc. tr. 2nd bte., as diss. v. fine-x-line flakes, and as narrow, irregular, discontinuous vults. consisting of these flakes mixed w/ gte. & ser. Fault: 538-539E'		530-540'	15	15	30	5	-1	—	A2268
531-538'	100															1-2	540-550'	Fine-granular gte., totally silicified rock; 0.5-1% py, texture similar to that in intervals above; irreg. clots of cpv up to 7 mm. in dia.; abundant gte. vults. Faults: 543-545', 546.6-550'		540-550'	50	5	-10	20	-1	—	A2269
538-545'	100															0.5-1	550-560'	Same as 540-550' to 557.5'		550-560'	35	15	-10	5	-1	—	A2270
545-548'	100															0.5-1	560-570'	Same as 530-540' to 560' from 557.5'		560-570'	75	2	20	5	-1	—	A2271
548-550'	100															0.5-1	570-580'	Entire interval badly shattered		570-580'	45	4	50	10	-1	—	A2272
550-552'	20															1-2	580-590'	Same as 570-580', exc. more gte., gte. vults. Entire interval is badly shattered		580-590'	50	26	-10	-5	-1	—	A2273
552-555'	100														6-7	590-595'	Quartz-sericitized INTRODUCTION BRECCIA; for full description, refer to description of spot core sample, 460-464', preceding p. 4; Sulfides increase from 2% at top of interval to 9% at base of interval, avg. 6-7%.	*** Fault: 581-582E, 584-585E	590-595'	15	6	-10	-5	-1	—	A2274	
555-557'	100														8-9	595-602'	Same as 570-580', except color is mottled lt. gray, lt. brownish-gray, & greenish-gray (traces of diss. chl. & 2nd bte.); a few small, irreg. patches, up to 2 mm. in dia., of intense blue, translucent, fairly hard mineral; spec. hem. appears again		595-602'	15	4	10	5	-1	—	A2275	
557-567'	100														7-8	602-610'	Same as 580-590', except ~ 0.5-1% of the blue mineral, also though, it occurs in narrow, irreg. discontinuous vults., and in irregular clots up to 10 mm. in dia.		602-610'	20	15	-10	5	-1	—	A2276	
567-572'	100														0.5-1	610-620'	Same as 540-550'		610-620'	20	16	10	-5	-1	—	A2277	
572-577'	100														11-12	620-630'	Quartz-sericitized INTRODUCTION BRECCIA (see 460-464') Fault same to 615B; a little of the blue mineral (580-590'), as diss. grains up to 3 mm. in dia.		620-630'	60	2	-10	25	-1	—	A2278	
577-583'	100														7-8		Same as 610-620', a few fragments of gte-ser. schist ?? or extremely well-foliated gte-ser. rock										
583-592'	100																										
592-599'	100																										
599-605'	100																										
605-612'	100																										
612-618'	66																										
618-623'	100																										
623-628'	100																										
628-635'	100																										



BEAR CREEK MINING COMPANY - *NORTHWEST* DISTRICT

Project Buckskin SouthwestHole No. BSW-7

Page 6 Of 13  
From 630' To 770'

[illegible]



SUPPLEMENT

780-784': HORNBLende BIOTITE DACITE (possibly QUARTZ LATITE) POPHYRY, lt. greenish to brownish-gray, strongly altered; original rock texture consists of 40-45% phenocrysts (plagioclase, biotite, hornblende, and quartz) embedded in a very fine-crystalline groundmass; plagioclase phenocrysts are stubby, anhedral to subhedral, and average 2 mm. in dia. or length (<1-4 mm. range); biotite books (avg. ~5%) are subhedral-euhedral, and average 2 mm. in dia.; hornblende prisms (1-2%) are subhedral to euhedral, and average 1 x 2 mm. in cross-section (longitudinal section); quartz phenocrysts (<1%) are anhedral, rounded, commonly embayed, and average <1 mm. in dia. \* plag. 50-55%

Bte (primary) and hbl. are totally altered to fine-crystalline aggregate of dk. brown and biotite and dark grayish-green chlorite — these 2 minerals also sparsely replace plagioclase phenocrysts and matrix feldspar; large amounts of silica have apparently been added to the groundmass; rock is laced w/ translucent quartz veins up to 1 cm. (avg 1 mm.) in width; pyrite and chalcopyrite occur as ① diss. anh. grains, <0.1-2 (avg ~0.3) mm. in dia., ② as grains and selvages in quartz veins, ③ singly or together as microveinlets w/ quartz

note: additional presence of slightly reddish-gray metallic mineral, texturally similar to pyrite & cpy (energitte? tetrahedrite?)



## BEAR CREEK MINING COMPANY - NORTHWEST DISTRICT

Logged By J.B. Hulen

Project Buckskin Southwest

Hole No. BSW-7

Page 7 Of 13  
From 770' To 910'

Core Recovery		Graphic Geology	Graphic Alteration							Graphic Mineral (%)							Vol % Sulf Est	GEOLOGY		Thin Section Location & Description	Analyses							Sample Nos.	
Size & Interval	% Rec		silica	ser.	clay	qtz. w/lt.	chl.	2 <sup>nd</sup> bte	hem.	calcite	py	sp	MoS <sub>2</sub>	gh	sp	Interval		Description	Interval		ppm Cu	ppm Mo	ppm Pb	ppm Zn	ppm Ag	ppm Au			
764- 774'	100																770- 780'	INTRUSION BRECCIA, same as 720-730'; Entire interval badly shattered		770- 780'	* 0.13%	32	10	70	1	—	A2293		
774- 784'	100																	780- 784'	HORNBLENDE BIOTITE DACITE PORPHYRY *Refer to supplement*		780- 784'	* 0.11%	20	10	60	-1	—	A2294	
784- 790'	100																	784- 790'	same as 780-784', exc. intrusion breccia below 788'		784- 790'	* 0.16%	22	80	65	2	—	A2295	
790- 795'	100																	790- 800'	same as 780-784', exc. total sulfide % has dropped to 1-2%		790- 800'	* 0.19%	29	10	65	1	—	A2296	
795- 804'	100																	800- 810'	same as 780-784', except strong qtz veining.		800- 810'	* 0.18%	41	20	90	2	—	A2297	
804- 809'	100																	810- 820'	810-815': same as 780-784' 815-820': INTRUSION BRECCIA, qtz-chlorite alteration, w/ ~4% of the reddish-gray metallic described under 780-784 supplement.		810- 820'	* 0.12%	32	10	85	2	—	A2298	
809- 815'	100																	820- 830'	820-827': same as 815'-820' 827-830': same as 780-784'		820- 830'	* 0.34%	19	210	490	8	—	A2299	
815- 824'	100																	830- 840'	same as 780-784', exc. no 2nd bte. below 834' (just qtz-chl.)		830- 840'	* 0.35%	5	60	370	4	—	A2300	
824- 834'	100																	840- 850'	same as 780-784'; pyritic texture obscure, but apparent; 2 <sup>nd</sup> bte to 846'; just qtz-chl. alt. (w/ a little spec. hem) from 846' to 850'; intense qtz veining.		840- 850'	* 0.21%	44	10	125	2	—	A 20101	
834- 840'	100																	850- 860'	same as 780-784'; pyritic texture obscure; sp has definitely increased; intense qtz veining.		850- 860'	* 0.42%	20	30	455	4	—	A 20102	
840- 845'	100																	860- 870'	same as 850-860'		860- 870'	* 0.39%	26	10	100	3	—	A 20103	
845- 850'	100																	870- 880'	same as 850-860'		870- 880'	* 0.24%	17	10	85	2	—	A 20104	
850- 860'	100																	880- 890'	same as 850-860' Faults: 880'-885' 886-887' 887'-888'	Reddish-gray metallic mineral (refer to 780-784 supplement) moderately abundant — is this tetrahedrite or enargite?		880- 890'	* 0.35%	15	10	115	2	—	A 20105
860- 866'	100																	890- 900'	same as 850-860' Faults: 890-890.5' 890.5-891.5' 891.5-892.5'	Abund. 2 <sup>nd</sup> bte.		890- 900'	* 0.45%	20	10	105	3	—	A 20106
866-869 869-871'	100																	900- 907'	same as 850-860' above 903.5', w/ 0.1' qtz-sp. vi. - 901-903.5' Below 903.5': qtz-ser. fault breccia w/ clasts of dacite ppx & an unknown fine-grained ppx & abund. calcite		900- 907'	* 0.30%	38	100	280	12	—	A20107	
871-872 872-875'	100																	907-920	REFER TO PAGE B		907-920							A20108	

2447 34 40 163 7



BEAR CREEK MINING COMPANY - *NORTHWEST* DISTRICT

## Project

Hole No. BSW-7

BSW-7

Page 8 Of 13  
From 910' To 1050'

Page 8 Of 13  
From 910' To 1050'

[illegible]



Logged By J. B. Hulen

Project Buckskin Southwest Hole No. BSW-7

Core Recovery		Graphic Geology	Graphic Alteration										Graphic Mineral (%)					Vol % Sulf Est	GEOLOGY		Thin Section	Analyses						
Size & Interval	% Rec.		silic.	ser.	clay	qtz. vln.	chl.	calcite	2nd Esp.	py.	cpv.	mosa.	gm.	sp.	gypsum	Interval	Description		Location & Description	Interval	PPM Cu	PPM Mo	PPM Pb	PPM Zn	PPM Ag	PPM Au	Sample Nos.	
1049-1055'	↑																1050-1060'	SAME AS 1040-1050', exc. some of the clasts below 1059' contain a little 2nd bte.		1050-1060'	480	1	20	65	2	—	A20122	
1055-1062'																	1060-1070'	Below 1065'; SAME AS 1040-1050' exc. w/ a few vlns. of waxy white clay. Above 1065': QUARTZ-FELDSPAR PPY DIKE, qtz.-chl.-bte. anh.		1060-1070'	800	28	20	65	1	—	20123A	
1062-1066'																	1070-1081'	SAME AS 1040-1050'		1070-1081'	430	20	20	30	1	—	20124A	
1066-1071'																	1081-1088'	SAME AS 1040-1050'		1081-1088'	610	30	30	90	2	—	20125A	
1071-1075'																	1088-1100'	HORNBLANDITE DIKITE PACITE PPY, altered, lt. brownish-gray; mafics altered to bte.-clay-ser (bte. predominates) aggregate; feldspars altered to qtz-ser-clay aggregate; cut by narrow irreg. vlns. of qtz (some rose) granular qtz-ksp (up to 10 cm. wide), & qtz-calcite; a few stellite masses, up to 0.5 mm. in dia., of radiating, acicular, dull black xls. (TOLUENALINE?) A few waxy white clay vlns.; sulfides texturally similar to 780-784' - some	supergene enrich-ment, w/ develop-ment of sooty ex-lines, esp. on qtz.	1088-1100'	665	20	10	60	1	—	20126A	
1075-1080'																	1100-1110'	SAME AS 1088-1100', exc. 1107'-1108': rose qtz - py-cpy vein 1108'-1109': clay-seritized 1109'-1110': w/ vlns lined w/ exh. qtz xls. up to 2 mm. long		1100-1110'	870	28	30	30	3	—	20127A	
1080-1084'																	1110-1120'	SAME AS 1088-1100', exc. a little diss. & vln. chl., v. little 2nd ksp.		1110-1120'	*	0.12%	30	10	60	1	—	20128A
1084-1091'																	1120-1126'	SAME AS 1110-1120'		1120-1126'	865	41	10	40	2	—	20129A	
1091-1098'	NYME																1126-1140'	INTRUSION BRECCIA, qtz-seritized, v. lt. gray; heavy AFV-qtz vlns. badly shattered; Fault: 1135-1139'		1126-1140'	*	0.13%	57	140	130	5	—	20130A
1098-1100'																	1140-1150'	same as 1126-1140', exc. a little 2nd bte & mod. chl appear at 1155'; interval is badly shattered. A few narrow, irreg. lt. brownish-pink dolomite-py-cpy-sp vlns.		1140-1150'	*	0.11%	36	10	35	3	—	20131A
1100-1103'																	1150-1159'	same as 1126-1140', exc. a little 2nd bte & mod. chl appear at 1155'; interval is badly shattered. A few narrow, irreg. lt. brownish-pink dolomite-py-cpy-sp vlns.		1150-1159'	*	0.10%	25	10	50	4	—	20132A
1103-1107'	100																1159-1170'	QUARTZ-FELDSPAR PPY, too badly altered for determination of orig. comp. by direct mic. examination; 25-30% alt. feldspar phen., anh.-subh., avg. 2 mm. length or dia.; 1-2% qtz phen., anh., avg. 1 mm. dia.; aphanitic to v. f. xline quizes; hybrid alt. → qtz-ser-chl-bte in roughly equal prop.; relict mafic aites mostly alt. to chl-bte. aggr.; some diss. & vln. calcite.	Tr. 2nd Esp w/ qtz as vlns.	1159-1170'	*	0.17%	76	10	35	2	—	20133A
1107-1112'																	1170-1180'	same as 1159-1170', exc. a little less bte. bte, here as above, occurs as ① diss. scattered, fine-xline flakes ② in microvlns w/ or w/o qtz & 2nd Esp. ③ as fine-xline chl.-bte. aggregates replacing original mafics.		1170-1180'	*	0.16%	45	10	35	2	—	20134A
1112-1117'																	1180-1190'	same as 1159-1170', exc. appearance of a few gypsum vlns. up to 5 mm. wide. py. begins to resemble PACITE PPY, but unsure as yet. pyritic texture more apparent.		1180-1190'	*	0.11%	28	10	40	1	—	20135A
1117-1122'																												
1122-1127'																												
1127-1130'																												
1130-1135'																												
1135-1138'																												
1138-1141'																												
1141-1146'																												
1146-1151'																												
1151-1156'																												
1156-1160'																												
1160-1165'																												
1165-1169'																												
1169-1174'																												
1174-1179'																												
1179-1184'																												
1184-1189'																												
1189-1199'	↓																											



Logged By J. B. Hulen

Project Buckskin Southwest Hole No. BSW-7

[illegible]



ALTERATION  
SUPPLEMENT

## Project

Buckskin Southwest

Hole No. BSW-7[illegible]







ALTERATION  
SUPPLEMENT

Project Buckskin Southwest Hole No. BSW-7

Page 13 Of 13  
From 1610' To 1715'

irreg. grs  
up to 0.5  
th dia.

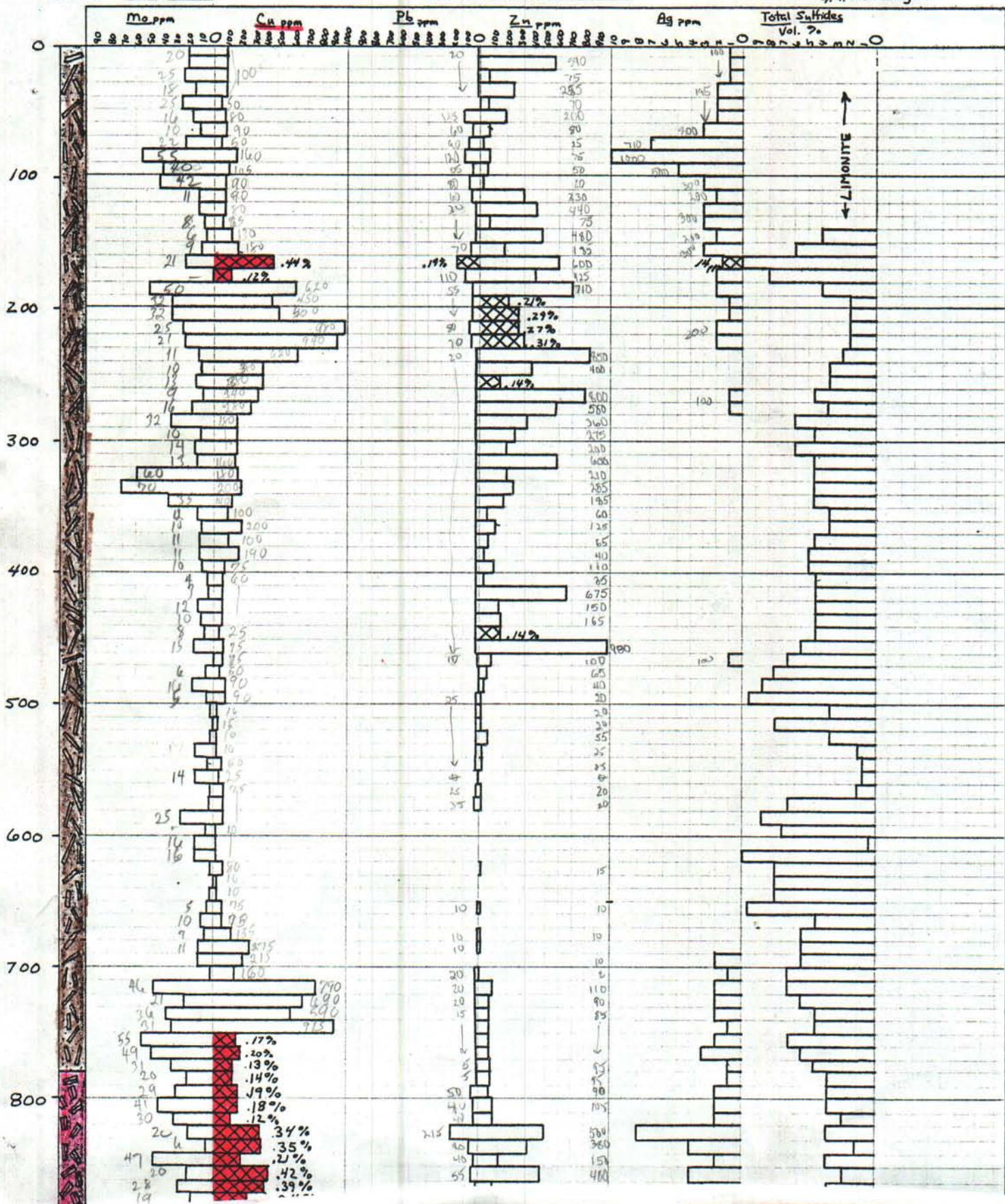


DOUGLAS CO. NEW.

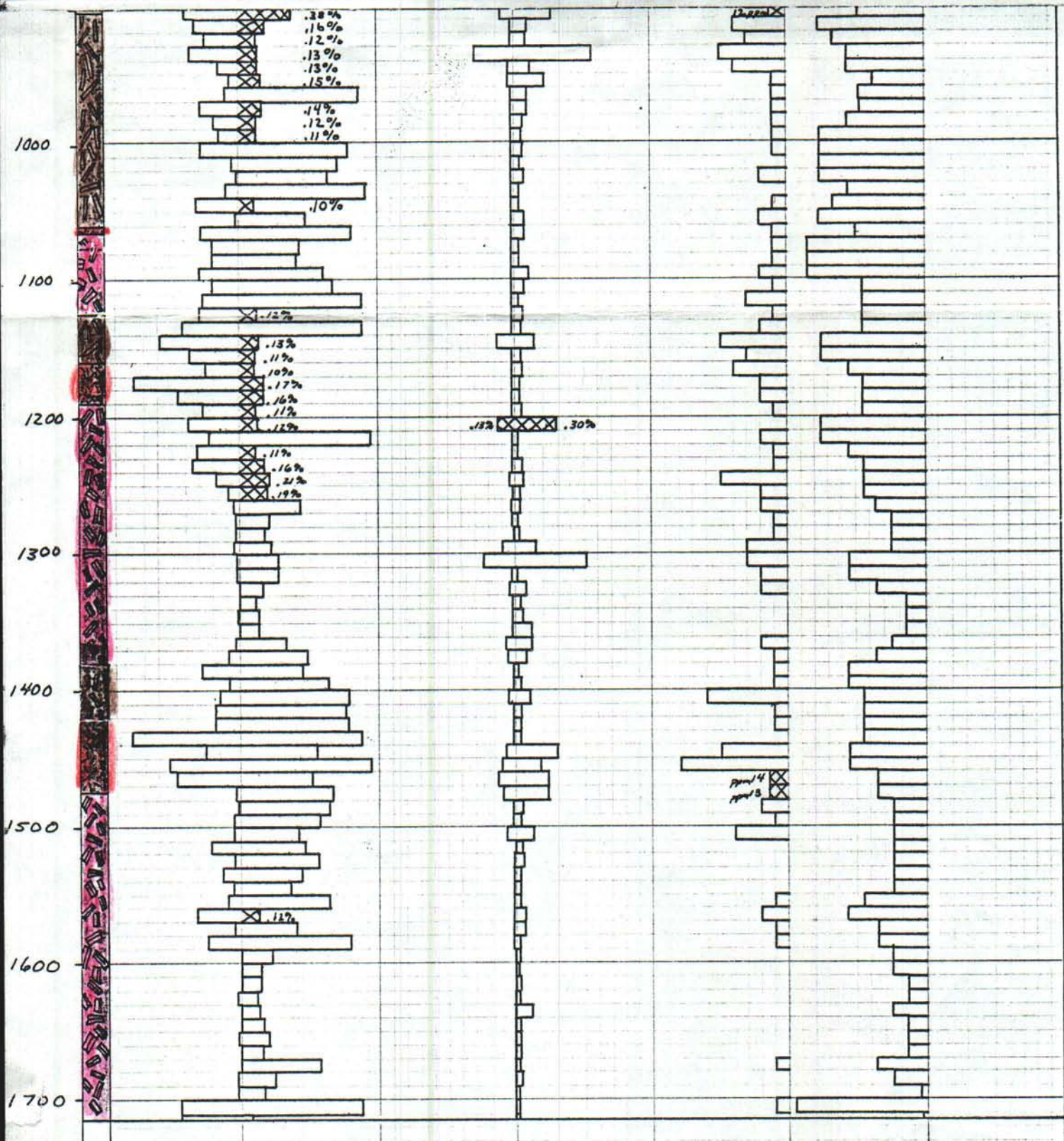
Code 005-01-0411

START \_\_\_\_\_ COMPLETED \_\_\_\_\_

DATA: G. R. Heinemeyer







BSW-7  
bottom half



516' ~ .16

88' of .35

at 820'

Drill hole BSW-7 (TD 1,715 feet), the deepest of three holes completed at Buckskin Southwest, was drilled to test for the presence of a deep primary porphyry copper deposit, vertically zoned beneath a quartz-sericite "lithocap". The hole penetrated, beneath 750 feet of quartz sericite rock above a 510-foot zone of complexly altered dacite porphyry, volcanic breccia, and quartz-feldspar porphyry averaging 0.16% Cu as chalcopyrite (Figure 5). All rock types within this zone are heavily quartz-veined and altered with the addition of various amounts and combinations of biotite, chlorite, sericite and quartz. Chlorite-quartz-biotite alteration is most often associated with relatively higher copper grades; chlorite-quartz-biotitized dacite porphyry between 820 feet and 907.7 feet, for instance, averages 0.35% Cu.



xxxx 89510

July 22, 1977

R. C. Babcock, Jr.  
District Manager  
Bear Creek Mining Company  
Box 9, Dishman Branch  
Spokane, Washington 99213

Dear Russ:

The sample numbers for hole BSW-7 at the Buckskin project are listed below as per our phone conversation last week. Also listed are the sample numbers for hole BSW-9 which we did not discuss but I would appreciate it if these pulps were also made available for Mr. Bamford's analyses. Could you please send the pulps to Mr. Bamford in Salt Lake City at your convenience?

We believe the analysis of these pulps will be useful in delineating a porphyry target at Buckskin, and appreciate your cooperation.

Also enclosed for your files is Bamford's report on the sulfide and limonite study he conducted on the entire Buckskin project this spring.

Pulp Sample Numbers:	BSW-7	A2126 - 2170
		A2261 - A2300
		A20101 - A20188
	BSW-9	28317 - 28399
		28400 - 28450
		28651 - 28660

Sincerely,

William M. Oriel  
Geologist

pb  
Enc.



## Buckskin Project

BSW-7

Item 15

~~THE GOLDFIELD CORPORATION~~

100' Composites

~~GETCHELL MINES DIVISION~~

Month/ of \_\_\_\_\_ 19\_\_\_\_

Journal Voucher No. \_\_\_\_\_

[illegible]



BSW 6

(35)  
Item 15

ROCK DESCRIPTION	GEOCHEMISTRY		
	Interval	Cu (ppm)	Mo (ppm)
0-10' ALLUVIUM. Same as 0-20' in BSW-3 (fig. 7).	0-10'	NO SAMPLE	
10-70' TUFFACEOUS SILTSTONE. Same as 20-60' in BSW-3 (fig. 7).	10-20'	50	1
	20-30'	55	2
	30-40'	25	-1
	40-50'	10	-1
	50-60'	15	-1
	60-70'	15	2

Figure 10. Summarized rotary drill hole log, BSW-6.

#### DISCUSSION AND GEOLOGIC HISTORY

Geologic history of the southwestern Buckskin Range can be hypothetically reconstructed from field evidence in the following sequence:

- (1) Extrusion of eugeosynclinal, Triassic intermediate lavas.
- (2) Burial and regional thermal metamorphism of the Triassic intermediate volcanic sequence, possibly accompanied by gentle, open folding. Such thermal metamorphism, of lower greenschist grade in the southwestern Buckskin Range, resulted in extensive alteration of primary mafic minerals and plagioclase to chlorite and/or epidote and calcite-dolomite. Rocks thus altered closely resemble hydrothermally propylitized varieties.
- (3) Emplacement of irregular plutons of fine-crystalline granodiorite, probably during Early Cretaceous (or possibly Late Jurassic) time, with attendant contact metamorphism of Triassic metavolcanic rocks.



BSW 5

(35)  
Item 15

ROCK DESCRIPTION	GEOCHEMISTRY		
	Interval	Cu (ppm)	Mo (ppm)
0-10' ALLUVIUM. Same as 0-20' in BSW-3 (fig. 7).	0-10'	NO SAMPLE	
10-60' TUFFACEOUS SILTSTONE. Same as 20-60' in BSW-3 (fig. 7).	10-20'	65	2
	20-30'	90	6
	30-40'	50	-1
	40-50'	30	-1
	50-60'	50	3

Figure 9. Summarized rotary drill hole log, BSW-5.



Log BSW 4

(35)  
Item 15

ROCK DESCRIPTION	GEOCHEMISTRY		
	Interval	Cu (ppm)	Mo (ppm)
0-10' ALLUVIUM. Same as 0-20' in BSW-3 (fig. 7).	0-10'	NO SAMPLE	
10-85' TUFFACEOUS SILTSTONE. Same as 20-60' in BSW-3 (fig. 7).	10-30'	10	-1
	30-40'	15	5
85-90' HORNBLENDE BIOTITE DACITE PORPHYRY, lt. greenish-gray.	40-50'	30	9
25-30% subh.-euh. plagioclase phenocrysts, avg. 2 mm. in length,	50-60'	30	5
9-10% subh.-euh. biotite books, avg. 2 mm. in diameter, 1-2% subh. hornblende phenocrysts, avg. 2 mm. in length, and 1% anhedral, rounded, translucent	60-70'	30	-1
qtz. phenocrysts, avg. 2-3 mm. in diameter, embedded in an aphanitic groundmass. Hornblende and biotite phenocrysts are partially altered to chlorite and/or epidote aggregates.	70-80'	35	1
	80-85'	30	1
	85-90'	35	1

Figure 8. Summarized rotary drill hole log, BSW-4.

ROCK DESCRIPTION	GEOCHEMISTRY		
	Interval	Cu (ppm)	Mo (ppm)
0-10' ALLUVIUM. Same as 0-20' in BSW-3 (fig. 7).	0-10'	NO SAMPLE	
10-60' TUFFACEOUS SILTSTONE. Same as 20-60' in BSW-3 (fig. 7).	10-20'	65	2
	20-30'	90	6
	30-40'	50	-1
	40-50'	30	-1
	50-60'	50	3

Figure 9. Summarized rotary drill hole log, BSW-5.



#### DRILLING

Between late fall 1970 and early spring 1971, six short air rotary holes, for a total of 830 feet, were drilled at Buckskin Southwest to perfect 32 lode claims. Summarized drill hole logs for the six holes are presented in figures 5 through 10. Drill holes BSW-4, 5, and 6 — drilled subsequent to completion of the Buckskin Southwest geologic map (pl. 2) — are collared in Quaternary pediment gravels a few hundred feet west of the western edge of the area mapped (pl. 2).



Log BSW 3

(35)  
Item 15

ROCK DESCRIPTION	GEOCHEMISTRY			
	Interval	Cu (ppm)	Mo (ppm)	Ag (ppm)
0-20' ALLUVIUM. Pebbles, cobbles and boulders of all pre-alluvial rock types exposed in the southern Buckskin Range, with minor quartz monzonite from the Pine Nut Range, to the west.	0-20'	NO SAMPLE		
	20-30'	45	5	-1
	30-40'	40	-1	-1
	40-50'	35	6	-1
	50-60'	45	9	3
20-60' TUFFACEOUS SILTSTONE, semi-consolidated, white to lt. gray. Tr. manganese oxide locally present along fractures. Tr. shredded plant debris toward base of interval.	60-70'	45	9	3
	70-80'	60	5	2
	80-90'	70	3	2
	90-100'	85	7	1
60-118' HORNBLLENDE ANDESITE PORPHYRY, lt. greenish-gray where unoxidized. 15-17% stubby, anh.-subh. plagioclase phenocrysts, avg. 1-2 mm. in length, and 2-3% anh.-subh. hornblende phenocrysts, avg. 1 mm. in length or diameter set in an aphanitic groundmass. Less than 0.5% diss., fine-xline., anh. <u>py</u> below 100'. From 60' to 100', <u>py</u> is partially to totally oxidized to powdery, yellowish-brown goethite.	----- <u>Base of Oxidation</u> -----			
	100-110'	60	5	2
	110-118'	50	3	1

Figure 7. Summarized rotary drill hole log, BSW-3.



ROCK DESCRIPTION	GEOCHEMISTRY			
	Interval	Cu	Mo	Ag
0-200' HORNBLLENDE QUARTZ LATITE PORPHYRY, strongly altered, light greenish-gray. Relict texture almost totally obscured by alteration. Hematitic leached capping from 0 to 120'. Porphyry in this interval is altered to a fine-xline. <u>qtz.-ser. aggregate, w/2-3% diss., fine-xline, anh.-subh. magnetite, 2-3% diss., fine-xline., anh.-subh. secondary biotite, 1% diss., fine-xline., anh.-subh. chlorite, 3-5% narrow (less than 2 mm. in width), banded (epithermal?), transl. stockwork qtz. vnlt., 1-2% goethite (as diss., fine-xline., anh. grains and, less commonly, as narrow, stockwork vnlt.), and 4-7% powdery, dark reddish-gray hematite, as diss., fine-xline., anh.-subh. pseudomorphs after primary sulfides and as narrow (less than 1 mm. wide) stockwork vnlt. in association with one or more of the minerals goethite, pitch limonite, quartz, turquoise, and "neotocite". Zone of supergene enrichment from 120' to 150'. From 120' to 140', rock is texturally and compositionally similar to that cut between 0 and 120', but contains 0.5-1% dark gray, sooty cc, as diss. fine-xline., anh.-subh. pseudomorphs after primary sulfides. From 140' to 150', rock contains 4-5% diss., fine-xline. py, partially replaced by dk. gray sooty cc. (0.5-0.7%). Primary sulfide zone from 150' to 200'. Rock is texturally and compositionally similar to that cut between 0 and 120' but contains 4-7% diss., fine-xline. anh.-subh. py, and less than 0.2% diss., fine-xline. anh. cpy.</u>	0-10'	0.12%	20	6
	10-20'	990	23	10
	20-30'	0.12%	16	7
	30-40'	0.11%	11	3
	40-50'	0.14%	15	4
	50-60'	0.23%	17	5
	60-70'	0.21%	11	3
	70-80'	0.23%	19	2
	80-90'	0.17%	19	3
	90-100'	970	16	17
	100-110'	0.15%	41	11
	110-120'	0.16%	8	11
	----Base of Oxidation----			
	120-130'	0.41%	1	11
	130-140'	1.20%	10	6
	140-150'	0.29%	20	4
	--Base of Supergene Enrichment--			
	150-160'	515	15	5
	160-170'	0.12%	13	7
	170-180'	820	26	6
	180-190'	815	16	4
	190-200'	690	22	2

All values reported in ppm unless otherwise noted.

Figure 6. Summarized rotary drill hole log, BSW-2.



BSW 2

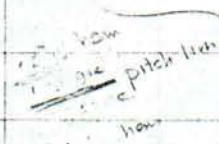
(35) Item 15

At surface - minor bleached, <sup>v. et.</sup> greenish-brown slaty mica along a few fracs, part. diss. — may have been chl. or bte. — also cracked mori. along fracs., minor diss. Cu-mineral (green, amorphous)

BSW 2

0-10' Ex. is primarily granular gtz. w/ minor clay (lt. greenish-gray to lt. brownish-gray), poss. some sericite; a few gtz. veinlets,  $\leq 1$  mm. in width; 7-9% hematite, — occurs as minute diss. grains & as veinlets (irreg.)  $\leq 1$  mm. in width.; ~1% goethite, mostly along fracs (some pitch limonite), but some as pseudomorphs/sulfides, generally  $\leq 0.5$  mm. in dia.; tr. unknown lt. green Cu(?) mineral as thin crusts along fracs.; poss. tr. cc along fracs. \*\* hem. appears dk. metallic gray on fresh surface — streaks intense brick reddish-brown.; poss. tr. MnO (Cu-bearing?) along fracs

10-20' Same as above, exc. a few more gtz. veinlets (up to 2 mm wide — <sup>some</sup> appear epithermal [cockade structure])



one frag. shows this relationship

; some of the hem post-dates introd. of gtz. veinlets ;

% goethite slg. increasing.; a few veinlets of translucent bluish-gray clay present ( $\leq 1$  mm. wide)

20-30' Same as above

30-40' Same as above, exc. several frags. of bte. decite spg., fresh black bte., feldspars argillized. — some reflect gtz. phenos.; minor diss. gre. pseudos; % hem dropping to ~5% in rx. other than Dp.



BSW2

40-50' - same rock type, but app. of several chunks  
w/ < 1% diss. hem.; fewer gtz. veinlets present

50-60' - same, except a few crusts of metallic-appearing,  
dk gray mineral along fracs - streak med. brown to black  
(may be cc.) — hint of relict pyritic texture in a few  
frags. — few gtz. veinlets present.

60-70' - same — a few veinlets of <sup>v. fine</sup> granular hem. present  
up to 15 mm. wide

70-80 "

80-90 "

90-100' same rx. type, exc. most frags. more intensely  
blackish-gray; a bit more clay in rx.; 2-3% diss. hem.,  
most appear blackish - but a few grains are ~~diss.~~  
bluish-gray metallic. — goethite same as 0-10'

100-110' same - a few more pcs. w/ almost no minls.

110-120' same - exc. 35-50% of chips contain < 1% casts/pseudos -  
also chips w/ up to 1% cc(?) (dk. gray metallic)

120-130' same w/ appearance of thoroughly argillized Dp  
some gtz. rx. frags. contain up to 1% cc(?)



130-140' >75% of sample is argillized & chloritized. Dp  
w/ ~1-1.5% cc (may be microcline) and <0.1% brilliant orange-red  
amorphous mineral (cinnabar, realgar?); embedded bte  
books are converted to dk. green chlorite; rest of  
sample is granular gtz. rock, also w/ 1% diss cc(?)

140-150' granular gtz-clay-(chl) rock, v. fine-gr., lt.-med.  
slty. greenish-gray; no relict texture discernible;  
~3% diss. v. fine-gr. (<0.5 mm.) py - minor py along  
tiny irreg. fracs.; a few small white clay veinlets.  
~0.3-0.5% diss. dk. gray metallic mineral similar in  
texture to py.

150-160 - Same as above, exc. less chl. (?) more gtz.  
poss. tr. 2nd bte.

160-170 - same as above

170-180 - "

180-190 - " exc. 2nd bte appears to be increasing.



BSW/

100' Average assays

(35)

Item 15

	Cu	Mo	Pb	Zn	Ag	
0-100	294	13	NA	NA	1.6	2
200	678	12	NA	NA	4.1	4
300	283	6.	NA	NA	1.9	2
400	919	19	29	154	2.5	2 <sup>1/2</sup>
500	998	10	19	136	1.5	1
600	294	2	86	335	1.8	2
700	652	9	43	180	2.6	2 <sup>1/2</sup>
800	871	30	23	109	2	2
900	806	14	32	115	3.8	4
1000	616	13	33	131	1.4	1
1100	537	19	35	66	0.8	1
1200	658	24	23	43	0.5	-
1313	293	22	24	67	0.3	-

700-710 = .7 Ag



Drill hole BSW-1 (TD 1,313 feet - 1974) was drilled in hopes of intersecting a thicker and higher grade interval of copper mineralization than was encountered in BSW-7 (1972). Drilling encountered 450 feet of quartz-sericitized volcanic flows and breccias with a 100-foot zone of 0.16% Cu between 350 and 450 feet. At 450 feet, a dacite porphyry sill, which persisted to a depth of 690 feet and is evidently the source of the copper mineralization at Buckskin Southwest, was intercepted. Beneath this sill and extending to the bottom of the hole, drilling encountered 600 feet of strongly altered (varying degrees of quartz-sericite-chlorite) volcanics which contained seven spotty, 10-foot intervals of 0.10 to 0.16% Cu. Nearly all of these occurrences in the lower volcanic section are directly associated with narrow dacite porphyry sills. Again, as was noted in the drilling results of BSW-7, the highest copper values are closely associated with quartz-chlorite-biotite altered volcanics and relatively unaltered dacite sills directly beneath.



Log

BSW 1

(35)  
Item 15

## ROCK DESCRIPTION

0-292' INTRUSIVE BRECCIA, intensely qtz.-sericitized, lt. greenish to brownish-gray. Relict fragmental texture almost totally obscured by alteration. Base of oxidation at about 110'. Base of supergene enrichment at about 150'. Where unoxidized and unenriched, rock is uniformly a fine-xline., qtz.-ser. aggregate w/4-7% diss., fine-xline., anh.-subh. py, 0.1-0.3% diss., fine-xline., anh. cpy, 1-3% diss., fine-xline., lt.-med. grayish-brown, anh.-subh. andalusite, tr. diss., fine-xline. anh. magnetite and hematite, and a few narrow (less than 1 mm.) transl. qtz. and lt.-med. grayish-brown dolomite (or ankerite) vnlt. Where supergene enriched (110-150'), py and cpy are partially replaced by dark gray, sooty cc (avg. 0.1-0.2%). Within zone of oxidation (0-110') ser. is partially replaced by white to lt. bluish-gray kaolinite: Sulfides in this interval have been thoroughly oxidized, with resultant formation of lt. yellowish-brown jarosite (as stockwork vnlt., often admixed with qtz., and as fine-xline. aggregates partially filling sulfide casts), and local traces of goethite, pitch limonite, and hematite: A few, narrow (less than 1 mm.), powdery, light grayish-blue-green turquoise vnlt. are erratically distributed throughout this interval.

## GEOCHEMISTRY

Interval	Cu (ppm)	Mo (ppm)	Ag (ppm)
0-10'	245	28	1
10-20'	470	14	4
20-30'	515	3	1
30-40'	400	1	-1
40-50'	365	16	-1
50-60'	305	14	1
60-70'	170	12	1
70-80'	55	15	2
80-90'	80	8	3
90-100'	340	17	3
100-110'	390	14	2
-----Base of Oxidation-----			
110-120'	965	10	2
120-130'	0.22%	13	1
130-140'	655	6	1
140-150'	715	8	3
--Base of Supergene Enrichment--			
150-160'	270	16	9
160-170'	350	16	6
170-180'	465	15	7
180-190'	380	12	8
190-200'	395	11	4
200-210'	390	14	3
210-220'	345	7	3
220-230'	155	1	1
230-240'	430	1	1
240-250'	375	8	2
250-260'	215	6	1
260-270'	165	1	1
270-280'	380	13	4
280-292'	235	10	3

All values reported in ppm unless otherwise noted.

Figure 5. Summarized rotary drill hole log, BSW-1.



S P O K A N E      O F F I C E

Page 1 of

Dip

E

Total Depth

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DEPTH & SIZE	Graphic Geology	ALTERATION					MINERALIZATION					EST. TOT. SULF. Vol. %	GEOLOGY		NOTES	CORE REC.	ANALYSES							Sample No.
		Sil	Ser	Clay	Chl	Zn	Py	Ms	Other	Int.	Description		Int.	Cu			Mo	Pb	Zn	Au	Ag			
355														UNUSUAL BRECCIA & VOLCANIC FLOWS CONT. - Original texture still observed by attention.	-357' TS	100%	350 to 360	10%	26	20	70	-.1	2	26207
365														353-359' - Abundant quartz veining (barren) 6 per foot 359' - 1" wide barren fr. vein. 359-363' - Zone of silty secondary biotite 363' - Zone of very weak alteration - original texture of rock is more 372' clearly visible. There is porphyry ophitic with most of the silica phases altered to sericite. Rock type is probably a little intermediate, i.e. most of the drill core previously logged would seem to be of a TYPE PORPHYRY with some brecciated bands. Some sparse secondary biotite is also present.	-370.5' TS	100%	360 to 370	11%	23	20	130	-.1	2	26208
375														372' - Degree of Silicification increases again.		100%	370 to 380	14%	19	30	130	-.1	4	26209
385														376' - Zone of quartz flooding and 7mm. gr. vein with fine grains of <u>Ms</u> 376-378' Core badly broken. Cpy may be intergrown with fine grained pyrite? 384' - Zone of large secondary biotite blebs.		100%	380 to 390	17%	42	30	195	-.1	6	26210
395														388' - 1' zone of extremely intense clay sericite alteration with abundant secondary biotite.		100%	390 to 400	860	10	20	120	-.1	2	26211
405														396' - Evidence of several stages of quartz veining - one is truncated and another cuts the truncated wht. Abundant secondary biotite blebs.		100%	400 to 410	13%	15	30	150	-.1	3	26212
415														402-408' - Possible fragmental intrusive rock in fine grained silicified volcanic member 1 to 6 mm quartz units with magnetite rims. Also traces of Cpy and <u>Ms</u> . Blebs of secondary biotite. A member of the quartz units are truncated and offset by chlorite - pyrite white 402' - Abundant secondary biotite NOTE: Intrusive appearing zones in the core with abundant secondary biotite are becoming more common. These zones are not necessarily fragment but could be ductile alteration obscuring original. Possibly hydrothermal contact of volcanic rock with intrusive rock. 412-414' - Apparent ductile of <u>Pt</u> & <u>Chry</u> 417-421' - Moderate quartz veining - Considerable magnetite also, with this veining. The area of silty, yellow, argillaceous, irregular zones of abundant secondary biotite blebs.		100%	410 to 420	16%	10	30	170	-.1	3	26213
425														Core badly broken			420 to 430	26%	32	40	210	-.1	3	26214



DEPTH & SIZE	Graphic Geology	ALTERATION		MINERALIZATION		EST. TOT. SULF. Vol. %	GEOLOGY		NOTES	CORE REC.	ANALYSES							Sample No.	
		Silic	Sericite	Clay	Qtz		Py	Fe			Int.	Description	Int.	Cu	Mo	Pb	Zn		Au
430								<u>VOLCANIC ROCK CONT. (Aikard)</u> - numerous zones of intrusive crystalline dike-like approaching contact 3 quartz veins per foot (lower) Considerable secondary brt. (strong) magnetite asax. w/ quartz units Most quartz veins are barren			430 to 440	.31%	29	20	145	-.1	4	26215	
440								<u>QUARTZ DIORITE PORPHYRY</u> 20-25% white phenocrysts (.25-2mm across) of plagioclase 10% phenocrysts of altered (chlorite) hornblende (.5-1.5mm across) 10% phenocrysts of partially altered biotite (.5-2mm across) 55% groundmass which contains considerable quartz and feldspar <u>ALTERATION</u> - Hornblende phenocrysts are strongly altered to chlorite and the feldspar phenocrysts are weakly altered to clay-sericite. Biotite grains are very weakly sericitized.	440.5'- to 443.3'	0%	440 to 450							2	26216
450								<u>MINERALIZATION - NONE!</u> 440.5' - 451' - Interval of alternating contacts between VOLCANICS and QUARTZ DIORITE PORPHYRY. QUARTZ DIORITE PORPHYRY disminutely cuts VOLCANICS (i.e. truncate structure and quartz veinlets in VOLCANICS). HOWEVER VOLCANICS contains abundant fragments of definite intrusive rock. They may represent the pre-volcanic intrusive whose fragments were incorporated in the overlying volcanic flow (not intrusive breccia). Then flow cut by Quartz Diorite Porphyry.		Tr.	450 to 460	90	-1	10	125	-.1	-1	26217	
460								minor fractures with clay-chlorite filling 461.5' 3mm quartz vein - contains minor pyrite Magnetite disseminated magnetite has become apparent.		Tr	460 to 470	165	4	10	115	-.1	-1	26218	
470								469' 1-2mm quartz vein with trace pyrite & trace cpy 471.5' - Conspicuous Cpy, Py, quartz vein (1mm wide) Numerous hairline veins with py, cpy		Tr	470 to 480	50	-1	10	70	-.1	-1	26219	
480								476' - Qtz, cpy vein. 476.5' - Fracture - vein surface with py & cpy - Also possible secondary K-spar. MINOR EPIDOTE? is associated with most Qtz, cpy, py units. Weak clay-sericite alteration becomes apparent - Core badly fractured most alteration (clay-sericite) associated with fracturing.		Tr	480 to 490	45	-1	10	90	-.1	-1	26220	
490								486.5' - 2-4mm open space quartz vein with pyrite and a trace of m.s.s., cpy and sphalerite. 488' - 1" wide fault zone.		Tr	490 to 500	70	4	10	115	-.1	-1	26221	
500								493' 3-6mm open space quartz vein with weak pyrite. Pyrite replaces occasional minor mineral adjacent to quartz vein. 495-498' - Zone of weak clay-sericite alteration - associated with extensive fracturing slight silicification evident in a few zones. Original texture is disrupted partially. Disseminated pyrite is present. Silicification is mostly in the form of quartz float on adjacent to fault. Minor brecciation is also present along fault. A quartz vein is also present. Clay fault zone with thickens present. Pyrite mostly subradial on the flanks of the fault zone.		Tr	500 to 510	110	2	10	125	-.1	-1	26222	



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DEPTH & SIZE	Graphic Geology	ALTERATION										MINERALIZATION										EST. TOT. SULF. Vol. %	GEOLOGY		NOTES	CORE REC.	ANALYSES								Sample No.																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																			
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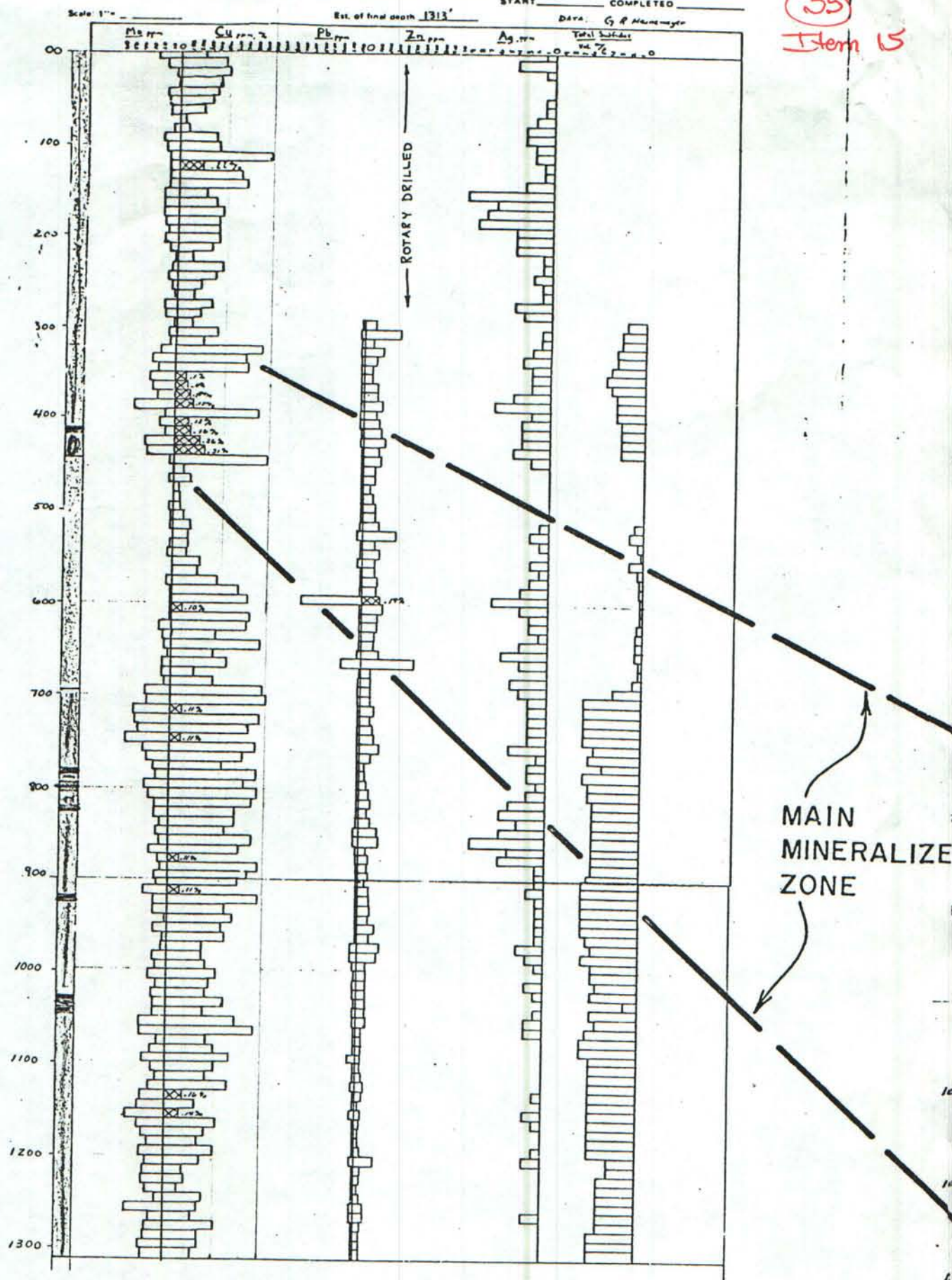
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35  
Item 15



### EXPLANATION

- |                  |   |
|------------------|---|
| Qal              | Alluvium  |
| Qol              | Older Alluvium  |
| Ts               | Semi-consolidated tuffaceous sedimentary rocks  |
| Trt              | Rhyolitic ash-flow tuff   |
| Tqlp             | Biotite, Hornblende, Quartz latite porphyry flows   |
| Kdp <sub>1</sub> | Hornblende biotite (Kdp <sub>1</sub> - syn-mineral)   |
| Kdp <sub>2</sub> | Dacite porphyry (Kdp <sub>2</sub> - post-mineral)   |
| Qz               | Quartz - Feldspar porphyry (Intrusive)<br>(Indeterminate original composition)              |
| Kgrd             | Hornblende granodiorite   |
| Int              | Intermediate Metavolcanic Rocks<br>(Predominantly flow breccias) (Including metaquartzites) |



SCALE 1"=20'  
START \_\_\_\_\_  
REMARKS \_\_\_\_\_

LOGGED BY lmo  
COMPLETION 1972

COORDINATES \_\_\_\_\_  
DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
BEARING \_\_\_\_\_

PROJECT \_\_\_\_\_  
ELEVATION \_\_\_\_\_

Page 1 of \_\_\_\_\_  
DRILL HOLE # BSW 1  
TOTAL DEPTH 1313  
INTERVAL 294 to 430

ASSAYS

LITH.

GEOLOGY

Rock Type, Alteration, Mineralization, Structure

ALTERATION

PERVASIVE VEIN

(35)  
Item 15

290

306

10

20

30

40

50

60

70

80

90

400

10

20

30

MV<sub>u</sub>

MV<sub>u</sub>

FPA?

MV<sub>u</sub>

MV

294 - 369 MV + MV<sub>u</sub> ungt. fine grained volcanic w- 2-5% <sup>vtg</sup> diorite py - light grey color  
Alt - chl - ~~or~~ <sup>or</sup> ~~ser~~  
- some ptg veining - uncondensed  
- 319 - some brown hot spring deposits?

369 - 373 ± FPA (MV?) possible contact at 369 - but can't see it at ~ 373 where it apparently  
grades back into MV or MV<sub>u</sub>. Reminds FPA in BSW 7 but not certain  
perhaps it's a cooling unit

373 - 430 MV as before

Foliation - 35-40°

ser closed

M

W?

M

M

M

1-2



SCALE \_\_\_\_\_  
 START \_\_\_\_\_  
 REMARKS \_\_\_\_\_

LOGGED BY \_\_\_\_\_  
 COMPLETION \_\_\_\_\_

COORDINATES \_\_\_\_\_  
 DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
 BEARING \_\_\_\_\_  
 DIP \_\_\_\_\_

PROJECT \_\_\_\_\_  
 ELEVATION \_\_\_\_\_

Page 2 of \_\_\_\_\_  
 DRILL HOLE # BSW 1  
 TOTAL DEPTH \_\_\_\_\_  
 INTERVAL 430 to 570

%	ASSAYS	LITH.	GEOLOGY Rock Type, Alteration, Mineralization, Structure	ALTERATION	
				PERVASIVE	VEIN
41		40	430-440 Mvu as before py-cpx veins 438 - hem. also	sericite	clay
18%		40	440-441 BLP as described below	M	
		40	441-443 Mvu as above		
		450	443-570 BLP - Biotite late Porphyry (half way between BLP of BSW 7 & PBL of BSW 7) ctc ~ 10° to CA Biot 1-2 1/2 mm - 3-5% Fsp 15-25% 1-7mm possibly some small hmb? < 1mm rare gtz phos Alt - weakly sericitized plagi - matrix fresh looking scattered chl after Biot Ming - < 1/2% py + no veins - py all on fractures. 476 & 482 - another phase of the porphyry or possibly late pre-crystalline movement in the xtal matrix - anywhere 1/2 to 1" zones of lighter porphyry cut across foliation 444 - dominant foliation ~ 35° + much zone 20° to CA in opposite sense.	W W	
		80	CPX veins w/ K sp some hmb present up to 1"		
		90			
		500			
		10			
		20			
		30			
		40			
		550			
		60			
		70			
		80			
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		540			
		550			
		560			
		570			



GEOLOGY		ALTERATION	
Rock Type, Alteration, Mineralization, Structure		PERVASIVE	
		VEINS	
70	570-693 as before but more qtz veining → strong stockwork 589-590		
80	Alt - thin pink schreger - possibly Kspas. Some chl schreger along fractures		
90	Structure - numerous micro faults - offset 1/4 to 1" of qtz etc veins. Fault dip 20-35° N40°E		
600	600		
10	10		
20	Texture becomes slightly finer grained - more brot + fewer feldspar		
30	30		
40	Very few qtz veins		
650	650		
60	60		
70	70		
80	80		
90	684-693		
700	693-710 MV - 1/2 g silicified volcanics - py - vfg dis. - 3-5%		



SCALE \_\_\_\_\_  
START \_\_\_\_\_  
REMARKS \_\_\_\_\_

LOGGED BY \_\_\_\_\_  
COMPLETION \_\_\_\_\_

COORDINATES \_\_\_\_\_  
DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
BEARING \_\_\_\_\_

DIP \_\_\_\_\_  
ELEVATION \_\_\_\_\_

PROJECT \_\_\_\_\_

Page 4 of \_\_\_\_\_  
DRILL HOLE # B5W1  
TOTAL DEPTH \_\_\_\_\_  
INTERVAL 710 to 851

ASSAYS

LITH.

GEOLOGY

Rock Type, Alteration, Mineralization, Structure

ALTERATION

PERVASIVE VEIN

ser. Al Ag Si Q Cal

710

710-778 MV greenish grey if g. Qtz veins more common 3-4/foot up to 7mm wide  
strong may throughout volcanics except in sericitized & silicified zones.

200

MV

30

40

750

ser

60

MV

70

80

Blp

778-780.5 Blp (?) Strongly sericitized & bleached - can't see contacts - rock broken & filled  
w/ waxy calcite on bottom side - so probably faulted in. MV is sericitized also  
on entire side

90

MV

780.5-819 MV as before

6

Blp

795-6" Blp dike c/c 3.5°

800

Blp

799 - " " w/ inclusions of MV

Blp

801-802 " " dark green same chl as veins

10

MV

814 - steep fault? - displaces MV against a foliose intense-looking rock - not massive because veins  
816 other side of fault of different orientation on either side end abruptly at fault which is filled by  
18" Qtz veins. no gouging or clay.

14

MV

16

MV

20

Blp?

819-821? Blp dike altered strongly - ser w/ some superimposed silica-sericite along a Qtz-calves.

821-827 Bxa or fault Bxa - too altered to discern

30

MV?

827-834 Bxa - volcanic matrix - subvolcanic intrusion? or fault bxa? can't tell. could even be a  
slow bxa.

40

MV

834-846 f.g. slightly porphyritic bleached andesite? - appears to be the matrix for the bxa above.  
resembles a finer grained FPA.

50

MV

846-851 Bxa - green space - filled w/ waxy calcite & dolomite?

50

MV?







ASSAYS				LITH.	GEOLOGY Rock Type, Alteration, Mineralization, Structure	ALTERATION						
						PERVASIVE				VEIN		
						Sr	Chl	Ans	Sil		Cal	gyp
1-2				990	990-1027							
					MVb (FPA matrix??) as before							
					Ser, Elongated fags - 35° dip							
				1000								
				10								
				20								
1-90				30	1027-1028 BLP dike hem on fract							
					sericite mod to strong							
					1028-1034 mvb as before							
				40	1034-1042 BLP - sericitized							
					can't see etc.							
				50	1042-1055 faulted - sericitized - argillaceous							
3-4												
				60	1055-1089 mvb silicified & sericitized							
				70								
2-3				80								
				90	1089.5-1090 - BLPg - strongly chloritized - fuzzy gm - silicified?							
					reminds BLP but I don't think it is.							
					1091-1091.6 - " as above							
				1100	1091-1100 mvb as before							
					Some steep gyp veins							
					1100-1114 FPA as before							
				10	1114-1130 mvb as before							
				20								
				30								



SCALE \_\_\_\_\_  
START \_\_\_\_\_  
REMARKS \_\_\_\_\_

LOGGED BY \_\_\_\_\_  
COMPLETION \_\_\_\_\_

COORDINATES \_\_\_\_\_  
DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
BEARING \_\_\_\_\_  
DIP \_\_\_\_\_

PROJECT \_\_\_\_\_  
ELEVATION \_\_\_\_\_

Page 7 of 8  
DRILL HOLE # 135W1  
TOTAL DEPTH \_\_\_\_\_  
INTERVAL 1130 to 1270

ASSAYS

LITH.

GEOLOGY

Rock Type, Alteration, Mineralization, Structure

ALTERATION

PERVASIVE VEIN

1130

1130-1185±

Mvb as before - more gyp - light gray color - Ser? or S.I-Ser

40

1150

60

70

80

90

1200

1206±

FPA or possibly MV - gradual etc

some alteration - 35°

same mvk mixed in as at 1261

Alt - ? bleached - no map - some Ser - & occasional silica-pervasive

Stray gyp

20

30

40

1250

60

Fault 1260-1261 zone

70



SCALE \_\_\_\_\_  
START \_\_\_\_\_  
REMARKS \_\_\_\_\_

LOGGED BY \_\_\_\_\_  
COMPLETION \_\_\_\_\_

COORDINATES \_\_\_\_\_  
DRILLER \_\_\_\_\_

LOCATION \_\_\_\_\_  
BEARING \_\_\_\_\_  
DIP \_\_\_\_\_

PROJECT \_\_\_\_\_  
ELEVATION \_\_\_\_\_

Page 8 of 8  
DRILL HOLE # BSW1  
TOTAL DEPTH 1313  
INTERVAL 1270 to 1313

ASSAYS

LITH.

GEOLOGY

Rock Type, Alteration, Mineralization, Structure

ALTERATION

PERVASIVE VEIN

1270

1270-1302

MV & MVB - on before

80

90

1300

1302-1317 FPA? or MV

10

moly

Seems like we've gone thru a sequence of flows & flowbas stacked on top of one another w/ the occasional sill & dike of much younger age.) FPA is prob a coarser grained volcanic that may also sill out. Gyp has increased steadily & is strongest at bottom of hole. Only moly found is at bottom of hole on a qb-moly hauling fracture.

If this hole was ever to be interesting it would have to be deepened drastically. We are too high in the system. (and probably off target to boot.)

lmo



NORTHWEST DISTRICT  
Summary Drill Hole Log

Name BUCKSKIN SOUTHWEST

Code 005-01-0411

DOUGLAS CO., NEV.

DRILL HOLE NO. BSW-1 PAGE      OF     

BEARING      DIP 90°

LOCATION: N.      E.     

COLLAR ELEV. ≈ 5560'

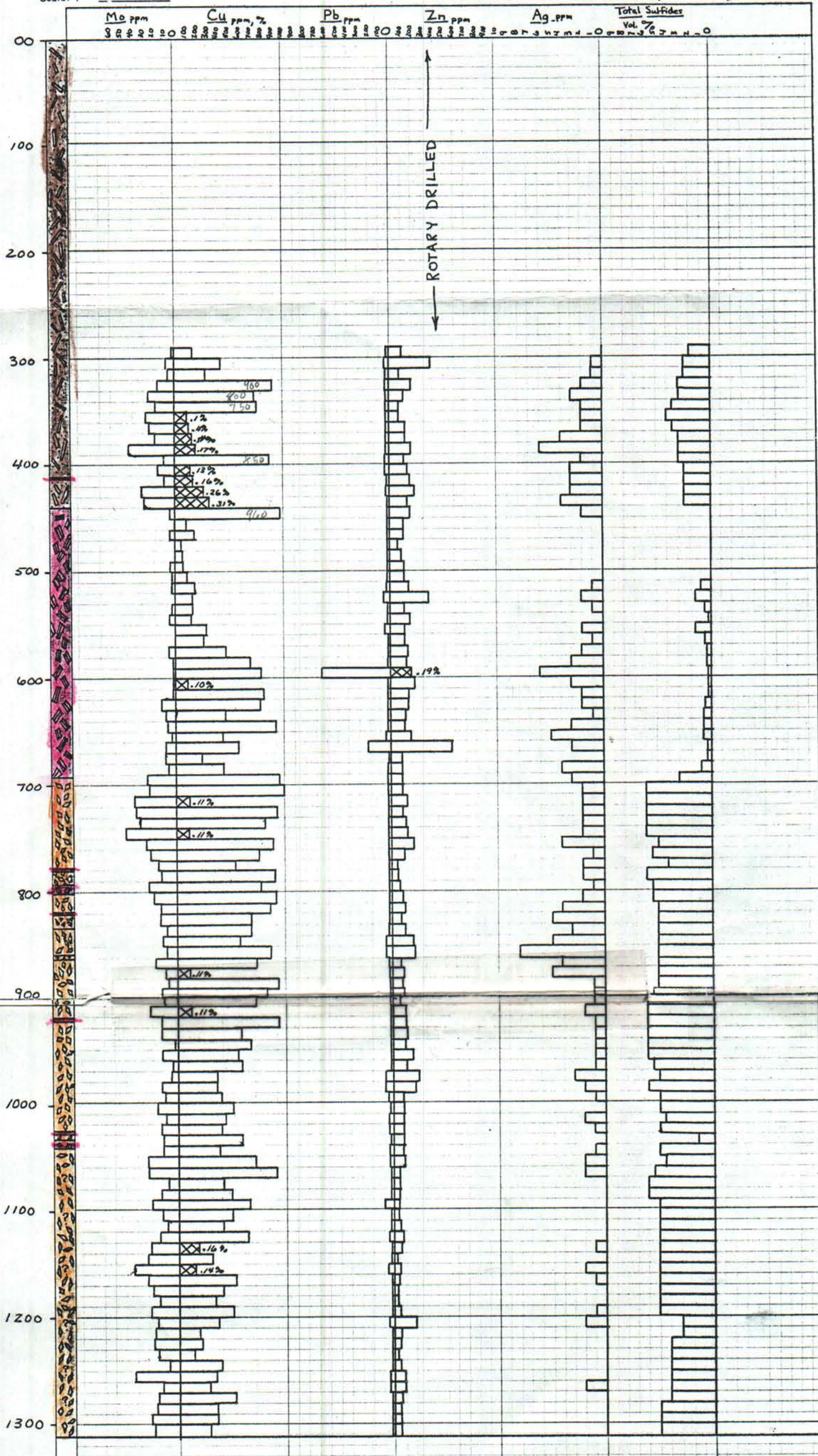
TOTAL DEPTH      BY     

START      COMPLETED     

Scale: 1"=

Est. of final depth 1313'

DATA: G. R. Heinemeyer





BSW#1

0-10' Qtz - ss. rock:

0.7-1% diss. dk. gray, partially metallic(?) mineral

irreg. grains < 0.5 mm. dia. — streak

mostly dk. grayish-red, but some relict

gray metallic mass — ~~some~~ may be almost

completely oxidized cc. (hem. w/ tiny cc. relicts)

minor turquoise and/or variscite

→ minor kaol. in body of rx; montmorillonite along fracs.

minor — ~2-5% pitch quartzite along fracs; dk. reddish-

brown to black — often reniform incrustations.

→ 2-5% casts, most < 0.5 mm. in dia. — most irreg.

some partially filled w/ bright yellowish-brown jarosite  
↓ or often ~~sub~~ translucent.

10-20' ~45-55% of sample is translucent dk. bluish-gray kaolinite

w/ ~5% diss. irreg. grains — <sup>dk.</sup> dark bluish-gray metallic

mineral same as above — use more of the grains

appear to be relict cc. — rest of sample same

as above

20-30' clay - ss. rx. w/ minor Qtz. — minor pitch <sup>reddish</sup> pyrite (same

as 0-10' [prob. < 0.5%]); dk. gray metallic mineral decrea-

sing (prob. ~0.5-0.7%) → prob. new addition is

clay - ss. rx w/ 2-5% silf. casts, partially filled w/ jar. (same

as 0-10')



BSW #1

30-40' - new appearance - clay rock, lt-med, gray - traces  
of reddish-brown - poss. a few relict feldspar  
phenocrysts - rock is probably altered rhyodacite  
intrusive ~~by~~ <sup>to</sup>; ~0.5% pitch limonite along fracs. -  
r/p (same as 0-10'); < 0.5% metallic gray mineral  
(same as 0-10')

40-50' - same as 30-40', exc. ~1-1.5% bluish-gray metallic  
mineral; irreg. grains up to 1 mm. in dia. -  
in hem.; rest is comprised of frags - of 10-40'

50-60' - gtz-se. <sup>rx</sup> w/minor clay - 3-5% sulf. casts - same as  
0-10' - < 0.3% metallic gray mineral (same as 0-10')  
rest of sample - (frags - 10-50')

60-70' same as 50-60' exc. only tr. of med. reddish-gray mineral  
prob. ss. > gtz.

70-80' same as 60-70', exc. 2-3% sulfide <sup>(casts)</sup> & a few tiny  
gtz. v. small < 0.5 mm. in dia.



80-90': same as 70-80' - exc. only tr. cc(?)

tr. qtz. veinlets (same as 70-80')

90-100': same as 80-90', exc. 5-7% sulfide casts - most  
partially filled w/ jarosite<sup>+</sup> - but pubs n.o.s.  
slightly hematitic; \* perhaps limonite  
becoming slightly more goethitic (up to 25% G)

100-110' same as 90-100' except a few clumps of 5% sulfide casts  
up to 1.5% dk. gray metallic (same distribution as  
irr. sulf. casts)

110-120' appearance of ~4% py, 0.5-0.7% cc(?)  
some chunks w/ up to 3% dk. gray metallic

120-130' <sup>SPF 814</sup> mottled clay-(qtz)-chlorite rock w/ 2-3% diss.  
<sup>mottled dk. bluish gray to greenish  
lt. gray</sup>

anh.-subh. py, 0.5-0.7% py along tiny fractures

(very dist.) ~0.5% diss dk. gray metallic minrl. ? (consp.)

pieces easily w/ knife

130-140' same as above

140-150' same as 120-130'



150-160' clay-(sl.?) -qtz. rock w/ ~5% py, 0.3-0.5% cc(?)

anyway dk slt, brownish to bluish gray metallic drs, irreg,  
clot up to 1.5 mm. dia.

160-170' same as above exc. uniform ~0.5% cc(?)

170-180' same as above exc. 1 dot cpy ~0.5 mm. dia.  
(anhedral). & prob ~0.5-0.7% cc(?)

180-190' same as above

190-200' same above except a few pc. clau-grt. chlor.  
rich - a few chips of dk. gray metallic mineral  
w/ reddish streak

200-210' same as above, except ~0.3% metallic mineral



(35)

Item 15

RECONNAISSANCE GEOLOGY  
OF THE  
BUCKSKIN RANGE, NEVADA

Donald M. Hudson  
December, 1977



(35)  
Item 15

RECONNAISSANCE GEOLOGY  
OF THE  
BUCKSKIN RANGE, NEVADA

Donald M. Hudson  
Continental Oil Company  
Minerals Department  
Reno District  
December, 1977





Figure 1. Looking generally northwest at several low angle Basin and Range faults in the southern Buckskin Range. Note that the attitude of the of the faults vary from east dipping to west dipping. Explanation of symbols is given on plate 1.



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| 4 | Low angle Basin and Range faults in the northern Buckskins                 | page 23 |
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### Plates

- |    |  |           |
|----|--|-----------|
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| 1B | Generalized stratigraphic column of the Buckskin Range | "         |
| 1C | Intrusive column of the Buckskin Range                 | "         |
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## SUMMARY

Approximately 18,000 feet of stratigraphic section is exposed in the Buckskin Range in west-central Nevada. Metavolcanic rocks, carbonates, and clastic sediments of the Upper Triassic Oreana Peak Formation and carbonates and sediments of the Lower Jurassic Gardnerville Formation are overthrust by andesitic metavolcanic rocks of the Lower Jurassic Artesia Sequence and dacitic to quartz latitic metavolcanic rocks of the Lower Jurassic to Lower Cretaceous Churchill Canyon Sequence. The Lower Mesozoic rocks are intruded by Jurassic granodiorite and quartz monzonite.

Oligocene ash-flow tuffs of the Guild Mine Tuff, Weed Heights, Singatse Tuff, Nine Hill Tuff, and Eureka Canyon Tuff unconformably overlie the Mesozoic rocks. Miocene andesites, Plio-Pleistocene lake sediments, and Quarternary alluvium unconformably overlie the ash-flow tuffs. Numerous varieties of Miocene andesitic to dacitic dikes and stocks intrude the older rocks of the Buckskin Range.

Gently dipping Basin and Range faults which underlie steeply westward dipping Tertiary ash-flow tuffs and Mesozoic metavolcanic rocks, were originally steeply dipping normal faults. The strata were probably tilted westward by the low angle faults and/or by later steeply dipping normal faults.

Most of the rocks in the Buckskin Range were hydrothermally altered to propylitic, sericitic, and aluminophyllic alteration facies by Miocene alteration events and to a lesser extent by Mesozoic alteration events.



## INTRODUCTION

### Location:

The Buckskin Range is in Douglas and Lyon Counties in west central Nevada, approximately 90 road miles southeast of Reno and 10 miles west of Yerington. The Buckskin Range lies between the Pine Nut Range to the west and the Singatse Range to the east, in Townships 13 and 14 north, Ranges 23 and 24 east (Figure 2). Dirt roads, which are periodically graded, ring the range and a paved road approaches the range at the Minnesota Mine from U.S. 95 to the east. Very few roads penetrate the range. The climate is arid with very few trees and mostly low brush. The range has relatively low relief with a maximum elevation of 6,915 feet.

### History and Previous Work

According to some reports the Buckskin district was discovered in 1904 by a prospector who found gold in veins on or near the present "Red Top" patented claims. The mine proved to be unsuccessful. Pyrophyllite was explored at the Blue Danube Mine, but was impure and was not mined. Iron ore was mined from the Minnesota open pit mine from 1944 to 1966 and produced nearly \$17,000,000.

Exxon, Anaconda, Bear Creek Mining Co., Phelps Dodge, and Conoco have drilled the southern and eastern portions of the range (drill hole locations shown on Plate 3). Anaconda currently holds unpatented claims on most of the southern and eastern Buckskins. Conoco holds claims in the south central Buckskins. There are several small blocks of unpatented claims held by prospectors in the northern Buckskins.

Published geologic mapping was done by Moore (1969) and Proffett (1977) and there is an unpublished thesis by Castor (1972). Bear Creek Mining Co.





Figure 2. Location map of the Buckskin Range and the surrounding region in west-central Nevada (scale 1:250,000). Outlined area is area shown on plate 1.



and Anaconda have mapped portions of the Buckskins but their maps have proven to be largely inaccurate due to the complex Tertiary history of the Buckskin Range and the Yerington area. The Conoco holdings and immediately surrounding area was mapped in detail by Conoco geologist W. M. Oriel in 1976. John Geisman of Stanford University is mapping the range as a thesis project, with support from Anaconda.

#### Methods of Investigation

The Buckskin Range was mapped at a scale of 1:12000 using blown up U. S. G. S. Como and Wellington 15' quadrangle topographic maps. Aerial photographs at 1:24000 and 1:6000 were mapped on for most of the area. About 120 thin sections and several X-ray diffraction patterns were used in defining stratigraphy, mineralogy, and alteration mineralogy.

The igneous rock classification of Travis (1955) is used in this report. The Tertiary time scale used by Proffett and Proffett (1976) and the Mesozoic time scale of Van Eysinga (1970) are used in this report.

#### Acknowledgement

Ken Howard of the Anaconda Co. gave permission to map on Anaconda claims. Edward C. Bingler of the Nevada Bureau of Mines was extremely helpful in correlating Mesozoic and Tertiary stratigraphy. Bill Oriel of the Conoco staff was very helpful in orientation to the Buckskin Range and in constructive criticism during the course of the work.



## STRATIGRAPHY

### Mesozoic Rocks

#### Oreana Peak Formation (R o)

The base of the stratigraphic section exposed in the Buckskin Range, along Churchill Canyon (Plate 1), is the Upper Triassic (Norian) Oreana Peak Formation of Noble (1962). About 1400 feet of Oreana Peak is exposed in two main units.

The lower unit (R op) consists of at least 1200 feet of interfingering beds of volcanic rocks and conglomerate (R opv) and limestone (R opl) in beds a few inches to 25 feet thick. The volcanics consist of felsic pyroclastics, finely varved water-lain tuffs, tuffaceous sediments, and dense aphanitic basalts with elongate amygdules up to 20 cm long and 2 cm thick. Coarse conglomerates and massive grey limestones are interbedded with the volcanics.

The overlying carbonate unit (R oc) of the Oreana Peak Formation (Upper carbonate member of Noble, 1962) contains about 200 feet of massive-to-medium bedded grey limestone with 1/2 to 3 inch, (often pinkish) siltstone partings. The limestone locally contains abundant fossil fragments. It is occasionally oolitic and locally dolomitic.

#### Gardnerville Formation (Jg)

The Gardnerville Formation (Noble 1962) of uppermost Triassic and Lower Jurassic age, conformably overlies the Oreana Peak Formation. One hundred and fifty to 1200 feet of Gardnerville Formation is exposed in the northern Buckskins (Plate 1). The sequence is composed primarily of grey, locally pyritic, ashy(?), calcareous siltstone interbedded with thin green-grey locally pyritic shale (Jgs). The siltstones break parallel to bedding and



contain impressions of ammonites. The Gardnerville contains several discontinuous lensoidal or patchy massive grey limestones that locally contain "fossil hash" and are locally oolitic. The limestones are from 1 to 20 feet thick. Occasionally the limestones contain thin interbeds of crossbedded calcarenite.

Contact metamorphism of the Gardnerville Formation by adjacent Mesozoic and Tertiary intrusions converts the siltstones to a dense dark grey hornfels with a fine varved "chert" appearance. The hornfels fractures across bedding. The limestones are strongly recrystallized though rarely marbleized. The break between hornfels and apparently unmetamorphosed Gardnerville can be very abrupt.

#### Artesia Sequence

The Artesia Sequence of probable uppermost Lower Jurassic age is named for Artesia Mountain in the central Buckskin Range. About 2000 feet of the Artesia Sequence is exposed south of Churchill Canyon (Plate 1) while up to 6000 feet of the Artesia Sequence is exposed in the southern Buckskins. The Artesia Sequence lies in thrust contact with underlying rocks. The majority of the Artesia Sequence is composed of dark lavender, red-lavender, and greenish andesitic to dacitic volcanic flows (Jav) with occasional volcanic breccia and conglomerate-sandstone interbeds. The volcanics are typically porphyritic with 10 to 40% phenocrysts up to 3 mm in length; however, they generally are less than 1 mm long. Phenocrysts of pyroxene (?), hornblende, biotite and weakly to strongly zoned plagioclase (An 55-35) occur, with rounded quartz phenocrysts up to 0.5 mm in some flows in the upper portion of the sequence. Some flows completely lack phenocrysts.

The aphanitic groundmass contains plagioclase, hornblende, biotite and rare quartz, where the groundmass is not recrystallized by metamorphism or alteration.



Felsic breccias (Jafb) occur strataconformably in about the middle of the Artesia Sequence in the southern Buckskins (Plate 1). The white-to-buff felsic unit which is up to 2000 feet thick, is often flow banded and contains rounded felsic fragments up to 20 cm in diameter. The composition of the unit may be rhyolite or quartz latite.

A thick sandstone and conglomerate unit (Jas) occurs in the Upper Artesia Sequence in the central and southern Buckskin Range (Plate 1). About 500 feet of the clastic unit is present in NE 1/4, Sec. 3, T13N, R23E and it thins to the north, south, and east. It is apparently absent in the northernmost Buckskins. The clastic unit is coarsest in the west central Buckskins with clasts up to 15 cm. The clast size becomes progressively finer as the unit thins. The unit contains rounded clasts of Artesia volcanics grading upward into feldspathic sandstone in usually well defined graded beds 1 cm to 1 meter thick.

A series of andesitic to dacitic volcanic flows up to 1200 feet thick (Fulstone Formation of Anaconda terminology) overlies the clastic unit in the central Buckskin Range. These flows thin rapidly to the north and south and are completely absent in the northern Buckskins. These flows are essentially the same composition as the rest of the Artesia Sequence except that the phenocrysts are up to 5 mm in length.

Thick, discontinuous, nearly strataconformable quartz "reefs" are a distinctive part of the Artesia in the central and southern Buckskins (Plate 3). Many of the quartz "reefs" contain breccias or have protobrecciation. Original textures are usually obscured. The quartz "reefs" rarely occur above the thick clastic unit (Jas). They probably represent epigenetic selective silification associated with local hydrothermal activity.



The Artesia Sequence is regionally metamorphosed to the greenschist facies containing epidote, sericite, calcite, chlorite and clays with slight albitization of the plagioclase.

#### Churchill Canyon Sequence (Jcc)

The Churchill Canyon Sequence, named for exposures just south of Churchill Canyon (Plate 1), conformably overlies the Artesia Sequence. Possibly as much as 3000 feet of Churchill Canyon Sequence is exposed in the Buckskin Range. A thin conglomerate unit (Jccs) usually lies along the contact between the Artesia Sequence and the Churchill Canyon Sequence. The conglomerate is 1 to 3 feet and rarely up to 7 feet thick. It contains clasts of both Artesia and Churchill Canyon lithologies.

Lithologically, the Churchill Canyon Sequence consists of dacitic to quartz latitic flows and crystal rich tuffs (Jccv) with associated porphyritic intrusions (Jcci) of essentially the same composition. The exposed portion of the sequence contains primarily ash-flow tuffs in the northernmost Buckskins and dominantly flows in the southern and central Buckskins. The ash-flow tuffs and flows interfinger in the north-central Buckskins. Both rock types are present throughout the range although structural and erosional complications expose only a small portion of the Churchill Canyon Sequence in the southern Buckskins. Quartz phenocrysts up to 10 mm in diameter are distinctive of the Churchill Canyon Sequence. The quartz phenocrysts are usually strongly resorbed, rounded and rarely bipyramidal. Not all flows and intrusions contain quartz phenocrysts.

The ash flow tuffs are usually tan to greenish light brown, containing 20 to 60% phenocrysts. Phenocrysts include 0 to 4% hornblende, 1 to 4% biotite, 0 to 10(?) sanidine, 1/2 to 7% quartz. The remaining phenocrysts



are plagioclase (generally AN35-25) with weak oscillatory zonation. Crystals are usually broken and often bent. Flattened pumice fragments up to 5 cm long are often seen in thin section. Tuff breccias occur in the sequence and are composed of up to 50% deformed fragments of tuff that are strongly welded into a tuff breccia.

The volcanic flows of the Churchill Canyon Sequence are grey to grey-green, and generally are thin bedded (5 to 30 feet thick). Occasional conglomerate or breccia interbeds up to 1-1/2 feet thick are present between some flows. The flows contain 5 to 40% subhedral to anhedral, often broken phenocrysts of 1 to 5% biotite, 0 to 5% oxyhornblende(?), 0 to 4% quartz, weakly zoned plagioclase (generally AN30-40) and occasional large pink sanidine. Phenocrysts up to 1 cm and rarely 3 cm in length with hial (sharp break in size between phenocrysts and groundmass) texture are set in an aphanitic groundmass of plagioclase, sanidine(?), and quartz microlites with possible glass.

Intrusions (Jcci) which are very similar in texture and composition to the Churchill Canyon volcanic rocks intrude the Churchill Canyon Sequence. They are difficult to distinguish from the flows except where crosscutting relationships exist. Where intrusives and volcanics are indistinguishable they are mapped as Jccu (Plate 1).

The Churchill Canyon Sequence is regionally metamorphosed to the greenschist facies containing calcite, sericite, epidote, clinozoisite, chlorite and quartz with slight albitization of the plagioclase.

Both the Artesia Sequence and Churchill Canyon Sequence are probably correlative to the Double Springs Formation of Noble (1962). Ammonites collected from the Double Springs Formation near Topaz Lake indicate an



uppermost Lower Jurassic age of the Double Springs Formation (H. F. Bonham, oral comm.). Age dates on hornblende obtained by Castor (1972) from Churchill Canyon Sequence type rocks in the Pine Nut Range yielded dates of 146 m.y. and 124 m.y. The age dates would indicate an Upper Jurassic to Lower Cretaceous age for the Artesia and Churchill Canyon Sequences. However, the dates may have been reset by subsequent intrusions and/or by metamorphism. Although the correlation of the Artesia and Churchill Canyon Sequences with the Double Springs Formation is tentative, an uppermost Lower Jurassic age for the sequences is used in this report.

#### Cenozoic Rocks

The Oligocene is represented by a series of ash-flow tuffs and tuffaceous sediments which are thoroughly discussed by Proffett and Proffett (1976) from the Singatse Range. Only brief descriptions are given here to note local variations in the tuff units.

#### Guild Mine Tuff (Tgmt)

About 1000 to 2000 feet of brown to lavender buff, moderately welded, crystal rich ash-flow tuff of the Guild Mine Tuff unconformably overlies the Mesozoic rocks in the Buckskin Range. The Guild Mine Tuff contains 20 to 35% phenocrysts, up to 2.5 mm in length, of quartz, sanidine, plagioclase, and 1 to 3% biotite. Rare fragments of petrified wood occur at the base of the Guild Mine Tuff, overlain by 5 to 25 feet of black vitrophyre. The vitrophyre is overlain by 700 to 900 feet of moderately welded and occasionally densely welded tuff with moderately abundant small pumice fragments. The top of the Guild Mine Tuff consists of 200 to 300 feet of buff colored vapor phase tuff with abundant large pumice fragments and frequently large irregular blocks of deep red jasper.



#### Weed Heights Tuff (Twht)

Two hundred and fifty to 300 feet of lavender to reddish brown, moderately welded, moderately crystal rich ash-flow tuff of the Weed Heights Tuff conformably overlies the Guild Mine Tuff. The Weed Heights Tuff contains 5 to 25% phenocrysts including plagioclase, sanidine, quartz, and 1 to 2% biotite. Abundant large white pumice fragments are distinctive of the unit. There appears to have been a very short cooling break between the Guild Mine Tuff and the Weed Heights Tuff in the Buckskin Range as evidenced by the lack of intervening sedimentary units which are present in the Singatse Range (Map Units 4 and 5 of Proffett and Proffett, 1976).

Fifty to 75 feet of rhyolitic sediments overlie the Weed Heights Tuff and are mapped with it. The lower portion is usually bright red while the upper portion is yellowish green. The uppermost sediments contain abundant fragments of petrified wood and sparse leaf impressions.

#### Singatse Tuff (Tst)

About 2500 feet of brown to red-brown, locally lavender, strongly to moderately welded, crystal rich ash-flow tuff of the Singatse Tuff conformably overlies the Rhyolitic sediments. The Singatse Tuff contains 30 to 45% phenocrysts of plagioclase, quartz, sanidine, hornblende, and 3 to 6% biotite up to 5 mm in length. Hornblende, abundant biotite, large phenocrysts, and its massive nature are distinctive of the Singatse Tuff. The unit contains sparse pumice fragments and 1 to 4% lithic fragments except local zones slightly above the base where the unit contains up to 30% lithic fragments. Five to 10 feet of black vitrophyre occurs locally at the base.



#### Nine Hill Tuff (Tnht)

About 100 feet(?) of deep lavender to buff, poorly welded to non-welded, crystal poor ash-flow tuff and tuff breccia of the Nine Hill Tuff of E. C. Bingler (oral comm.) conformably overlies the Singatse Tuff. The Nine Hill Tuff is part of the Bluestone Mine Tuff of Proffett and Proffett (1976). The tuff contains less than 5% phenocrysts of plagioclase, quartz, sanidine, and trace quantities of biotite. White flattened pumice fragments up to 20 cm across are distinctive of the unit. Locally, the Nine Hill Tuff is a tuff breccia.

#### Eureka Canyon Tuff (Tect)

Fifty feet(?) of bright red-orange, poorly to non-welded, crystal poor tuff of the Eureka Canyon Tuff of E. C. Bingler (oral comm.) conformably overlies the Nine Hill Tuff. The Eureka Canyon Tuff is also part of the Bluestone Mine Tuff of Proffett and Proffett (1976). The tuff contains sparse pumice fragments and less than 5% phenocrysts of plagioclase, quartz, sanidine, and trace biotite.

#### Andestic Volcanics (Tv)

Miocene andesitic volcanic flows and flow breccias overlie the Oligocene ash-flow tuffs in angular unconformity. More than 1000 feet of andesite is present in north Churchill Canyon while only small exposures crop out on the east flank of the Buckskin Range. More than 5000 feet of andesite is present north of Lincoln Flat (Proffett and Proffett, 1976). Many of the flows are oxyhornblende and augite rich andesitic porphyries (Tvhpa) and probably were derived from a volcanic intrusive center of the same composition and texture in the northern Buckskins (see discussion on Tertiary intrusive rocks).



A few flows of possibly younger oxyhornblende biotite andesite porphyry (Tvhba) crop out south of the Minnesota Mine.

#### Lake Sediments (QTs)

At least 250 feet of pliocene to Pleistocene lacustrine sediments occur on the west flank of the Buckskin Range. These lie with angular unconformity on Miocene to Mesozoic rocks. The lacustrine sediments are white and rarely green, finely to thinly bedded shales, siltstones, and fine grained quartz arenites.

#### Older Alluvium (Qoal)

Fifty or more feet of coarse unsorted, unconsolidated conglomerate of older Quarternary alluvium lie conformably or slightly unconformably on the lake sediments. The conglomerate contains moderately rounded pebbles to boulders. Most of the clasts are granitic material that is foreign to the Buckskins with a lesser amount of locally derived material.



## INTRUSIVE ROCKS

### Mesozoic Intrusions

#### Granodiorite (Jgd)

Fine to medium-grained, equigranular, phaneritic granodiorite is probably the oldest intrusion exposed in the Buckskin Range. It may be equivalent to the Black Mountain Granodiorite in the Singatse and Wassuk Ranges (E. C. Bingler, oral comm.). The granodiorite contains crystals up to 1 mm in length but they are usually 0.7 mm or less. It contains 1 to 3% biotite, trace hornblende, 8 to 12% interstitial quartz, 10 to 15% orthoclase and the remainder unzoned plagioclase locally varying from An<sub>35</sub> to An<sub>45</sub>. South of the Minnesota Mine a small body of coarser grained granodiorite (Jgdc) intrudes the fine granodiorite. The rock is essentially the same mineralogically except that it contains about 10% biotite and 1 to 2% hornblende with crystals up to 5 mm.

#### Diorite

Several exposures of black, fine-grained diorite that may be older or younger than the granodiorite crop out along Churchill Canyon (Plate 1). The diorite contains 4 to 10% augite, 1 to 2% quartz, trace orthoclase, and Plagioclase (An<sub>45-55</sub>) in very weakly zoned crystals.

#### Quartz Monzonite (Jqmp)

A pinkish brown quartz monzonite porphyry intrudes the granodiorite just east of the Minnesota Mine (Plate 1). Subhedral plagioclase (oligoclase?) and biotite phenocrysts up to 5 mm long occur in seriate texture with groundmass minerals up to 0.4 mm. Overall the rock contains 1 to 3% biotite, about 15% quartz, 50 to 55% plagioclase, 30 to 35% orthoclase, 1 to 2% magnetite,



and trace pyrite. The quartz monzonite porphyry is cut by abundant quartz-magnetite-pyrite veins. A coarse-grained quartz monzonite breccia (Jqmb) containing clasts of quartz monzonite porphyry in an aphanitic igneous matrix is exposed just south of the Minnesota Mine (Plate 1).

#### Microgranodiorite (KJmgd)

A small body of buff colored subequigranular microgranodiorite intrudes the Artesia Sequence west of Fulstone Spring No. 1 (Plate 1).

Plagioclase is up to 1 mm long, but the crystals are mainly less than 0.4 mm. The rock contains 15 to 20% anhedral quartz, 10 to 15% anhedral orthoclase, and subhedral to anhedral unzoned plagioclase, (An35-40). The exact age is unknown except that it is upper Mesozoic.

#### Churchill Canyon Intrusives (Jcci)

The Churchill Canyon Intrusives are discussed under Churchill Canyon Sequence.

#### Latite Porphyry Dikes (KJlp)

Latite porphyry dikes intrude the older Mesozoic rocks throughout the Buckskin Range. They contain subhedral 1 to 3 cm K-feldspar phenocrysts with small hornblende, biotite, plagioclase, and about 1% quartz phenocrysts in an aphanitic groundmass. Several dikes in the northern Buckskins are essentially identical in texture and mineralogy but contain about 5% quartz phenocrysts (Kqlp). The dikes are similar in composition to Churchill Canyon Sequence intrusions but they are distinctive enough in appearance to map separately. The latite porphyry dikes could be a later phase of Churchill Canyon intrusive activity, based upon intrusive relationships, but may be considerably younger.



### Tertiary Intrusions

Numerous Tertiary dikes and stocks are present in the Buckskin Range. Relative ages are given where known but otherwise the Tertiary intrusions are assumed to be Miocene based upon intrusive relationship and age dates. Several intrusives in the southern Buckskins are not described here and are described by W. M. Oriel (1976).

### Biotite Pyroxene Diorite (Tbpd)

A few small stocks of dense, black, fine-grained diorite crop out on the north end of the Buckskin Range (Plate 1). The diorite contains 5 to 7% shreddy, greenish-brown biotite less than 0.05 mm in length, usually in magmatic reaction with augite, 7 to 10% subhedral augite up to 2 mm long, less than 2% quartz, weakly zoned subhedral plagioclase (An50-55) up to 2 mm in length, and about 1% magnetite. The rock shows strong foliation in thin section but it is not evident in hand sample.

### Hornblende Pyroxene Diorite Porphyry (Thdpd)

Phaneritic diorite porphyry and locally porphyritic diorite underlies the north end of the Buckskin Range and intrudes the biotite pyroxene diorite. Phases of the diorite vary widely in texture and composition and are generalized here. The rock contains 10 to 40% Phenocrysts of hornblende that are commonly 5 to 10 mm in length but range up to 3 cm in length. Plagioclase phenocrysts are up to 5 mm in length and rarely reach 10 mm. The phenocrysts grade downward in size in a seriate texture to a groundmass generally 0.01 to 0.5 mm. The diorite contains 5 to 25% euhedral to subhedral green to green-brown zoned hornblende with numerous inclusions, 3 to 7% subhedral augite, 2 to 3% quartz, oscillatory normally zoned plagioclase



An50-40), and trace sphene and apatite. The diorite is often foliated with subparallel plagioclase and hornblende phenocrysts.

Numerous andesitic dikes and small stocks related to the hornblende pyroxene diorite porphyry intrude the northern portion of the Buckskin Range. They have variable compositions and highly variable textures. These have been generalized into 2 groups: those with less than 25% phenocrysts (Thppa), and those with 25 to 50% phenocrysts (Thpap). The rocks are similar in composition to the hornblende pyroxene diorite porphyry but have an aphanitic groundmass of plagioclase, augite, and a small amount of quartz, and phenocrysts of plagioclase, hornblende and usually augite. These intrusives are green to greenish-grey and many have weakly developed columnar jointing.

Several intrusive breccias (Timb) crop out in Sec. 22, T14N, R23E. These have a matrix of hornblende pyroxene porphyritic andesite with angular to highly rounded pebbles to boulders of hornblende pyroxene porphyritic andesite, hornblende pyroxene andesite and diorite porphyry, various ash-flow tuffs, and Mesozoic metavolcanic rocks as well as granitic rocks not exposed in the Buckskin Range.

The hornblende pyroxene andesite and diorite porphyries, hornblende pyroxene porphyritic andesites, and intrusive breccias form a large Miocene intrusive complex in the northern Buckskin Range. This diorite-andesite intrusive complex is Miocene because dikes of the complex intrude Tertiary ash-flow tuffs across Basin and Range faults. Some of the faults have post-intrusive displacement, however. The intrusives may have volcanic equivalents in the Lincoln Flat Andesite (17 to 10 m. y.), Proffett and Proffett (1976) but may be younger, because most of the Basin and Range faults cut by the intrusives do not show later movement.



#### Biotite Hornblende Andesite Porphyry (Tbhap)

Several dikes of coarse grained biotite hornblende andesite porphyry intrude the older rocks in and around the Minnesota Mine (Plate 1). The andesite contains 45 to 50% phenocrysts of euhedral biotite up to 1.2 cm across in books up to 1 cm thick, euhedral hornblende up to 5 mm long and subhedral plagioclase to 1 cm in length. The phenocrysts are in hial texture with an aphanitic groundmass of glass, plagioclase, microlites, and quartz. The rock contains 7 to 10% biotite, 2 to 3% hornblende, 2 to 4% quartz, 5 to 10% brown glass, and weakly oscillatory normally zoned plagioclase (An 43-38). The biotite hornblende andesite porphyry is apparently younger than the sheared hornblende pyroxene andesite porphyry exposures in the Minnesota Mine.

#### Porphyritic Hornblende Dacite (Tphd)

An intrusive porphyritic hornblende dacite crops out extensively in the southern Buckskin Range (Plate 1). The unit contains 15 to 25% phenocrysts of plagioclase and hornblende up to 1 cm and rarely up to 3 cm in length, although locally as much as 40% phenocrysts may be present. The plagioclase usually has epidotized cores. Locally the dacite contains rounded quartz "eyes" up to 3 mm (Tphdp) and locally occurs as an autobreccia (Tphdb). One to 2% biotite occurs locally in the dacite in Sections 2, 3, 10, and 11, T13N, R23E (Plate 1). The rock contains 2 to 8% hornblende, up to 10%(?) K-feldspar, 0 to 7% quartz, and weakly zoned plagioclase (An 30-35?). The groundmass contains K-feldspar and plagioclase microlites with or without quartz and hornblende.

Conflicting evidence indicates possibly both Tertiary and Mesozoic ages for the porphyritic hornblende dacite. In exposures in Sections 11, 12, 13,



and 14, T13N, R23E, (in Conoco's project area) the dacite lies along low angle Basin and Range faults with very altered Tertiary ash-flow tuffs above the dacite and highly altered Mesozoic metavolcanic rocks below. The dacite is weakly propylitized. The difference in alteration suggests a late Tertiary age for the dacite in these outcrops. In Sections 2 and 3 N1/2 Sec. 15, NW1/4 Sec. 11, and NE1/4 Sec. 10, T13N, R23E (Plate 1), the dacite is locally highly sericitized, intruded by highly altered Tertiary dikes (N1/2, NE1/4, Sec. 10, T13N, R23E), and apparently may not intrude the Tertiary ash-flow tuffs. This evidence, along with its similar appearance to the intrusives of the Churchill Canyon Sequence, indicates a Mesozoic age for these porphyritic hornblende dacite occurrences. Although the dacite is, for the most part, probably an intrusive portion of the Churchill Canyon Sequence, some of the dacite may be Tertiary. The Tertiary and Mesozoic dacites are indistinguishable in the field because they have identical composition and textures.

#### Hornblende Biotite Andesite Porphyry (Thbap)

Several intrusions of grey hornblende biotite andesite porphyry crop out in the southern Buckskins (Plate 1). Phenocrysts make up 25 to 35% of the rock with 4 to 6% hornblende with thin magnetite rims, 4 to 5% biotite, and oscillatory zoned subhedral to anhedral plagioclase (An 42-48). The phenocrysts are up to 1 cm in length and occur in an aphanitic groundmass of plagioclase, 1 to 4% mafics, up to 4% quartz, 1 to 2% magnetite and trace apatite. The rock generally is propylitized and rarely is sericitized. A large mass of hornblende biotite andesite porphyry in Sec. 2, T13N, R23E (Plate 1) may be either extrusive or intrusive although contact relationships, where seen, seem to indicate that it is extrusive.



#### Biotite Hornblende Sanidine Dacite Porphyry (Tbhsdp)

Several biotite hornblende sanidine dacite porphyry dikes crop out on the southern tip of the Buckskin Range (Plate 1). Phenocrysts make up to 60% of the rock, in seriate texture. They consist of subhedral oscillatory normally zoned plagioclase (An<sub>52-39</sub>) up to 10 mm long, rounded sanidine up to 5 mm making up 2 to 3% of the phenocrysts, 4 to 6% biotite up to 5 mm across, and 4 to 6% strongly zoned oxyhornblende up to 4 mm long with magnetite rims. The aphanitic groundmass contains 1 to 2% sanidine, plagioclase microlites, and minor amounts of glass.

#### Hornblende Andesite Porphyry (Thap)

Dikes of hornblende andesite porphyry intrude along faults in the southern Buckskins (Plate 1). The dark green andesites contain 30 to 40% phenocrysts, in seriate texture, composed of 30 to 35% hornblende up to 10 mm long and 10 to 20% subhedral plagioclase up to 15 mm in length.

#### Basalt (Tb)

A few black, dense, aphanitic basalt dikes crop out in the southern Buckskins (Plate 1). These are probably the youngest intrusives in the range but their exact age is uncertain.



## STRUCTURE

The Buckskin Range has two distinct types and ages of deformation: Mesozoic thrust faulting and Cenozoic Basin and Range faults. The Basin and Range faults are of two types: Older gently dipping faults which may have been originally steeply dipping, and younger steeply dipping faults.

### Mesozoic Deformation

A thrust fault is exposed in several locations in the northern Buckskin Range (Plate 1), bringing the Artesia Sequence in contact with the Gardnerville Formation. Evidence of this thrust is abundant. Assuming that the Artesia and Churchill Canyon Sequences are correlative with the Double Springs Formation, at least 8000 feet of stratigraphic section are absent from the Buckskins. They consist of the upper Gardnerville Formation, Preachers Formation, Veta Grande Formation, all of which are present in the southern Pine Nut Range (Noble, 1962). The upper Gardnerville, 150 to 300 feet of Ludwig limestone, about 600 feet of gypsum, and up to 1400 feet of Preachers Formation (which are the uppermost exposed Mesozoic units in the Singatse Range) are present near Ludwig (E. C. Bingler, written comm.) but are absent from the Buckskins.

Further evidence of the thrust fault lies in thickness and textures in various exposures of the Gardnerville Formation. About 150 to 200 feet of Gardnerville lies between the Oreana Peak Formation and the Artesia Sequence (Figure 3) in the middle of Sec. 11, T14N, R23E (Plate 1). A limestone unit of the Gardnerville is highly sheared and recrystallized with clasts of Artesia Sequence and Churchill Canyon Sequence intrusives occurring within the limestone several feet below the contact with the Artesia Sequence. Just to the south, in Section 11 (Plate 1) along strike in the Gardnerville Formation, at least 800 feet of Gardnerville Formation is exposed below the



Artesia, also along a sheared contact. Elsewhere, where the contact between the Gardnerville and Artesia is exposed, the rocks within the Gardnerville below the contact with the Artesia are usually strongly sheared.

The Buckskin Thrust postdates most of the Churchill Canyon Sequence and is probably uppermost Lower Jurassic to Middle Jurassic in age. In the vicinity of the Minnesota Mine a few Churchill Canyon Sequence affinity intrusions cut the Gardnerville. Although none of these intrusions were observed crosscutting the Buckskin Thrust, they appear to have been a pulse of magmatism following the emplacement of the thrust sheet. Assuming that the intrusions in the lower plate do represent a late pulse of Churchill Canyon Sequence type rocks, the thrust would then probably be uppermost Lower Jurassic.

The direction of thrusting is unknown as well as its areal extent. It is not present in the southern Pine Nut Range, however, a structural discontinuity similar to that in the northern Buckskin Range has been found in the northern Pine Nut Range southeast of Carson City (E. C. Bingler, oral comm.). A thrust contact there is as yet unproved.

Several faults with small displacement, probably of Mesozoic age, occur in the vicinity of the Minnesota Mine (Plate 1). They originally were normal faults with an east-west trend, except one just south of the Minnesota Mine, which was a reverse fault. Tertiary westward tilting has rotated the strata that the faults displace so that the originally dip slip faults now resemble east-west trending strike slip faults.

There may have been some pre-Tertiary folding in the Buckskin Range. The Churchill Canyon and Artesia Sequences lie in an angular unconformity



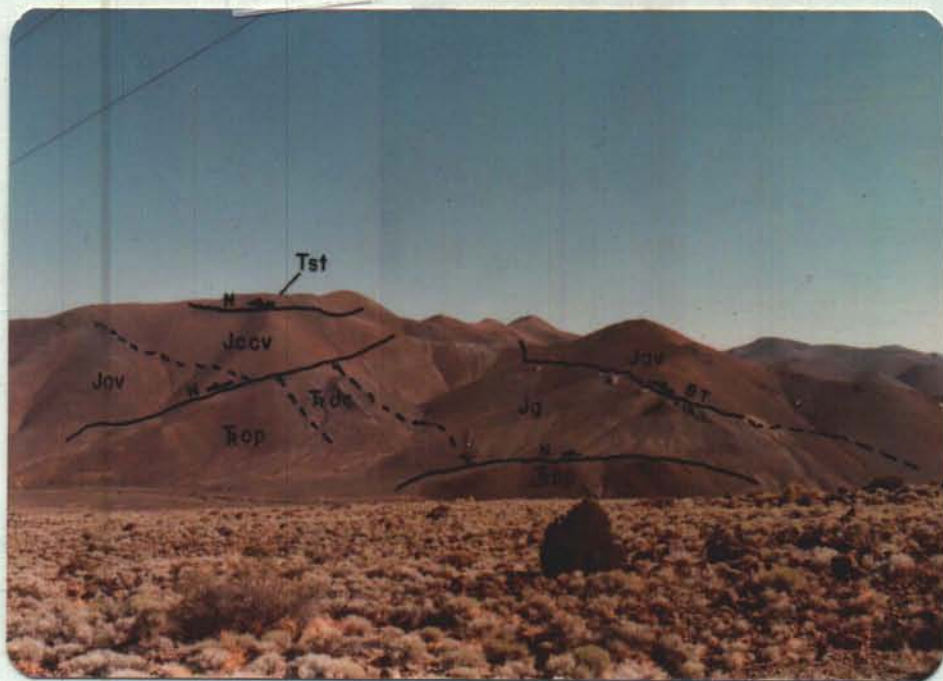


Figure 3. Looking south at the north end of the Buckskin Range showing the Buckskin Thrust (BT) and low angle Basin and Range faults (N).



Figure 4. Looking northwest at low angle Basin and Range faults and accompanying drag folds in Tertiary ash-flow tuffs in the northern Buckskin Range.



of about  $30^{\circ}$  with the overlying Guild Mine Tuff near Fulstone Spring No. 1 (Plate 1). The Guild Mine Tuff overlies the Artesia Sequence in the SW 1/4 Sec. 26, T14N, R23E (Plate 1). Elsewhere, the Guild Mine Tuff overlies highly variable thicknesses of the Churchill Canyon Sequence. This evidence implies pre-Tertiary folding but could be explained by Tertiary erosion with highlands to the west and a river valley to the east (Proffett and Proffett, 1976).

#### Tertiary Deformation

By far, the dominant structural elements in the Buckskin Range are low angle Basin and Range faults. These faults generally dip 5 to 15 degrees east although faults dipping up to 7 degrees west and 25 degrees east also occur (Plate 2). The low angle faults are generally easily recognizable where they displace Mesozoic rocks because of the lack of distinctive marker horizons. All of these faults have normal dip slip movement. The faults are generally subparallel with a northerly strike and frequently intersect each other. In cross sections drawn parallel to strike, the faults are generally crescent shaped with the concave surface up. Individual fault blocks are 100 to perhaps 1500 feet thick and 1000 to 15,000 feet long parallel to strike.

The average dip of the Tertiary ash-flow tuffs is 75 to 90 degrees to the west, except adjacent to the low angle Tertiary faults where intense drag folding occurs. This dip clearly indicates drastic rotation of the Buckskin Range and the ash-flow tuffs serve as good control for fault displacement. Offsets as small as 50 feet can be clearly seen, with maximum observed displacements of 6000 feet, all with displacement directed to the east.



Intense drag folding is usually found adjacent to the low angle faults, and often is accompanied by thick fault gouge. Drag folding is usually difficult to observe in the Mesozoic rocks but, in NE1/4, Sec. 11, T14N, R23E (Plate 1), limestones of the Oreana Peak Formation are folded up to 90 degrees and they lie parallel to the fault where the Churchill Canyon and Artesia Sequences are dropped down against the Oreana Peak. The best examples of drag folding occur where the highly competent Singatse Tuff is faulted against the incompetent Weed Heights, particularly in the NE 1/4, Sec. 26, T14N, R23E (Figure 4). While the dip of the Singatse may vary from 90 degrees at a distance from the fault to about 70 degrees west adjacent to it, the Weed Heights changes dip from about 70 degrees west far from the fault to 3 degrees west adjacent to the fault, with intense, closely spaced slickensides developed, completely grinding up the pumice fragments. Up to 100 feet of fault gouge is exposed by Anaconda bulldozer cuts in the southern Buckskins and it is not uncommon to find boulders of Mesozoic rocks in fault gouge between Tertiary ash-flow tuffs.

Proffett (1972, 1977) envisions progressive westward tilting of fault blocks by rotation on younger faults in the Yerington area (Figure 5). According to Proffett's model, the gently dipping faults which underlie the steeply dipping Tertiary volcanic rocks were originally steeply dipping Basin and Range faults. In the vicinity of the Buckskin Range, the only younger east dipping faults which could account for the rotation of the range are the frontal faults on the east flank of the Pine Nut Range. Although displacements along these faults are younger than the low angle Basin and Range faults, their displacements probably do not exceed 10,000 feet between the Pine Nuts and the Buckskins. A steep frontal fault on the



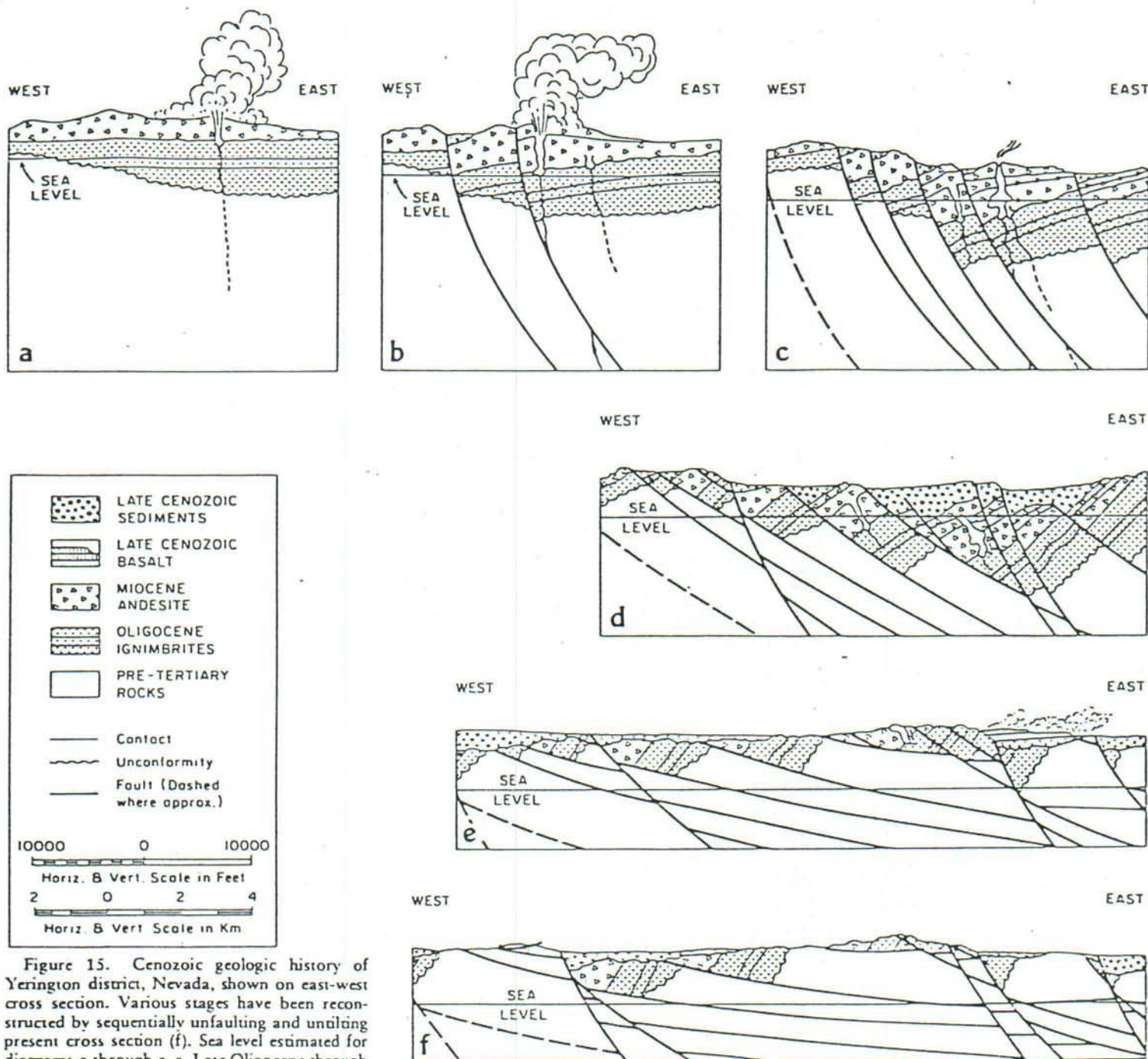


Figure 15. Cenozoic geologic history of Yerington district, Nevada, shown on east-west cross section. Various stages have been reconstructed by sequentially unfaulting and untilting present cross section (f). Sea level estimated for diagrams a through e. a, Late Oligocene through early Miocene; deposition of ignimbrite sequence followed by beginning of deposition of hornblende andesite, 18 to 19 m.y. ago. b, Early Miocene, 17 to 18 m.y. ago; continued eruption of andesite and beginning of Basin and Range faulting and tilting. c, Early Miocene, 17 m.y. ago; end of andesite volcanism and continued Basin and Range faulting with erosion of up-faulted blocks. d, Early and middle Miocene, 11 to 17 m.y. ago; continued faulting, tilting, erosion of ranges, and deposition of sediments in basins. Early faults rotated to such a flat dip that they are no longer in favorable orientation for continued faulting. New steeper faults form to take their place. e, Eight to 11 m.y. ago; all early faults tilted to gentle dips and inactive, but faulting and tilting continues on newly formed steep faults; erosion of ranges and deposition of sediments and locally basalt in basins. f, Present time; continued faulting and tilting along steep range-front faults, erosion of ranges and deposition in basins.

Figure 5. Cenozoic structural history of the Yerington area from Proffett, 1977.



east flank of the Buckskins may have a few thousand feet of displacement near the south end of the Range but the displacement is probably minor along the north-east flank of the Range. Basin and Range faults elsewhere in the Basin and Range Province usually do not appear to have intense drag folding and thick gouge zones associated with them. All of this suggests that Proffett's model may be incomplete, however, no better model can be offered at this time. Extensive, drastic westward tilting and crustal extension in the Yerington area, however, cannot be disputed.

A few west dipping normal faults are present in the northwestern Buckskins (Plate 1) but appear to be minor elements in the structure of the range. One of the faults is intruded by a Tertiary hornblende pyroxene andesite porphyry dike with extensive breccia adjacent to the dike. Most of the breccia is composed of fragments of Mesozoic rocks, including granitic rocks not exposed in the Buckskins. This is probably an intrusive breccia without igneous matrix.

#### Unresolved Structural Problems

A major disconformity exists between the Artesia Sequence and underlying metavolcanic rocks in the southeastern Buckskin Range. The rocks of the lower sequence (Tv, Plate 1) consist of andesitic volcanic flows with small phenocrysts and occasional "chicken track" andesite. These rocks may be correlative with a Lower Triassic metavolcanic sequence exposed in McConnell Canyon in the Singatse Range (Figure 2). These metavolcanic rocks are older than the Lower Jurassic Artesia Sequence, since they are intruded by granodiorite which is provisionally correlated with the Lower Jurassic granodiorite of Black Mountain. Foliations in the lower sequence rocks generally



strike east-west and dip to the north. The Artesia Sequence strikes approximately north-south and dips to the west. The contact between the Artesia Sequence and the lower metavolcanic sequence may be part of the Buckskin Thrust but insufficient exposure prevents determination of the type of contact.

The large area of Churchill Canyon Sequence exposed in the northwestern Buckskins (Plate 1) undoubtedly does not represent a single Basin and Range fault block, based upon the tectonic style in the Buckskin Range. Lack of stratigraphic control and poor outcrop inhibited the mapping of low angle Basin and Range faults which certainly are present.

The large area of Tertiary hornblende biotite andesite porphyry in Sec. 1 and 2, T13N, R23E (Plate 1) may be a fault block with a low angle fault beneath it (Plate 2, Cross Section G-G'). A bulldozer trench along the contact on the ridge in NW1/4 Sec. 1, T13N, R23E reveals a low angle fault contact. This fault relationship could not be definitely traced around the block of hornblende biotite andesite porphyry (Thbap), however. A low angle Basin and Range fault under the porphyry is indicated, but not proven.

The fault block of Tertiary ash-flow tuff and intrusions north of the Minnesota Mine (Plate 1) appears to have been emplaced following the intrusion of the Tertiary hornblende pyroxene andesite porphyry dikes (Cross Section B-B'). No Tertiary dikes cut the low angle fault on the west side of the block and dikes near the fault have locally intense shearing. Additionally, the tuffs and some of the intrusions are strongly altered adjacent to the fault while the limestones to the west of the fault contain only very weak recrystallization. This strongly suggests that the Tertiary ash-flow tuffs and intrusions were faulted into place after intrusion and alteration.



A major fault paralleling the trend of Churchill Canyon (about N25E) (Plate 1) has been shown by Moore (1969) to explain the much lower elevation of Miocene volcanics north of the Canyon in contrast to the higher elevation of Mesozoic rocks south of the Canyon. No evidence was found to support a large fault but rather there may have been much deeper erosion to the north of the Buckskin Range before the Miocene volcanics were deposited. Regionally, the lineament through Churchill Canyon appears to be a major break between Tertiary volcanic rocks to the north and Mesozoic rocks to the south and may in fact be a fault. Insufficient exposure prevents confirmation.



## ALTERATION

### Propylitic Alteration

Most of the rocks exposed in the Buckskin Range have varying degrees of propylitization. Propylitic alteration is difficult to determine in the Mesozoic rocks, because they have been metamorphosed to the greenschist facies, except where they are more sericite and epidote rich. The Tertiary ash-flow tuffs are highly propylitized in the southern and northern Buckskins near Tertiary intrusions and become less altered in the west-central Buckskins. Alteration minerals are clays, calcite, sericite, and minor quantities of epidote. The tuffs, particularly the Singatse Tuff, often become whitish-lavender when propylitically altered. Tertiary intrusive rocks in the northern Buckskins are generally not propylitized although a few Tertiary hornblende pyroxene andesite porphyry dikes north of the Minnesota Mine are moderately propylitized. The only Tertiary intrusives that are not altered in the southern Buckskins are the basalt, the biotite hornblende sanidine dacite porphyry, the hornblende andesite porphyry and the exposure of hornblende biotite andesite porphyry in Sec. 22, T13N, R23E.

### Sericitic Alteration

Sericitic alteration is present in much of the Buckskin Range, and is particularly abundant in the east and south. Several types of sericitic alteration were used in field identification: Strong sericitic (ss, Plate 3) where the rock is completely altered to sericite and quartz with little clay; moderate sericitic (ms, Plate 3) where there is partial sericitization with little clay; strong sericite-clay (ssc, Plate 3) where the rock is completely altered to sericite with relatively abundant clay; moderate sericite-clay



(msc, Plate 3) where the rock is partially altered to sericite and clay; and weak sericite-clay (wsc, Plate 3) where the rock is slightly altered to visible sericite and clay. The dominant clay present is kaolinite, although some montmorillonite may occur. The Mesozoic rocks are sericitized over much of the range while sericitization in Tertiary rocks is more restricted. A few of the Tertiary hornblende pyroxene andesite porphyry intrusions are highly sericitized and are presumably the earliest phases of hornblende pyroxene andesite and diorite porphyry intrusive activity because adjacent hornblende pyroxene andesite and diorite porphyry intrusives are fresh. The Tertiary ash-flow tuffs in the northern Buckskins are highly sericitized, particularly adjacent to the Tertiary hornblende pyroxene diorite porphyry intrusions. Sericitic alteration is also widespread in the Tertiary ash-flow tuffs in the southern Buckskins except at the extreme southern end of the Range.

#### Aluminophyllic Alteration

Two areas of aluminophyllic alteration were found in the Buckskin Range. Aluminophyllic alteration, a subtype of advanced argillic alteration, is defined as an assemblage consisting of quartz, ordered kaolinite group minerals, pyrophyllite, andalusite, diaspore, corundum  $\pm$  topaz and fluorite. The area around the Blue Danube Mine contains predominately quartz, pyrophyllite, and diaspore with some andalusite and topaz. An assemblage of quartz, pyrophyllite and diaspore was found just west of the Minnesota Mine. In both cases extreme recrystallization of the original rock accompanied aluminophyllic alteration.

#### Silification

Strong silification (si, Plate 3) is largely restricted to the Artesia



Sequence except for a small area in the Jurassic granodiorite in Sec. 24, T13N, R23E. In general, the silification in the Artesia Sequence is nearly strataform with lensoidal quartz "reefs" which locally crosscut bedding. The silicified bodies often contain up to 15% sericite.

#### Tourmaline Breccias

Several small outcrops of tourmaline breccia (Tbx, Plate 3) were found in the Buckskins. They are usually ovoid in outcrop and generally less than 30 feet in length. The tourmaline breccias are generally peripheral to major centers of alteration and contain weakly altered rock and/or quartz fragments in a matrix of tourmaline. Often veinlets of tourmaline radiate from the breccias into the country rock.

#### Biotite veinlets

One exposure of coarse biotite veinlets was found in nearly fresh Artesia Sequence in the center of Section 1, T13N, R23E. The biotite veinlets are one to 30 mm wide and irregular in trend with biotite flakes up to five mm across.

#### Age of Alteration

Most of the alteration in the Buckskin Range appears to be Miocene in age because the Tertiary ash-flow tuffs and intrusions are altered. The silification in the Artesia Sequence and some of the associated alteration is Mesozoic. An altered and mineralized intrusive complex in the southern Buckskin Range, one dike of which has been dated at 15.7 m.y. (Oriel, 1976), has a pervasive alteration halo in the surrounding Mesozoic rocks. This, together with alteration associated with Miocene intrusions in the northern Buckskins, indicates there is a Miocene alteration overprint on most of the Mesozoic rocks in the Buckskin Range.



## MINERALIZATION

### Limonites

Goethite occurs in varying concentrations, both as disseminations and fracture coatings. Goethite is present throughout areas with sericitic and aluminophyllic alteration and in silicified bodies. There appears to be little correlation in most areas between the intensity of alteration and the quantity of goethite.

Jarosite occurs in a number of areas as both disseminated grains and fracture coatings. The jarosite is impure and is mixed with goethite and hematite. Only areas where jarosite exceeds 50% of the total limonites are outlined on Plate 4; jarosite is indicated by the letter "J" where its concentration is greater than 25%, (Plate 4).

Strong lavender hematite was found in a few areas but it never exceeded about 20% of the total limonite content. Brick red hematite stains are common near some fault zones. A few veins of black specular hematite (SH, Plate 4) were found in the eastern south-central Buckskins. It occurs in veins up to two feet wide, occasionally accompanied by quartz, and as thin fracture fillings. The significance of the specular hematite is unknown.

### Minnesota Mine

A magnetite-copper skarn in the limestone unit of the Oreana Peak Formation was mined at the Minnesota Mine. Massive magnetite with stringers and blobs of pyrite and rarely chalcopyrite replace the limestone parallel to bedding. Essentially pure magnetite-pyrite bodies up to 30 feet wide are exposed in the mine with smaller bodies of magnetite-pyrite stringers in



the limestone. The limestone is strongly recrystallized with a weak development of epidote, clinozoisite(?), tremolite-actinolite, and some pyroxenes. Some marble is locally present. Numerous Mesozoic and Tertiary dikes are exposed in the mine. The Tertiary dikes are fresh and have no spatial relationship to the ore. Mesozoic dikes of Jurassic quartz monzonite porphyry also seem to have no direct correlation to the ore. Ore often occurs on one side of a quartz monzonite dike and recrystallized limestone occurs on the other. The rocks are complexly sheared by low angle Basin and Range faults. The Jurassic quartz monzonite porphyry contains quartz veins just east of the mine. Pyrite and magnetite occur both in the veins and disseminated in the porphyry. The quartz monzonite porphyry thus appears to be the mineralizing intrusion. Weak copper staining is often present where the Jurassic granodiorite comes in contact with limestone of the Oreana Peak and Gardnerville Formations in the vicinity of the Minnesota Mine.



## GEOCHEMISTRY

Rock chip geochemical samples were collected from outcrop or sub-outcrop in the Buckskin Range. Ten to 14 small chips were collected from about a 20 foot diameter area from outcrops and from float of local origin for each sample.

Geochemical background (in ppm) for various unaltered rock types is summarized as follows, based upon relatively few samples (refer to Plate 5).

	Tertiary Volcanics	Tertiary Intrusions	Mesozoic Volcanics	Mesozoic Sediments	Mesozoic Intrusions
Cu	-10	10 to 30	5 to 15	35 to 60	10 to 20
Mo	1 to 4	-1 to 5	1 to 2	-1	1 to 4
Zn	10 to 75	40 to 100	30 to 70	20 to 160	40 to 100
Ag	-1	-1	-1	-1	-1
As	-10	-10	-10	-10	-10

The Singatse Tuff tends to be higher in zinc (45 to 75 ppm) than the other tuffs (25 to 40 ppm) but is similar in its content of other elements. Altered tuffs are about the same or lower in Cu, Mo and Zn than fresh tuffs.

The area where Anaconda has drilled intensively (S1/2 of S1/2, Section 36, T14N, and N1/2, Section 1, T13N, R23E, Plate 3) is, in general, relatively lower in Cu, Mo and Zn and higher in As than the surrounding areas.

The Tertiary hornblende diorite porphyry related intrusions are relatively zinc rich and this is often reflected in adjacent wall rocks (77-BR-568, 579, 660, 679, etc., Plate 5).

Sample 77-BR-345 (157996) (Plate 5) is very high in Cu, Zn, and Ag.



The sample comes from vapor-phase Guild Mine Tuff with no apparent surrounding alteration. The sample may have been mislabeled as the rock does not have the appearance of any mineralization.



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W. M. Oriel  
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Continental Oil Company  
Minerals Department  
Reno District Office  
May, 1977



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NOTE: The Plates noted with \* may be found in the roll which accompanies this report.



## SUMMARY

The Buckskin property is a large intensely altered, mineralized, and geologically young porphyry copper prospect that covers nearly 3.75 square miles. Several generations of exploration began in 1967, but only on small parcels on the fringes of the area. Conoco drilled hole BC-1 to 1994 feet in late 1976 with encouraging geochemical results. The results of drilling, geochemistry, induced polarization surveys, geologic mapping and sampling combine to give at least two strong porphyry copper target areas and probably a third (Figures 1, 10, and 11).

Induced potential surveys at Buckskin indicate a large area of sulfide mineralization approximately 9000 feet wide by 12,000 feet long, containing a centrally located elongated zone of apparently more abundant sulfides (Figure 20). Drilling results and limonite mapping largely confirm this pattern. Pyrite is the predominant sulfide in all drill holes, occurring as very fine grained disseminations and in veinlets rarely two inches wide and commonly 1/32" to 1/8" wide. The pattern indicates several geologically permissible models involving a general decrease in volume percent sulfides at depths approximating the limits of the technique (2500-3000 feet).

The large system of mineralization at Buckskin is undoubtedly related to the mineralized later Tertiary porphyry dikes that Bear Creek has dated at  $15.7 \pm 1.5$  million years in BSW-9. Conoco's mapping indicates a similar age for these dikes. Older Mesozoic(?) phaneritic quartz monzonites crop out on the east flank of the range, however, and are uniformly high in copper. Therefore, it is possible to speculate that there are two widely



separated ages of mineralization at the prospect, but age dates are needed to establish the relationships. If this interpretation is true, then the deep IP anomaly on the down thrown side of the eastern Basin and Range fault could become a separate target for mineralization similar in age to the Yerington deposits.

Because of the nature, age and pattern of the bulk of the mineralization and the geologic environment, it is reasonable to interpret the present erosion surface at Buckskin as the top of a porphyry system that is either upright or tilted westerly, possibly as much as 45 degrees.

The first target area (Figure 1) has the highest probability of (1) being centered over the top of the system and (2) approaching the target closest to the surface. It is between holes BC-1 and BSW-7 (Figure 10) and is strongly altered to clay and sericite and weakly veined with quartz and clay. A breccia pipe and a small pebble dike occur near this area also. Hole BSW-7 encountered 510 feet of 0.16% copper beneath a zinc halo. The copper may have been in the intermediate potassic zone, as may occur at Red Mountain, Arizona (Corn, 1975). The target could be below BSW-7 or to the east because of regional tilting and faulting. It is probable that the target is greater than 1500 feet deep and possibly is much deeper.

The second area should be drilled either contemporaneously or following the drilling of the first target. Structural complexities permit the second target to be located east of hole BC-1.

The third target is on the east flank of the prospect. Uniform copper mineralization in older tilted quartz monzonites near a mineralized Tertiary dike swarm provide intriguing geologic possibilities of either (1) separate but overlapping mineralized episodes or (2) downfaulted mineralized quartz



monzonites along the east range. The depth of burial after faulting, and the weaker alteration products overlying the area are reasons for the lower priority of this target. The values of copper in these eastern rocks, however, are very high for the intensity of alteration and in comparison to the other areas of the prospect.



## RECOMMENDATIONS

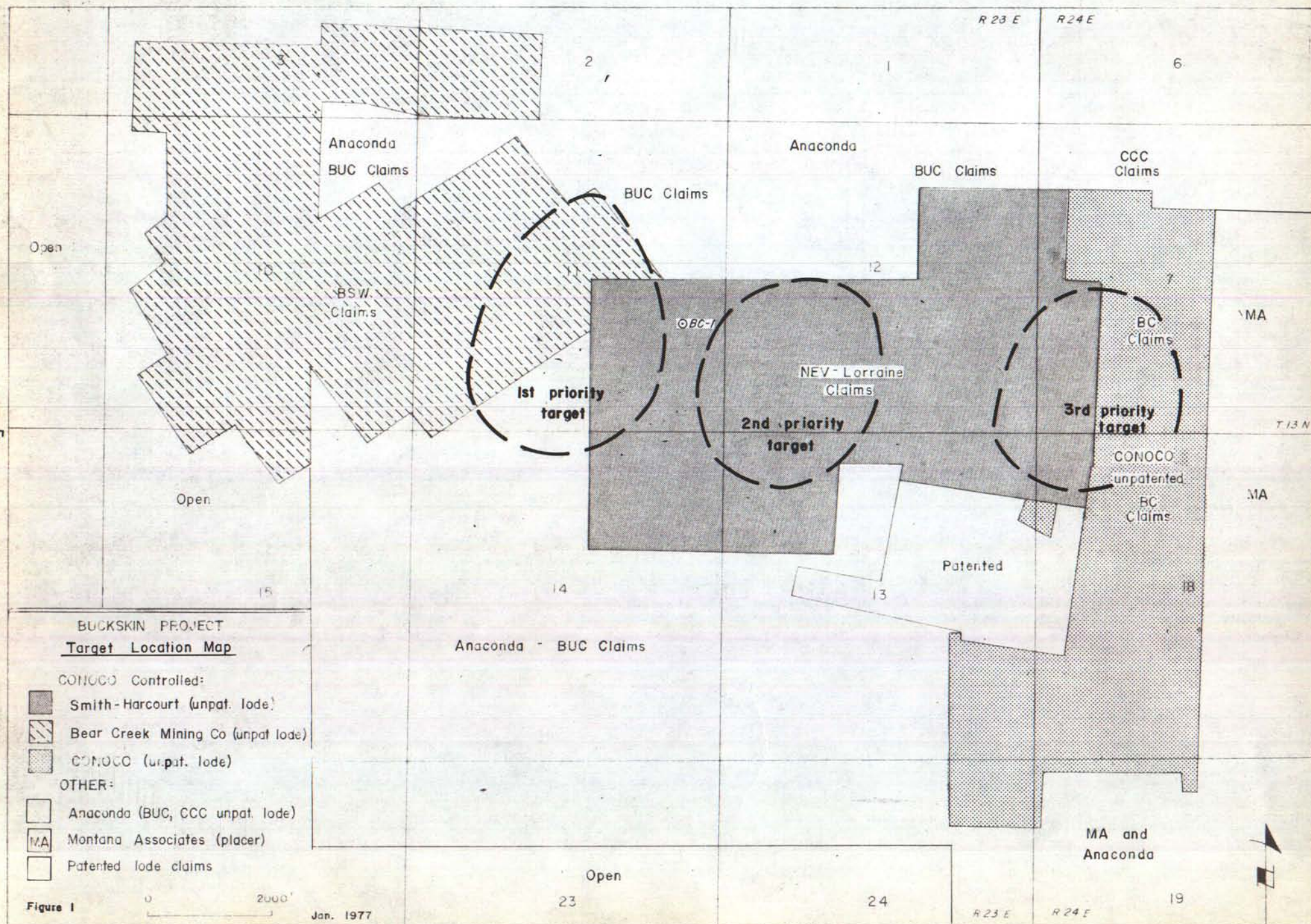
It is recommended that Conoco drill two holes in the first target area to depths of at least 2000 feet in 1977. The area should not be abandoned until at least one hole has been drilled to approximately 4000 feet.

The second target area could be drilled consecutively with the first target area, finances permitting. Otherwise it should be drilled following the results of the primary drilling target.

The third target area deserves consideration in the future, but need not be acted upon promptly. Age dates of the older phaneritic intrusives would meanwhile be useful for geologic interpretation.

Obligations to the contract with Bear Creek Mining Company require that we drill at least 1500 feet upon the BSW claims during 1977. These obligations can be met by drilling within the primary target area as recommended.







## INTRODUCTION

### Location

The Buckskin district is in Douglas County in west central Nevada, approximately 95 road miles southeast of Reno and 10 miles west of Yerington (Figures 2 and 3). The prospect is in Township 13 North, Ranges 23 and 24 East, in the south central portion of the Buckskin Range. It is easily accessible by dirt roads that are periodically graded throughout the year. The climate is arid and only a few juniper trees exist in the area. The range has relatively low relief and a maximum elevation of 6915 feet.

### Property and Ownership

Conoco controls 128 contiguous unpatented lode claims containing approximately 2300 acres spanning the Buckskin Range from east to west. Forty-seven (47) claims are controlled through a joint venture with Bear Creek Mining Co. (BCMC) and 43 are under option from two local residents, Smith and Harcourt. Conoco staked 38 unpatented claims on open ground at the eastern edge of the range as a result of IP data and permissive geology (Figure 4).

Anaconda controls unpatented claims on the north and south margins of our claim block and Montana Associates, a small group of local investors, controls unpatented placer claims adjoining Conoco on the east.

All claims are on public domain lands controlled by the Bureau of Land Management.

A small group of patented claims surround the Buckskin mine at the southeastern portion of our claim group. They are not of interest to



Conoco at the present time.

#### Holding Costs and Obligations

Bear Creek Joint Venture. Conoco must have spent \$200,000 by December 13, 1981 to earn 51% interest. Conoco also must drill 1500 feet yearly upon the BCMC claims (BSW group) until the 51% has been earned.

Smith-Harcourt. An expenditure of \$500,000 in annual payments and/or royalty payments to the owners before 1992 will purchase 100% of the 43 "Nev-Lorraine" claims.

Conoco is responsible for assessment obligations for the Bear Creek and Smith-Harcourt claims as well as our own 38 unpatented "BC" claims, on which no assessment work is required until September 1, 1978. No obligations or special requirements exist, other than those imposed upon us by the BLM.

Total payments and expenditures in 1976 were as follows:

Smith-Harcourt: \$39,001.60 paid to owners and applicable to the \$500,000 end price.

BCMC: Conoco spent \$58,178.36 in drilling and other exploration activities which is applicable toward the \$200,000 buy-in price.

#### History and Previous Work

According to some reports the Buckskin district was discovered in 1904 by a prospector who found gold in veins on or near the present "Red Top" patented claims. Subsequently, copper was found below the oxidized gold zones and considerable excitement arose. This resulted in the building of a townsite and the construction of pumping and metallurgical plants. Very little ore was produced, and most activity ceased in 1907 except for the



floating of large amounts of stock. The district was rediscovered and re-evaluated in the early 1930's, and as late as 1974 a head frame and mill building were erected. Drilling, however, apparently found no ore and the site is now abandoned.

A more successful mining operation was carried out on combined iron-copper ore in the northern part of the Buckskin range. Iron ore was mined from the Minnesota open pit mine from 1944 to 1966, producing nearly \$17,000,000, which is 99.5% of the total values of the district. Exploration and drilling have been carried out sporadically in the district since 1966 and Anaconda currently holds claim to much of the north-eastern part of the Buckskin range. The only published geologic mapping was done by Moore (1969) and recently by Proffett (1977). Both Bear Creek Mining Co., and Anaconda mapped portions of the project area, but the scale and detail were largely inadequate, due to the intense degree of weathering and the complex Tertiary history of the Buckskin Range.

The Yerington, or Weed Heights open pit copper mine has been operating since 1952. It is a porphyry copper mine associated with Mesozoic intrusives. It contained about 200 million tons of about 0.6% copper. Other porphyry copper deposits near Yerington include the Mickey Pass (Ann Mason), Bear, Airport, and Lyon (Pumpkin Hollow) (Figures 3 and 5).

#### Recent Exploration

Exploration for porphyry copper ore at the Buckskin prospect apparently began in 1966 or 1967 when Phelps Dodge acquired large acreage, including the "Nev" claims, in the eastern part of the property. They drilled holes BUE-1 and BUE-2 to depths of less than 1000 feet on the east flank of the prospect, and BUL-9 to 734 feet in the central portion of the prospect.



Anaconda had control of the "Nev" group until the autumn of 1975. They drilled two holes within our claim group (DH 281 and 289) and several holes bordering our property about which we have no information. Anaconda dropped the property before a large land payment was due.

In 1969, Bear Creek Mining Co., staked claims on the west flank of the range around a turquoise showing. They mapped the prospect and took some rock chip samples which were anomalous. Nine holes were drilled by BCMC, five of which were diamond drill holes centered around their geochemical anomalies, and four others were shallow rotary holes drilled in 1970 in the alluvium to the west. Two of their deep holes, BSW-1 and BSW-7, intersected anomalous mineralization and alteration. Because of their geologic interpretation they tried to follow the mineralization unsuccessfully to the west with drill hole BSW-9.

#### Present Program

Conoco's involvement is partially the result of the fortuitous availability of both the Bear Creek claims and the "Nev-Lorraine" claims. The Conoco staff had the ability to recognize the potential of a large complex sulfide system that had been looked at only in bits and pieces previously.

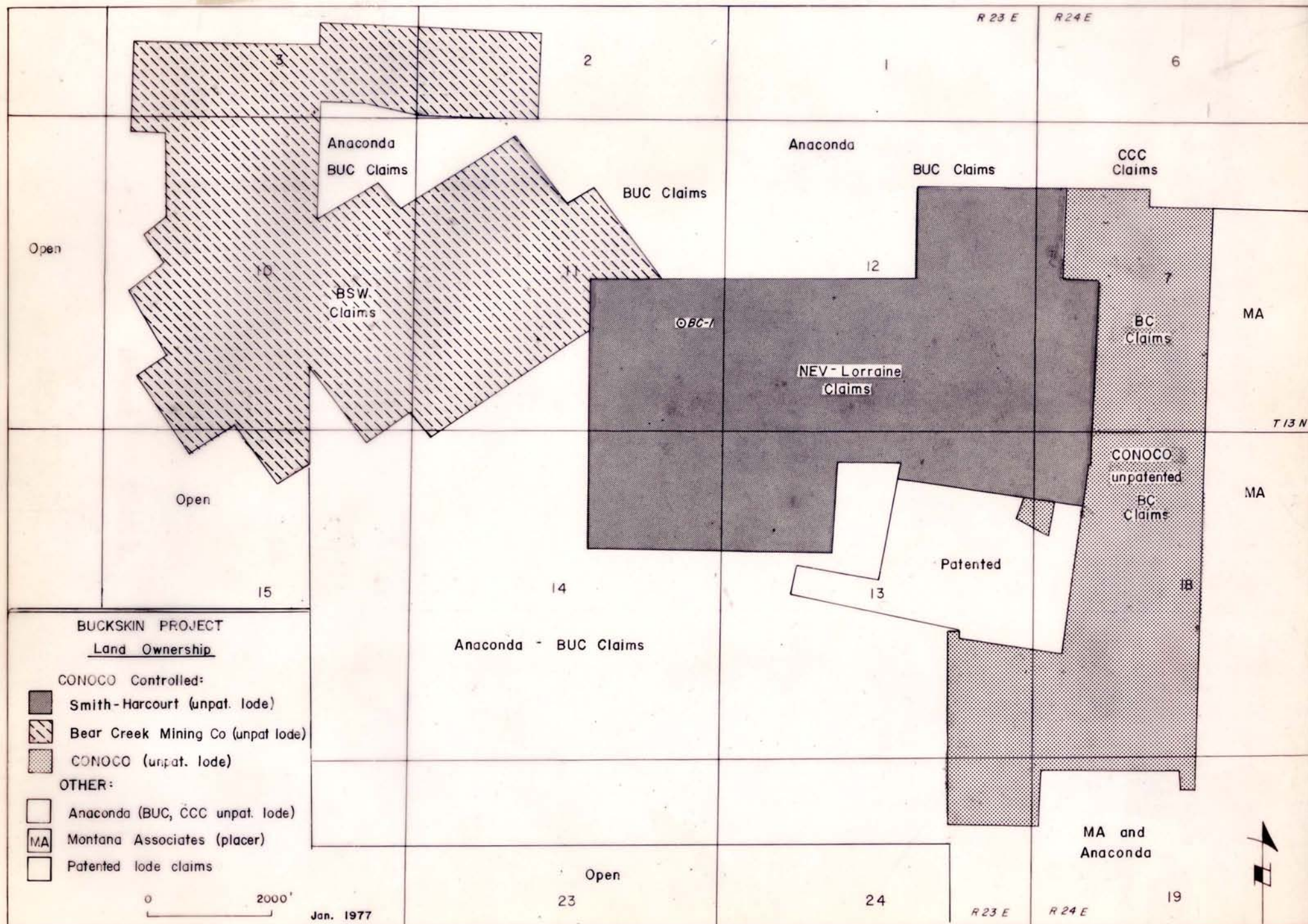
When the property became available in late 1975 and early 1976, we secured our interests with letters of intent and began a large soil geochemical program to outline the anomalous areas. Detailed geologic mapping and extensive IP surveys were conducted in the summer and fall of 1976, and Conoco began drilling hole BC-1 in September, to a depth of 1994 feet. This report is a summary and interpretation of the mapping, geophysics, geochemistry and drilling conducted during 1976.



#### Acknowledgements

Mr. P. H. Kirwin was able to act quickly to examine and acquire the two halves of the property when they became available in late 1975. Bear Creek Mining Co., and Kennecott Research have been very cooperative and helpful since we acquired the property. Aubrey J. Beck assisted mapping for several weeks on the Buckskin area and Karen Kling and Pat Brioady helped with drafting and design of the plates and figures. The Conoco staff was helpful in its constructive criticism.







## GENERAL GEOLOGY

The Buckskin prospect is in the southern half of the Buckskin Range at the western edge of the Great Basin Province and near the western margin of the Walker Lane structural zone. It is bounded by a moderately steep east dipping fault on the east flank of the range but is dissimilar to classical basin ranges because it was formed by progressive late Tertiary westward tilting. Crustal extension on the margins of the province possibly caused the tilting (Proffett, 1977). According to Proffett (1977), westward tilting and extension ceased about 11 million years ago and is an important consideration in exploring for mineralized intrusives older than 11 million years.

Metamorphosed Mesozoic volcanic and sedimentary rocks make up the underlying country rock (Moore, 1969). They are intruded by intermediate intrusive rocks in the Yerington area, which are mostly Jurassic in age and often are mineralized. There are extensive mineralized late Tertiary porphyritic dikes and sills locally at Buckskin. They intrude voluminous late Oligocene ash flow tuffs that overlie the Mesozoic volcanics. The porphyry dikes and sills at Buckskin have not been recognized previously for their true age or significance.



## PROSPECT GEOLOGY

### Mesozoic Volcanic Rocks

The oldest rocks exposed in the vicinity of the Buckskin prospect are undifferentiated Triassic and Jurassic metavolcanic rocks and metasedimentary rocks. They consist of interlayered sandstones, limestones, shales, and felsic and intermediate metavolcanic rocks (Castor, 1971) in the Northern Buckskin Range. In the southern portion of the range, the rocks are primarily andesitic flows, metavolcanic breccias, and tuffs, with rare discontinuous lenses of siltstones and arkoses. Some of the andesites are probably semiconcordant intrusive sills in the volcanic pile. The volcanics all are similar in composition, however, and they are difficult to correlate over small distances even with detailed mapping, because of metamorphism, the superimposition of hydrothermal alteration and Cenozoic faulting. Only the volcanic breccias in the west side of the prospect are distinctive enough to correlate with any confidence. Most of the older rocks have undergone weak regional greenschist-facies metamorphism, causing chloritization and epidotization of the andesites. Because these minerals are principle constituents of the propylitic hydrothermal alteration assemblage surrounding typical porphyry copper deposits, some uncertainties can arise when mapping alteration effects.

Numerous field names were used in mapping the volcanics. They were mostly textural descriptions used to correlate rocks throughout the prospect. Undifferentiated Mesozoic metavolcanics (Mvu) is the dominant variety and consists predominantly of non-descript, very fine grained, weakly foliated andesites. The andesites are probably flows



that possessed a weak primary foliation that formed roughly parallel to the ground surface during deposition (pseudo-bedding) that has been enhanced by alteration effects. Mvu was also applied to any metavolcanic rock whose original texture was destroyed by surface weathering or intense alteration.

Fine grained andesite (Mva) is the most abundant volcanic rock other than the undifferentiated Mesozoic volcanics. The plagioclases are generally recognizable in these rocks as opposed to the undifferentiated volcanic rocks. The andesite is probably a less altered rock of equivalent composition.

Metavolcanic breccia (Mvb) consists of elongate clasts several millimeters to rarely a foot long in a fine grained foliated volcanic matrix. The best exposures are in the western portion of the range where they are traceable for thousands of feet. Here they dip moderately to steeply westward and are 300 to 500 feet thick. The foliation in the surrounding volcanics closely mimics the strike of the breccia unit, lending strong support for a primary origin of the foliation.

Silicified volcanics (Mvs) occur throughout the prospect but tend to be concentrated in the western portion of the prospect. Although they were mapped on the photo, they were not shown as specific units on the printed geological map because of doubts about their origin as primary depositional features. They may be included on an alteration map in the future because most of the evidence indicates that they may have been formed by silica metasomatism. The silicious rocks (Mvs) cap hills and form ridges because of their resistance to erosion. They are generally 90% to 100% silica, and some may be replacements of originally porous



silicious volcanic rocks. Several outcrops show distinct relict crystals which may be feldspars. The silicified volcanic rocks form discontinuous ridges parallel to foliation but occasionally they are parallel to fractures that cross-cut the pseudo-bedding foliation. In at least one case, the silicification appears to be an altered sedimentary sandstone.

Several breccia units (Mvsb) appear to be the result of silicification of a tectonic breccia.

Several other textural and compositional varieties of volcanics were mapped. Most, however, are merely un-noteworthy variations that are to be expected in a thick volcanic pile. A few varieties such as fine quartz rhyolite (fqr), and fine latite porphyry with quartz (flpq), may have been felsic tuffs or crystal tuffs, but they are rare. Other textural and mineralogical varieties are as follow:

Mvfpa	Fine porphyritic andesite - usually foliated with greater variation in feldspar sizes than other andesites.
Mvpfa	Porphyritic, strongly foliated andesite - uncommon.
Mvha	Hornblende andesite.
Mvhap	Hornblende andesite porphyry - $\geq 25\%$ pheno.
Mvpha	Porphyritic hornblende andesite - $\leq 25\%$ pheno.

Some of the andesitic rocks can be traced for short distances and some appear to be sill-like intrusive bodies. This is not surprising in a volcanic pile, nor is it surprising that they are difficult to recognize after metamorphism and subsequent hydrolytic alteration.

#### Mesozoic(?) Intrusives

The oldest mapped intrusive rocks are exposed on the east flank of the



prospect. Fine quartz monzonite (Fqm) is the major map unit. Other textural varieties are speckled quartz monzonite (Sqm), fine grained monzonite (Fqm) and medium grained quartz monzonite (Mqm). They are mostly fine grained phaneritic quartz monzonites with slightly varying amounts of mafic minerals and quartz phenocrysts. Although several more varieties exist in drill holes, they are not clearly recognizable on the surface and will not be discussed separately. No contacts can be found between units on the surface and the variations are probably the normal ones found within many intrusive bodies. The intrusive rocks are predominately altered to a moderately pervasive chlorite - calcite - sericite assemblage within the prospect area.

The Mesozoic(?) intrusives are a small pluton which has been tilted approximately 90 degrees westward and which has been truncated by recent Basin and Range faulting (Figures 10 and 11).

#### Mesozoic(?) Dikes

Several steep dikes, termed coarse latite porphyry (Clp), coarse hornblende - biotite - latite (Chblp) and coarse latite porphyry with quartz (Clpq) intrude the phaneritic quartz monzonite and the Mesozoic volcanics. The dikes are undoubtedly from the same intrusive event but possess minor textural and compositional variations. They are characterized by large, one-half to one inch, K-feldspar phenocrysts with hornblende and biotite in varying amounts and sizes. Smaller ( $\leq 4$  mm) quartz phenocrysts are visible occasionally, but they are always less than one percent in hand specimen. Some of the field name variations are the result of destruction of the fabric by strong sericitization. All other porphyritic dikes in the prospect area are younger and do not possess the characteristic



large feldspars.

#### Tertiary Ash Flow Tuffs

An extensive erosion surface was formed during early Tertiary or possibly Late Cretaceous time, and large volumes of ash flow tuff were later deposited on it. The basal Tertiary deposits, which are probably less than 30 million years old (Proffett, 1976), were rhyolitic ash flow tuffs. They were partially eroded before great volumes of Oligocene to Miocene age ash flow tuffs were deposited (Proffett, 1972, 1976, 1977). Only the basal rhyolites (Trt and Tqrt) and the lower Mickey Pass tuff (Tbt) are preserved at the Buckskin prospect. They dip steeply westward and are locally vertical as a result of progressive regional tilting and faulting. The rhyolitic tuffs are pervasively sericitically altered, and weakly pyritized ( $\pm 1\%$ ). They have been intruded by younger mineralized porphyritic dikes that are associated with most of the hydrolytic alteration and mineralization in the area.

The ash flow tuffs are definitely pre-mineral, although they are not a good host for mineralization or alteration. They thus set parameters for the age of the porphyry system as well as helping decipher the structural problems at Buckskin.



### Tertiary Igneous Rocks - Mineralized

Mineralized porphyritic dikes, plugs, and sills occur throughout the prospect and largely post-date the faulting. One dike (Fbhp) was dated at about 15 million years old. Detailed mapping (Plate 2) shows that the mineralized intrusive rocks intrude late Oligocene to Miocene ash flow tuffs and cross-cut the tilted Basin and Range faults (see structural section). They are latitic to dacitic in composition with varying amounts of phenocrystic quartz, hornblende, and biotite. Strong pyritization and sericitic, argillic and propylitic alteration affect the dikes and the country rocks enclosing them (see alteration section).

Some compositional and textural trends occur from east to west. Phenocrystic quartz-rich dikes occur almost exclusively in the east. Dikes of biotite latite quartz porphyry with hornblende (Blqph) form a dike swarm at the east flank of the range and dikes of biotite hornblende porphyry with quartz (Bhpq) and fine quartz biotite porphyry (Fqbp) are also concentrated in the east. Phenocrystic quartz-poor rocks (Fbhp, Hbp and Bhp) are dominant in the central and western portions of the district. The dikes on the west tend to be fine grained also and the biotites are less stacked than in the dikes on the east and central parts of the prospect.

Fine biotite quartz porphyry (Fbqp) is perhaps the oldest of the mineralized Tertiary porphyries. It occurs only on the east flank of the range as several small dikes up to 100 feet wide and 600 feet long. In thin section, 2-3 mm rounded quartz phenocrysts and biotite books 2 mm thick and 4 to 5 mm long are set in a felted groundmass. Microlites of plagioclase and K-feldspar are normally  $\frac{1}{2}$  to 1 mm long. Hornblende



( $\pm 0.2$  mm) is widespread in the groundmass and sphene and apatite are common accessories.

Biotite-hornblende porphyry (Bhp) and hornblende-biotite porphyry (Hbp) are compositional variations of the same rock. They occur as small dikes and irregular sills up to 2500 feet long and 600 feet wide in the central area. Stacked biotite phenocrysts 3 to 5 mm wide characterize both these rocks. The hornblende-biotite porphyry megascopically has more large hornblende crystals, however. Feldspars are as large as 5 mm but are subdued by alteration. Quartz is rare in both these rocks with the possible exception of biotite-hornblende-porphyry with quartz (Bhpq).

Biotite-hornblende-porphyry with quartz (Bhpq) is similar to the above rocks except for containing 5% quartz phenocrysts and having more hornblende in thin section. A gradational contact to Bhp was mapped in the field but should be re-examined this summer.

Fine-grained biotite-hornblende porphyry (Fbhp) is compositionally very similar to biotite-hornblende porphyry (Bhp) and may be texturally gradational into it. It can be distinguished by the much smaller biotite phenocrysts that are distinctly less stacked. These dikes occur exclusively in the west portion of the district and have been dated by K-Ar methods at  $15.7 \pm 1.5$  million years in hole BSW-7 (Nielsen, 1975).

Biotite, latite quartz-porphyry with hornblende (Blqph), is a common rock in the eastern area and is probably one of the youngest of the mineralized porphyries. Large, rounded and embayed quartz phenocrysts 3 to 6 mm long, stacked biotite books 3 to 5 mm high, and large plagioclase and feldspar phenocrysts 4 to 5 mm long are characteristic for field identification. Quartz phenocrysts are not common and can be lacking altogether



in thin section. One or two in a fist-sized specimen is typical, and smaller, one mm crystals are common.

Fine-foliated hornblende quartz dacite porphyry (Fhqdp) occurs as small east-west trending dikes in the central and eastern parts of the property. They are very young and cut most of the mineralized porphyries except Blqph with which it doesn't come in contact. These rocks are very dark and fresh looking on the surface, and at first they were thought to be post-mineral. Thin sections reveal very fine disseminations of up to 5% pyrite and greater alteration than previously recognized.

Quartz biotite hornblende dacite porphyry (Qbhdp) forms dark dikes that are weakly chloritized but apparently are not mineralized with sulfides. These are among the latest dikes as determined through mapping.

#### Breccias

Breccias are limited to the western side of the prospect. A small pipe-like tectonic breccia containing porphyry and volcanic fragments in a rock flour matrix crops out about 500 feet northeast of hole BSW-1 (Plate 2). Float mapping indicated its dimensions as approximately 400 feet by 100 feet. A small pebble dike containing rounded silicified fragments of metavolcanics crops out east of the breccia. It is about six inches to a foot wide and dips 52 degrees to the east.

A small outcrop of breccia with a questionable porphyritic matrix with biotite hornblende porphyry and fine grained phaneritic(?) igneous fragments occurs southwest of hole BUE-9. It is highly altered to sericite and clay, and the nature of the matrix is difficult to distinguish.



### Tertiary(?) Lake Sediments

Slightly westward tilted lake sediments fringe the western margins of the prospect. Their gentle dip implies they are probably 11 million years old or younger.

### Quaternary Alluvium

Quaternary alluvial deposits can be separated into older and younger products of erosion. The old (Qoal) deposits are distinguished by granitic cobbles foreign to the Buckskin area. Recent alluvium fills valleys and covers gentle slopes.







## STRUCTURE

Proffett (1972, 1976, 1977) studied the Cenozoic structure of the Yerington area in detail for several years while working for Anaconda. He concluded that Basin and Range faulting accompanied regional westward tilting from Late Oligocene through at least early Miocene (Figure 13). This resulted in steeply dipping recent Basin and Range faults, mostly on the east flanks of ranges, but strongly tilted older inactive Basin and Range faults that are so gently dipping that they were earlier presumed to be gravity slide blocks.

Conoco's detailed mapping strongly supports Proffett's work. There are several older Basin and Range faults in the Buckskin area possessing gentle dips in the order of 20 to 30 degrees east (Figures 10, 11, and 13). They are easily recognized and traced where erosional remnants of the Oligocene ash flow tuffs are preserved. Faults within the metavolcanic country rocks, however, are extremely difficult to locate in the immediate vicinity of the prospect. Tertiary mineralized porphyry dikes have intruded along and across most faults and cut the ash flow tuffs (Plate 2, Figure 14).

Numerous other smaller presumed fault blocks which are defined best by juxtaposition of foliations within the metavolcanics are present within the large tilted Basin and Range fault blocks. Many of the bounding faults are only presumed to be present because of the lack of outcrop. The foliations in the Mesozoic country rocks dip moderately to steeply west and strike north-south except in the eastern quarter of the prospect where they strike east-west and dip mostly north 45 to 65 degrees. There is no apparent folding and the postulated structural feature responsible for the



large change in foliation does not fit well into the Cenozoic pattern of faulting. It is possible that it is a pre-Tertiary structure.

Most of the metavolcanics apparently were not disturbed by faulting until after the Oligocene ash flow tuffs were deposited. Foliations mimicking pseudo-bedding in the volcanics are nearly parallel to the ash flow tuffs that were deposited horizontally upon them. In the southwest part of the prospect the old pre-Oligocene erosion surface can be traced and hills and gullies preserving older ignimbrites are recognizable.

Some gentle folding of the metavolcanics can be seen in the breccia units on the west flank of the range (Plate 2). The fold axes trend west-northwest nearly coincident with the outcrop appearance of turquoise. The folding is also subparallel to the moderate and strong IP anomalous trends, suggesting it may be involved in some way with the mineralization event. It is probable that the turquoise is formed late as a result of supergene activity controlled by the fold axes.

An apparent thrust fault dipping gently west at the south-central edge of the property has emplaced Mesozoic volcanics over Tertiary ash flow tuffs. From present information it is not possible to determine if it is a true thrust fault or a gravity slide from the higher hill to the east. There does not appear to be an extensive gouge zone developed although much of the fault is covered by alluvium.

The trends of the faults are mostly north-south, northeast and west-northwest. The northerly trending faults are the traces of the old eastward, steeply dipping Basin and Range faults which are now tilted to the west so that the fault plane dips gently to the east (Figures 11, 12, and 13). Because of the trough-like nature of the old Basin and Range faults,



they usually curve into north-west trending boundaries on the south ends of the blocks and trend to the northeast on the north end of the blocks (Proffett, 1977). Often these blocks are terminated by later northeast and northwest trending faults (Figure 11).

Most of the later northeast and west-northwest faults displace Mesozoic(?) dikes and possibly some Tertiary dikes in a right lateral sense. In a series of west-northwest trending faults which dissect a Mesozoic (Clp) dike, some of the interior fault blocks have moved their segments of the dike right laterally farther to the east than adjacent faults and thereby locally producing an apparent left lateral sense of motion of several hundred feet.



Tilting and reduction of dip and curvature of early normal fault (fault of Fig. 8) by tilting due to displacement on later normal fault. Note that vertical displacement as well as tilting are due to hanging-wall block adjusting in shape to fill a potential void resulting from horizontal extension between hanging wall and footwall. Hanging wall could have adjusted to its new shape by development of antithetic faults instead of tilting, as has clearly been shown by Hamblin (1965), but in Yerington district tilting was dominant.

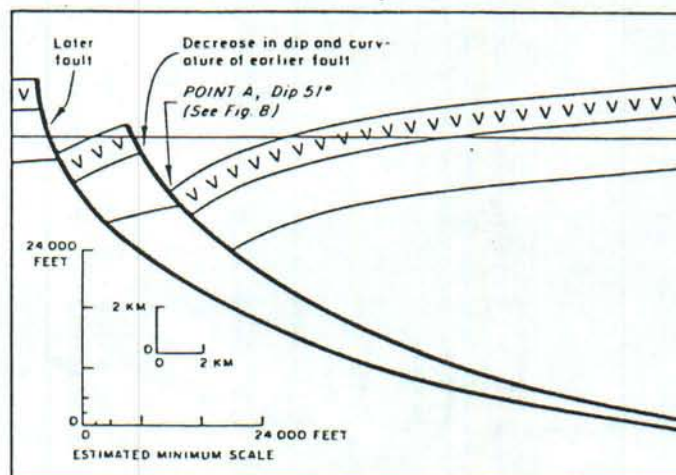
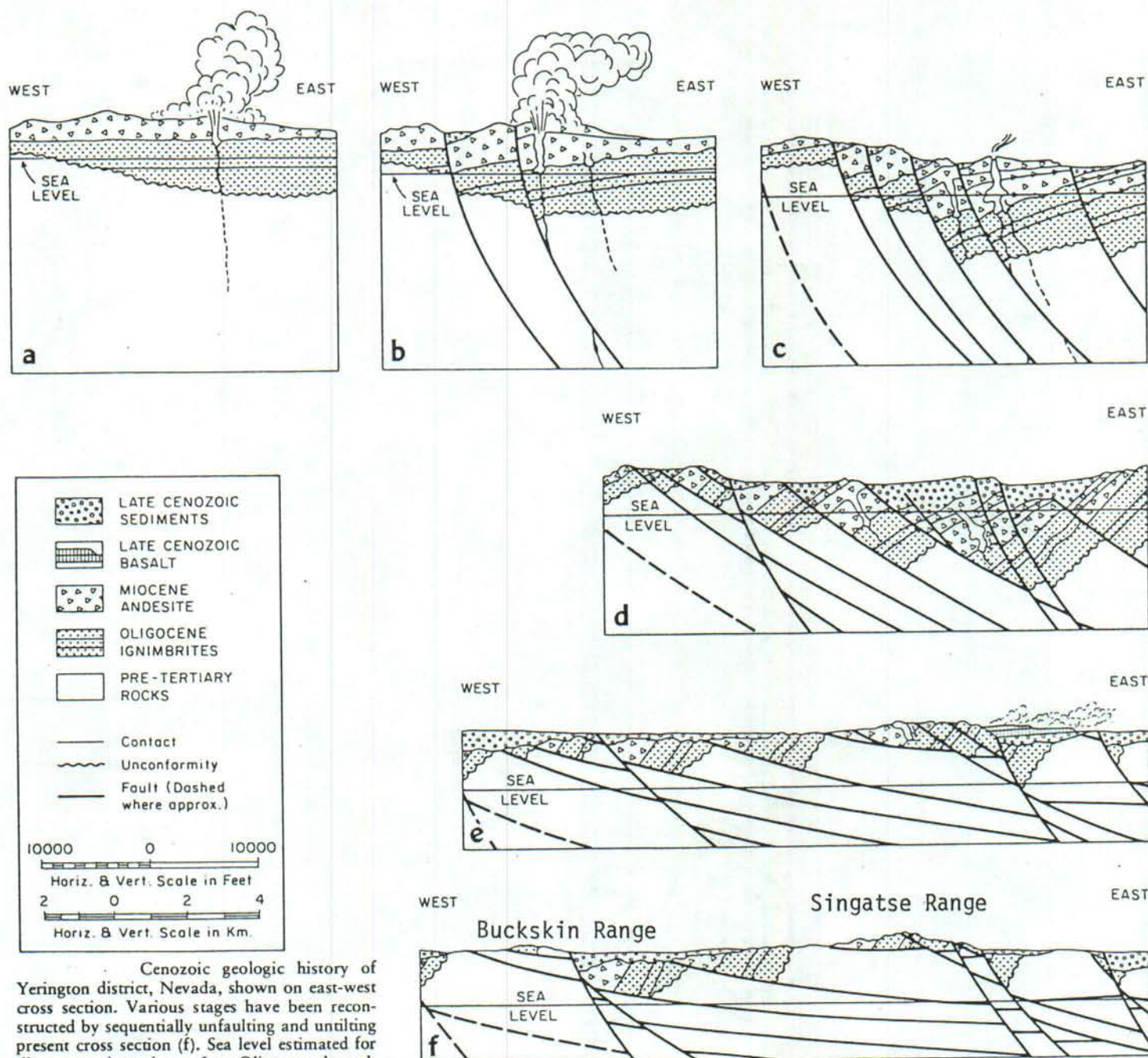


Figure 12

Tilting of Basin and Range faults  
From Proffett, 1977



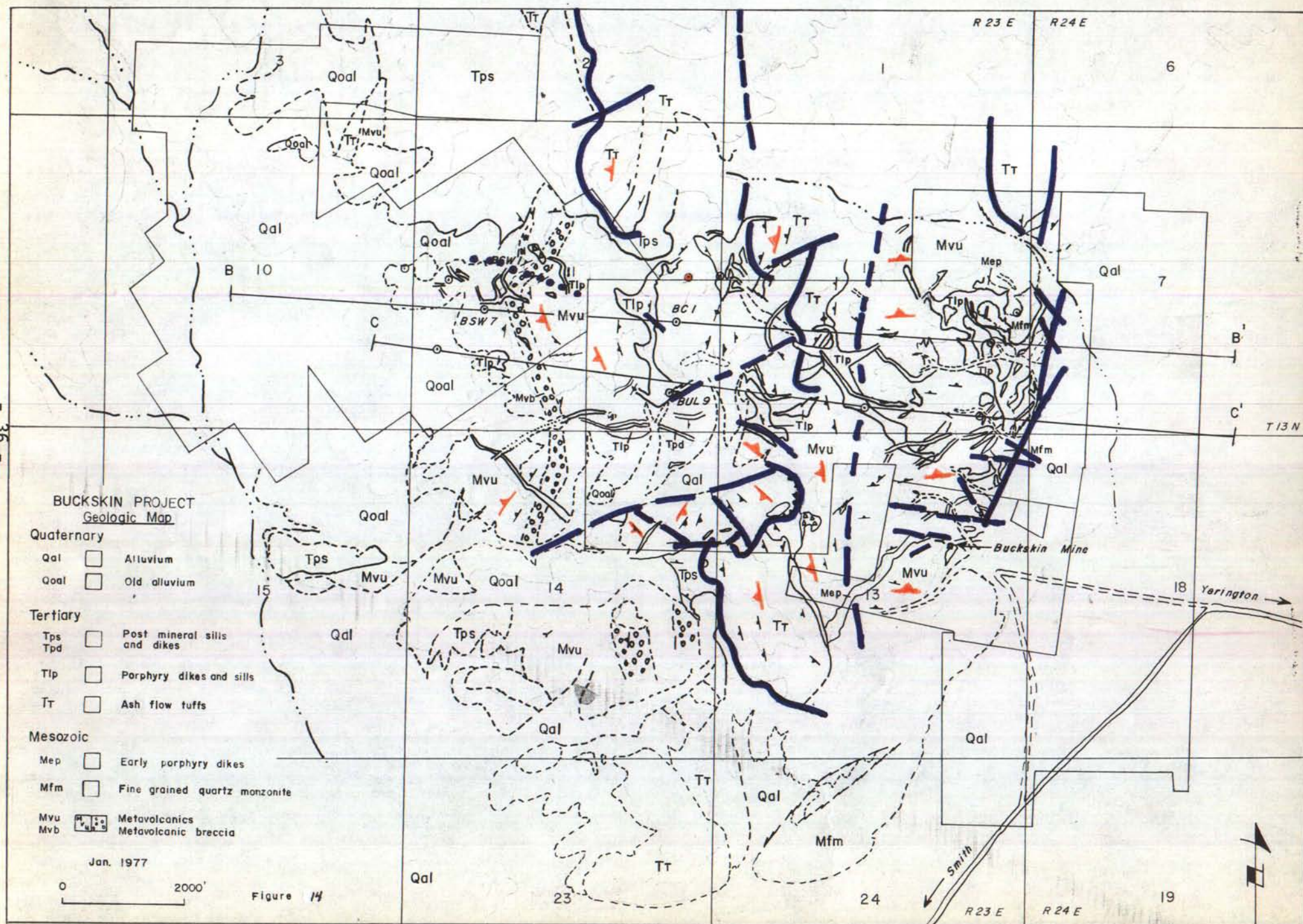


Cenozoic geologic history of Yerington district, Nevada, shown on east-west cross section. Various stages have been reconstructed by sequentially unfauling and untilting present cross section (f). Sea level estimated for diagrams a through e. a, Late Oligocene through early Miocene; deposition of ignimbrite sequence followed by beginning of deposition of hornblende andesite, 18 to 19 m.y. ago. b, Early Miocene, 17 to 18 m.y. ago; continued eruption of andesite and beginning of Basin and Range faulting and tilting. c, Early Miocene, 17 m.y. ago; end of andesite volcanism and continued Basin and Range faulting with erosion of up-faulted blocks. d, Early and middle Miocene, 11 to 17 m.y. ago; continued faulting, tilting, erosion of ranges, and deposition of sediments in basins. Early faults rotated to such a flat dip that they are no longer in favorable orientation for continued faulting. New steeper faults form to take their place. e, Eight to 11 m.y. ago; all early faults tilted to gentle dips and inactive, but faulting and tilting continues on newly formed steep faults; erosion of ranges and deposition of sediments and locally basalt in basins. f, Present time; continued faulting and tilting along steep range-front faults, erosion of ranges and deposition in basins.

Figure 13

Cenozoic Structural History  
From Proffett, 1977







## ALTERATION

Field observations defined the following pervasive mineral assemblages which appear to form alteration zones: sericite, strong clay-sericite, weak clay-sericite, mixed sericite-chlorite and chlorite-epidote. Field alteration studies of weathered outcrop are of necessity general and occasionally unreliable. Thin section studies of 112 surface and 134 drilling samples tend to confirm the patterns established in the field at Buckskin, however.

The sericite zone covers most of the prospect. It extends to the southwest and the northeast beyond our property line but has a relatively sharp north-south trending boundary with the mixed sericite-chlorite (propylitic) zone on the east (Figure 19). This boundary is interrupted in the north and west portions of the prospect by patches of clay-sericite (argillic to advanced argillic) and mixed sericite-chlorite. The feldspars and the groundmass in this zone are typically 80% or more replaced by sericite. Accessories such as sphene and apatite are relatively unaltered. Strong surface leaching of the sulfides is everywhere apparent within the sericite zone. The limonites are predominately jarosite even where the initial sulfide content was low.

The largest clay-sericite alteration zone is in the west. Sericite and low birefringence clays are pervasive in this zone. Rarely the K-feldspars are only 50% altered. Some silicification of the groundmass occurs and quartz and clay veins are most common in these localities. The surface boundaries of this alteration zone were defined by the presence of sparse clay and alunite veining. Preliminary X-ray diffraction data by Mr. Bamford



indicate that the western and largest area of strong clay-sericite may represent advanced argillic alteration as defined by the presence of natroalunite, diaspore and pyrophyllite. The central elongate strong clay-sericite zone centered about hole BC-1 probably does not represent advanced argillic alteration. The clay veins here are metahalloysite, and alunite is apparently not present. The clay veins are probably supergene products.

The weak clay-sericite alteration has been studied in hand specimen only. It is found in the ash flow tuffs of the Mickey Pass sequence and is defined by a greater percentage of clay than is present in the strong clay-sericite zones and yet weaker intensity of alteration. Rock type may be an important factor in this zone because the original tuff apparently contained a great deal of clay. It is not immediately obvious what the significance of this zone is.

The mixed sericite-chlorite zone forms most of the eastern part of the prospect and possibly more of the central portions of the prospect than previously was recognized. On the surface it is characterized by weak to strong sericite mixed intimately with weak to moderate chlorite both microscopically and megascopically. Thin sections reveal a chlorite - sericite - calcite - epidote alteration assemblage in both dikes and volcanic country rock. The intensity or thoroughness of the alteration varies considerably from about 15 volume percent chlorite - calcite - sericite to greater than 90%. Mapping of the intensity from thin sections has not been undertaken at this time.

Propylitic alteration is found in the northern and southern portions of the property in late Tertiary sills and plugs. It also occurs in relatively fresh looking Mesozoic andesites. Regional metamorphism probably accounts for the propylitic assemblage in the older volcanics. The alteration



appears to be deuteric in the younger sills and dikes and is probably unrelated to the overall hydrothermal system which produced the other suites of minerals.

Bear Creek drilled five diamond drill holes in the western part of the prospect. A moderate to strong pervasive sericite - clay - quartz  $\pm$  chlorite and calcite assemblage occurs in BSW-1 to about 600 feet depth. This assemblage surrounds the highest copper interval in the hole from 320 to 450 feet. Minor secondary biotite with chlorite alteration of mafics occurs in a dike at about 660 feet and again in a smaller dike at 820 feet. It is uncommon and does not appear to be related to higher copper values. The first dike also has strong sericite and occasionally moderate chlorite and calcite alteration to about 700 feet. Strong quartz-sericite (phyllic) alteration continues to about 1100 feet and decreases in intensity to the bottom of the hole at 1313 feet. Traces of calcite and chlorite alteration are also present near the bottom of the hole.

There is no core from the surface to 464 feet in BSW-7 but Bear Creek logs indicate there is mixed sericite - chlorite pervasive alteration. Extreme to very strong pervasive quartz - sericite - pyrite (phyllic) alteration occurs from about 400 feet to 780 feet. Pervasive biotite - chlorite - sericite (intermediate potassic) alteration from 780 feet to 860 feet coincides with the highest copper values. Late chlorite ( $\pm$ calcite) veinlets are prominent in this zone and are associated with most opaques in detail. Overall, however, there is a strong association of copper and the secondary biotite - sericite - chlorite assemblage. Strong quartz-sericite occurs from 860 to about 1500 feet and then gives way to weak to moderate chloritization to the bottom at 1715 feet. Some quartz veins and



numerous late chlorite - calcite  $\pm$  pyrite veinlets are common throughout the hole. Gypsum veinlets occur from 700 feet to the bottom of the hole.

Hole BSW-9 encountered moderate to strong pervasive chlorite and sericite alteration both above and below the higher copper interval, from 750 to 1050 feet. This alteration assemblage is associated with high zinc values as it also is in drill holes BSW-7, BSW-1 and BC-1. The higher copper interval contained moderate to strong pervasive sericite - chlorite and siderite(?) alteration but no secondary biotite as in BSW-7.

Conoco's hole BC-1 encountered moderate to strong pervasive sericite - chlorite - calcite alteration in the top 1000 feet of the hole associated with strong lead-zinc mineralization. Clay veins (metahalloysite), quartz - calcite - chlorite veinlets and late chlorite veinlets also are common. The clay veins may be a surface weathering product and may not be representative of hypogene alteration. Epidote - chlorite - calcite - actinolite alteration is strong and pervasive from 1000 to 1650 feet with occasional massive epidote lenses occurring to the bottom of the hole. Zones of strong quartz sericite alteration, one inch to several feet wide, overlap the epidote - chlorite - amphibole alteration from 1650 feet to 1994 feet accompanying an increase in lead and zinc and total sulfides. Vein and disseminated tourmaline occur at 1942 feet.

Overall the hole has mostly a propylitic alteration assemblage with increasing amphibole, tourmaline(?) and strong quartz - sericite zones near the bottom of the hole. Mr. Bamford's X-ray diffraction of the 100 feet composite pulps from hole BC-1 show the presence of orthoclase from 1200 to 1960 feet. His total X-ray results are shown below.



Interval of composite pulp in feet	Quartz	Sericite	Chlorite	Calcite	Amphibole	Orthoclase
20 - 100	x	x				
100 - 200	x	x				
200 - 300	x	x	x			
300 - 400	x	x	x	x		
400 - 500	x	x	x	x		
500 - 600	x	x	x	x		
600 - 700	x	x	x	x		
700 - 800	x	x	x	x		
800 - 900	x	x	x	x		
900 - 1000	x	x	x	x		
1000 - 1100	x		x	x	x	?
1100 - 1200	x	x			x	?
1200 - 1300	x	x	x		x	x
1300 - 1400	x	x	x	x	x	x
1400 - 1500	x	x	x	x		x
1500 - 1600	x	x	x	x	x	x
1600 - 1700	x	x	x	x		x
1700 - 1800	x	x	x	x		x
1800 - 1900	x	x	x	x	x	x
1900 - 1960	x	x	x	x		x

Holes Bue-1, Bue-2, D-281, and D-289 on the eastern portion of the property all have alteration assemblages roughly corresponding to surface samples. Pervasive mixed sericite - chlorite - calcite  $\pm$  epidote alteration occurs in varying proportions and mostly with weak to moderate intensity. No definite trends of intensity or proportion of alteration minerals can be determined at this time.

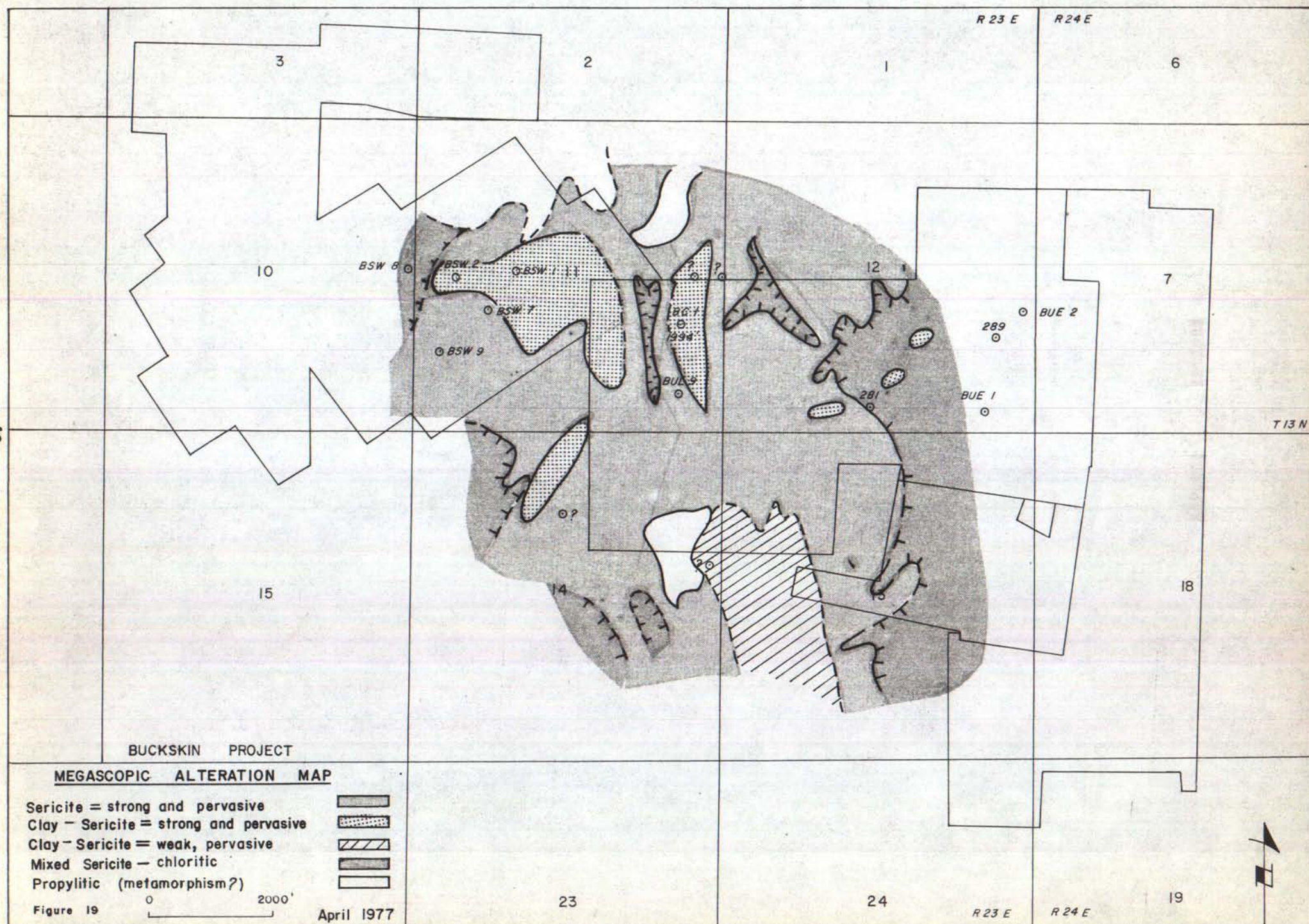
Although the extreme eastern area shows no clear trends the holes in the west, including hole BC-1 show rough alteration zoning trends. Hole BSW-7 encountered a sericite - chlorite zone (propylitic?) high in the hole associated with a strong zinc anomaly. This alteration passes into a very strong quartz - sericite - pyrite zone which gives way to a chlorite -



biotite - sericite assemblage associated with the highest hypogene copper values yet drilled on the property. This assemblage is by no means the "core" potassic zone of a classic porphyry system, but may instead be the outer margins of what may be termed an intermediate potassic zone such as is found at Red Mountain, Arizona (Corn, 1975). Alteration decreased below the chlorite - biotite zone (intermediate potassic) to a mixed sericite - chlorite (propylitic) assemblage again. This possibly indicates that the hypogene system may be slightly tilted and therefore plunging, perhaps steeply, to the east; or possibly the alteration is following a structural zone of weakness that is tapping the main system to the east.

Holes BSW-9 and BSW-1 lack the extreme quartz-sericite alteration and the chlorite - biotite assemblage of BSW-7. Therefore it appears that the system does not extend to the north, west or southwest, and hence must lie east or southeast of BSW-7. Surface alteration mapping also indicates an area east or southeast of BSW-7, centered near the advanced argillic alteration zone, as the primary target area (Figure 19). Hole BC-1 has an expanded sericite - chlorite (propylitic) zone with increasing amounts of amphibole and orthoclase near the bottom. This could indicate that either the hole was drilled very high over the top of the system or obliquely down the side of the target. If the latter is the case, then BC-1 would have to be northerly (north, northeast, or northwest) of the target due to geological, geophysical and geochemical evidence.







## LIMONITE CAPPING

Surface mineralization consists predominately of limonites that formed after pyrite. A few patches of less weathered rocks bearing fresh disseminated pyrite occur randomly throughout the prospect. Figure 35 is a volume intensity summary map of surface limonites. It depicts only my visual estimates of "moderately" abundant limonites on fractures, combined with a jarosite content generally greater than 75% of the total limonite volume. Its shape and size agree fairly well with the moderate IP response, which represents about four to five percent pyrite in drill holes (see discussion on Induced Polarization).

Hematite stains are scattered everywhere, but in a manner that is not amenable to mapping. Typical "lavender" hematite after chalcocite can be seen in a few places and drilling in holes BSW-1, 2, 7 and 8 indicates there is a small amount of secondarily enriched copper beneath the surface staining. "Brick red" hematite stains are also common and usually are associated with small fractures or fault zones. This color and type of occurrence is not related to copper mineralization.

Jarosite is typical of the thoroughly leached zones overlying moderate to strong pyritization and usually occurs in areas of sericitic and/or argillic alteration, possibly because leaching is favored by these alteration products.

Goethite stains on fractures are found in less altered rock indicating generally less disseminated pyrite than the area of predominately jarositic limonites. Goethite is more abundant in the propylitically altered volcanics on the east flank of the prospect.



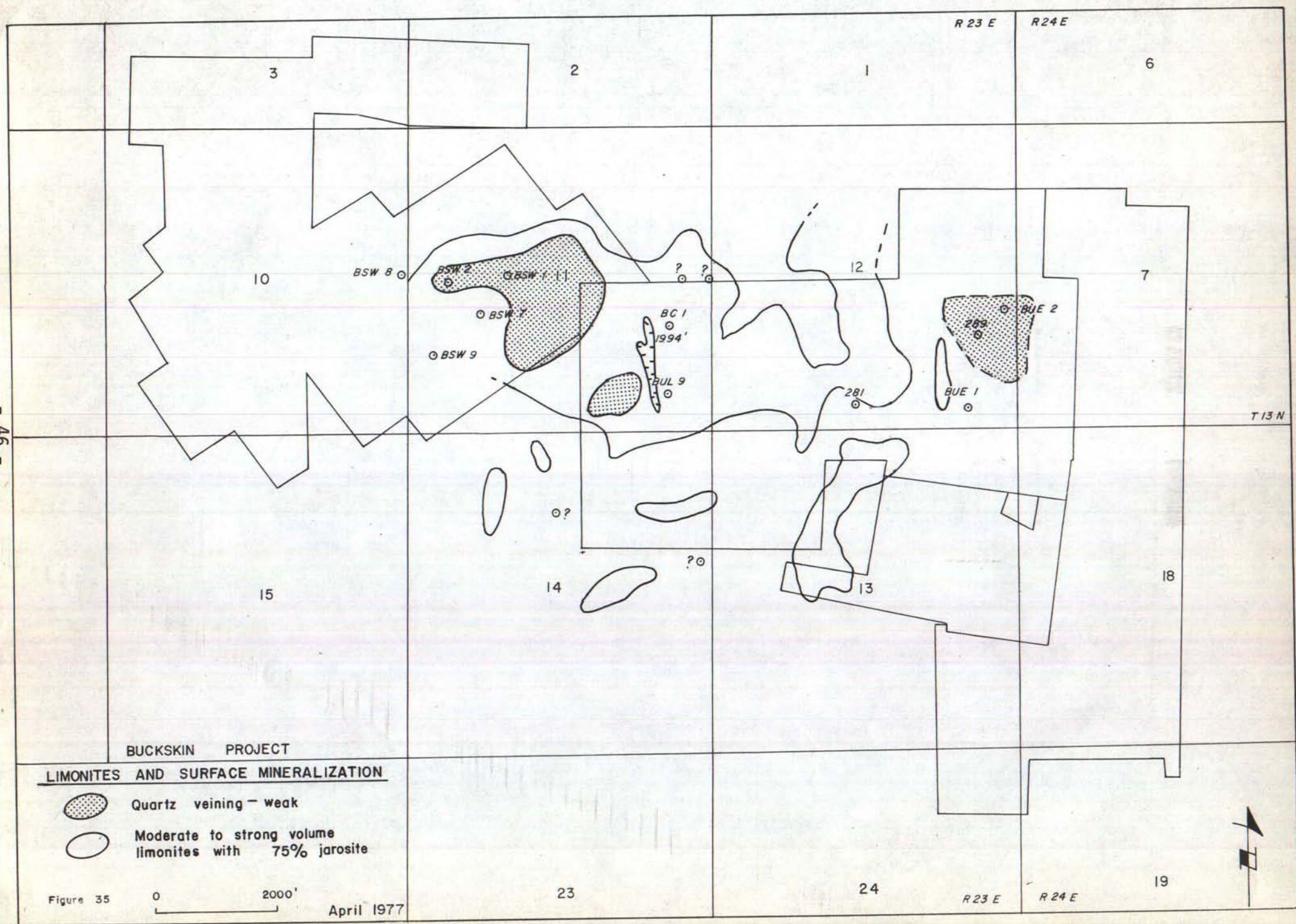
All three limonites can be found on separate fractures in practically the same hand specimen, but they are characteristically fairly pure. Jarosite is usually 90% pure (estimated) and only occasionally has more than 10 to 20% blended hematite and/or goethite. The same is true for the hematite stains. Goethite may have more mixtures of jarosite and hematite, but it is more difficult to estimate the impurities in goethite.

Turquoise veinlets are abundant in roughly linear west-northwest trending zones on the eastern and western sides of the property. They do not quite join in the center of the prospect (Plate 8).

Outcrops showing malachite, azurite and "neotocite" (Mn-Cu oxide) are sparse. These minerals can be found along small veinlets and fractures. They occur in weakly altered and mineralized fringe areas where strong leaching has not taken place. Their presence is not related to, or representative of, soil geochemistry.

Small quartz veinlets invade the Mesozoic volcanics in the vicinity of the Bear Creek holes BSW-2, 7, and 1, and locally on the east (Figure 35). They are not abundant and do not form the typical stockwork veining of many porphyry copper systems.







## GEOCHEMISTRY

### Soil Geochemistry

Soil samples were collected over a large area for geochemical analysis early in 1976 as a first step in defining the Buckskin system. To this end the program was entirely successful.

Approximately five hundred samples were collected on a 600 by 300 foot grid and assayed for copper, molybdenum, lead, zinc, and silver. The property has only a thin to non-existent soil horizon overlying a few inches to several feet of decomposed bedrock. The samplers tried to consistently collect the finer material from the top of the decomposed bedrock. Quaternary alluvium was also sampled and serves as a rough guide to background values. Contour intervals were selected by comparison of frequency histograms of the values of each element.

All elements have closely similar outer limits defining the geochemically anomalous areas. Only the distribution of the higher values changes significantly within this pattern. It should be noted that the geochemically anomalous values do not coincide directly with the strongest alteration. Although the values generally define a large roughly east-west trending system, three areas (western, central and eastern) can be separately discussed.

The western anomaly was drilled by Bear Creek before 1973. It consists of a copper anomaly with weak zinc and silver anomalies. The outer or western part of the central lead anomaly just touches this area. Widespread thin veinlets of turquoise may account for the high copper values, but a thin blanket of supergene chalcocite 120 to 170 feet deep underlies the



surface anomaly.

The central anomaly is a large bulbous shaped area containing high values of lead and silver, moderate zinc and low copper. Its eastern boundary is roughly defined by decreasing values in the vicinity of a large tilted Basin and Range fault which drops Tertiary ash flow tuffs down against Mesozoic metavolcanics. These ash flow tuffs are sericitically altered but are apparently a poor host for mineralization. Drilling indicates the high lead-zinc values are representative of the underlying bedrock, but the anomalous zone may dive westerly beneath the western geochemical anomaly.

The eastern anomaly is characterized mostly by northeasterly trending linear zones of higher values of copper, lead, and zinc, within a broad area of moderately anomalous values. Lead is a possible exception in that its linear elements form a roughly arcuate zone. This linearity may in part be due to drainage. Most drainages, however, trend easterly and not northeasterly. The soil survey terminated to the east in alluvium parallel to the fault that bounds the range. Not much is presently known about the controls of mineralization in the east or exactly how it relates to the central and western soil anomalies.

Soil geochemistry has many deficiencies due to both sampling technique and dispersion effects by ground water and local drainages. It is, however, a relatively quick and easy way of defining large geochemically anomalous areas and is essentially an averaging technique.

#### Rock Chip Geochemical Sampling

Rock chip sampling is less subject to the many errors of soil sampling.



COTTON FIBER CONTENT

The samples, however, are more tedious to collect, especially where large amounts of land are covered by alluvium as at Buckskin.

Because of the timely availability of Mr. Bob Bamford, formerly Senior Research Geologist with Kennecott, it has been decided to retain him to conduct a rock chip limonite geochemical study of the Buckskin area. Mr. Bamford hopes to use the techniques and interpretive data developed by him at Kennecott in the last three years to predict the direction and depth to the core zone of the hydrothermal system. Mr. Bamford will not complete his program until after this report has been finished; unfortunately, we cannot discuss his methods and results at this time.



BUCKSKIN PROJECT  
Soil Geochemistry

60-139

140-259  
COPPER  
(ppm)

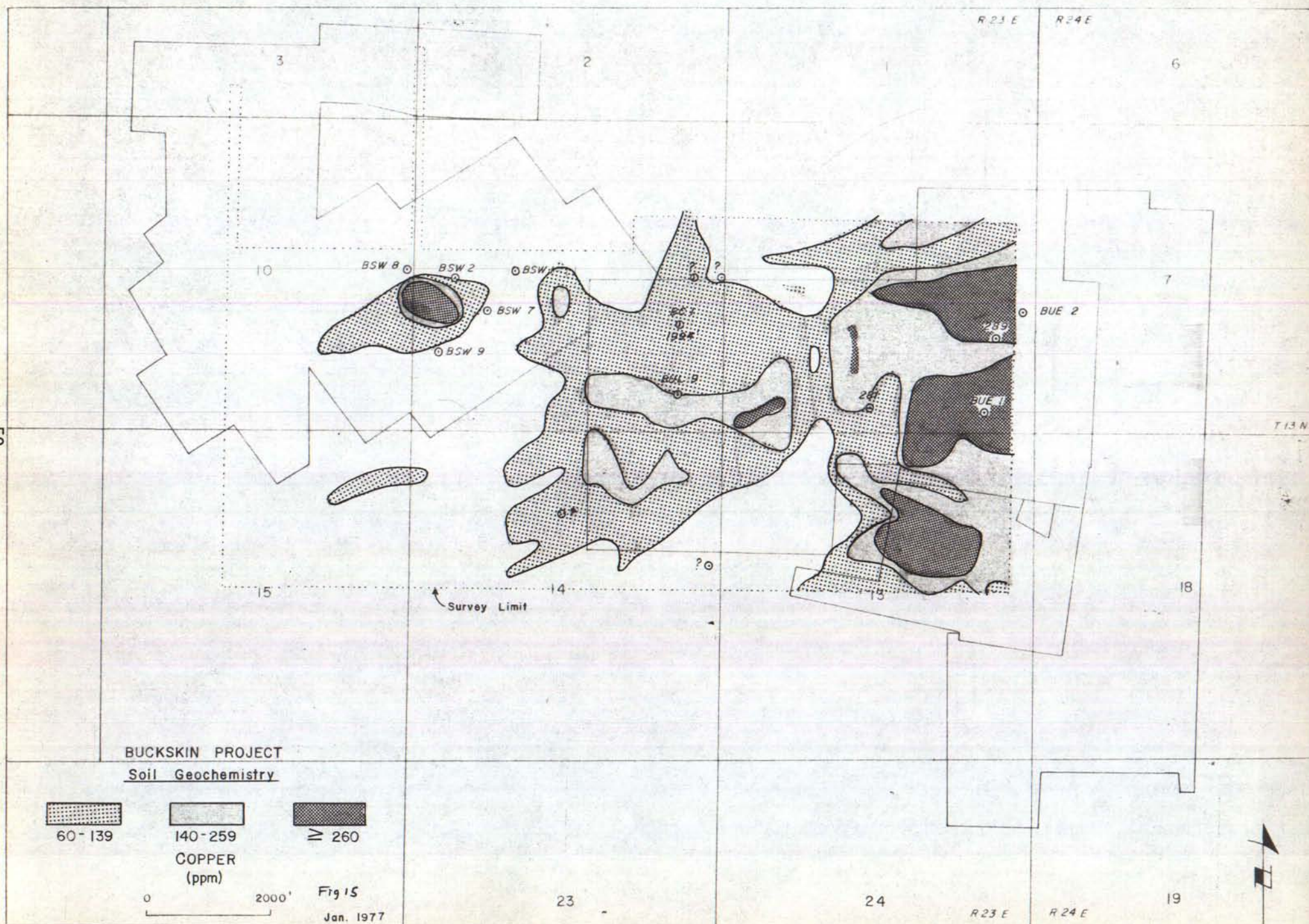
≥ 260

0 2000'

Fig 15

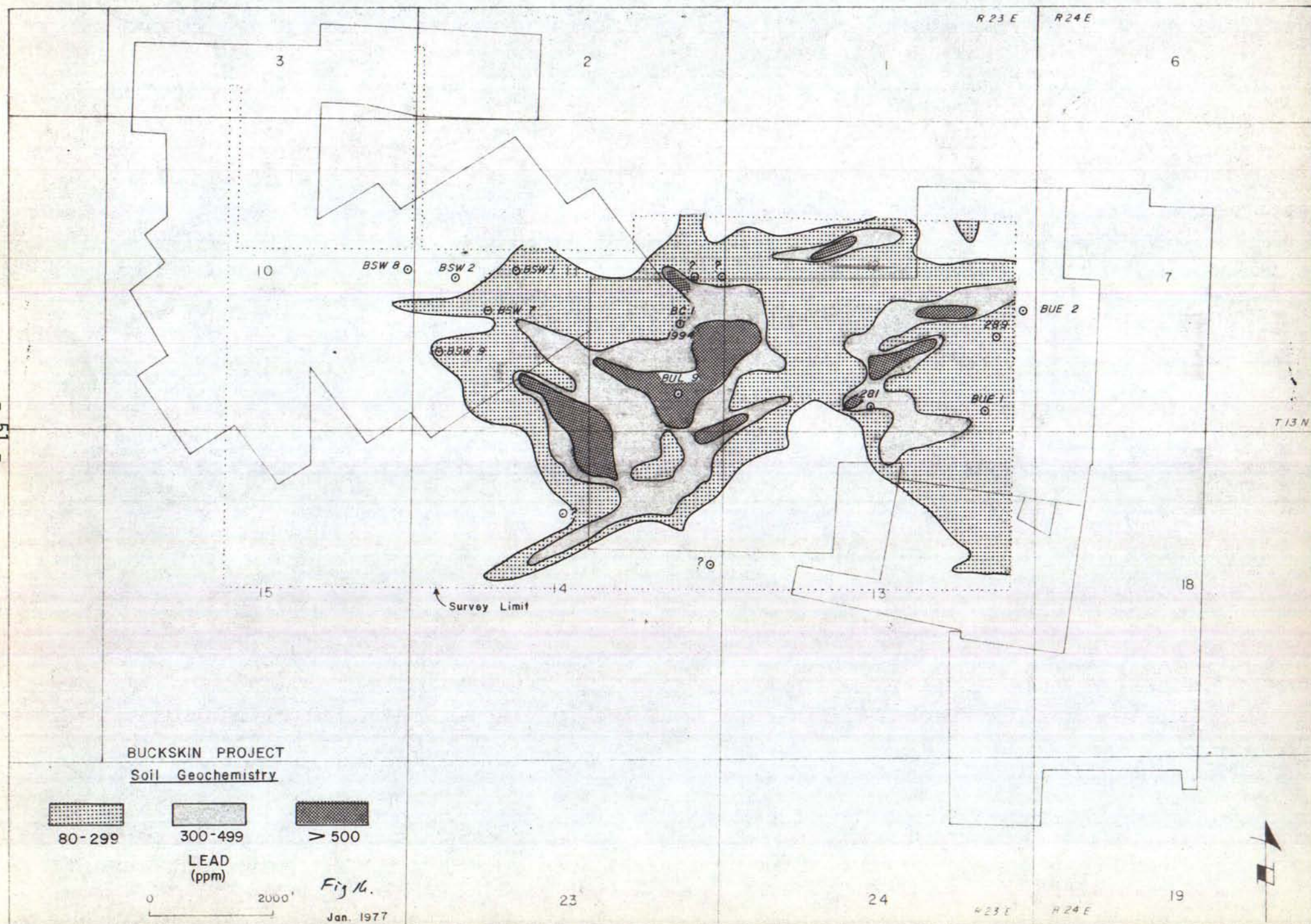
Jan. 1977

- 50 -



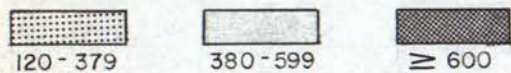


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BUCKSKIN PROJECT  
Soil Geochemistry

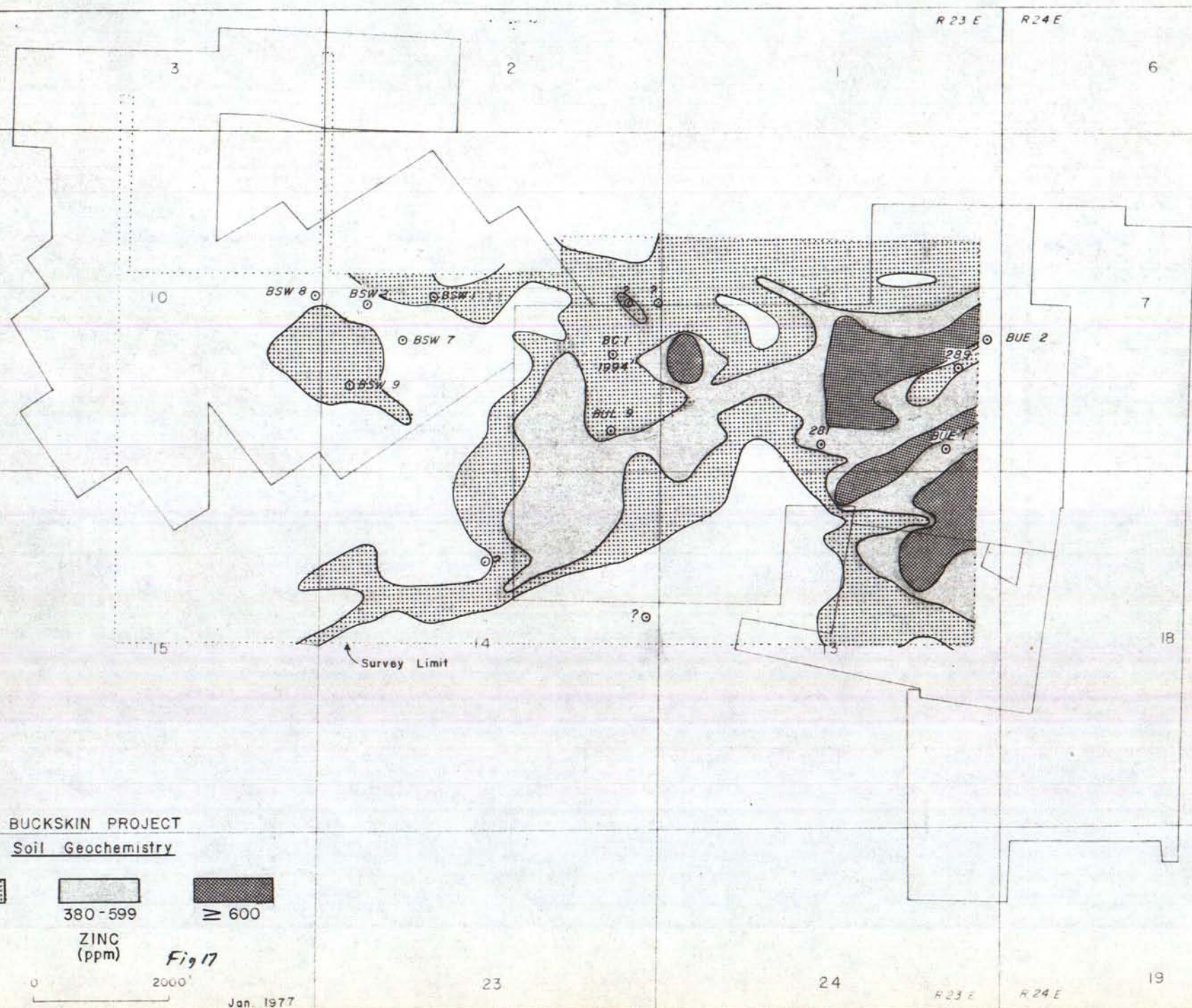


ZINC  
(ppm)

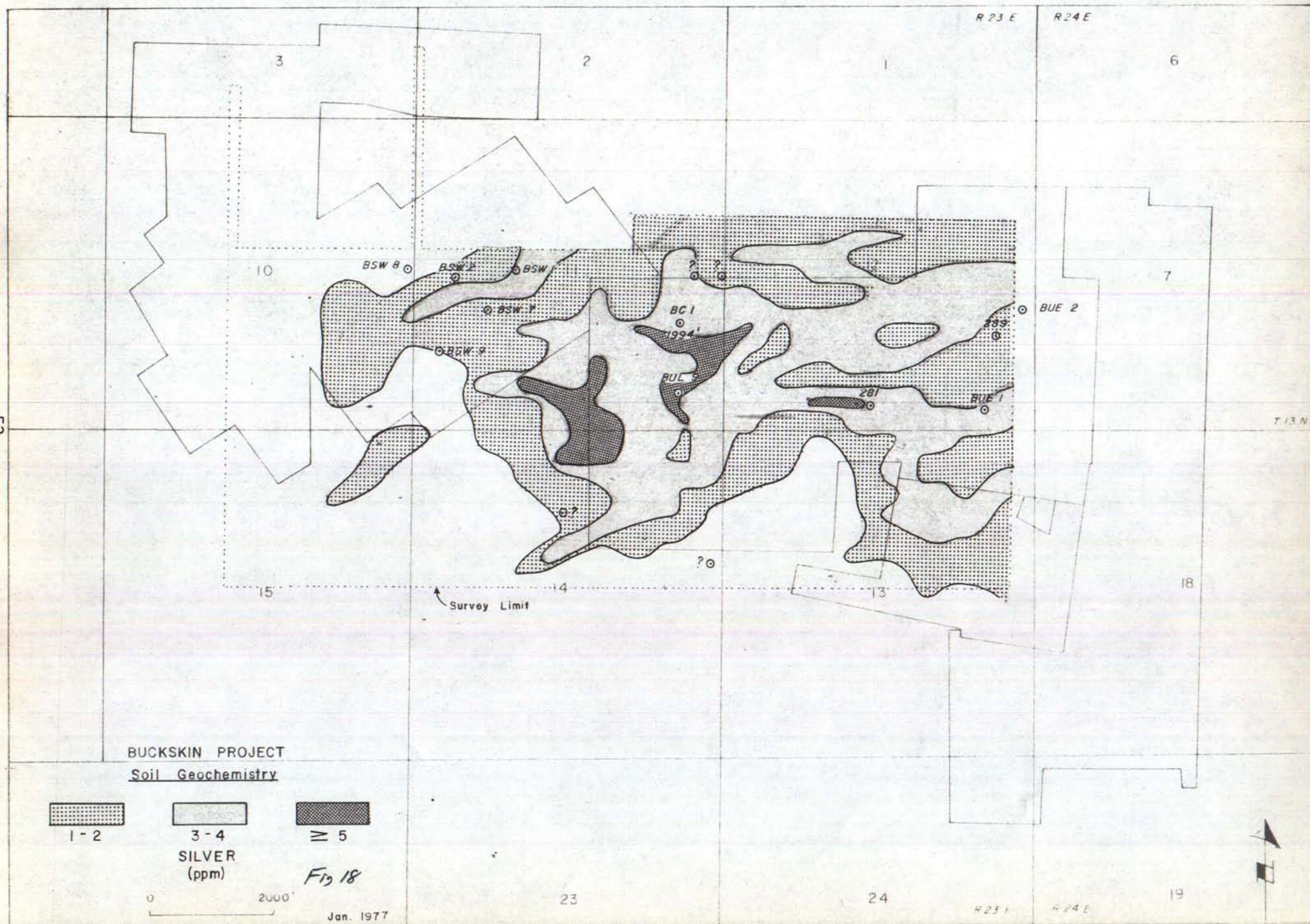
Fig 17

0 2000'

Jan. 1977









## GEOPHYSICS

### Induced Polarization (IP)

Conoco ran approximately 23 line miles of IP consisting of two east-west and six north-south lines in 1976 to complement and confirm IP data obtained from Bear Creek and Phelps Dodge. Lines 10W and 20N had 2000 foot dipole-dipole spacing; all the rest had 1000 foot dipole-dipole spacing. Mining Geophysical Surveys conducted the survey of the two east-west lines, and a Conoco crew collected the data for the north-south lines. Joe Anzman was hired as a consultant to interpret all the new data and the interpretation, which is summarized below, is based on his report and several hours of personal consultation.

The IP surveys are summarized on Figure 20 and Plate 13. The data indicate a large anomalous area 9000 by 12,000 feet elongated east-west and extending to at least 3000 feet in depth. Within this area is a "moderate" anomaly 2500 feet wide, 11,000 feet long and at least 2000 feet deep. The strongest anomaly extends 4000 feet east-west and is about 1000 feet wide and approximately 1500 feet deep. It is in a central position relative to the lesser anomalies and apparently has steep boundaries. The presence of sulfides as the source of the anomalies has been confirmed by drill holes.

Extensive discussions with Mr. Anzman pointed out several areas that were not clear in his report. The "strong" anomaly of Mr. Anzman's interpretation may or may not be significant to the interpretation of the entire system. Drill holes outside the "strong" anomaly (BSW-7 and Bue-9) have abundant sulfides but not necessarily of an amount or pattern to support



the existence of the strong anomaly as anything more than a statistical boundary.

His belief that the sulfides (IP anomalous area) decrease significantly at depth was not stated clearly in his text. The "anomalous" response decreases somewhere below a depth of 3000 feet and the "moderate" anomaly decreases at 2000 to 3000 feet in depth. Drill hole BC-1 confirms the presence of sufficient sulfides to cause the "moderate" anomaly to at 2000 feet. Although Mr. Anzman feels he can interpret data to about 2500 to 3000 feet with 2000 foot dipoles, at best there can be but a fuzzy picture at such depths. Therefore, he finds it permissible that the apparent sulfide-decrease could be the center of a porphyry system containing lower overall sulfides but a higher chalcopyrite to pyrite ratio, as in a classic "porphyry" system. Other interpretations are possible, but the lack of detailed geologic knowledge at depth allows only optimistic interpretations that are permissible with the known geology.

Line 20N, the east-west line which was run with 2000 foot dipoles, indicates a deep anomaly on the downthrown side of the Basin and Range fault on the east flank of the range. Conoco has staked this area and will direct future efforts to explain this anomaly (Figures 10 and 20).

The IP surveys show that a large sulfide system is present at the Buckskin prospect. Its limits are probably approximated by the anomalous area of Mr. Anzman's interpretation, as indicated by previous drilling. The sulfide system trends east-west, transects all local geology and structure, and appears to be decreasing at the depth limits of the IP method. It cannot be determined whether this decrease is permanent or local, or exactly what patterns the sulfides define at depth, but the



geology suggests that the Buckskin ground surface is near the top of a large, strong, porphyry-type sulfide system. The decreasing sulfides at depth may be the core of the system and the target area for copper.

Until more drilling has been accomplished, it must suffice to note merely that Conoco controls nearly all of the geophysically anomalous ground, especially the "moderate" and "strong" anomalies. Specific target location based solely upon IP is not feasible with the information presently available.

#### Local Airborne Magnetic Survey

Conoco flew an aeromagnetic survey over the Buckskin Range with a line spacing of 650 feet in the fall of 1976. No official interpretation was performed by a geophysicist. Joe Anzman looked briefly at the data while examining the IP results and concluded that it was not generally suggestive of any pattern relating to the known geology or the IP responses and hence was not very useful. The Conoco Reno geological staff has noted, however, that there is a large, quiescent area of relatively low magnetic values roughly correlative to the surface exposure of the mineralized rocks. It appears to trend northeasterly and is partially surrounded by small positive anomalies (Figure 21). Some of the high magnetic values are not readily explained because they occur over the altered Mesozoic volcanics instead of late Tertiary andesites. The high response in the southeast portion, however, may be coincident with relatively unaltered Mesozoic andesites.

#### Regional Airborne Magnetic Survey

A regional aeromagnetic survey was completed by the USGS in 1971



(Figure 22). It shows a roughly east-west trending, broad magnetic low which nearly completely contains the Buckskin prospect. It is difficult to determine if Conoco's detailed air mag would generally conform to this pattern without more lines to the west of our present survey. The Minnesota iron mine and the Pumpkin Hollow (Lyon) iron-copper district have definite positive magnetic expressions whereas the Yerington porphyry copper open pit is within the north end of a negative magnetic anomaly. Other deep porphyry copper deposits found by Anaconda lie on the margin of the same large magnetic depression.

In general both the local and regional negative magnetic anomalies at the Buckskin prospect are correlative and are probably caused by hydrothermal destruction of pre-existing magnetic minerals.



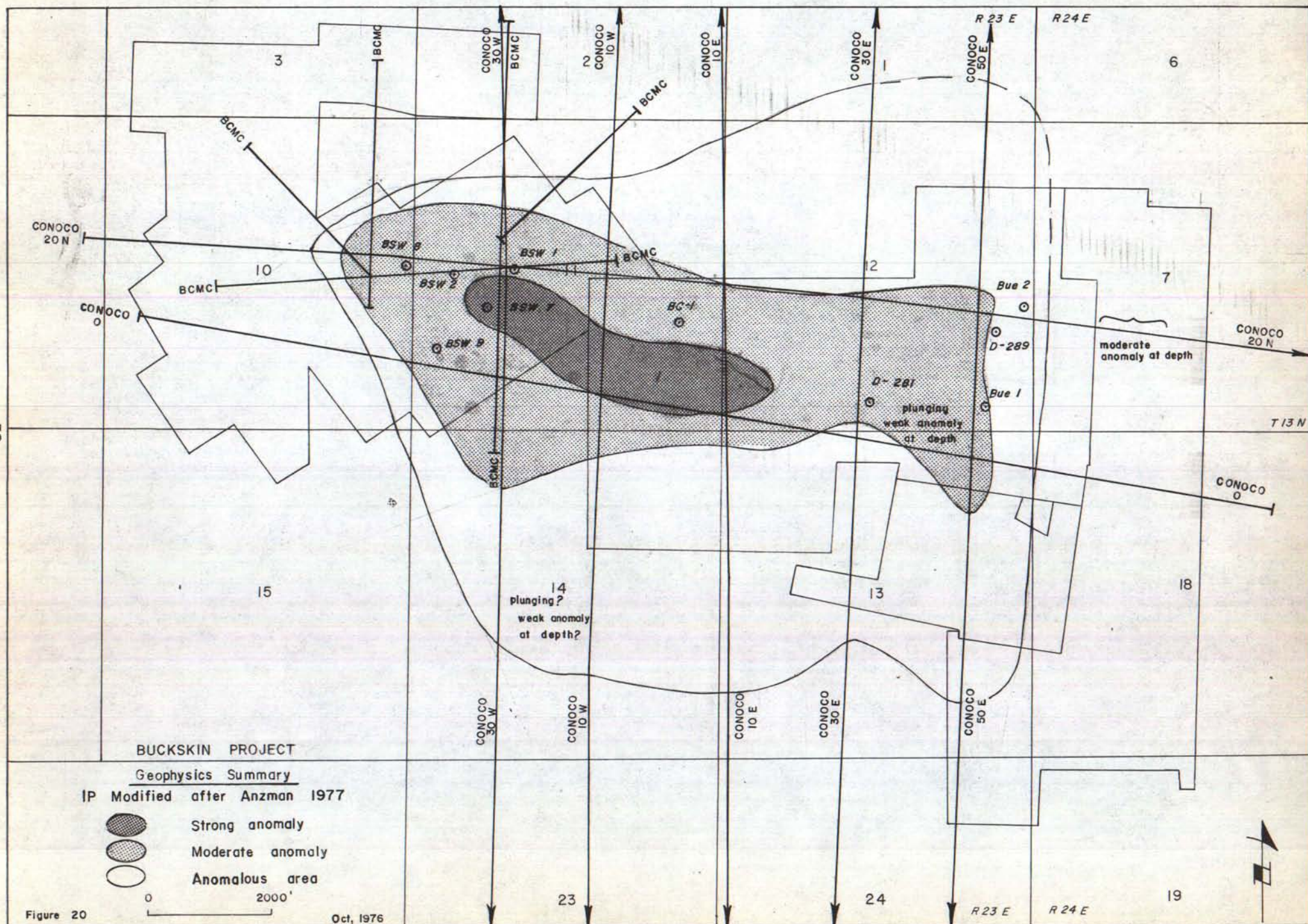
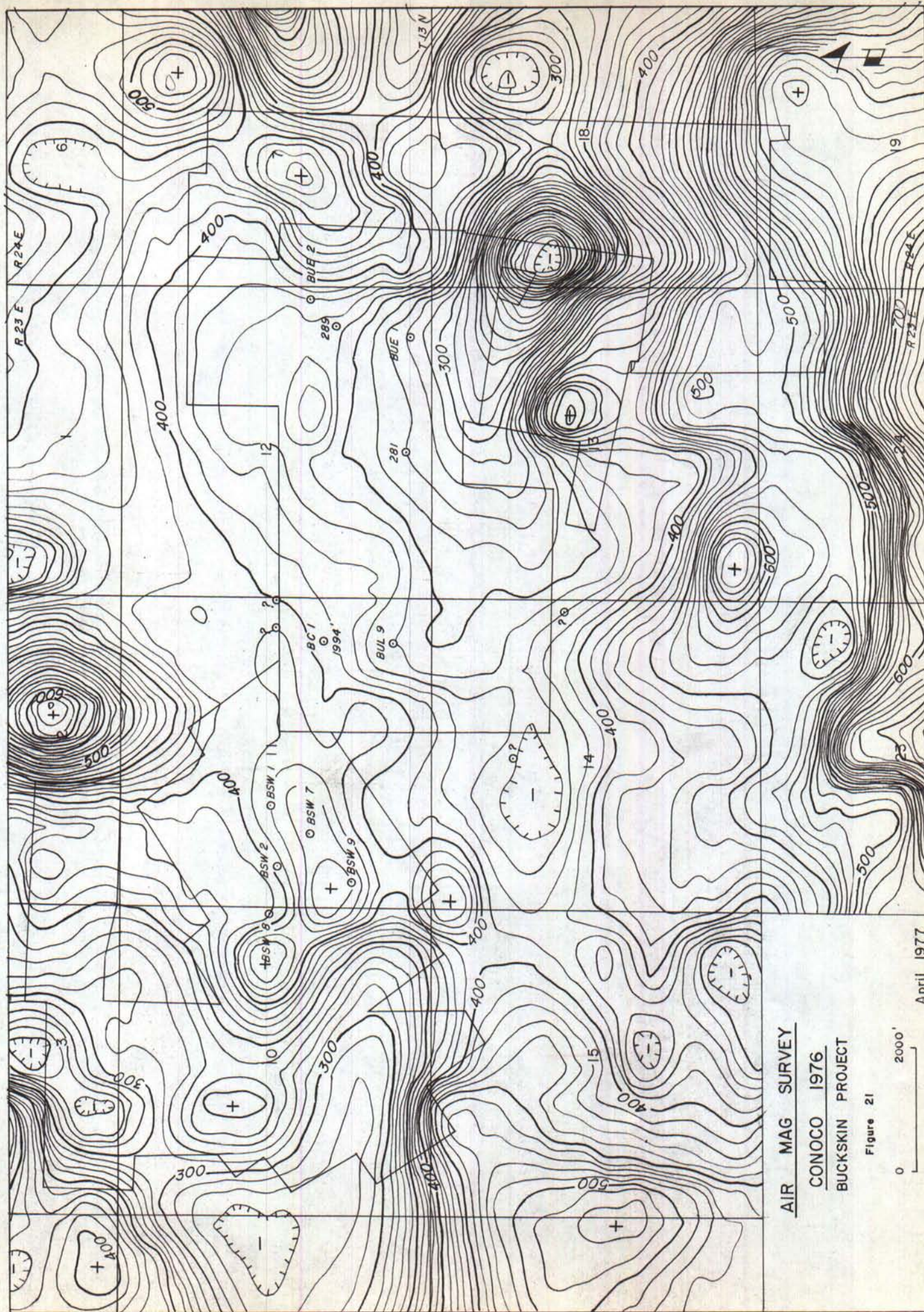
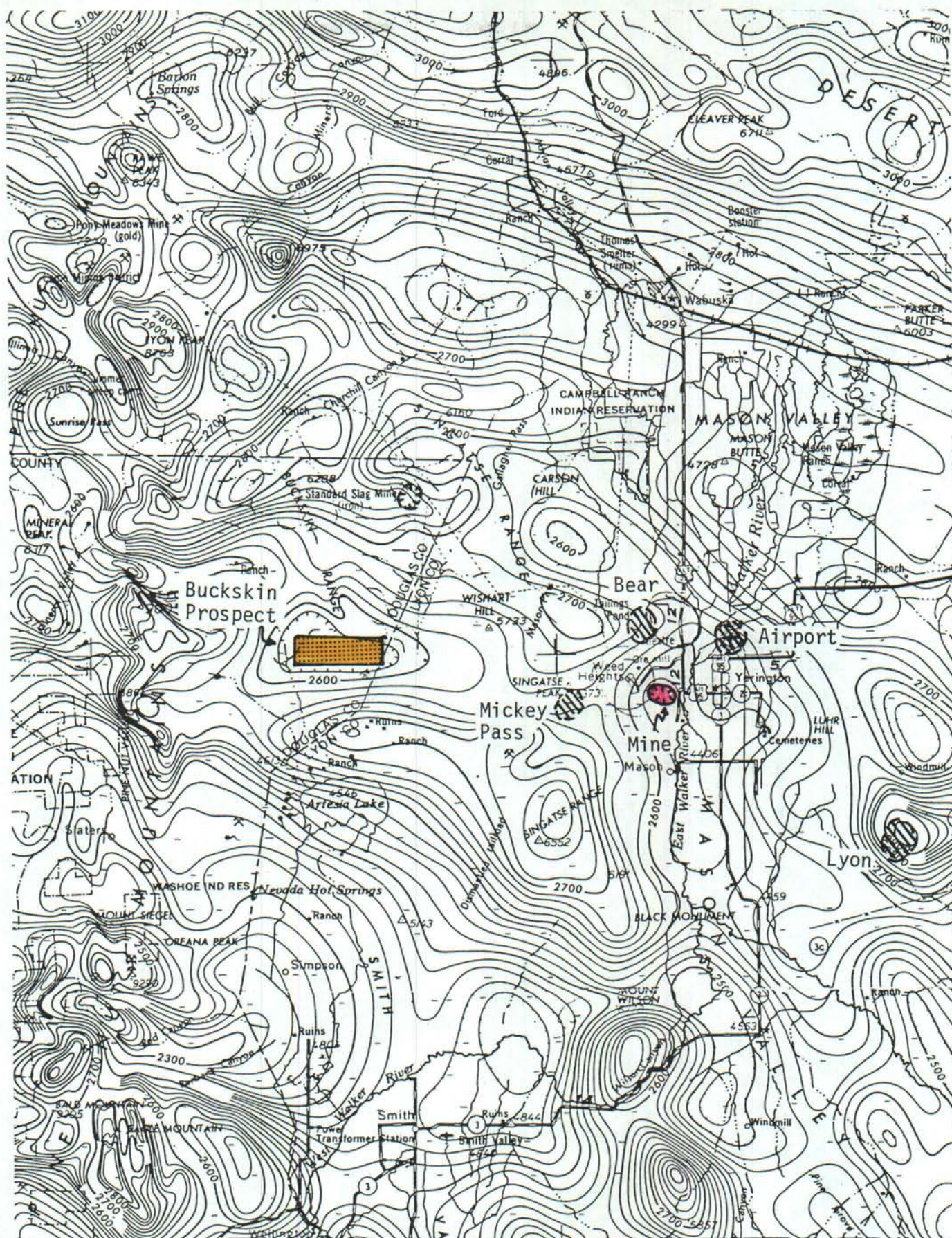


Figure 20









REGIONAL AIR MAG SURVEY  
BUCKSKIN REGION  
Scale = 1:250,000

Figure 22.

Taken from G.P. 751; USGS, 1971



## DIAMOND DRILL RESULTS

### Western Drilling

By far the best copper assays are in drill hole BSW-7 near the west side of the prospect (Figure 25). An interval of 510 feet averaged 0.16% copper from 750 to 1260 feet and 88 feet within that interval averaged 0.35% copper. Fine grained chalcopyrite is disseminated in faulted and broken metavolcanics. One small dike is within the ore zone, but most of the dikes are just below the high copper interval. Secondary biotite and sericite accompany the sulfide mineralization, but it is not my opinion or that of Bear Creek that the "core" or "potassic zone" of a porphyry system has been drilled here. Instead, the hole appears to be above the potential target. A small zinc anomaly (110 feet of 0.145% zinc) was drilled high above the copper interval. Although no core is now available from this specific interval, sphalerite accounts for the zinc elsewhere at Buckskin.

More than 0.1% copper was encountered in hole BSW-1 in two intervals from about 350 feet to 450 feet and 700 feet to 900 feet. The rest of the hole contained about 500 ppm copper. Lead, zinc, and silver were not as abundant in BSW-1 as in other western and central holes. Hole BSW-9 was drilled 1000 feet southwest of BSW-7 to follow up on the high grade interval in BSW-7. Bear Creek drilled there because they believed the mineralization in BSW-7 was related to a west dipping "intrusion breccia." Conoco's mapping indicates that this breccia is a volcaniclastic breccia that is conformable to the rest of the volcanics and only fortuitously acted as a good host for mineralization in BSW-7. BSW-9 did contain a



small 150 foot interval of about 0.1% copper in a fine biotite hornblende porphyry (Figure 29). A zinc anomaly similar to that in hole BSW-7 is also present high above this copper zone.

Bear Creek encountered a chalcocite enrichment blanket and holes BSW-8 and BSW-2 were drilled specifically to test for it. The best enrichment was about 30 feet of 0.64% copper in BSW-2, although holes BSW-1, 7, and 8 also encountered some small enrichment.

#### Central Area Drilling

Conoco's diamond drill hole BC-1 was completed December 20, 1976 at a vertical depth of 1994 feet. It was located in the northerly part of the central geochemical anomaly and in the center of the moderate IP anomaly (Figure 20). It intercepted anomalous but weak copper and molybdenum mineralization but had a very strong zone of lead and zinc from 70 to 1100 feet. Lead and zinc increase again from 1650 feet to the bottom.

	<u>Copper</u>	<u>Zinc</u>	<u>Lead</u>
200 feet - 1100 feet	202 ppm	1416 ppm	830 ppm
1100 feet - 1600 feet	262 ppm	284 ppm	143 ppm
1600 feet - 1994 feet	228 ppm	509 ppm	273 ppm

Sphalerite and galena are finely disseminated in the porphyry dikes but are almost entirely on fractures in the metavolcanics. Porphyry sills are near strong lead-zinc mineralization in most cases and are probably the local source. Sericite, clay veins, and weak epidote and chlorite are associated with the mineralization in the upper 600 feet of the hole, whereas in the lower portion of the hole the chlorite and epidote are more abundant and some potassium feldspar veinlets occur. Several large gouge



zones or faults are present high in the hole, but mineralization tends to overlap and generally disregard them. The hole appears to have intersected an outer lead-zinc halo in the propylitic zone, high and probably off to the side of the core of the "porphyry" system.

Mr. Robert Bamford, while working for Kennecott, recently conducted analyses of sulfide concentrates from 100 foot composite pulps from drill hole BC-1. His method involves analyzing for about 15 elements in a total sulfide concentrate. Kennecott's research indicates that the elemental distribution within the sulfides, combined with X-ray diffraction identification of alteration minerals, tends to reflect the hypogene zoning pattern of a porphyry copper deposit much better than whole rock analyses. The elements were chosen by Mr. Bamford from comparison data he developed by studying several porphyry systems in widely varying environments over a period of several years. His specific models are confidential information owned by Kennecott but the principles and patterns he developed are apparently his to use privately in a consulting capacity.

His results are presented in Figure 24 and Plate 16. He believes the data indicate that the hole was drilled toward the inner zoning of a porphyry sulfide system. This conclusion is drawn from the increasing gradients of bismuth, tellurium, tin and possible arsenic; and the decreasing gradients of lead, zinc and possibly silver (Figure 24). Copper is fairly uniform but because molybdenum seems to decrease downward rather than increasing, he believes we are not directly over the top of the target area but are off to one side. His X-ray diffraction analyses indicate a similar trend as shown by the presence of potassium feldspar near the bottom of the hole in an otherwise typically propylitic assemblage.



From our geologic mapping, he believes we could be skimming the top of a tilted system or obliquely penetrating the side of a nearly upright, but possible deep, sulfide system. The alteration pattern of the core at the bottom of hole BC-1 and the detailed geological mapping have indicated to me that the upright model is the more likely, given our present knowledge. Although future drilling plans at Buckskin may not change significantly, it was valuable work performed for us by Kennecott and the method could be very useful in the future because very little is known about the tops of porphyry systems and almost nothing is published. It is inherently a good idea and reasonable that Kennecott, with its large inventory of drilled porphyry copper systems, should have developed it first.

The results from BC-1 and the patterns of sulfides in holes BSW-7 and 9 indicate strong target potential in the area between BSW-7 and BC-1. The area east of BC-1 cannot be ruled out as a target area, however, because of our direct lack of knowledge about the system at depth and the complex structure.

#### Eastern Drilling

Phelps Dodge and Anaconda drilled four holes in the east portion of the prospect, all less than 1000 feet deep. Hole Bue-1 (Figure 31) was drilled to 838 feet and averaged about 0.05% copper except for a small supergene enriched zone of 0.1% copper and a 150 foot interval of chalcopryite that assayed approximately 0.15% copper. Lead and zinc decreased gradually until the last 28 feet where zinc increased significantly. Molybdenum also appeared to be increasing for the last 78 feet.



Bue-2 was drilled by Phelps Dodge in 1967 to a depth of 893 feet. It also encountered a 10 foot interval of enriched copper but otherwise averaged a little less than 0.1% copper in chalcopyrite for the entire hole. Mesozoic(?) quartz monzonites were the predominant host rocks. Pyrite increased to about 6% for the last 100 feet with traces of molybdenum and silver.

Anaconda drilled holes 281 and 289 in 1971 and 1972. Hole 289 was drilled to 958 feet in Mesozoic(?) quartz monzonites intruded by numerous porphyry dikes. Anaconda assayed only certain intervals, unfortunately, and these assays do not show any definite trends. The hole appears to average more than 0.05% copper, however.

Hole 281 was a short hole (446 feet) that averaged about 0.05% copper in Mesozoic metavolcanics.

The holes in the eastern portion are very intriguing. The Mesozoic(?) phaneritic quartz monzonites are uniformly high in copper and two holes show apparent increasing gradients in lead, zinc, and total sulfides. Late Tertiary dikes are well mineralized also and occasionally appear to control locally higher assays. Most of the mineralization in these holes, however, appears to be genetically related to the older quartz monzonites.



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(35)

Item 15

EXPLORATION OF ENHANCED SULFIDE MINERALIZATION  
AT THE BUCKSKIN PORPHYRY COPPER PROSPECT

DOUGLAS & LYON COUNTIES  
NEVADA

by  
William M. Oriel  
March, 1978



35

Item 15

EXPLORATION OF ENHANCED SULFIDE MINERALIZATION  
AT THE BUCKSKIN PORPHYRY COPPER PROSPECT

DOUGLAS & LYON COUNTIES  
NEVADA

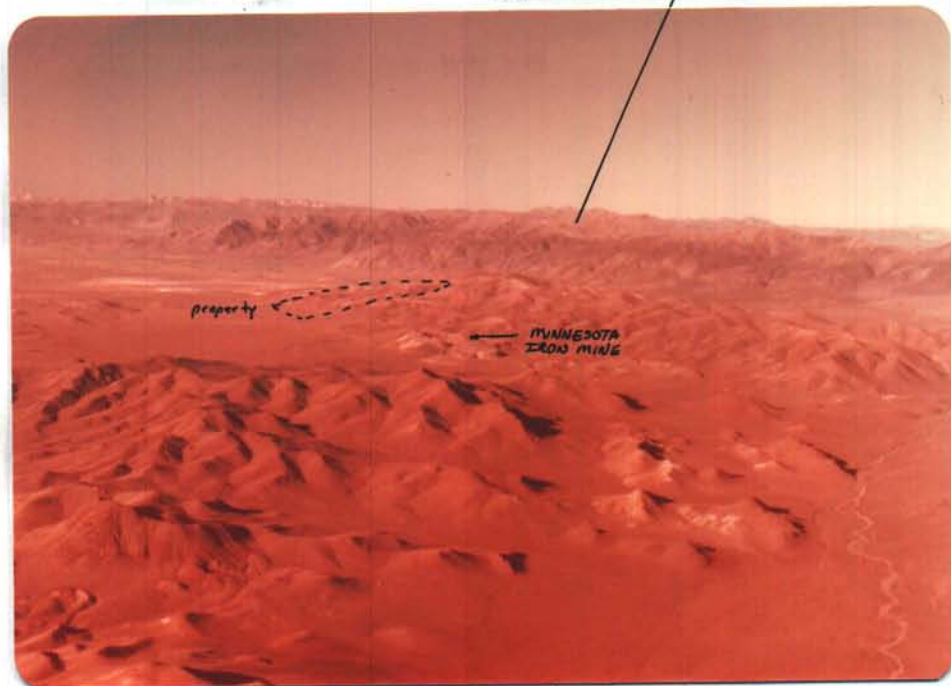
by

William M. Oriel

2nd Annual Report  
Reno District Metals Office  
Reno, Nevada  
March, 1978



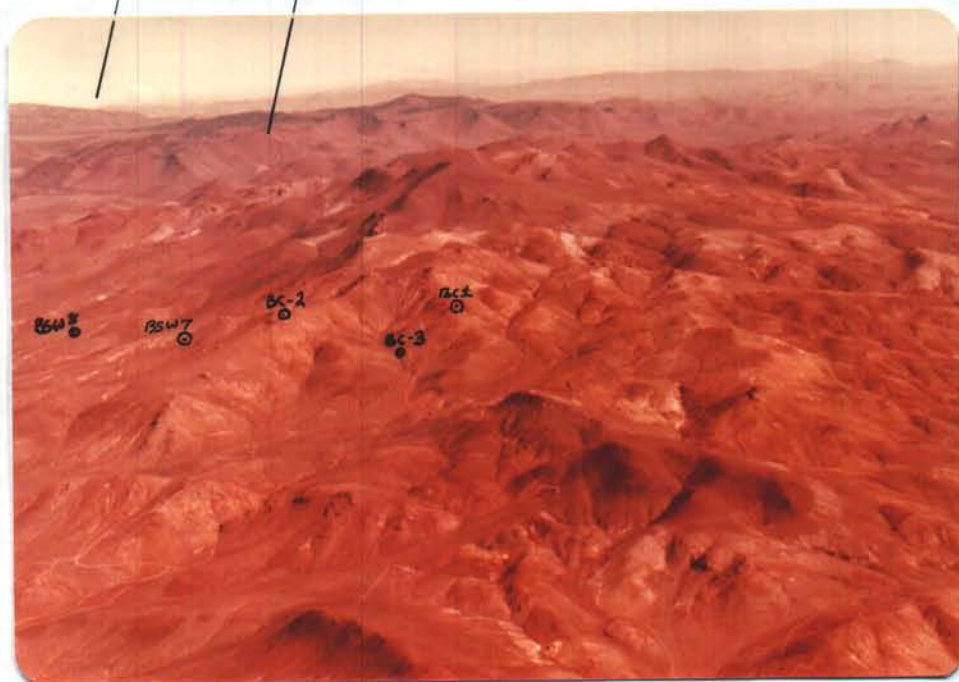
Pine Nut Mountains



Buckskin Range Looking Southwest

Reno

Northern Pine Nuts



Buckskin Range in Foreground - Looking North



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D-3	Lead	" "	" "
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D-5	Silver	" "	" "

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D-7	Moly	" "	" "
D-8	Lead	" "	" "
D-9	Zinc	" "	" "
D-10	Silver	" "	" "

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16	" " BC-3 - Conoco, 1977, 1" = 100'
17	" " BSW-7 - Bear Creek Mining Co.
18	" " BSW-1 - " " " "



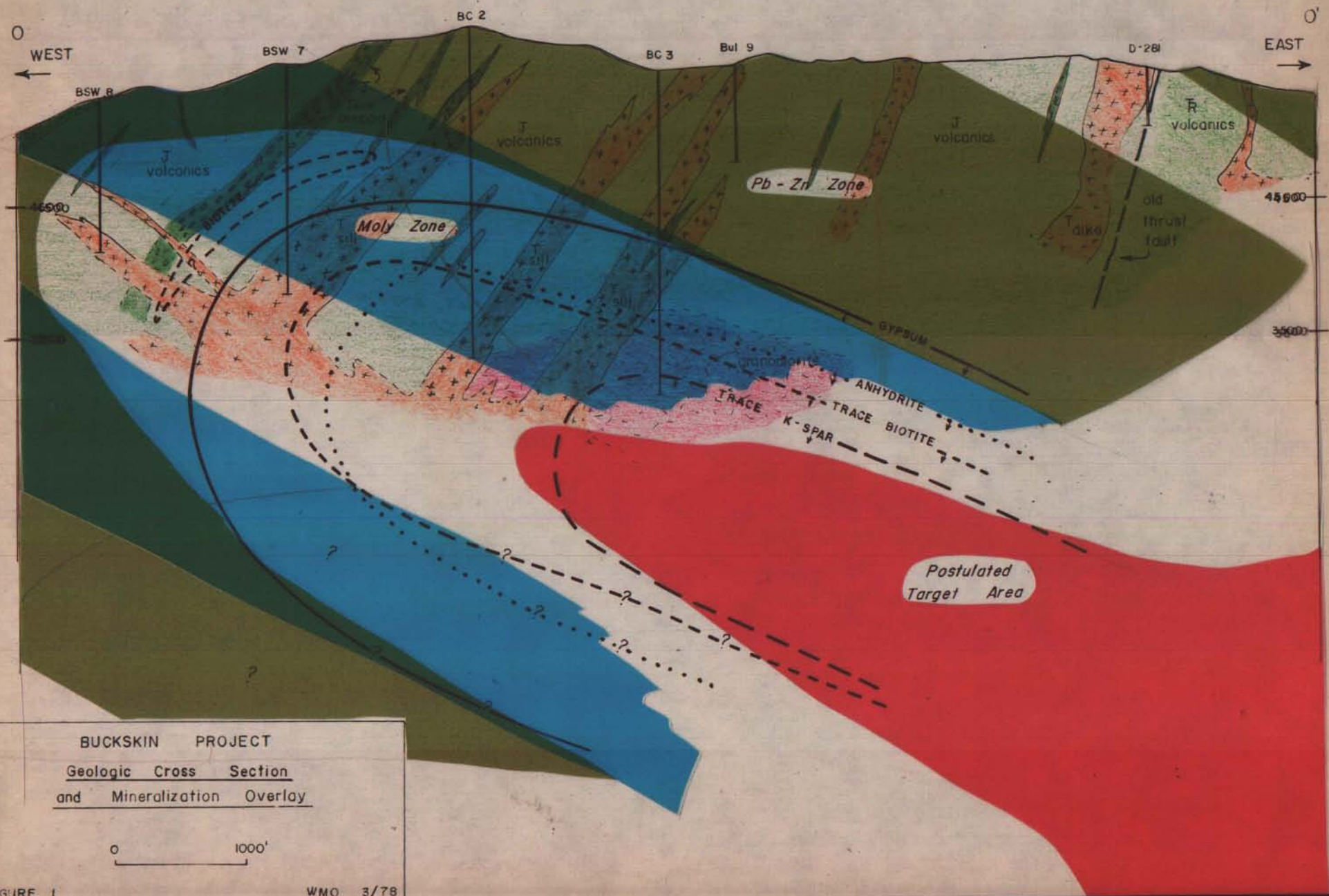
Plates (continued)Plate Number

19	Drill Log - Hole	BSW-2 - Bear Creek Mining Co.
20	" "	BSW-8 - " " " "
21	" "	BSW-9 - " " " "
22	" "	Bul-9 - Phelps Dodge
23	" "	Bue-1 - " "
24	" "	Bue-2 - " "
25	" "	D-281 - Anaconda
26	" "	D-289 - Anaconda

NOTE: The plates listed above may be found in the roll which accompanies this report, or may be requested from the Reno office.



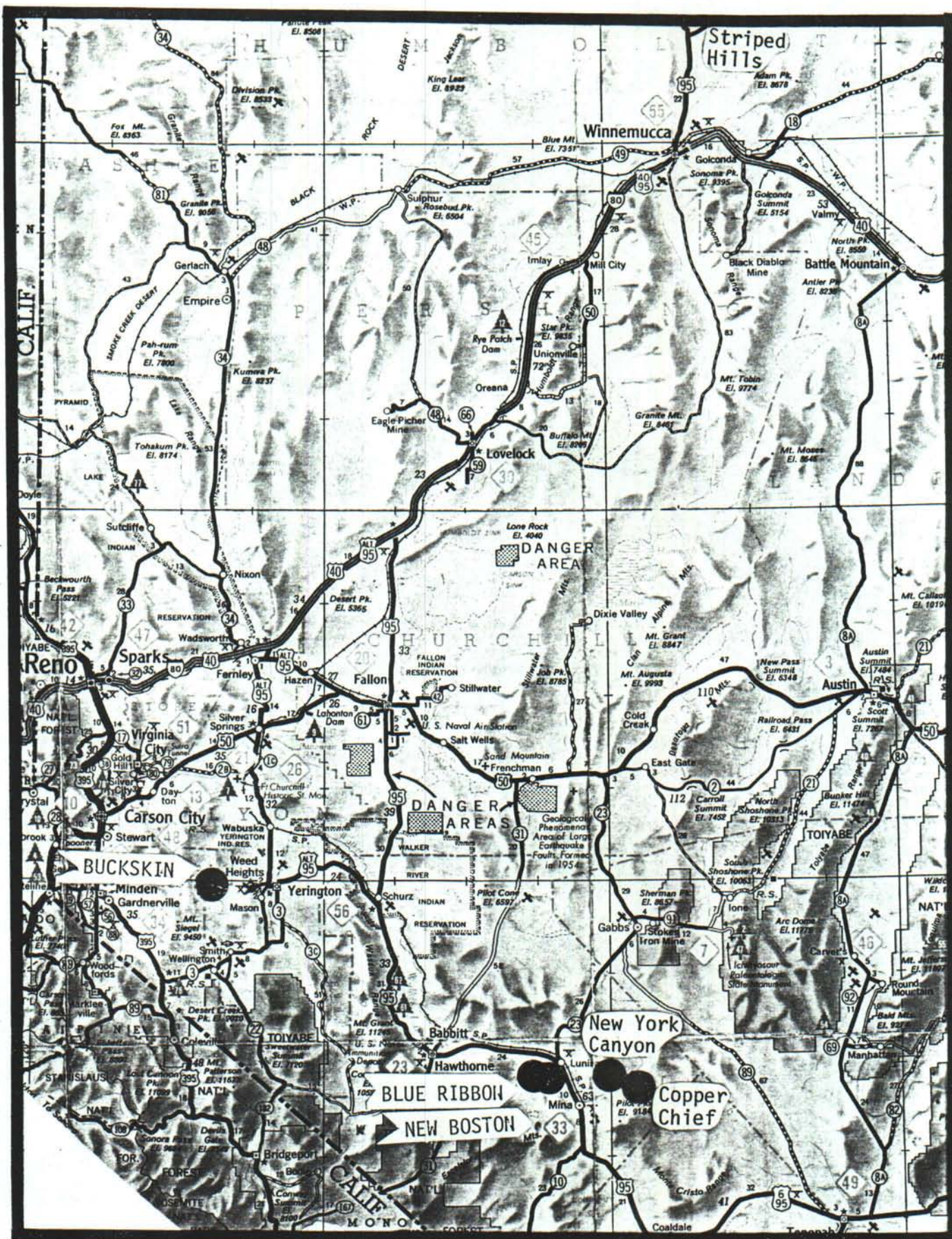
# BUCKSKIN PORPHYRY SYSTEM



BUCKSKIN PROJECT  
Geologic Cross Section  
and Mineralization Overlay

0 1000'





INDEX MAP - NEVADA HIGHWAY



## SUMMARY

The Buckskin project is 95 miles southeast of Reno, Nevada, in the Buckskin Range in Douglas County, Nevada, Township 13 North, Ranges 23 and 24 East. Conoco controls approximately 2400 acres on 128 contiguous unpatented lode mining claims partly through a joint venture with Bear Creek Mining Company.

After securing the properties early in 1976, Conoco conducted an extensive grid geochemical survey that successfully outlined a large hydrothermal system. Following this, all previously drilled holes were geologically logged and thin section studies undertaken. Detailed geologic mapping with emphasis on alteration and limonite maps proceeded in 1976, coincident with 23 line miles of IP and an airborne magnetometer survey. IP data outlined an anomaly 9,000 feet by 12,000 feet, mostly coincident with the geochemical soil anomaly. A separate deep IP anomaly was discovered east of our properties which resulted in the staking of the Conoco claims. Gravity profiles were conducted over this deep eastern anomaly in 1977 to determine the depth to bedrock.

Many altered and mineralized porphyritic and phaneritic dikes, sills, and plugs intrude broken and structurally separate blocks of strongly altered Triassic and Jurassic metavolcanics. The oldest and easternmost intrusives (Mesozoic granodiorites, etc.) have been tilted steeply westward by Basin and Range normal faults, and then were intruded by a swarm of altered and mineralized Tertiary porphyritic dikes and plugs of varying composition. Irregular Tertiary sills and dikes, and Mesozoic



metavolcanics in the central portion of the prospect are altered and leached to clay and sericite with limonite mixtures nearly obliterating rock textures. A series of small dikes and a small intrusive plug, which is mostly covered by alluvium, are present to the west, where Bear Creek has previously conducted some exploration. There is reason to speculate that the Tertiary plugs and sills originate from a common source between the western and eastern areas. They also show chemical and textural variations that indicate a common evolving source. Late post mineral dikes and sills intruded along apparent structural zones with a preferential east-west strike.

The soil geochemical survey conducted in the spring of 1976 showed three areas of interest. The westernmost anomaly is a small copper anomaly where Bear Creek drilled several holes in 1972 and 1973. The central anomaly contains high lead-zinc-silver values in the vicinity of several large sills. The eastern anomaly consists of higher copper, moly, zinc, and lead values overlying a complex swarm of porphyritic dikes which intrude Mesozoic granodiorites and quartz monzonites. It is separated from the central anomaly by Tertiary ash flow tuffs, which have been tilted, overturned, and dropped down to the east. The patterns of alteration and geochemistry derived from 1977 drilling (holes BC-2 and BC-3) indicate that the western anomaly is the uppermost exposed portion of a single tilted system that includes the central geochemical anomaly and possibly the eastern anomaly. The deep eastern IP anomaly is still speculative.

Hole BC-2 was drilled in the central lead-zinc-silver geochemical anomaly to a depth of 1994 feet late in 1976. It encountered strong lead-zinc mineralization in both sericitically and propylitically altered volcanic country rocks, which are cut by a few porphyry sills.



In 1977 holes BC-2 and BC-3 were drilled to 2500 and 2463 feet, respectively, and also intersected strong lead-zinc halo-type mineralization indicative of the fringes of a porphyry copper system.

The geochemical results of Conoco's drilling last summer were recently compiled and alteration in the two holes was studied in thin section. A series of cross sections were drawn to accommodate the new information from holes BC-2 and BC-3. Cross sections containing 100 foot averages of Conoco's analyses of Cu, Mo, Pb, Zn, and Ag show consistent and almost classical patterns for a large porphyry copper system. Lead and zinc form parallel zones of high geochemical values external to a strong molybdenum and a weak but downward increasing copper zone. Additionally, the alteration and mineralogical zoning parallel the geochemical zones in an equally classical pattern. Holes BC-2 and BC-3 show progressive increases in gypsum, anhydrite and secondary biotite toward the bottom of the holes. Secondary orthoclase and albite and anhydrite increase easterly toward and below hole BC-3 (Figure 1).

All Conoco data coincide to indicate that holes BC-2 and BC-3 were drilled into the top half of a porphyry copper system dipping 15 to 35 degrees easterly. The copper-rich zone was not entered but only approached by holes BC-2 and BC-3. Whether or not this system is strong enough to become an economic success now or in the future can only be determined by additional drilling.

The Reno office contracted Mr. Robert Bamford (formerly of Kennecott Research) to conduct a surface limonite analysis program and a sulfide concentrate analysis technique of our drill core. His technique is inherently correct and his data-base of 12 known and drilled Kennecott porphyry systems in various environments could be a valuable addition to



Conoco's exploration expertise. Buckskin provided a good opportunity to test his techniques because Conoco had good geologic and geochemical control of the system. We are still awaiting the geochemical analyses and zoning evaluation by Bamford on the additional holes at Buckskin. The analyses he has performed to date support Conoco's geochemical patterns and will likely provide additional supporting evidence and hopefully some clues about his geochemical zoning data base for other porphyry systems. His data will be integrated with all other geological and geophysical data.



## RECOMMENDATIONS

It is recommended that Conoco complete contract obligations to Bear Creek Mining Company by extending hole BC-2 another 1500 feet. Conoco should also drill two additional holes into the proposed target (Figure 1) eastward of hole BC-3 to depths approximating 4000 feet.

The eastern portion of the prospect should not be tested before 1978 drilling has been completed and evaluated. Ultimately, the eastern portion of the prospect should have at least two deep holes collared on bedrock and one hole in alluvium over the deep IP response before surrendering the property.



## INTRODUCTION

### Location

The Buckskin district is in Douglas County in west central Nevada, approximately 95 road miles southeast of Reno and 10 miles west of Yerington (Figure 2). The prospect is in Township 13 North, Ranges 23 and 24 East, in the south central portion of the Buckskin Range. It is easily accessible by dirt roads that are periodically graded throughout the year. The climate is arid and vegetation is sparse. The range has relatively low relief and a maximum elevation of 6915 feet.

### Property and Ownership

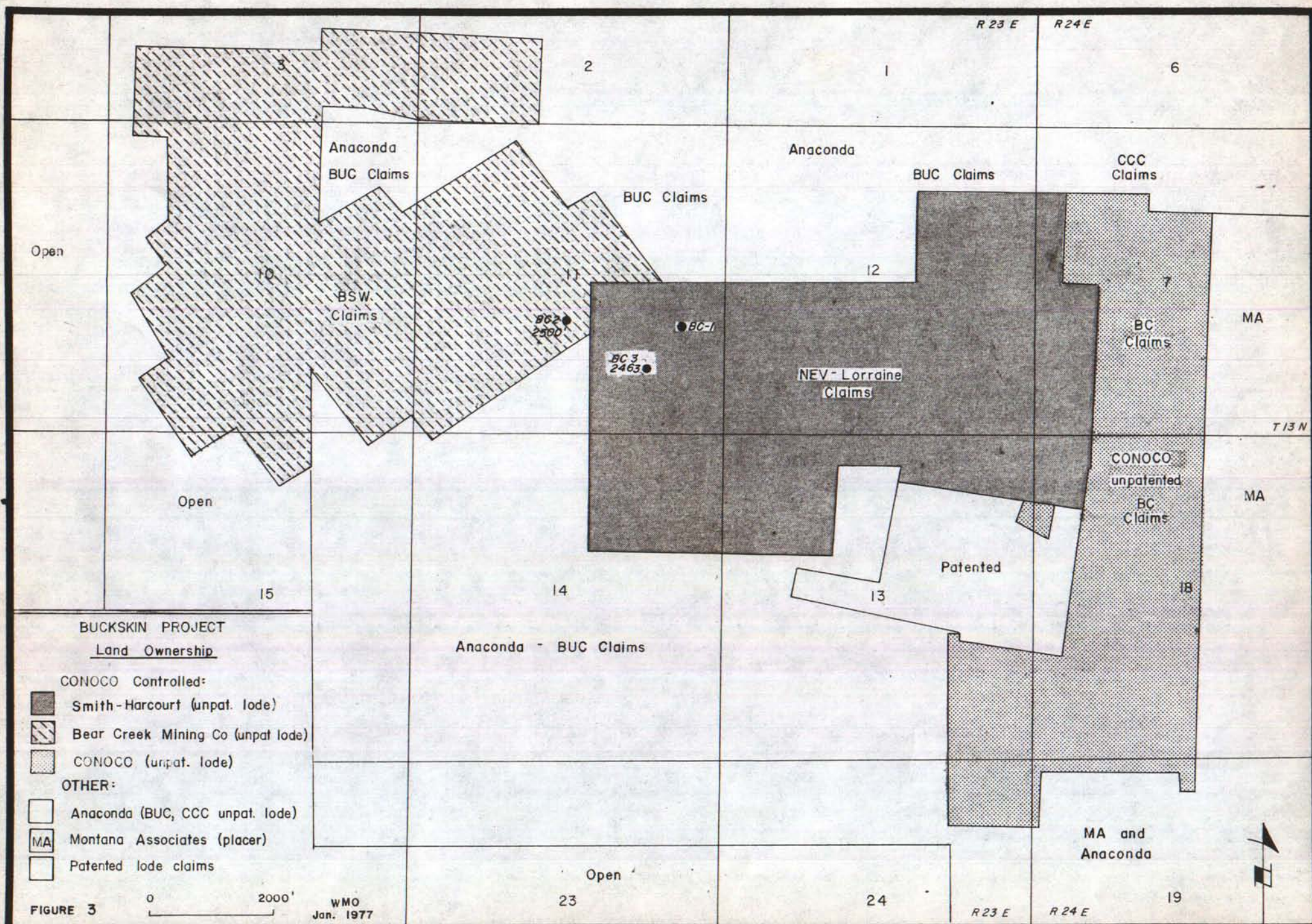
Conoco controls 128 contiguous unpatented lode claims containing approximately 2400 acres spanning the Buckskin Range from east to west (Figure 3). Forty-seven (47) claims are controlled through a joint venture with Bear Creek Mining Co. (BCMC) and 43 are under option from two local residents, Smith and Harcourt. Conoco staked 38 unpatented claims on open ground at the eastern edge of the range as a result of IP data and permissive geology.

Anaconda controls unpatented claims on the north and south margins of our claim block and Montana Associates, a small group of local investors, controls unpatented placer claims adjoining Conoco on the east.

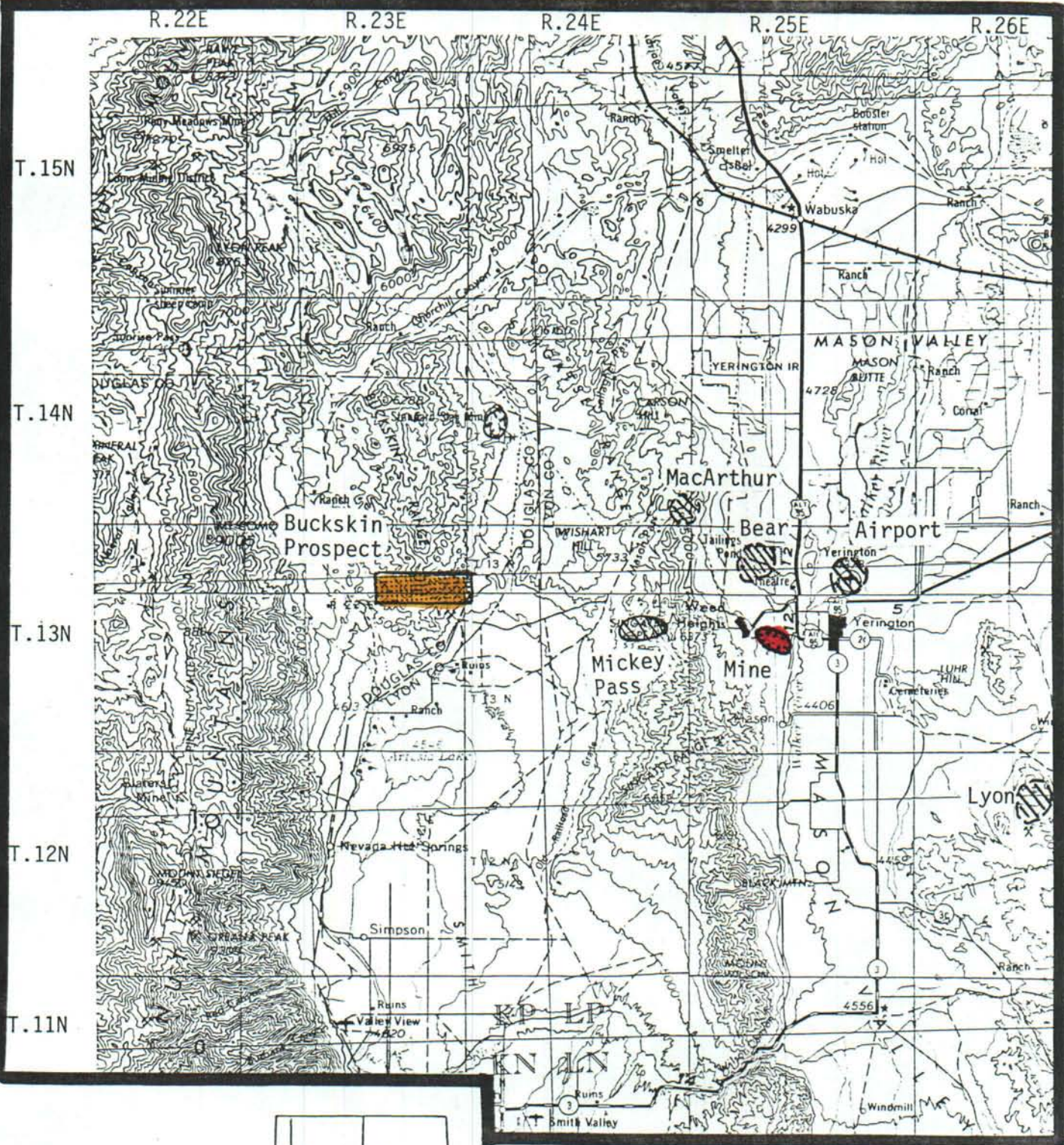
All claims are on public domain lands controlled by the Bureau of Land Management.

A small group of patented claims surround the Buckskin mine at the southeastern portion of our claim group. They are not of interest to









INDEX MAP - YERINGTON AREA  
TOPOGRAPHY

Scale 1:250,000

Figure 4



Conoco at this time; however, should an economic ore body be discovered, the claims would have to be acquired.

#### Holding Costs and Obligations

Bear Creek Joint Venture. Conoco must have spend \$200,000 by December 13, 1981 to earn 51% interest in Bear Creek's 47 claims. Conoco also must drill 1500 feet yearly upon the BCMC claims (BSW group) until the 51% has been earned. The cost of drilling the obligatory 1500 feet in 1978 and other geologic work should be sufficient to earn Conoco the 51% interest before year end.

Conoco has spent \$171,996.02 through 1977, applicable toward the \$200,000 buy-in.

Smith-Harcourt. An expenditure of \$500,000 in annual payments and/or royalty payments to the owners before 1992 will purchase 100% of the 43 "Nev-Lorraine" claims. Total payments and expenditures to Smith and Harcourt through 1977 were \$172,261.83, <sup>of \$55,300.12 is</sup> which ~~is~~ applicable to the \$500,000 end price.

A cost summary sheet is included as Appendix A at the end of this report.

Conoco is responsible for assessment obligations for the Bear Creek and Smith-Harcourt claims as well as our own 38 unpatented "BC" claims on which no assesment work is required until September 1, 1978. No obligations or special requirements exist, other than those imposed upon us by the BLM.

#### History and Previous Work

According to some reports, the Buckskin district was discovered in 1904 by a prospector who found gold in veins on or near the present



"Red Top" patented claims (Plate 2A). Subsequently, copper was found below the oxidized gold zones and considerable excitement arose. This resulted in the building of a townsite and the construction of pumping and metallurgical plants. Very little ore was produced, and most activity ceased in 1907 except for the floating of large amounts of stock. The district was rediscovered and re-evaluated in the early 1930's, and as late as 1974 a head frame and mill building were erected. Drilling, however, apparently found no ore and the site is now abandoned.

Pyrophyllite exploration at the Blue Danube Mine in the southwest portion of the range indicated that the pyrophyllite was impure and thus no commercial development occurred.

A more successful mining operation was carried out on combined iron-copper ore in the northern part of the Buckskin Range. Iron ore was mined from the Minnesota open pit mine from 1944 to 1966, producing nearly \$17,000,000, which is 99.5% of the total values of the district. Exploration and drilling have been carried out sporadically in the district since 1966 and Anaconda currently holds claim to much of the north-eastern part of the Buckskin Range. The only published geologic mapping was done by Moore (1969) and recently by Proffett (1977). Both Bear Creek Mining Co. and Anaconda mapped portions of the project area, and although Conoco has these maps, the scale and detail were inadequate, and the geology possessed gross errors and was mostly incorrect. Mapping was also hindered by the intense degree of weathering, alteration, and the complex structural history of the Buckskin Range.

Eight miles east of Buckskin, the Yerington, or Weed Heights, open-pit copper mine has operated since 1952, but is due to close in June, 1978. It is a porphyry copper mine which contained about 200 million tons of



0.6%± copper. Ore is associated with Mesozoic intrusives. Other porphyry copper deposits near Yerington include the Mickey Pass (Ann Mason), Bear, Airport, and Lyon (Pumpkin Hollow) (Figures 4 and 5). Mr. R. Babcock (V.P. - BCMC) has recently stated that 30 million tons of 2% copper were drilled at the Lyon deposit.

#### Recent Exploration

Exploration for porphyry copper ore at the Buckskin prospect apparently began in 1966 or 1967 when Phelps Dodge acquired large acreage, including the "Nev" claims, in the eastern part of the property. They drilled holes Bue-1 and Bue-2 to depths of less than 1000 feet on the east flank of the prospect, and Bul-9 to 734 feet in the central portion of the prospect. Anaconda had control of the "Nev" claim group until the autumn of 1975. They drilled two holes within our claim group (DH-281 and DH-289) and several holes bordering our property about which we have no information. Anaconda dropped the property in late 1975 before a large land payment was due.

In 1969, Bear Creek Mining Co. staked claims on the west flank of the range around a turquoise showing. They mapped the prospect and took some rock chip samples which were geochemically anomalous. Nine holes were drilled by BCMC, five of which were diamond drill holes centered around their geochemical anomalies, and four others were shallow rotary holes drilled in 1970 in the alluvium to the west. Two of their deep holes, BSW-1 and BSW-7, intersected anomalous mineralization and alteration. Because of their geologic interpretation, they tried to follow the mineralization to the west with drill holes BSW-8 and BSW-9. They never offset their best hole (BSW-7) to the east.



**Oxide and Sulphide Ore Mined, Grade of Ore, Copper Production,  
Waste Stripped, and Recovery of Copper at Anaconda Company's  
Yerington Mine from 1970 through First Six Months of 1975**

Year	Ore Processed, Tons	Percent Copper	Recovery, Percent	Waste Stripped, Tons	Total Copper Production, Tons
<b>SULPHIDE</b>					
1970	4,347,000	0.60	85.47	10,802,045	
1971	4,738,000	0.58	83.58	11,774,805	
1972	4,777,000	0.53	84.09	14,318,043	
1973	4,885,000	0.53	85.65	8,274,550	
1974	4,815,000	0.53	82.89	7,724,803	20,693 <sup>4</sup> (37,969) <sup>1</sup>
1975 <sup>2</sup>	2,403,000	0.52	78.41	4,000,000 <sup>3</sup>	
<b>OXIDE</b>					
1970	4,282,000	0.52	77.72		42,437 <sup>1</sup>
1971	4,514,000	0.51	79.51		42,541 <sup>1</sup>
1972	4,578,000	0.51	77.60		41,218 <sup>1</sup>
1973	5,673,000	0.36	72.83		35,835 <sup>1</sup>
1974	5,887,000	0.38	69.36		17,276 <sup>4</sup> (37,969) <sup>1</sup>
1975 <sup>2</sup>	2,995,000	0.37	67.86		18,997 <sup>1</sup>

1. Both oxide and sulphide ores. 2. First six months 1975. 3. Estimated. 4. First six months 1974

## Yerington district will be a major future porphyry copper producer

It's hard to take a choice.

Many call it the Tucson of the west. And there are many more that call it the Highland Valley of the west.

Yes, that's the Yerington, Nevada porphyry copper district.

There is only one mine in production now, Anaconda Company's Yerington pit and sulphide flotation mill at Weed Heights, which has been operating since 1954. It started out as an oxide mine, leaching plant, and tin can precipitation plant. The flotation mill was added as mining deepened into the primary sulphides. Recent production at the Yerington mine is shown in accompanying table. Total production has been over 100,000,000 tons, about 75 percent of which has been oxide.

Reserves at the Yerington mine proper as of June 30, 1975 were 27,000,000 tons assaying 0.33 percent acid soluble copper and 30,600,000 tons of sulphide assaying 0.48 percent.

There are five additional known deposits in the Yerington district as follows:

Ann Mason which is 3 miles west of the Yerington mine. The deposit does not crop out and is buried under 300 to

1,500 feet of alluvium and post ore cover. Drilling has indicated 495,000,000 tons of sulphides assaying 0.40 percent copper. Mining will probably have to be by underground methods. Recent drilling has discovered shallower and higher grade copper which could be mined by open pitting at a higher price for copper.

Airport does not crop out and was found by Phelps Dodge Corporation. Very little data has been published on this deposit. (500,000,000)

The Bear is 2 miles north of the Yerington mine, does not crop out, and is deeply buried. Deep drilling, only a few holes, indicates about 500,000,000 tons assaying 0.40 percent copper with molybdenum and precious metal contents. Much additional drilling and geological study must be made to define and sample this deposit.

Lyons, about 10 miles east of the Yerington mine, was discovered by the United States Steel Corporation using airborne geophysics which picked up an anomaly due to its magnetite content. This was another deposit which did not crop out. The Anaconda Company has been drilling under an agreement with United States Steel. Drill indicated re-

serves for underground mining are estimated at 25,000,000 tons assaying 0.40 percent total copper.

The MacArthur, four miles northwest of the Yerington pit, contains 13,000,000 tons assaying 0.43 percent total copper. About half the deposit can be economically mined by open pitting. Anaconda has made plans to mine the deposit and truck the ore to the Yerington leaching plant after the Yerington oxide has been mined out.

There's a lot of copper in the Yerington copper district. It is high cost copper and it's typical of where tomorrow's copper production will come from.

Figure 5

Yerington District Deposits  
Taken from World Mining Journal,  
July, 1976



### Present Program

Conoco's involvement is partially the result of the fortuitous availability of both the Bear Creek claims and the "Nev-Lorraine" claims. The Conoco staff recognized the potential of a large complex sulfide system that had been looked at only in bits and pieces previously.

When the property became available in late 1975 and early 1976, we secured our interests with letters of intent and began a large soil geochemical program to outline the anomalous areas in the spring of 1976. Detailed geologic mapping, including limonite and alteration and sulfide mapping, and extensive IP surveys were conducted in the summer and fall of 1976. Conoco began drilling hole BC-1 in September; total depth of the hole was 1994 feet. Holes BC-2 and BC-3 were drilled to 2500 feet and 2463 feet respectively, in 1977. This report is a summary and interpretation of the mapping, geophysics, geochemistry and drilling conducted through 1977.

### Acknowledgements

Mr. P. H. Kirwin was able to quickly acquire the two halves of the property when they both became available in late 1975 after he and Ron Long visited the property. Bear Creek Mining Co. (Spokane Office) and Kennecott Research have been very cooperative and helpful since we acquired the property. Aubrey J. Beck assisted mapping for three weeks on the prospect in 1976, and Don Hudson conducted important reconnaissance geologic mapping around the project during the summer of 1977. Karen Kling and Pat Brioady helped with drafting and design of the plates and figures. The Conoco staff, especially Mr. P. H. Kirwin and Ron Long, were helpful through constructive criticism. Mr. Robert Bamford also provided stimulating discussions and ideas.



## GENERAL GEOLOGY

The Buckskin prospect is in the southern half of the Buckskin Range at the western edge of the Great Basin Province and near the western margin of the Walker Lane structural zone (Figure 6).

Approximately 18,000 feet of stratigraphic section is exposed in the Buckskin Range. Triassic metavolcanic rocks are superseded by carbonate and clastic sediments of the Upper Triassic Oreana Peak Formation and carbonates and sediments of the Lower Jurassic Gardnerville Formation. These are overthrust by andesitic metavolcanic rocks of the Lower Jurassic Artesia sequence and dacitic to quartz latitic metavolcanic rocks of the Lower Jurassic to Lower Cretaceous Churchill Canyon sequence.

Oligocene ash-flow tuffs of the Guild Mine Tuff, Weed Heights, Singatse Tuff, Nine Hill Tuff, and Eureka Canyon Tuff unconformably overlie the Mesozoic rocks. Miocene andesites, Plio-Pleistocene lake sediments, and Quaternary alluvium unconformably overlie the ash-flow tuffs.

The Lower Mesozoic rocks are intruded by Jurassic granodiorite and quartz monzonite. Numerous varieties of Miocene andesite dikes, and mineralized dacitic and latitic porphyry dikes and stocks intrude the older rocks of the Buckskin Range, especially in the prospect area. The porphyry dikes and sills at Buckskin have not been recognized previously for their true age or significance.

Most of the rocks in the Buckskin Range were hydrothermally altered to propylitic, sericitic, or aluminophyllic alteration facies by Miocene alteration events and to a lesser extent by Mesozoic alteration events.



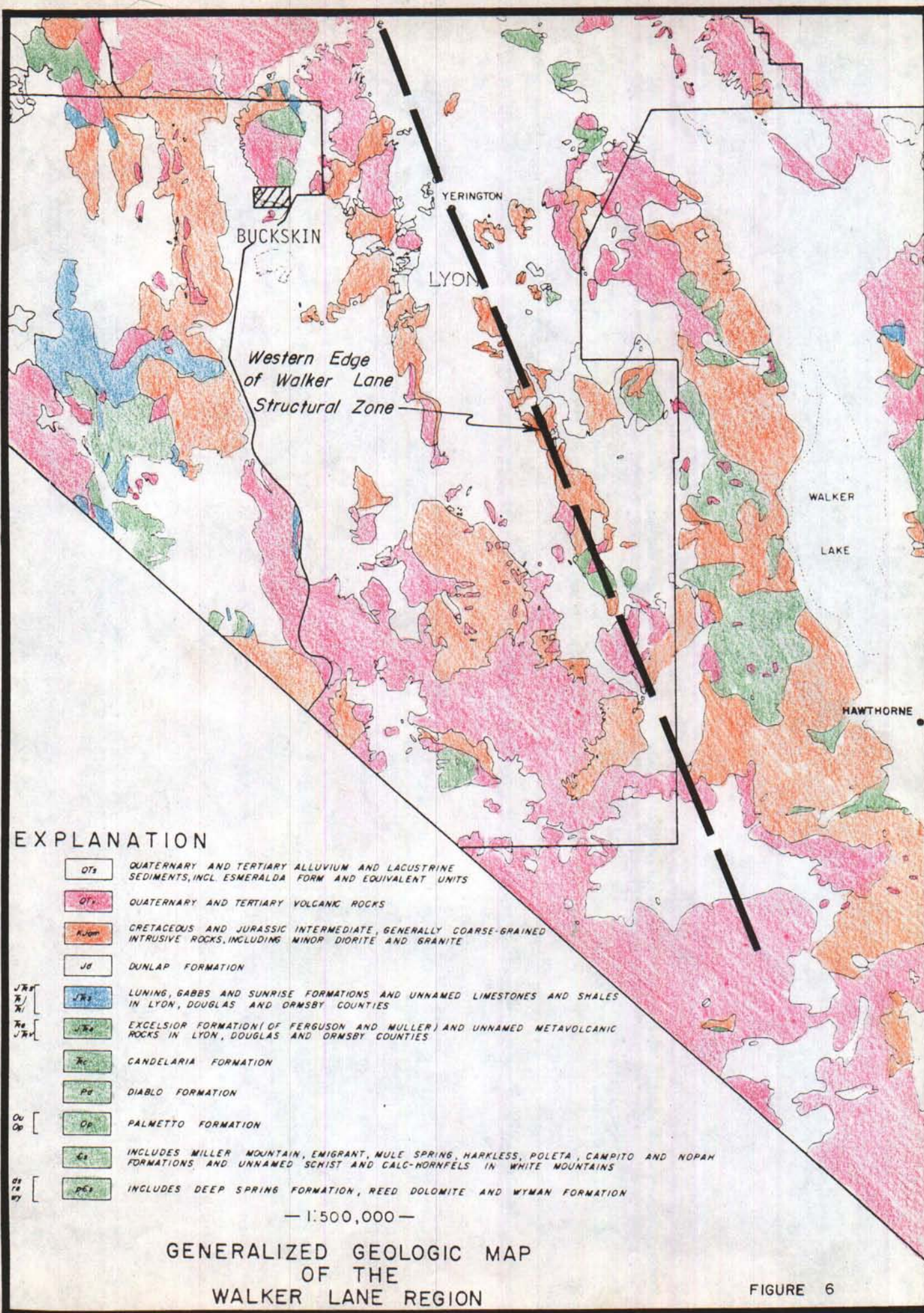


FIGURE 6



Structure is dominated by gently-dipping Basin and Range(?) faults which were originally steeply-dipping normal faults. These faults underlie steeply westward dipping Tertiary ash-flow tuffs and Mesozoic volcanic rocks. The strata were probably tilted westward by the low angle faults and/or by later steeply-dipping normal faults. According to Proffett (1977), westward tilting and extension began 18 m.y. ago and ended between 11 and 7 m.y. ago. The recognition of Tertiary structure is an important consideration in exploring for mineralized intrusives older than 11 million years.

Mesozoic thrust faulting is also documented in the northern Buckskin Range (Hudson, 1977) bringing the Artesia sequence in contact with the Gardnerville Formation, eliminating at least 8000 feet of stratigraphic section from the Buckskin Range. It is possible that a large structural unconformity identified in the eastern portion of the Conoco project area is also a Mesozoic thrust fault bringing Triassic metavolcanics into contact with Artesia group metavolcanics.



## PROSPECT GEOLOGY

STRATIGRAPHY

## MESOZOIC ROCKS

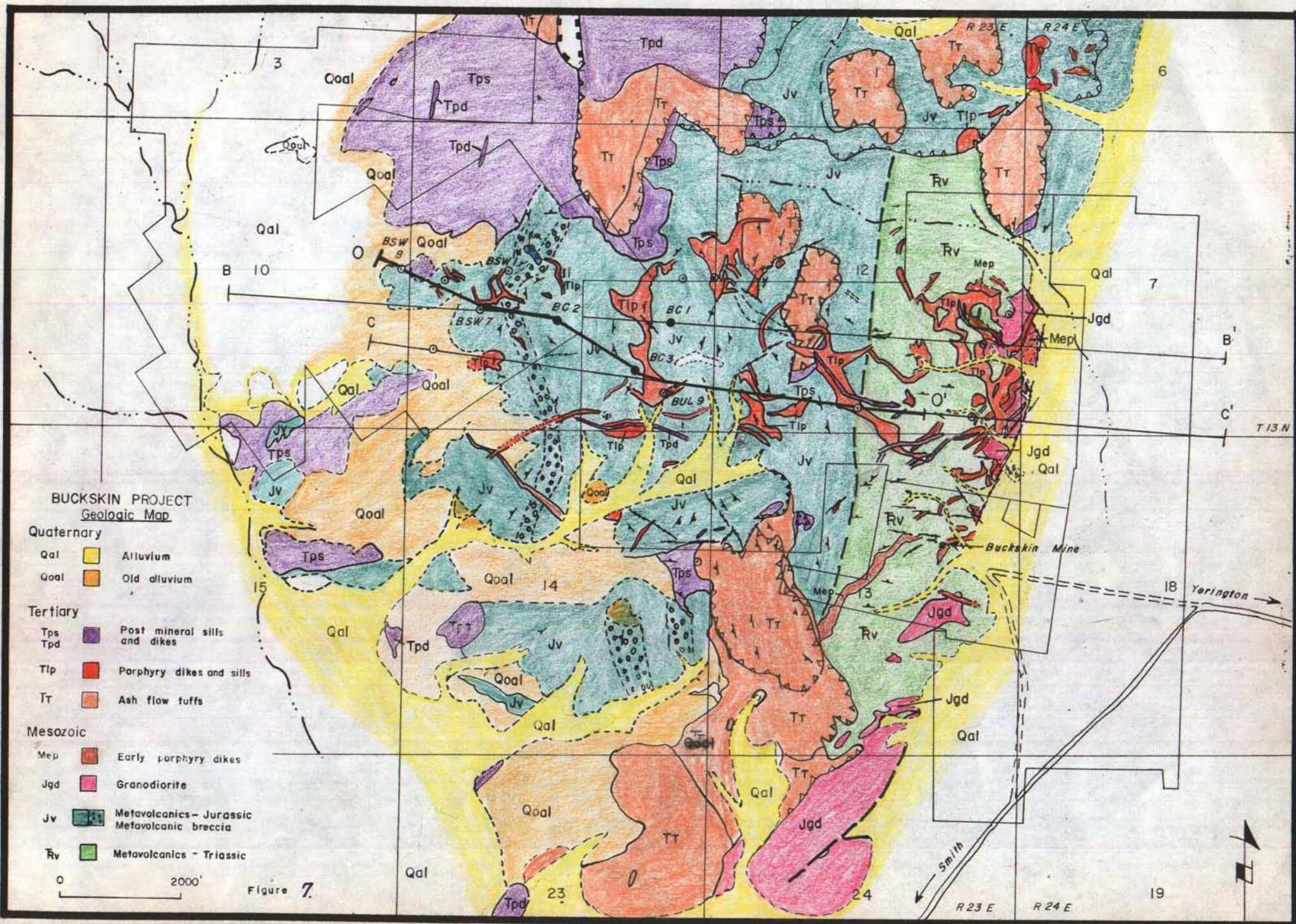
Triassic Metavolcanics (T v)

The oldest rocks exposed in the prospect vicinity as well as in the entire range are Triassic "chicken track" meta-andesites (T v) on the east flank of the prospect (Figures 7 and 8). These metavolcanics are tentatively correlated with a continuous section at nearby McConnel Canyon in the Singatse Range and are believed to underlie the Triassic Oreana Peak limestones. Approximately 1300 feet of felsic flows and angular breccias are missing from the Buckskin area. The andesites have mostly fine-grained, dark green groundmass with small, white feldspars occasionally forming a "chicken track" pattern. These rocks are older than the Lower Jurassic Artesia sequence, which they resemble because they are intruded by granodiorite which is provisionally correlated with the Lower Jurassic granodiorite of Black Mountain (Bingler, oral communication, 1977). Foliations in the lower sequence rocks generally strike east-west and dip to the north. The contact between the Artesia sequence to the west and the Triassic metavolcanic sequence may be part of the Buckskin thrust but insufficient exposure prevents exact determination.

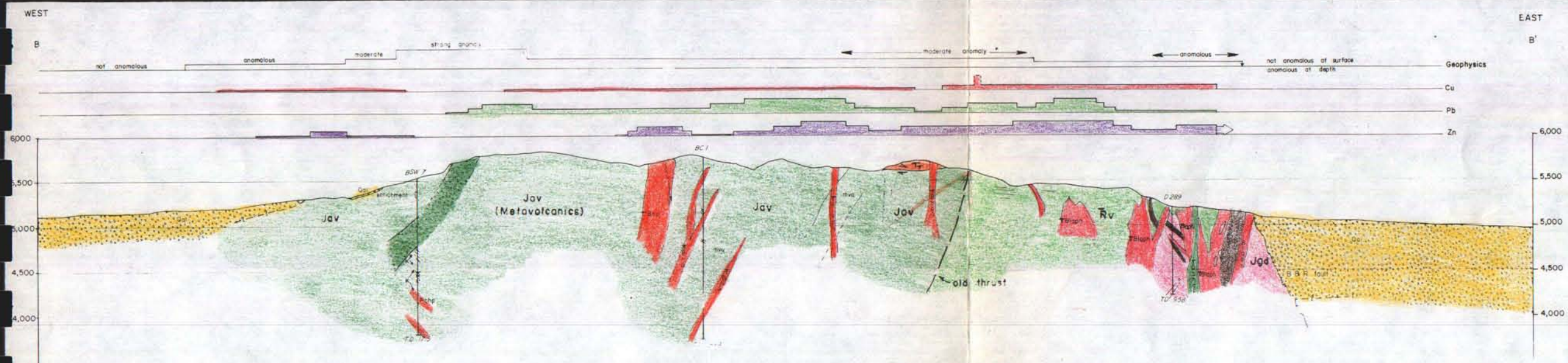
Triassic Sediments

The Triassic Oreana Peak Formation (T op) and Jurassic Gardnerville Formation (Jg) present in the northern part of the range (see Appendix B) are absent in the southern half of the range. The Buckskin thrust in the









BUCKSKIN PROJECT

GEOLOGIC CROSS SECTION B-B'

0 500'

Figure 8

WMO 3/77



northern Buckskin Range eliminates 8000 feet of stratigraphic section above the Gardnerville and may also be responsible for the missing section in the southern Buckskin Range.

#### Artesia Sequence (Jav)

The bulk of Conoco's claims are underlain by Lower Jurassic Artesia metavolcanics, which are up to 6000 feet thick in the southern Buckskin Range. The rocks strike northerly and dip predominately eastward,  $50^{\circ}$  to  $85^{\circ}$ . Artesia metavolcanics lie in thrust contact with underlying rocks where observed in the northern Buckskin Range. There is evidence to support the same conclusions in the prospect area (see Structure). The volcanics are repetitious andesitic to dacitic flows and occasional rhyolitic flow breccias within the prospect area. They are fine grained, foliated and usually weakly porphyritic (see Appendix B). They are also characterized by discontinuous "reefs" of silica which form resistant craggy outcrops. Extreme alteration in the prospect area usually destroys the fabric but in some cases enhances the foliation. Foliations are always parallel to small-bedded sandstone and arkose lenses and flow breccias, and represent primary pseudo-bedding features. Some of the andesitic volcanics are probably semi-concordant intrusive sills, in the volcanic pile, but are difficult to correlate, even over small distances with detailed mapping, because of similar composition, metamorphism, the super-imposition of hydrothermal alteration, and Cenozoic faulting.

#### Churchill Canyon Sequence (Jcc)

The Churchill Canyon sequence, named for exposures just south of Churchill Canyon (Plate 1), conformably overlies the Artesia sequence metavolcanics. In the prospect area, they lie 2000 feet west of Bear Creek's



most westerly diamond drilling and are not encountered in any drill holes. As much as 3000 feet of dacitic to quartz latitic flows, tuffs, and associated porphyritic intrusions is exposed in the Buckskin Range (Hudson, 1977). The exposed portion of the sequence contains primarily ash-flow tuffs in the northernmost Buckskin Range and dominately flows in the southern and central Buckskin Range. Quartz phenocrysts up to 10 mm in diameter are distinctive of the sequence, though not all flows and intrusions contain them (see Appendix B).

Both the Artesia sequence and Churchill Canyon sequence are probably correlative to the Double Springs Formation of Noble (1962). Ammonites collected from the Double Springs Formation near Topaz Lake indicate an uppermost Lower Jurassic age of the Double Springs Formation (F. G. Bonham, oral communication). Age dates on hornblende obtained by Castor (1972) from Churchill Canyon sequence type rocks in the Pine Nut Range yielded dates of 146 m.y. and 124 m.y. The age dates would indicate an Upper Jurassic to Lower Cretaceous age for the Artesia and Churchill Canyon sequences.

All of the preceeding volcanic units are regionally metamorphosed to the greenschist facies containing epidote, sericite, calcite, chlorite, and clays with slight albitization of the plagioclase. The units described were primarily mapped as textural varieties of andesites (Oriel, 1977) until the excellent regional mapping and correlation by Don Hudson in 1977.



## CENOZOIC ROCKS

The Oligocene is represented by a series of ash-flow tuffs and tuffaceous sediments which are thoroughly discussed by Proffett and Proffett (1976) from the Singatse Range. Only brief descriptions of the units immediately in the prospect vicinity are given here to note local variations in the units. More complete descriptions can be found in Appendix B.

Guild Mine Tuff (Tgmt)

About 1000 to 2000 feet of moderately-welded, crystal-rich ash-flow tuff of the Guild Mine Tuff unconformably overlies the Mesozoic rocks in a normal section in the Buckskin Range. It is a light brown to buff crystal-rich tuff consisting of 5 to 25 feet of basal vitrophyre with petrified wood fragments overlain by 200 to 900 feet of pumice fragment rich moderately-welded tuff and 300 feet of vapor phase tuff with rare blocks of red jasper. It is sericitically altered to a white or tan within the prospect area.

Weed Heights Tuff (Twht)

The Weed Heights Tuff conformably overlies the Guild Mine Tuff. It consists of 300 feet of lavender, moderately-welded, crystal-rich tuff topped by 50 to 75 feet of rhyolitic sediments. Abundant large white pumice fragments in the main tuff unit are distinctive.

Singatse Tuff (Tst)

About 2500 feet of brown to red-brown, locally lavender, strongly-to-moderately-welded, crystal-rich ash-flow tuff of the Singatse Tuff conformably overlies the rhyolitic sediments of the Weed Heights Tuff.



It is distinguished by its massive nature, normally reddish tint and the presence of sparse hornblende phenocrysts in its crystal-rich matrix.

It has been altered to a brown color within the prospect area.

Other Tertiary volcanics which are listed below are present in the Buckskin Range but not in the prospect area. They are described in Appendix B.

Nine Hill Tuff (Tnht) - 100 feet - overlies Singatse Tuff

(part of Bluestone Mine Tuff of Proffett and Proffett, 1969)

Eureka Canyon Tuff (Tect) - 50 feet - from E. C. Bingler - also part of Bluestone Mine Tuff.

Andesitic Volcanics (Tv) -

#### Lake Sediments (QTs)

At least 250 feet of Pliocene to Pleistocene lacustrine sediments occur on the west flank of the Buckskin Range. These lie with angular unconformity on Miocene to Mesozoic rocks and are tilted as much as 20° westward.

#### Older Alluvium (Qoal)

Fifty or more feet of coarse unsorted, unconsolidated conglomerate of older Quaternary alluvium lie conformably or slightly unconformably on the lake sediments. The conglomerate contains moderately-rounded pebbles to boulders. Most of the clasts are granitic material that is foreign to the Buckskin Range with a lesser amount of locally-derived material.

#### Ferricrete (Qfc)

Small patches of bright orange ferruginous breccia (limonite and manganese cemented alluvium) are developed along fault margins on the east



flank of the southeastern portion of the Buckskin Range. They both overlie and are locally overlain by recent alluvium (Qa1).



## INTRUSIVE ROCKS

## MESOZOIC IGNEOUS ROCKS

Jurassic Granodiorite (Jg)

Several textural varieties of Jurassic granodiorite intrude the Triassic "chicken track" metavolcanic (T<sub>rv</sub>) along the east flank of the prospect. The granodiorite is a fine to medium grained, equigranular, phaneritic intrusion and is the oldest exposed in the Buckskin Range (Jgd on Figures 7 and 9). It has been tentatively correlated with the Jurassic Black Mountain Granodiorite in the Singatse and Wassuk ranges (E. C. Bingler, 1977, oral communication). This correlation is also useful in defining the Triassic age and correlation of the unit.

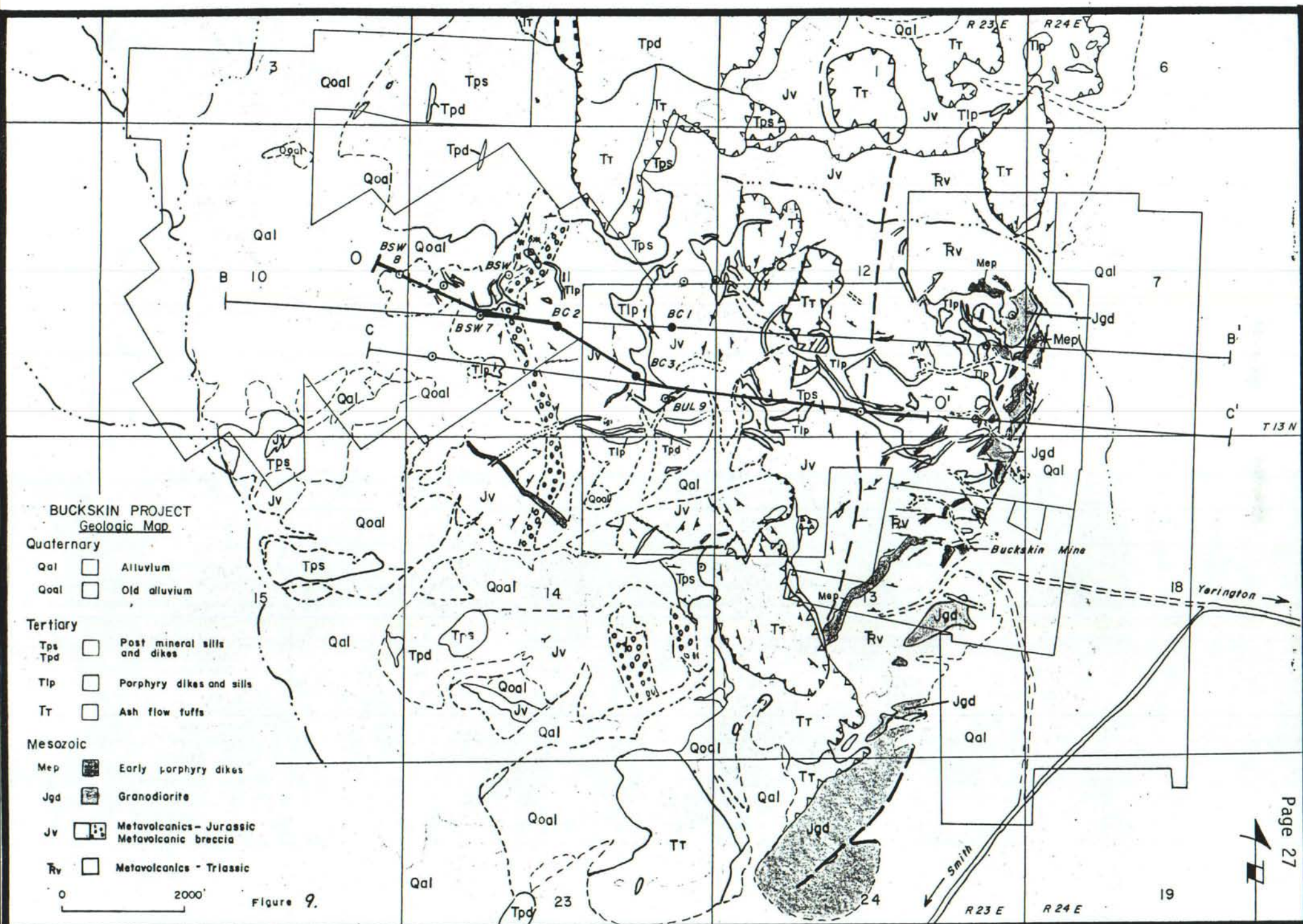
Previous field descriptions and drill log descriptions of these rocks were labeled as varieties of quartz monzonites and granites: i.e. fine quartz monzonite (fqm), speckled quartz monzonite (sqm), fine-grained granite (fg), medium-grained quartz monzonite (mqm), etc. These designations can still be found in early drill logs but should be regarded as textural variations of the Jurassic Granodiorite (Jgd).

The Jurassic granodiorite on Conoco's property appears to form a small discrete stock or plug that has been tilted westward approximately 80° and truncated by recent Basin and Range faulting. Some copper mineralization, discussed in later sections, may be associated with this stock.

Latite Porphyry Dikes (KJlp)

Mesozoic latite porphyry dikes intrude the older Mesozoic rocks throughout the Buckskin Range (Map on Figures 7 and 9). They are







characterized by large subhedral 3 to 7 cm K-feldspar phenocrysts and about 1% quartz phenocrysts in an aphanitic groundmass. Varying amounts of hornblende, biotite and plagioclase occur in the groundmass. Several dikes in the northern Buckskin Range are essentially identical in texture and mineralogy but contain about 5% quartz phenocrysts (Kq1p). The dikes are similar in composition to Churchill Canyon sequence intrusions (see Appendix B) but they are distinctive enough in appearance to map separately. The latite porphyry dikes could be a later phase of Churchill Canyon intrusive activity, based upon intrusive relationships, but may be considerably younger. All other intrusive rocks cut the latite porphyry dikes except the granodiorite mentioned previously.



## TERTIARY INTRUSIVE ROCKS

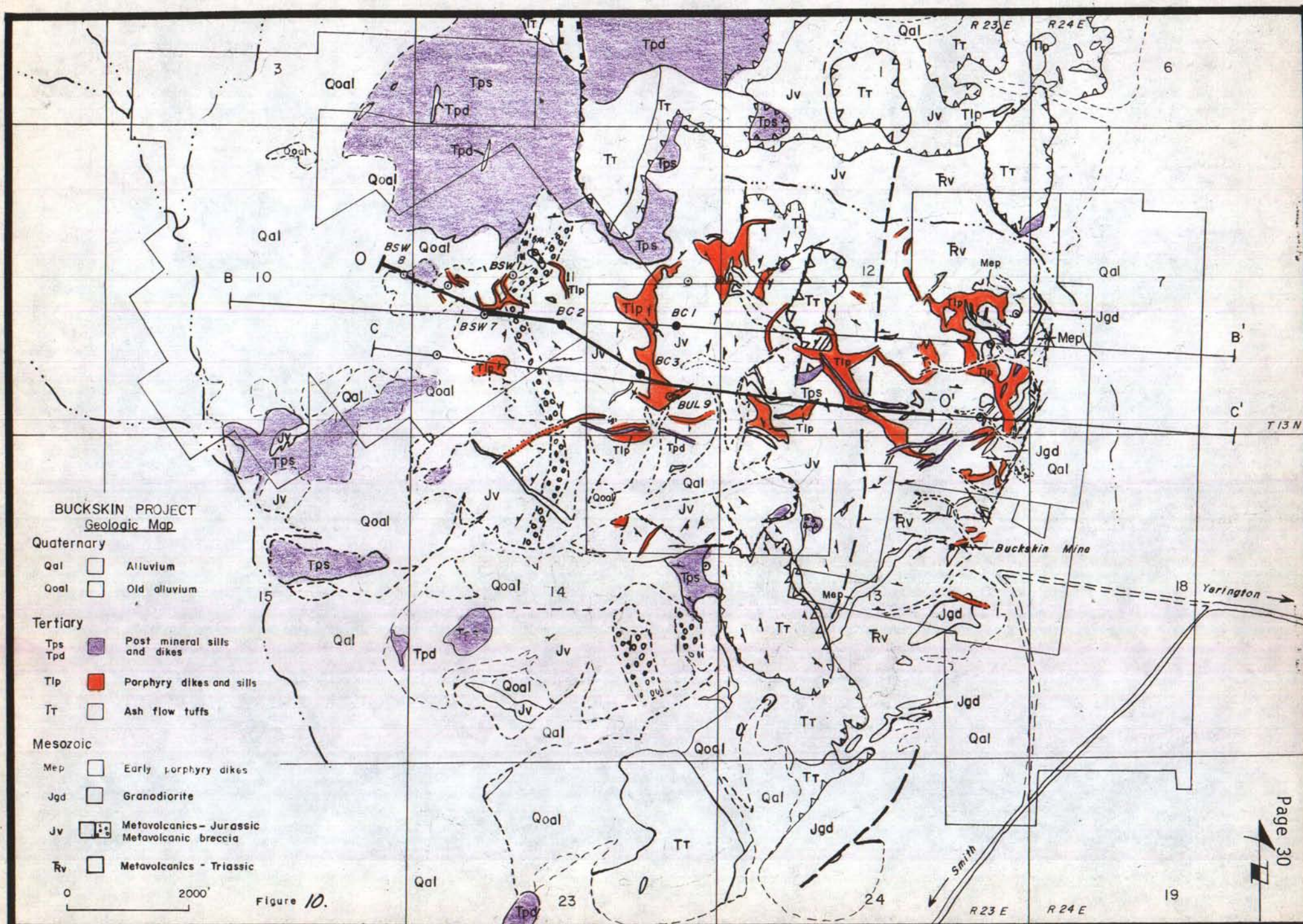
### Mineralized Igneous Rocks

Mineralized porphyritic dikes, plugs, and sills occur throughout the prospect and largely post-date the faulting (Tlp on Figure 10). One mineralized dike (TFbhp) was dated at about 16 million years by Bear Creek Mining Co. Detailed mapping (Plate 2 and Figure 7) shows that some mineralized intrusive rocks intrude late Oligocene to Miocene ash-flow tuffs and a few may cross-cut tilted normal faults (see Structure). They are latitic to dacitic in composition with varying amounts of phenocrystic quartz, hornblende, and biotite. Strong pyritization and sericitic, argillic and propylitic alteration affect the dikes and the country rocks enclosing them (see Alteration).

Some interesting east-west trending compositional and textural variations were recognized through detailed mapping. Closely-related Tertiary dikes are older, contain significantly more phenocrystic quartz and larger, more well developed stacked biotites in the east than in the west. The changes take place gradually over the entire length of the range and are possibly related to an evolving porphyry system subsequently tilted so that it is now nearly horizontal. The bottom of the system is in the east and the top is in the west. Although we are uncertain of the relationship of the parameters mentioned, other geochemical and mineralogical data strongly support a westward-tilted, elongate porphyry copper geochemical and geological system (discussed in Drilling and Geochemistry).

Fine biotite quartz porphyry (Tfbqp) is perhaps the oldest of the mineralized Tertiary porphyries. It occurs only on the east flank of the range as several small dikes up to 100 feet wide and 600 feet long.







In thin section 2-3 mm rounded quartz phenocrysts and biotite books 2 mm thick and 4-5 mm long are set in a felted groundmass. Microlites of plagioclase and K-feldspar are normally 1/2 to 1 mm long. Hornblende ( $\pm 0.2$  mm) is widespread in the groundmass, and sphene and apatite are common accessories.

Biotite-hornblende porphyry (Tbhp) occurs as small dikes and irregular sills up to 2500 feet long and 600 feet wide in the central area of the prospect. Stacked biotite phenocrysts 3-5 mm wide characterize this rock. Hornblende rarely is more abundant than biotite. Feldspars are as large as 5 mm but are visually subdued by alteration effects. Quartz phenocrysts are normally less than 1% of the rock, but gradational contacts with identical rocks containing as much as 5% quartz phenocrysts (Tbhqp) have been mapped on the surface.

Fine-grained biotite-hornblende porphyry (Fbhp) is compositionally very similar to biotite-hornblende porphyry (bhp) and may be texturally gradational into it. It can be distinguished by the much smaller biotite phenocrysts that are distinctly less stacked. These mineralized dikes occur exclusively in the west portion of the district. They have been dated by K-Ar methods at  $15.7 \pm 1.5$  million years in hole BSW-7 (Nielsen, BCMC, 1975).

Biotite latite quartz-porphyry with hornblende (Tblqph) is a common rock in the eastern area and is probably one of the youngest of the mineralized porphyries. Large, rounded and embayed quartz phenocrysts 3-6 mm long, stacked biotite phenocrysts 3-5 mm high, and large plagioclase and feldspar phenocrysts 4-5 mm long are characteristic for field identification. Quartz phenocrysts are not common; one or two in a fist-sized specimen is typical. This porphyry normally contains greater than 1% pyrite and is strongly altered to sericite, calcite and chlorite.



Fine-foliated hornblende quartz dacite porphyry (Tfhqdp) occurs as small, east-west trending dikes in the central and eastern parts of the property. These dikes are very young and cut most of the mineralized porphyries except Tblqph with which it doesn't come in contact. These rocks (Tfhqdp) are very dark and fresh looking on the surface, and at first they were thought to be post-mineral. Thin sections reveal very fine disseminations of up to 5% pyrite and greater alteration than previously recognized.

Quartz biotite hornblende dacite porphyry (Tqbhdp) forms dark dikes that are weakly chloritized and mineralized with pyrite. These are among the youngest dikes mapped.

Other Tertiary dikes occur within the prospect and elsewhere in the Buckskin Range. They are of minor significance, however, and their descriptions can be found in Appendix B or Hudson's reconnaissance report (1977).

### Breccias

Breccias are limited to the western side of the prospect. A small, pipe-like tectonic breccia containing porphyry and volcanic fragments in a rock flour matrix crops out about 400 feet northeast of hole BSW-1 (Plate 2). Float mapping indicated its dimensions to be approximately 100 feet by 400 feet. A small pebble dike containing rounded silicified fragments of metavolcanics crops out east of the breccia. It is about six inches to a foot wide and dips  $52^{\circ}$  to the east.

A small outcrop of breccia with a questionable porphyritic matrix containing biotite hornblende porphyry (Tbhp) and fine-grained phaneritic igneous fragments occurs southwest of hole Bue-9. It is highly altered



to sericite and clay, and the nature of the matrix is difficult to distinguish. The presence of breccia pipes and pebble dikes in the vicinity of drill hole BSW-1 supports all other data relating to a tilted porphyry system with its top near hole BSW-1 (Figure 7, Plate 2, and Appendix D).



## STRUCTURE

The Buckskin Range has two distinct types and ages of deformation: Mesozoic thrust faulting and Cenozoic Basin and Range faulting. The Basin and Range faults are of two types: older, gently-dipping to flat faults which may have been originally steeply dipping, and younger, steeply-dipping faults.

### Mesozoic Deformation

A thrust fault is exposed in several locations in the northern Buckskin Range (Plate 3), bringing the Artesia sequence in contact with the Gardnerville Formation. Evidence of this thrust is abundant (Hudson, 1977). Although the direction of movement is unknown, the fault brings Jurassic Artesia sequence rocks into contact over Gardnerville Formation sediments. It has been roughly dated from middle to uppermost Lower Jurassic age.

In the prospect area there is a large, structural discontinuity between east-west striking, north-dipping foliated Triassic volcanics on the east and north trending, west-dipping Jurassic Artesia metavolcanics on the west. The similarity of the volcanics and lack of outcrop have combined to make it impossible to locate the fault exactly or determine its subsurface orientation. The Triassic volcanics to the east all lie at a lower elevation than the Jurassic Artesia metavolcanics permitting a thrust type fault with Jurassic volcanics overlying Triassic volcanics. It may be the same thrust fault identified by Hudson (1977) in the northern Buckskin Range that has merely moved down section. Some bulldozer trenching is planned next summer to investigate the nature of this fault.



### Tertiary Deformation

The dominant structural elements in the Buckskin Range are low-angle Basin and Range faults. These faults generally dip 5 to 15 degrees, although some dip to 7° west and 25° east (Hudson, 1977). The low-angle faults are generally not recognizable where they displace Mesozoic rocks because of the lack of distinctive marker horizons, but are easily recognized and traced where erosional remnants of the Tertiary ash flows are preserved. All of these faults have normal dip slip movement and are generally subparallel with a northerly strike.

The average dip of the Tertiary ash-flow tuffs is 75 to 90 degrees to the west, except adjacent to the low-angle Tertiary faults where intense drag folding occurs. This dip clearly indicates drastic rotation of the Buckskin Range and the ash-flow tuffs serve as good control for fault displacement. Offsets as small as 50 feet and up to 6000 feet can be clearly observed; all displacement is to the east.

Proffett (1972, 1976, 1977) studied the Cenozoic structure of the Yerington area in detail for several years while working for Anaconda. He concluded that Basin and Range faulting accompanied regional westward tilting from Late Oligocene through at least early Miocene (Figures 11 and 12). This resulted in steeply-dipping recent Basin and Range faults, mostly on the east flanks of ranges, but strongly-tilted older, inactive Basin and Range faults that are presently so gently dipping that they were earlier presumed to be gravity slide blocks. Conoco's detailed and reconnaissance mapping strongly supports Proffett's work, but not necessarily his conclusions.

Another small thrust was mapped in the south central part of the prospect area which places Jurassic Artesia volcanics over near-vertical



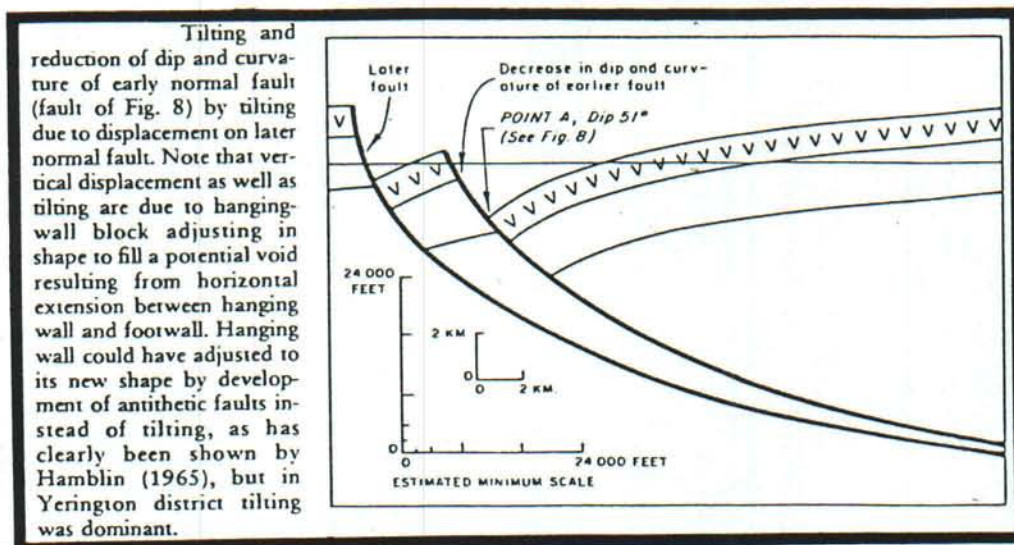


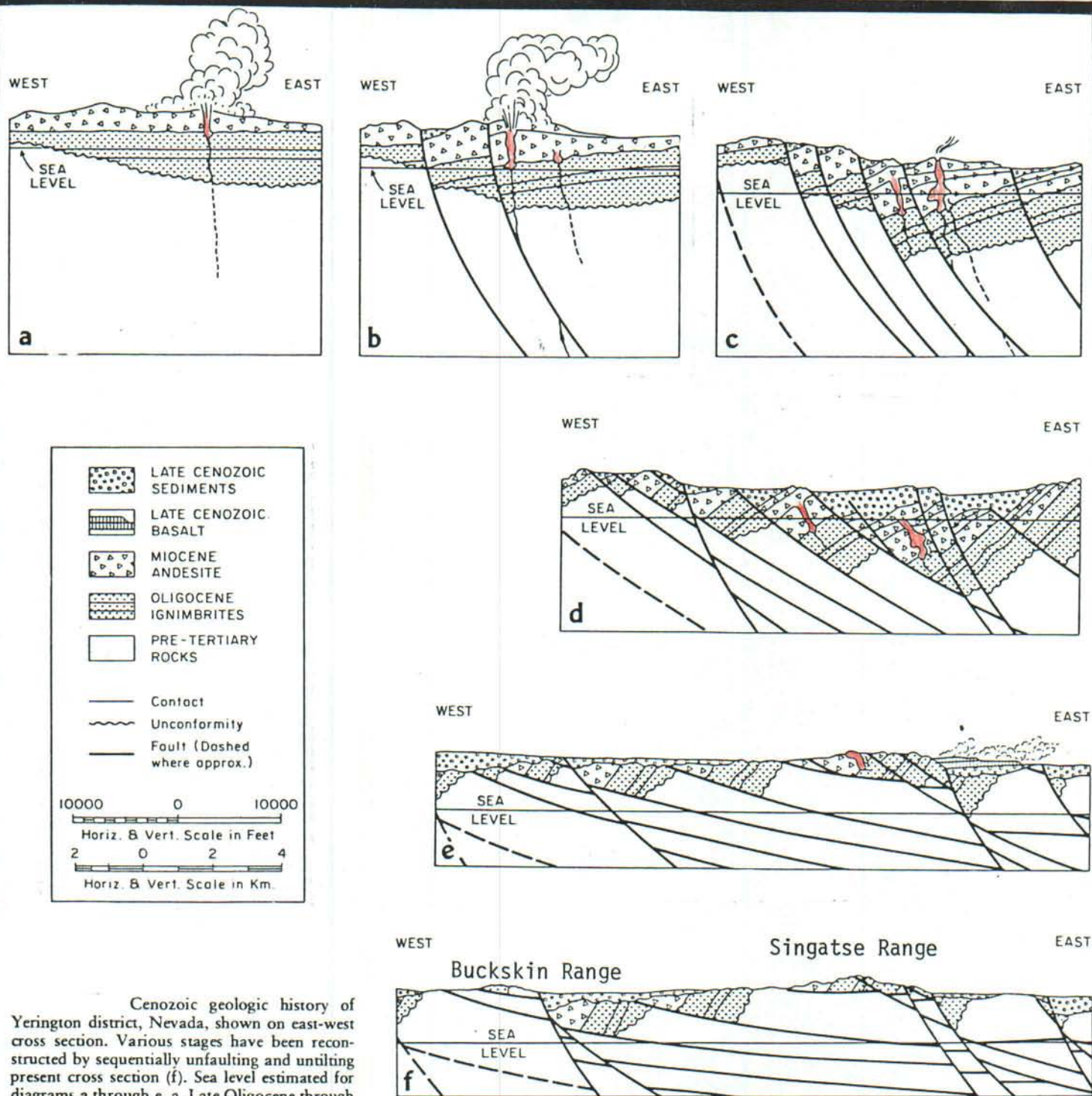
Figure 11

Tilting of Basin and Range faults  
From Proffett, 1977

dipping Tertiary ash-flow tuffs. The extent of this fault is uncertain but it is probably local. Tertiary mineralized porphyry dikes have intruded along and across some faults and locally cut the ash-flow tuffs (Figure 13, Plate 2).

Numerous other smaller presumed fault blocks which are defined best by juxtaposition of foliation within the metavolcanics are present within the large tilted Basin and Range fault blocks. Many of the bounding faults can only be presumed present because of the lack of outcrop. The foliations in the Mesozoic country rocks dip moderately to steeply west and strike north-south except in the eastern quarter of the prospect where they strike east-west and dip mostly north 45 to 65 degrees. There is no apparent folding and the postulated structural feature responsible for the large change in foliation does not fit well into the Cenozoic pattern of faulting. It is possible that it is a pre-Tertiary structure, perhaps the Buckskin thrust as discussed earlier.



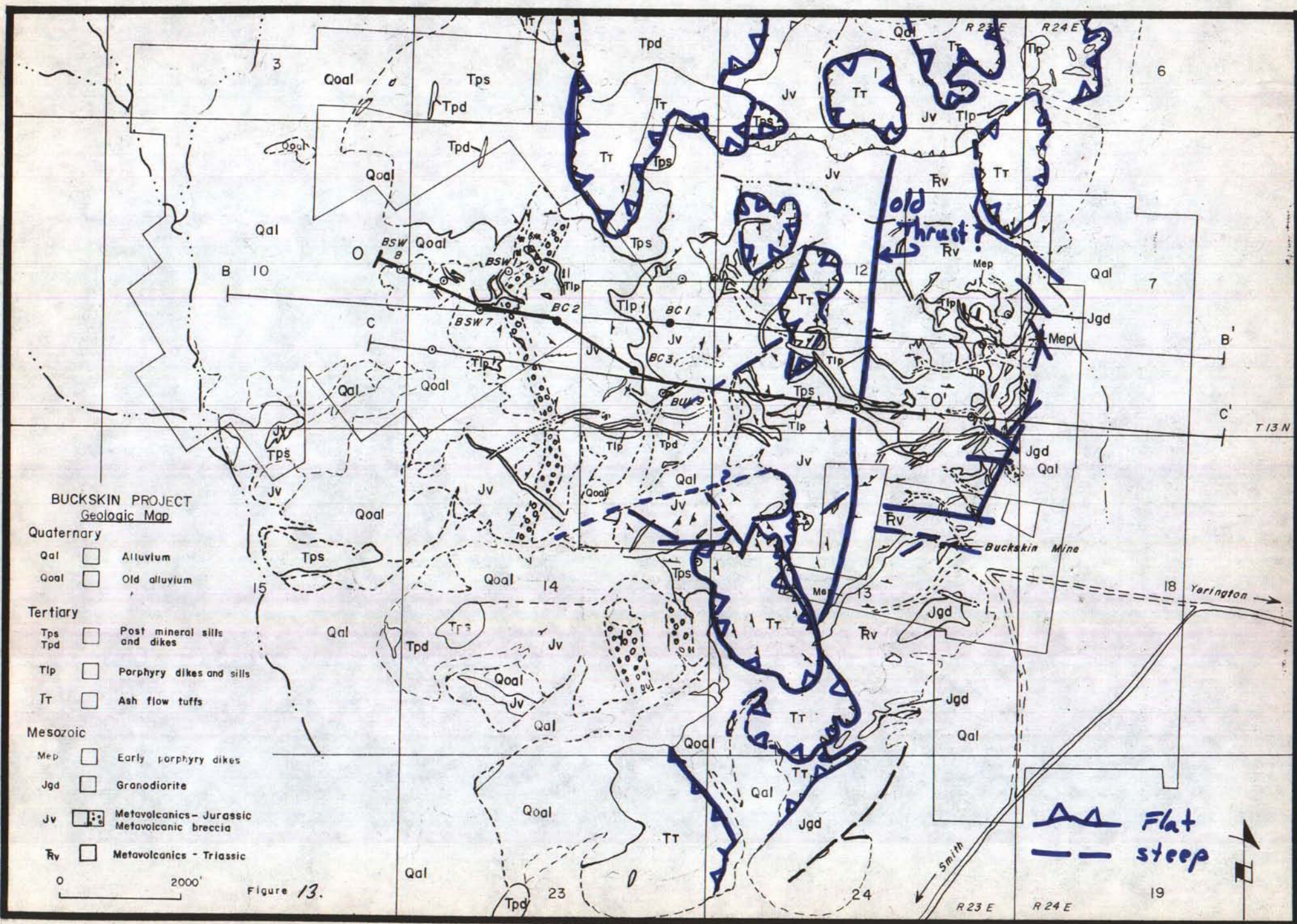


Cenozoic geologic history of Yerington district, Nevada, shown on east-west cross section. Various stages have been reconstructed by sequentially unroofing and untilting present cross section (f). Sea level estimated for diagrams a through e. a, Late Oligocene through early Miocene; deposition of ignimbrite sequence followed by beginning of deposition of hornblende andesite, 18 to 19 m.y. ago. b, Early Miocene, 17 to 18 m.y. ago; continued eruption of andesite and beginning of Basin and Range faulting and tilting. c, Early Miocene, 17 m.y. ago; end of andesite volcanism and continued Basin and Range faulting with erosion of up-faulted blocks. d, Early and middle Miocene, 11 to 17 m.y. ago; continued faulting, tilting, erosion of ranges, and deposition of sediments in basins. Early faults rotated to such a flat dip that they are no longer in favorable orientation for continued faulting. New steeper faults form to take their place. e, Eight to 11 m.y. ago; all early faults tilted to gentle dips and inactive, but faulting and tilting continues on newly formed steep faults; erosion of ranges and deposition of sediments and locally basalt in basins. f, Present time; continued faulting and tilting along steep range-front faults, erosion of ranges and deposition in basins.

Figure 12

Cenozoic Structural History  
From Proffett, 1977







Some gentle folding of the metavolcanics can be seen in the rhyolitic flow breccias on the west flank of the range. The fold axes trend west-northwest subparallel to the moderate and strong anomalous IP response.

The trends of the faults are mostly north-south, northeast and west-northwest. The northerly trending faults are the traces of the old eastward, steeply-dipping Basin and Range faults which are now tilted to the west so that the fault plane dips gently to the east (Figures 7, 11, 12, and 13). Because of the trough-like nature of these faults, they usually curve into northwest trending boundaries on the south ends of the blocks and trend to the northeast on the north end of the blocks (Proffett, 1977). Often these blocks are terminated by later northeast and northwest trending faults (Figure 13).

Most of the later northeast and west-northwest faults displace Mesozoic dikes and some Tertiary dikes in a right lateral sense. A series of west-northwest trending faults dissect a Mesozoic (KJLP) dike, into fault blocks that have moved right laterally. An apparent left lateral sense of motion is locally observed as a result of differential movement along the faults.



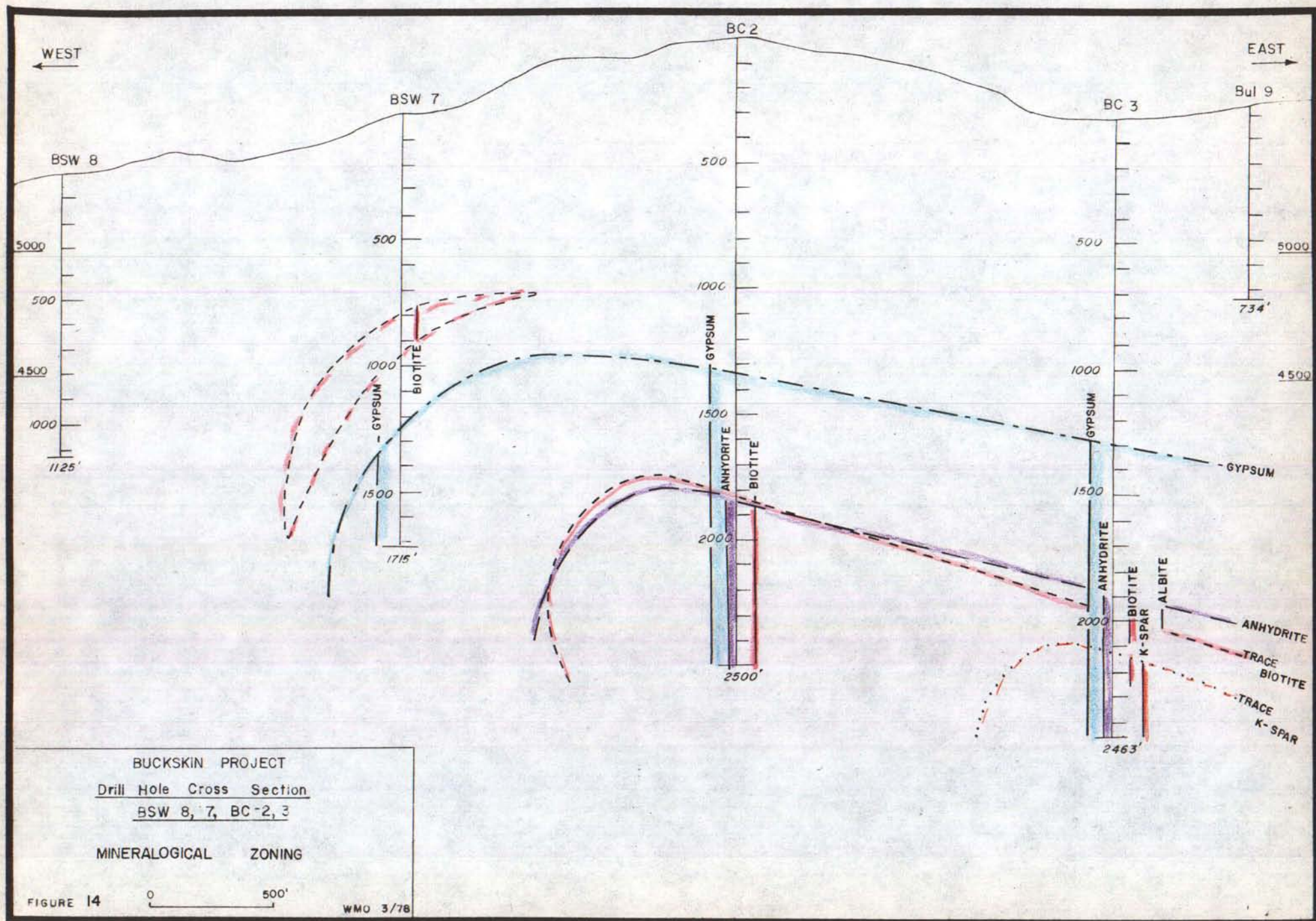
## ALTERATION

Surface and drill core thin section studies have developed patterns which largely conform to those typical of a porphyry system when westward tilting and erosion are accounted for (Figures 1 and 14). Thin section studies of Conoco's holes BC-2 and BC-3 (see Drilling) have especially formed a pattern and model that integrates all Conoco geology, geochemistry, limonites, geophysics and alteration data into a coherent consistent porphyry copper model (Figure 14 and Appendix D). Mineralogical gradients of gypsum, anhydrite, secondary biotite, and secondary K-spar form patterns indicating a porphyry copper system that is only approached at depth by holes BC-2 and BC-3. Pervasive sericite is present on the surface and in all holes with varying amounts of silica, clay, chlorite, epidote, calcite, and occasionally other products such as biotite, pyrophyllite, diaspore and rutile. Details of the holes drilled prior to 1977 can be found in last year's Buckskin report (Oriel, 1977).

Field observations defined the following pervasive mineral assemblages which appear to form mappable alteration zones at the surface: sericite, strong clay-sericite, weak clay-sericite, mixed sericite-chlorite and chlorite-epidote (Figure 15). Thin section studies of 112 surface and 245 core samples revealed conforming surface and subsurface alteration and mineralization patterns.

The sericite zone covers most of the prospect. It extends to the southwest and the northeast beyond our property line, but has a relatively sharp north-south trending boundary with the mixed sericite-chlorite (propylitic) zone on the east. The feldspars and the groundmass in this







BUCKSKIN PROJECT

MEGASCOPIC ALTERATION MAP

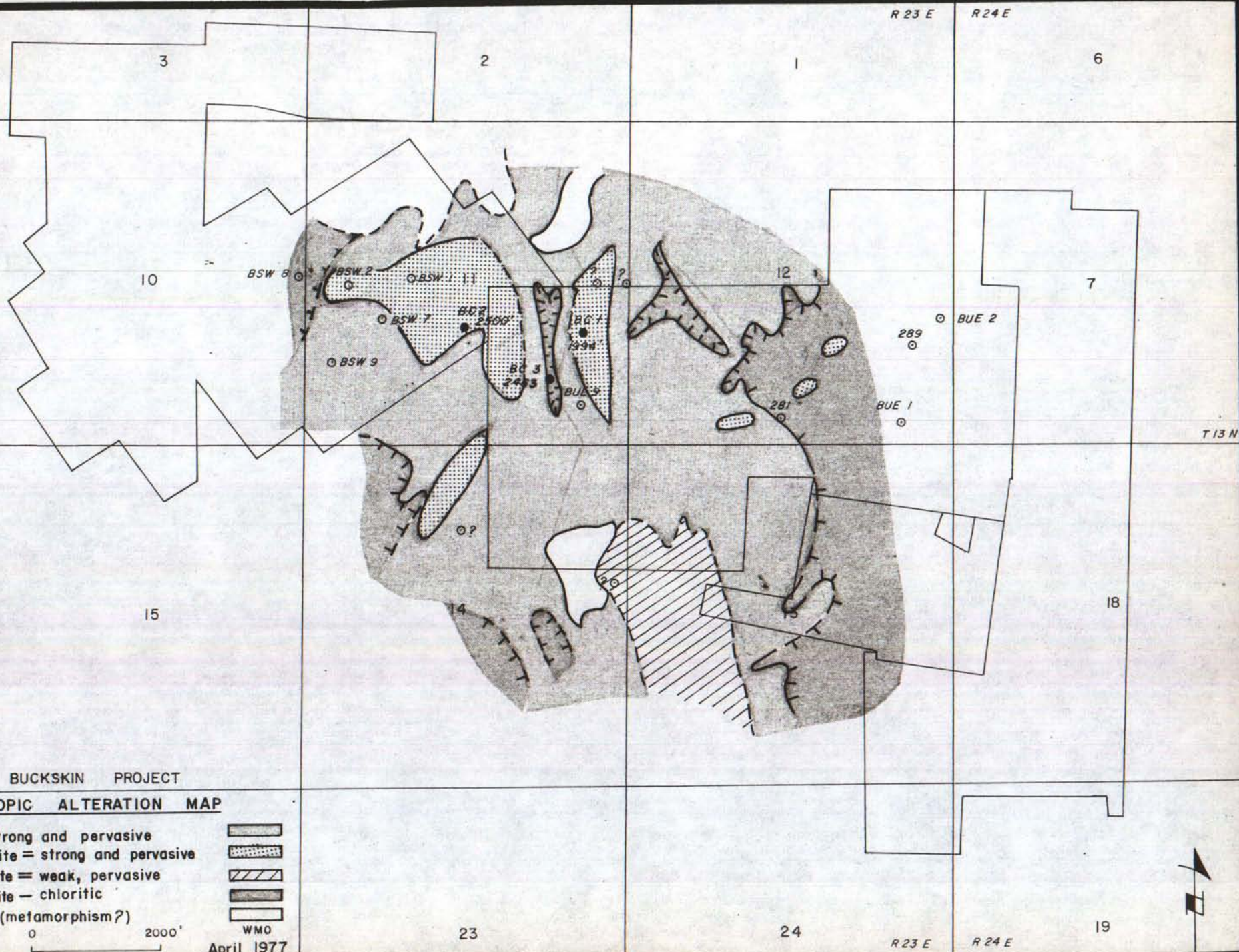
- Sericite = strong and pervasive
- Clay - Sericite = strong and pervasive
- Clay - Sericite = weak, pervasive
- Mixed Sericite - chloritic
- Propylitic (metamorphism?)



Figure 15

0 2000'

WMO April 1977





zone are typically 80% or more replaced by sericite. Accessories such as sphene and apatite are relatively unaltered. This zone is overlapped in the western portions of the prospect by patches of clay-sericite (argillic to advanced argillic) alteration.

The largest clay-sericite alteration zone is in the west. Sericite and low birefringence clays are generally pervasive in this zone. Some silicification of the groundmass occurs and quartz and clay veins are most common in these localities. The surface boundaries of this alteration zone were mostly defined by the presence of sparse clay veining, but thin section and X-ray diffraction data indicate that the western and largest area of strong clay-sericite occurrences may represent advanced argillic alteration as defined by the presence of natroalunite, diaspore and pyrophyllite. This mineralogic assemblage conforms well to the tilted porphyry copper model Conoco has developed. The central elongate strong clay-sericite alteration zone centered about hole BC-1 probably does not represent advanced argillic alteration. The clay veins here are metahalloysite, and are probably supergene products.

The weak clay-sericite alteration has been studied in hand specimen only. It is found in the ash-flow tuffs of the Mickey Pass sequence and is defined by a greater percentage of clay than is present in the strong clay-sericite alteration zones and yet weaker intensity of alteration. Rock type may be an important factor in this zone because the original tuff apparently contained a great deal of clay. The significance of this zone is not immediately obvious.

The mixed sericite-chlorite alteration zone forms most of the eastern part of the prospect. On the surface it is characterized by weak to strong sericite mixed intimately, both microscopically and



megascopically, with weak to moderate chlorite. Thin sections reveal a chlorite-sericite-calcite-epidote alteration assemblage in both dikes and volcanic country rock. The intensity or thoroughness of the alteration varies considerably from about 15% chlorite-calcite-sericite to total replacement. Mapping of the intensity of alteration from thin sections has not been undertaken at this time, but the pattern conforms well to a propylitic-type alteration zone around the sides and base of a tilted porphyry copper system (Figure 15 and Drilling section).

Propylitic alteration is found in the northern and southern portions of the property in late Tertiary sills and plugs. It also occurs in relatively fresh-looking Mesozoic andesites. Regional metamorphism probably accounts for the propylitic assemblage in the older volcanics. The alteration appears to be deuteric in the younger sills and dikes and is probably unrelated to the overall hydrothermal system which produced the other suites of minerals.



## LIMONITE MINERALOGY

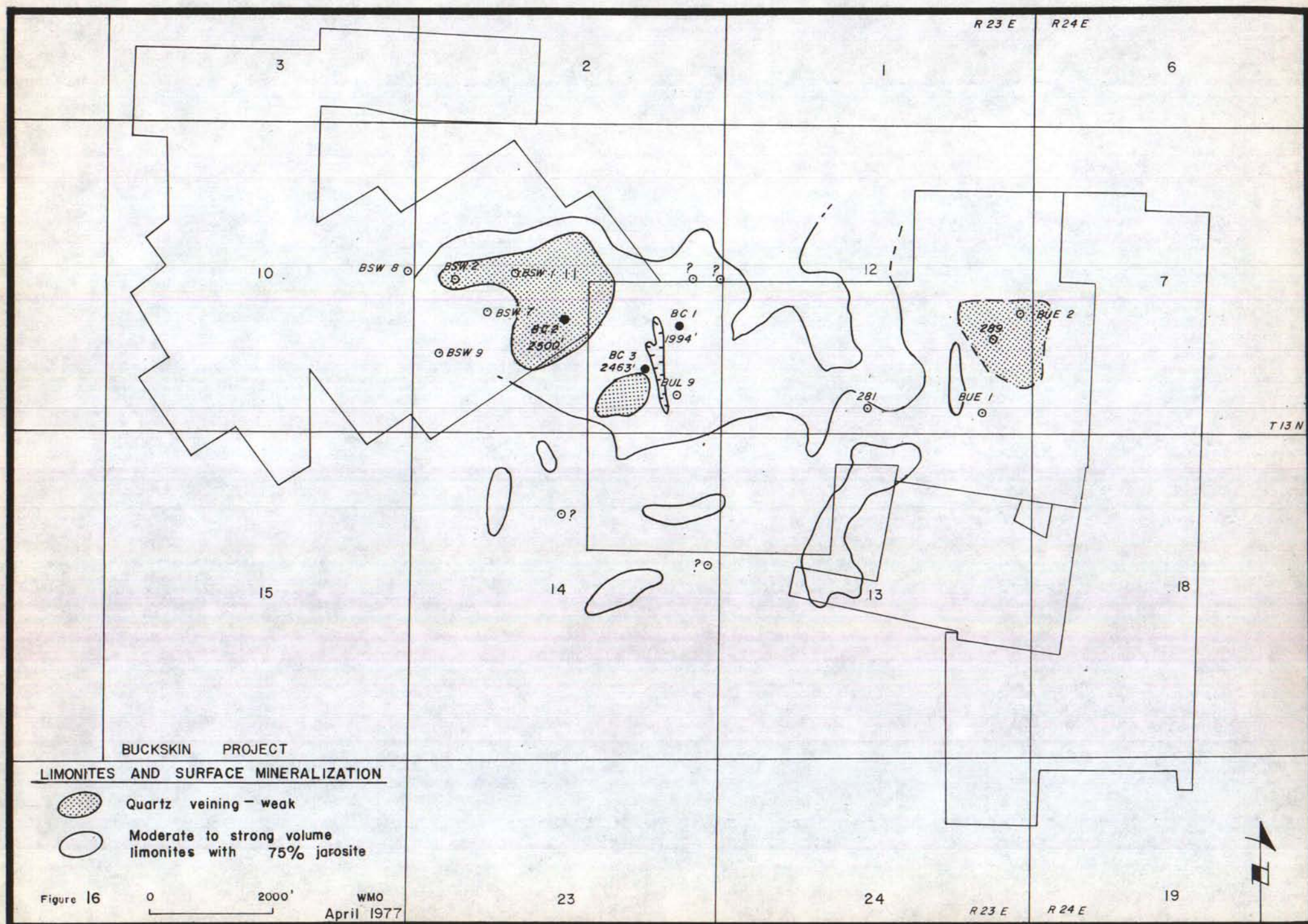
Surface mineralization consists predominately of limonites that formed after pyrite. A few patches of less-weathered rocks bearing fresh disseminated pyrite occur randomly throughout the prospect. Figure 16 is a volume intensity summary map of surface limonites. It depicts only visual estimates of "moderate" or "abundant" limonites on fractures, combined with a jarosite content generally greater than 75% of the total limonite volume. Its shape and size, however, agree remarkably well with the moderately anomalous IP zone (Figure 22), which represents about four to five percent pyrite in drill holes (see discussion on Induced Polarization).

Hematite stains are scattered everywhere, but do not form a mappable pattern. Typical "lavender" hematite after chalcocite can be seen in a few places, and drilling in holes BSW-1, 2, 7, and 8 indicates there is a small amount of secondarily enriched copper beneath the surface staining. "Brick red" hematite stains are also common and usually are associated with small fractures or fault zones. This color and type of occurrence is not necessarily related to copper mineralization.

Jarosite is typical of the thoroughly leached zones overlying moderate to strong pyritization and usually occurs in areas of sericitic and/or argillic alteration, possibly because leaching is favored by these alteration products.

Goethite stains on fractures are found in less-altered rock indicating generally less-disseminated pyrite than the area of predominately jarositic limonites. Goethite is more abundant in the propylitically-altered volcanics on the east flank of the prospect.







All three limonites can be found on separate fractures in practically the same hand specimen, but they are characteristically fairly pure. Jarosite is usually 90% pure (estimated) and only occasionally has more than 20% blended hematite and/or goethite. The same is true for the hematite stains. Goethite may have more mixtures of jarosite and hematite, but it is more difficult to estimate the impurities in goethite.

Turquoise veinlets are abundant in roughly linear west-northwest trending zones on the eastern and western sides of the property. They do not quite join in the center of the prospect (Plate 8).

Outcrops showing malachite, azurite and "neotocite" (Mn-Cu oxide) are sparse. These minerals can be found along small veinlets and fractures. They occur in weakly-altered and mineralized fringe areas where strong leaching has not taken place. Their presence is not related to, or representative of soil geochemistry.

Veinlets of pure jarosite up to 1/2 inch wide invade the Mesozoic volcanics in the vicinity of the Bear Creek holes BSW-1 and BSW-7.

Weak quartz veining areas have been included on the limonite map for convenience because the alteration map was too crowded. The western area of veining is in the vicinity of holes BSW-1 and BSW-7, which is reasonable for the top of the porphyry system near the axis as has been interpreted from other geologic and geochemical data.



## GEOCHEMISTRY

Soil Geochemistry

Soil samples were collected over a large area for geochemical analysis early in 1976 as a first step in defining the Buckskin system (Figures 17 through 21). To this end the program was entirely successful.

Approximately five hundred samples were collected on a 300 by 600 foot grid and assayed for copper, molybdenum, lead, zinc, and silver. The property has only a thin to non-existent soil horizon overlying a few inches to several feet of decomposed bedrock. The samplers tried to consistently collect the finer material from the top of the decomposed bedrock. Quaternary alluvium was also sampled and serves as a rough guide to background values. Contour intervals were selected by comparison of frequency histograms of the values of each element.

All elements have closely similar outer limits defining the geochemically anomalous areas. Only the distribution of the higher values changes significantly within this pattern. It should be noted that the geochemically anomalous values do not coincide directly with the strongest alteration. Although the values generally define a large roughly east-west trending system, three areas (western, central and eastern) can be separately discussed.

The western anomaly was drilled by Bear Creek before 1973. It consists of a copper and moly anomaly with weak zinc and silver anomalies. The outer or western part of the central lead anomaly just touches this area. Widespread thin veinlets of turquoise may account for the high copper values, but a thin blanket of supergene chalcocite 120 to 170 feet deep underlies the surface anomaly.



BUCKSKIN PROJECT  
Soil Geochemistry

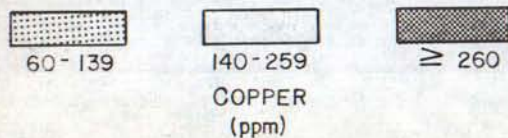


FIGURE 17

0 2000'

WMO  
Jan. 1977

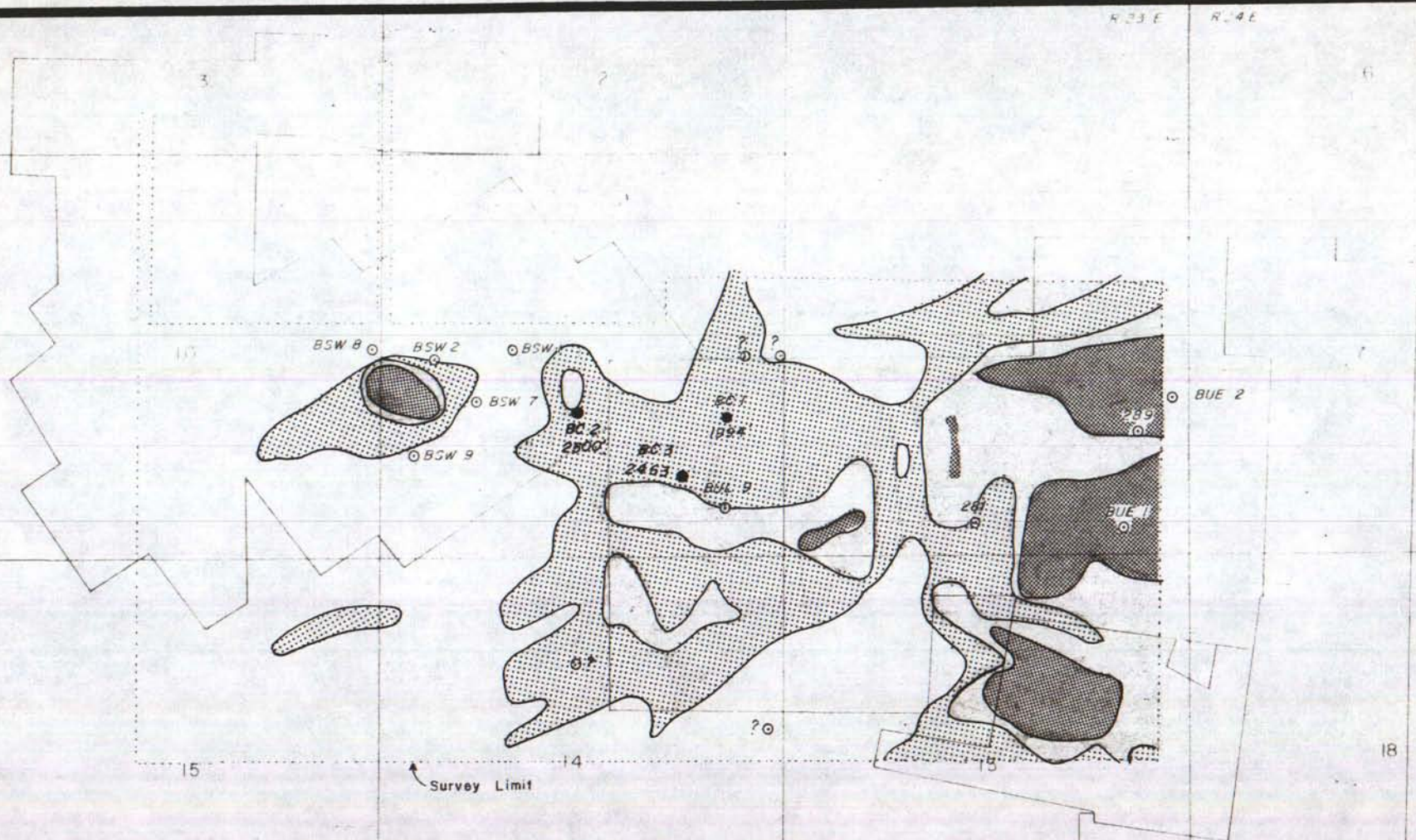
23

24

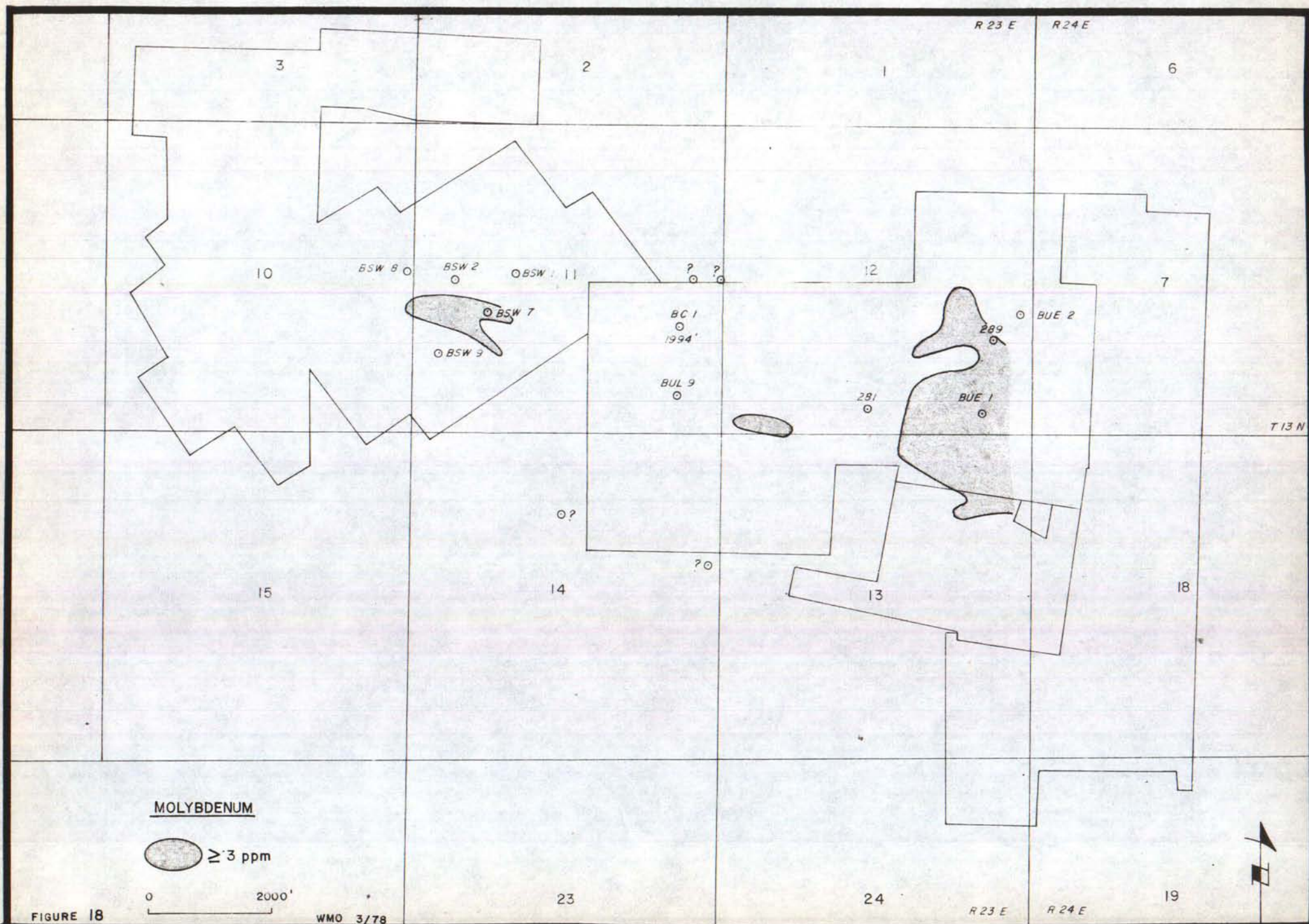
R 23 E

R 24 E

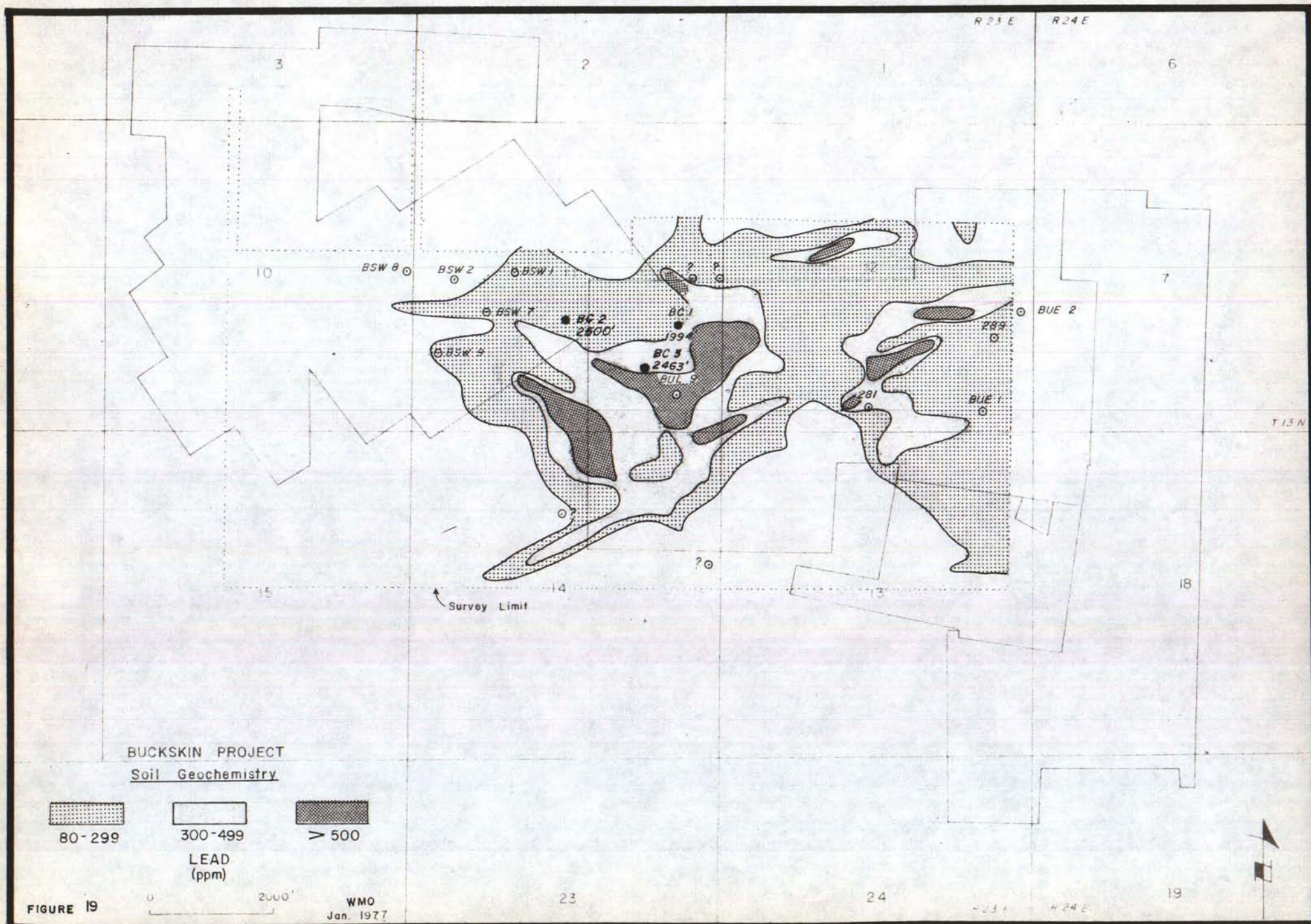
19













BUCKSKIN PROJECT  
Soil Geochemistry

120 - 379

380 - 599

≥ 600

ZINC  
(ppm)

FIGURE 20

0 2000'

WMO  
Jan 1977

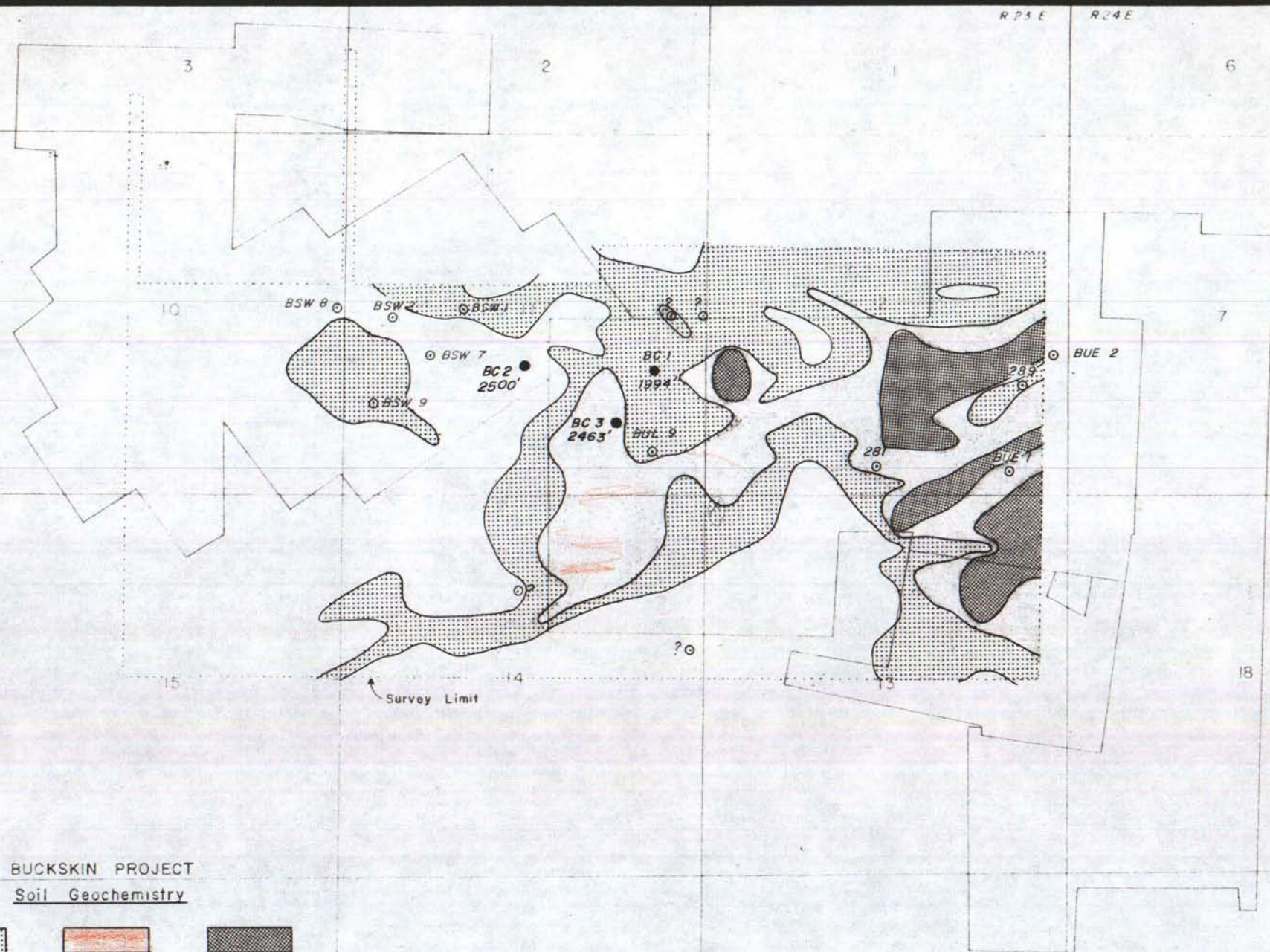
23

24

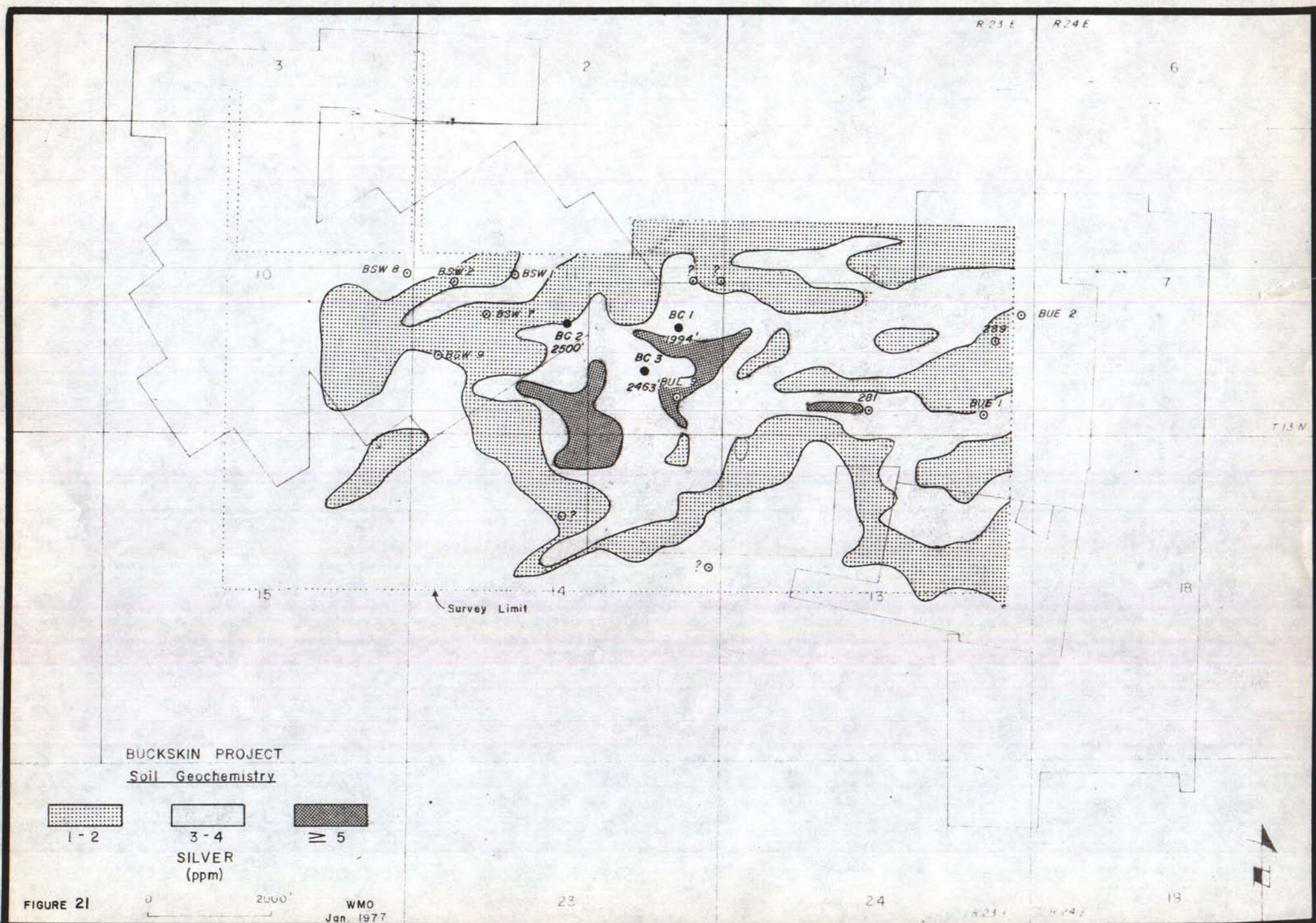
R 23 E

R 24 E

19









The central anomaly is a large bulbous-shaped area containing high values of lead and silver, moderate to high zinc and low copper. Its eastern boundary is roughly defined by decreasing values in the vicinity of a large tilted Basin and Range fault which drops Tertiary ash-flow tuffs down against Mesozoic metavolcanics. These ash-flow tuffs are sericitically altered but are apparently a poor host for mineralization. Drilling indicates the high lead-zinc values are representative of the underlying bedrock, but the anomalous zone may dive westerly beneath the western geochemical anomaly.

The eastern anomaly is characterized mostly by northeasterly trending linear zones of higher values of copper, moly, lead, and zinc, within a broad area of moderately anomalous values. Lead is a possible exception in that its linear elements form a roughly-arcuate zone. This linearity may in part be due to drainage. Most drainages, however, trend easterly and not northeasterly. The soil survey terminated to the east in alluvium parallel to the fault that locally bounds the range. Not much is presently known about the controls of mineralization in the east or exactly how it relates to the central and western soil anomalies.

The three areas can be considered to form a pattern consistent with the porphyry-type sulfide system outlined by drilling and mapping. The western copper anomaly is where the axis of the system apparently breaches the present surface. The central lead-zinc anomaly conforms to the appearance of an east-dipping lead-zinc halo peripheral to the axis. The eastern anomaly should be the extreme periphery of the system at depth. Complications of geology, older possibly-mineralized intrusions, and lack of drilling, prevent a ready analysis or comparison of the eastern portion of the district.



Soil geochemistry has many deficiencies due to both sampling technique and dispersion effects by ground water and local drainages. It is, however, a relatively quick and easy way of defining large geochemically anomalous areas and is essentially an averaging technique; and in this case, it successfully outlined the system and gave us a rough idea of where in the geochemical zoning the surface lay.



## GEOPHYSICS

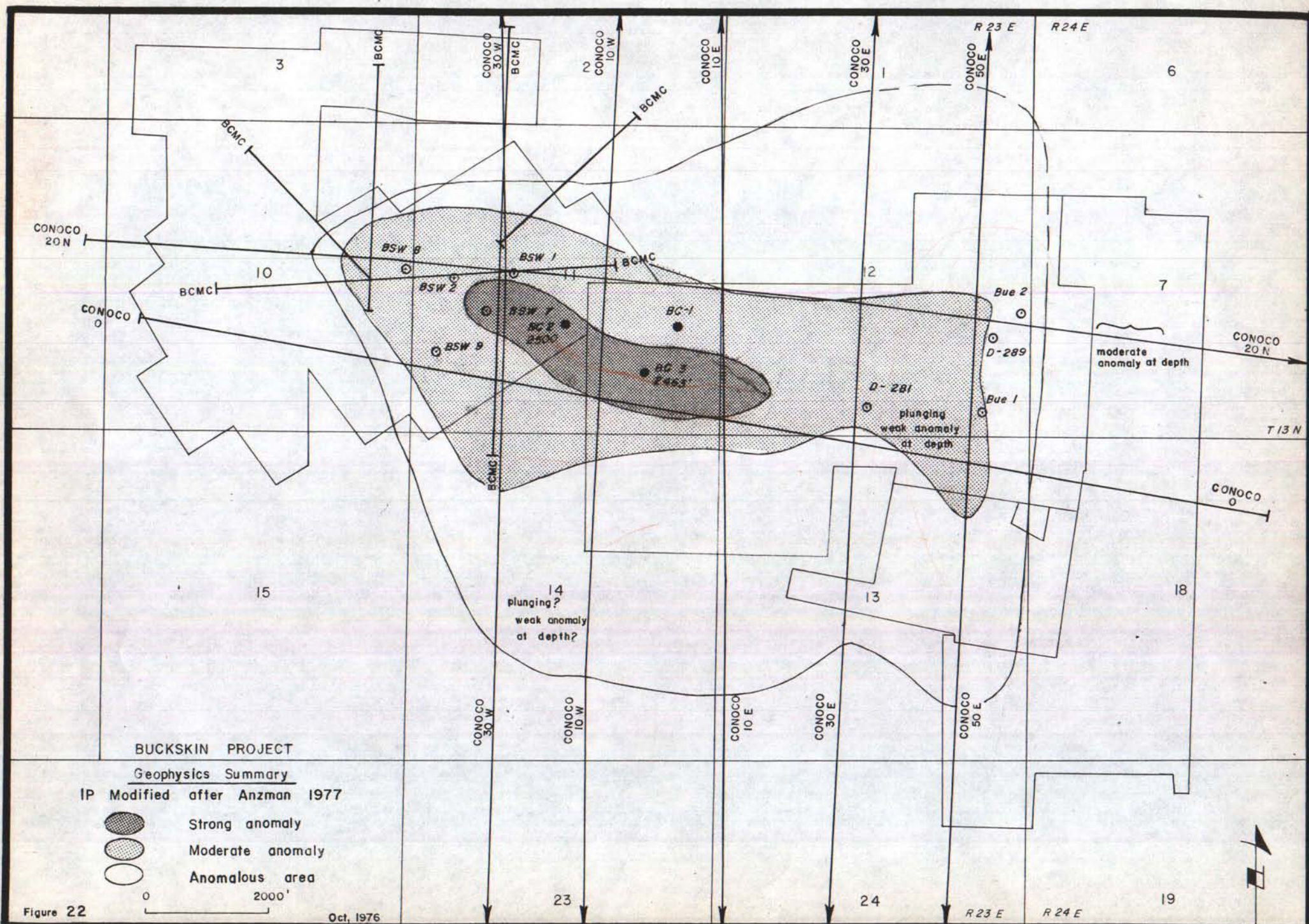
Induced Polarization (IP)

Conoco ran approximately 23 line miles of IP consisting of two east-west and six north-south lines in 1976 to complement and confirm IP data obtained from Bear Creek and Phelps Dodge. Lines 10W and 20N had 2000 foot dipole-dipole spacing; all the rest had 1000 foot dipole-dipole spacing. Mining Geophysical Surveys conducted the survey of the two east-west lines, and a Conoco crew collected the data for the north-south lines. Joe Anzman was hired as a consultant to interpret the data and Peter Kirwin and myself re-interpreted his results and made changes based upon drill hole sulfide content (Figure 22; Plate 13).

The data indicate a large anomalous area 9000 by 12,000 feet elongated east-west and extending to at least 3000 feet in depth. Within this area is a "moderate" anomaly 2500 feet wide, 11,000 feet long, and greater than 2000 feet deep. The strongest anomaly extends 4000 feet east-west and is about 1000 feet wide and approximately 1500 feet deep. It is in a central position relative to the lesser anomalies and apparently has steep boundaries. The presence of sulfides as the source of the anomalies has been confirmed by drill holes. Drilling indicates the strong anomaly represents greater than 5% total sulfides in what may be a "phyllic" shell of the classical porphyry modes. The moderate and weak anomalies reflect drilled sulfides of 2 to 4 percent and less than or equal to 2% total sulfides respectively.

The IP surveys show that a large sulfide system is present at the Buckskin prospect. Pyrite content of drill holes was used to refine the







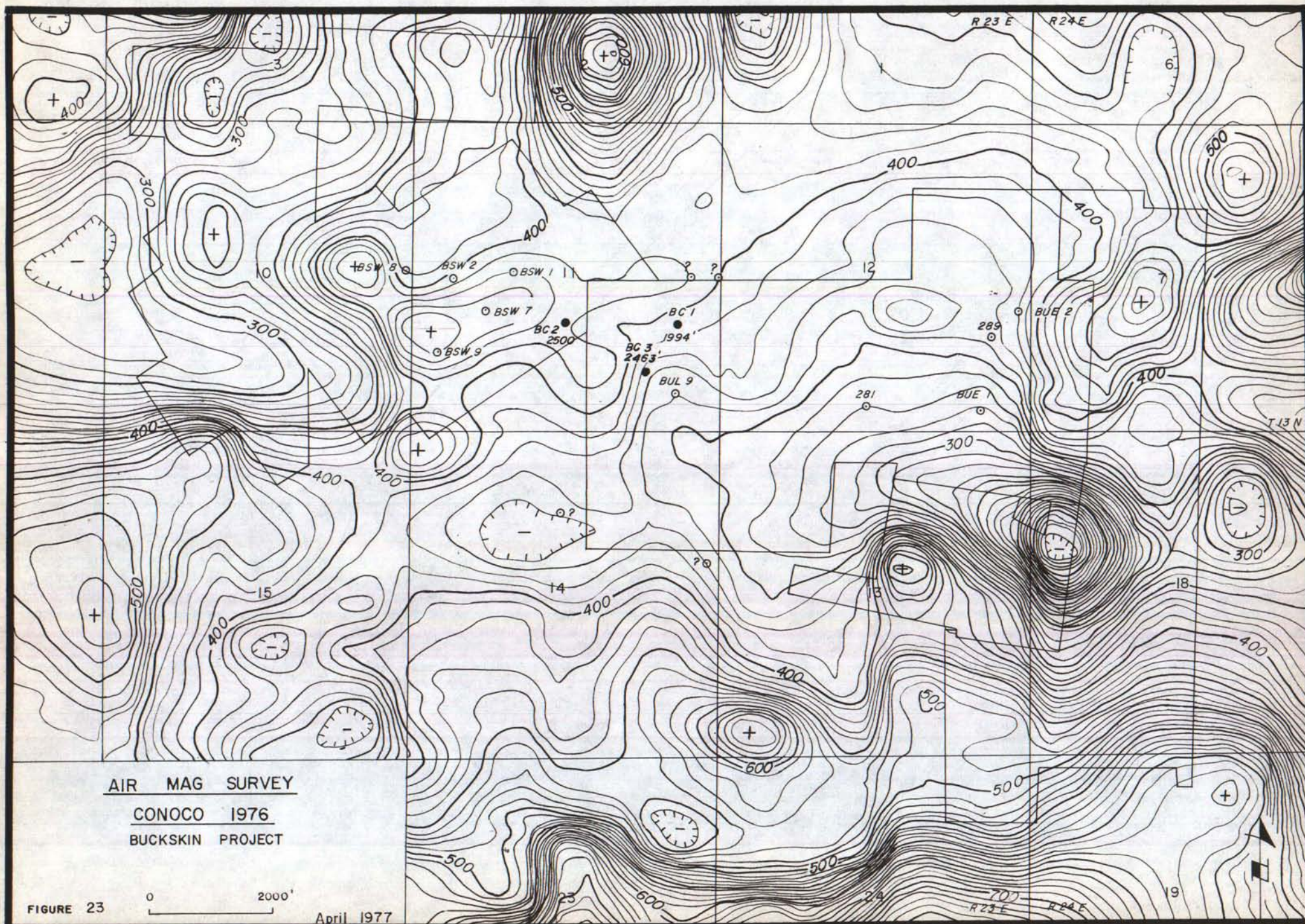
boundaries of the anomalous zone. The sulfide system trends east-west, transects all local geology and structure, and appears to be decreasing at the depth limits of reliable IP data. It cannot be determined whether this decrease is permanent or local, or exactly what patterns the sulfides define at depth, but the geology and drilling suggest that the western Buckskin ground surface is near the top of a large, strong, tilted, porphyry-type sulfide system. Within the decreasing sulfides at depth may be the core of the system and the target area for economic copper mineralization. This system plunges to the east coincident with the anomalous IP zones.

Line 20N, the east-west line which was run with 2000 foot dipoles, indicates a deep anomaly on the down-thrown side of the Basin and Range fault on the east flank of the range. Conoco has staked this area and will direct future efforts to explain this anomaly. Gravity profiles indicate that depth to bedrock is less than 1000 feet over this anomaly (Aiken, 1978).

#### Local Airborne Magnetic Survey

Conoco flew an aeromagnetic survey over the Buckskin Range with a line spacing of 650 feet in the fall of 1976 (Figure 23). No interpretation has been performed by any geophysicist. Joe Anzman briefly examined the data and concluded that it was not generally suggestive of any pattern relating to the known geology or the IP responses and hence was not very useful. There is, however, a large quiescent area of relatively low magnetic values roughly correlative to the surface exposure of the mineralized rocks. It appears to trend northeasterly and is partially surrounded by small positive anomalies. Some of the high magnetic values are not readily explained because they occur over







the altered Mesozoic volcanics instead of late Tertiary andesites. The high response in the southeast portion, however, may be coincident with relatively unaltered Mesozoic andesites.

#### Regional Airborne Magnetic Survey

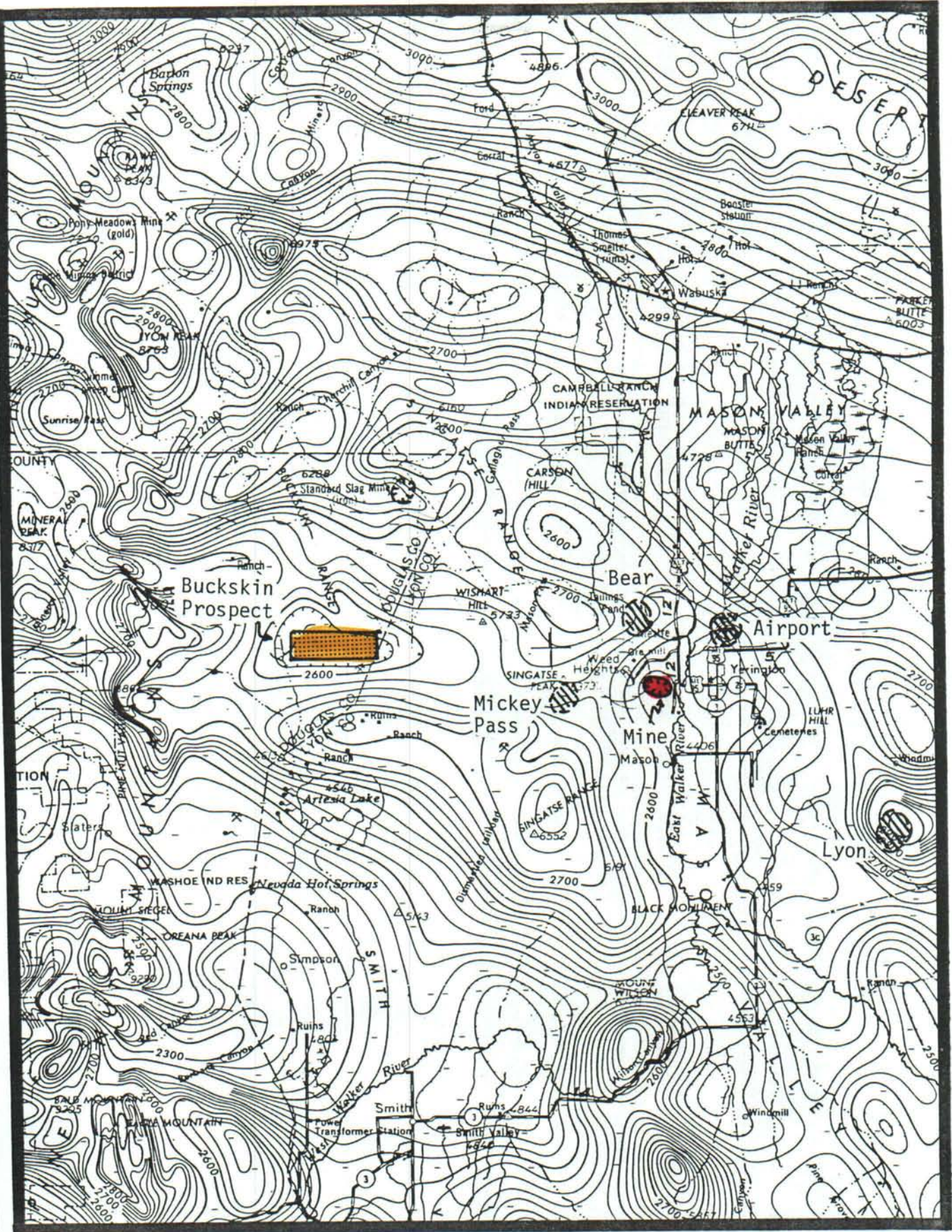
A regional aeromagnetic survey was completed by the USGS in 1971 (Figure 24). The Buckskin prospect is within a roughly east-west trending, broad magnetic low. It is difficult to determine if Conoco's detailed air mag would generally conform to this pattern without more lines to the west of our present survey. The Minnesota iron mine and the Pumpkin Hollow (Lyon) iron-copper district have definite positive magnetic expressions whereas the Yerington porphyry copper open pit is within the north end of a negative magnetic anomaly. Other deep porphyry copper deposits found by Anaconda lie on the margin of the same large magnetic depression.

In general both the local and regional negative magnetic anomalies at the Buckskin prospect are correlative and are possibly caused by hydrothermal destruction of pre-existing magnetic minerals.

#### Gravity

A limited Conoco gravity survey was performed late in 1977 over the deep IP anomaly on the east flank of the prospect. Results suggest that depth to bedrock over the anomaly "will not likely exceed 1000 feet" (Aiken, 1978).





REGIONAL AIR MAG SURVEY  
BUCKSKIN REGION  
Scale = 1:250,000

Figure 24.

Taken from G.P. 751; USGS, 1971



## DRILLING RESULTS AND INTERPRETATION

### Holes BC-2 and BC-3, Conoco, 1977

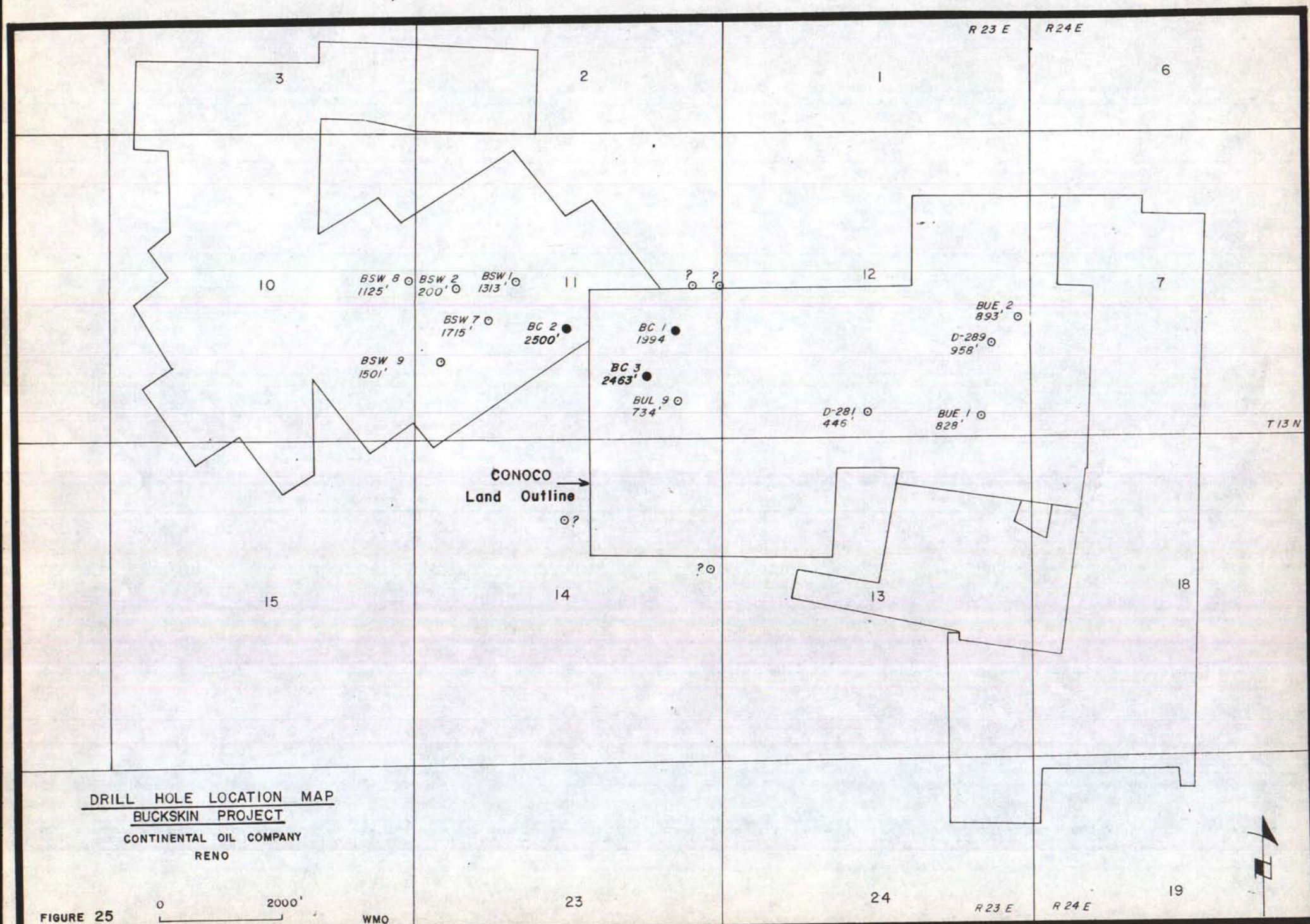
In 1977 Conoco drilled holes BC-2 and BC-3 to 2500 and 2463 feet respectively. Both intersected strong lead-zinc halo-type mineralization indicative of the fringes of a porphyry copper system (Figure 25; Appendix C).

### Geochemical Results

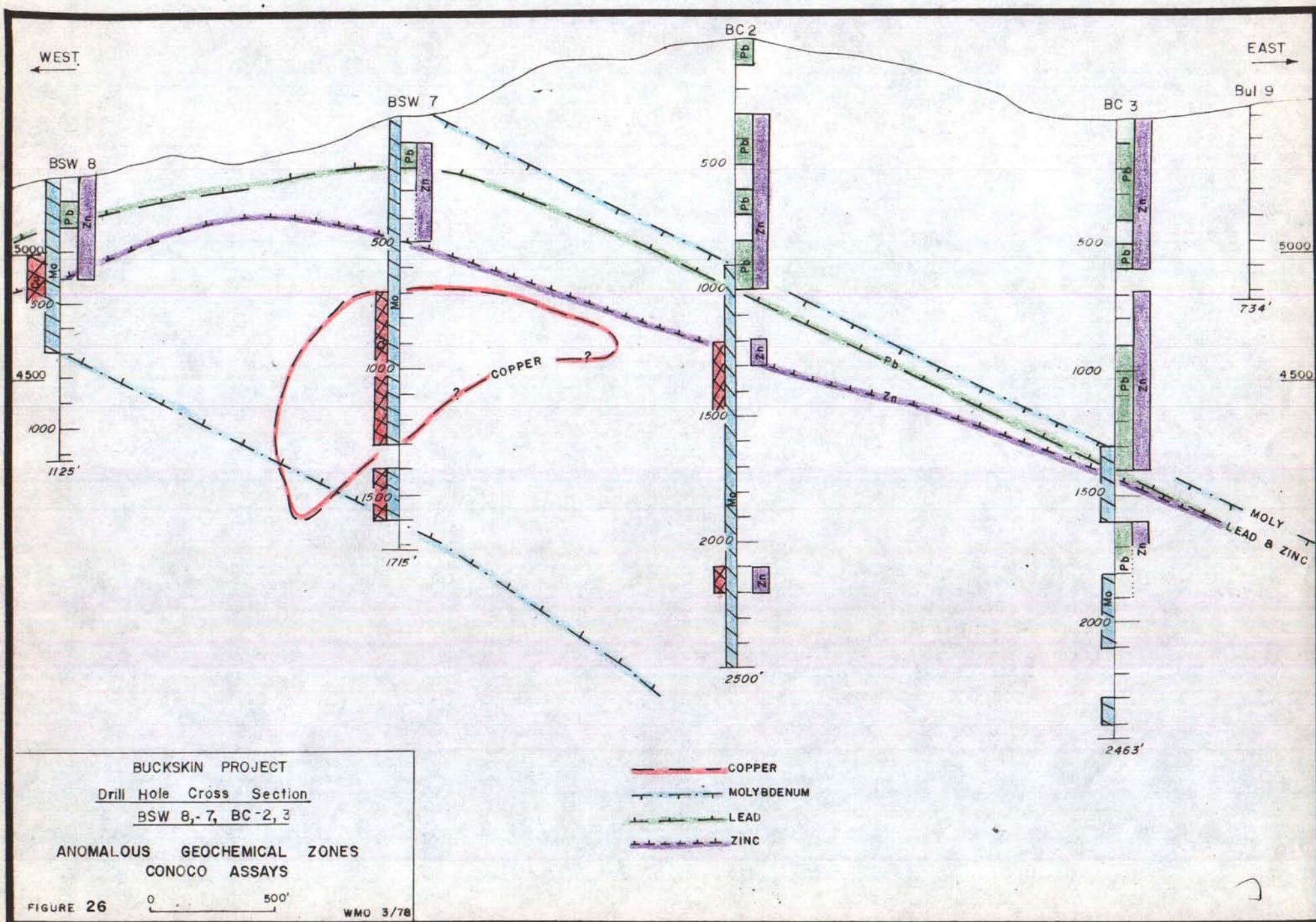
The geochemical results of Conoco's drilling last summer were recently compiled and alteration in the two holes was studied in thin section. A series of cross sections were drawn to accommodate the new information from holes BC-2 and BC-3 and BCMC holes BSW-7 and BSW-8 (Figures 26, 14, and 27). Cross sections containing 100 foot averages of Conoco's analyses of Cu, Mo, Pb, Zn, and Ag show consistent and almost classical patterns for a large porphyry copper system. Lead and zinc form parallel zones of high geochemical values external to a strong molybdenum zone and a weak but downward increasing copper zone. In BC-2 (Appendices C and D) total pyrite decreased from greater than 5% to 2-4% below 1900 feet. Copper inversely increased from less than about 300 ppm to as much as 1300 ppm for a 100 foot average assay below 1900 feet. Hole BC-3 (Appendices C and D) showed similar but lesser changes in pyrite vs. copper content.

Galena, spalerite and pyrite veins and veinlets occur sporadically throughout both holes, but are more common within the higher lead-zinc anomalies of both holes. A cross section through BCMC holes BSW-1, BSW-7, and BSW-9 (Figure 28) shows that the tip of the system may breach the present topography in the vicinity of hole BSW-1. This is indicated by

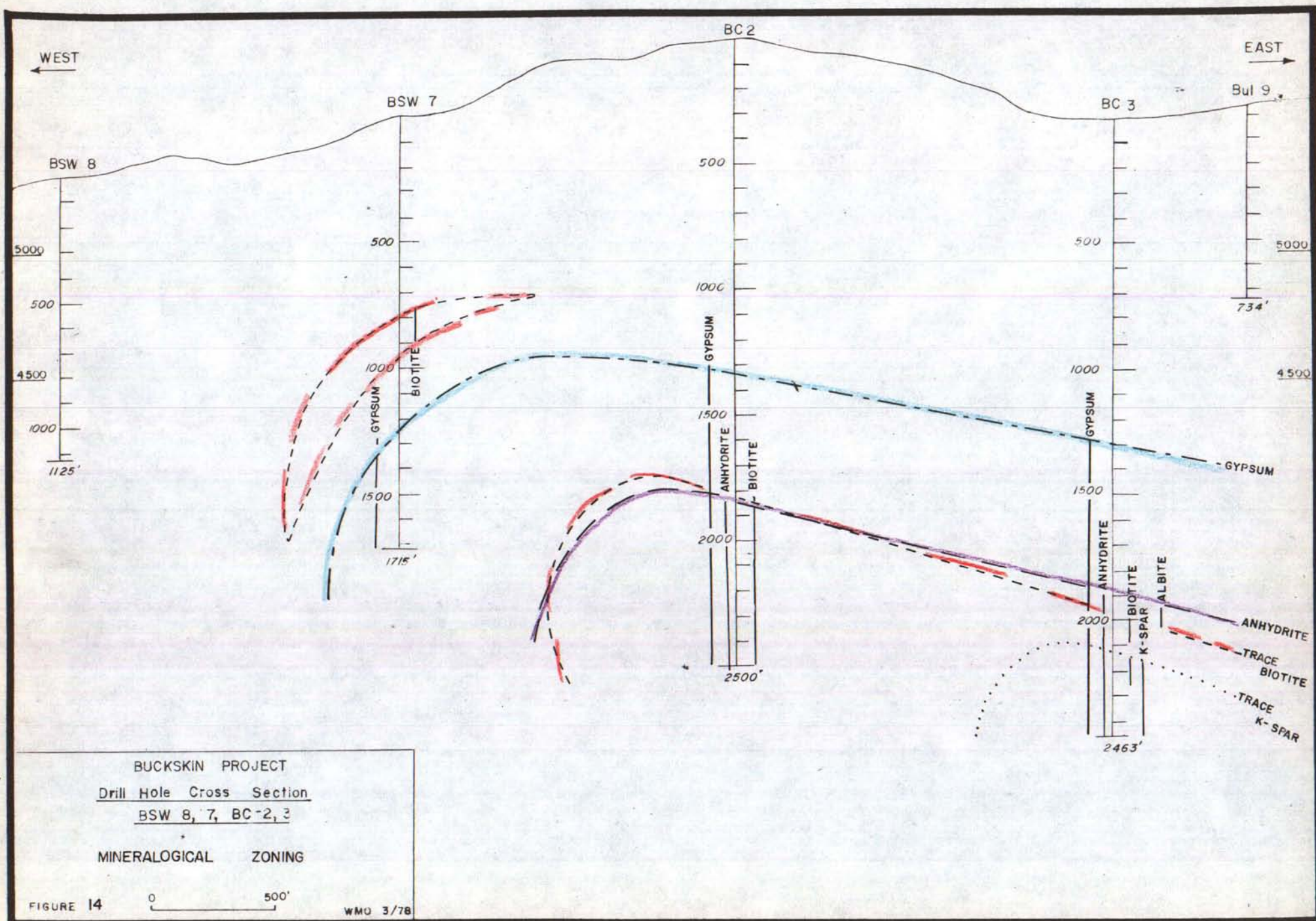




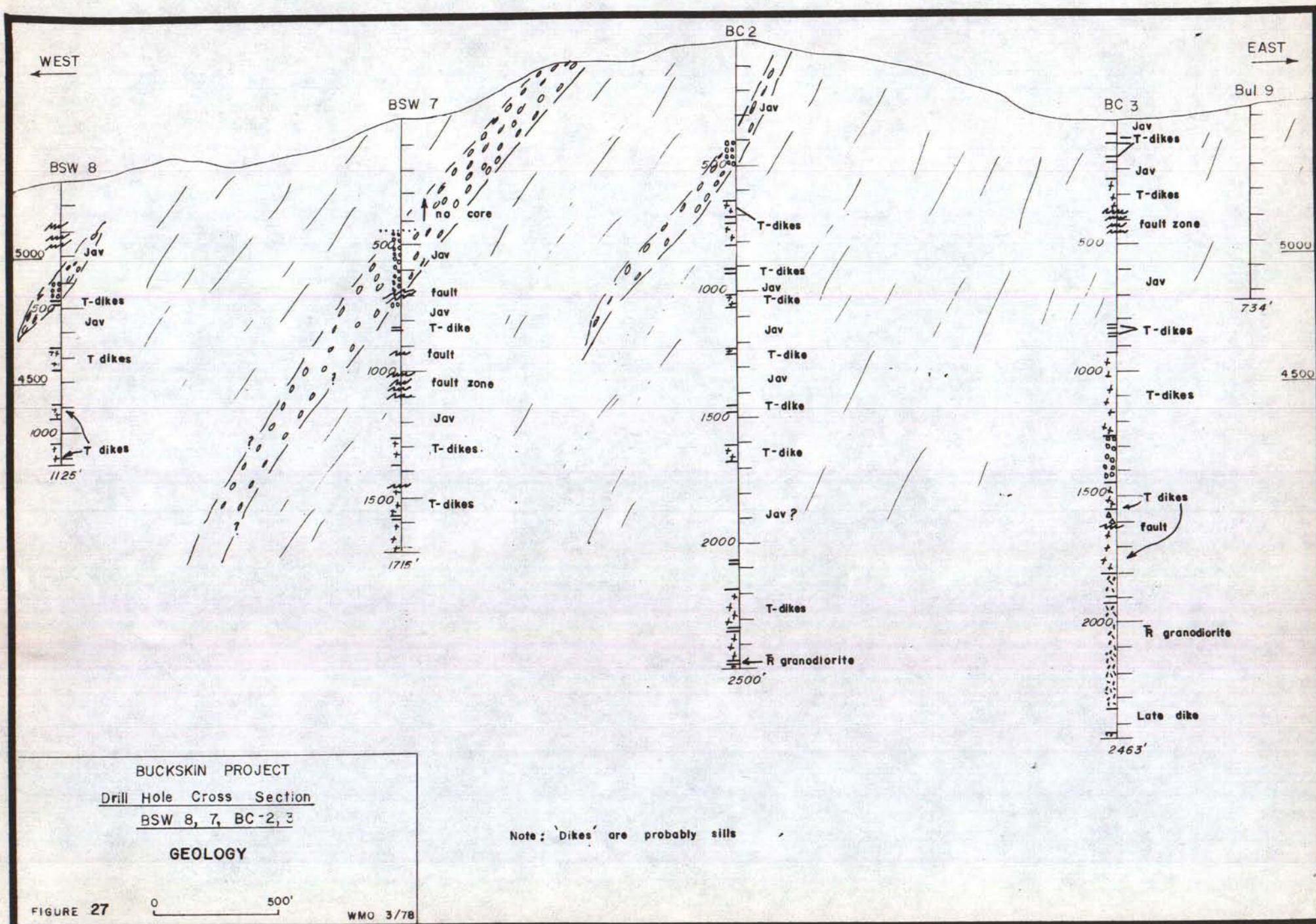




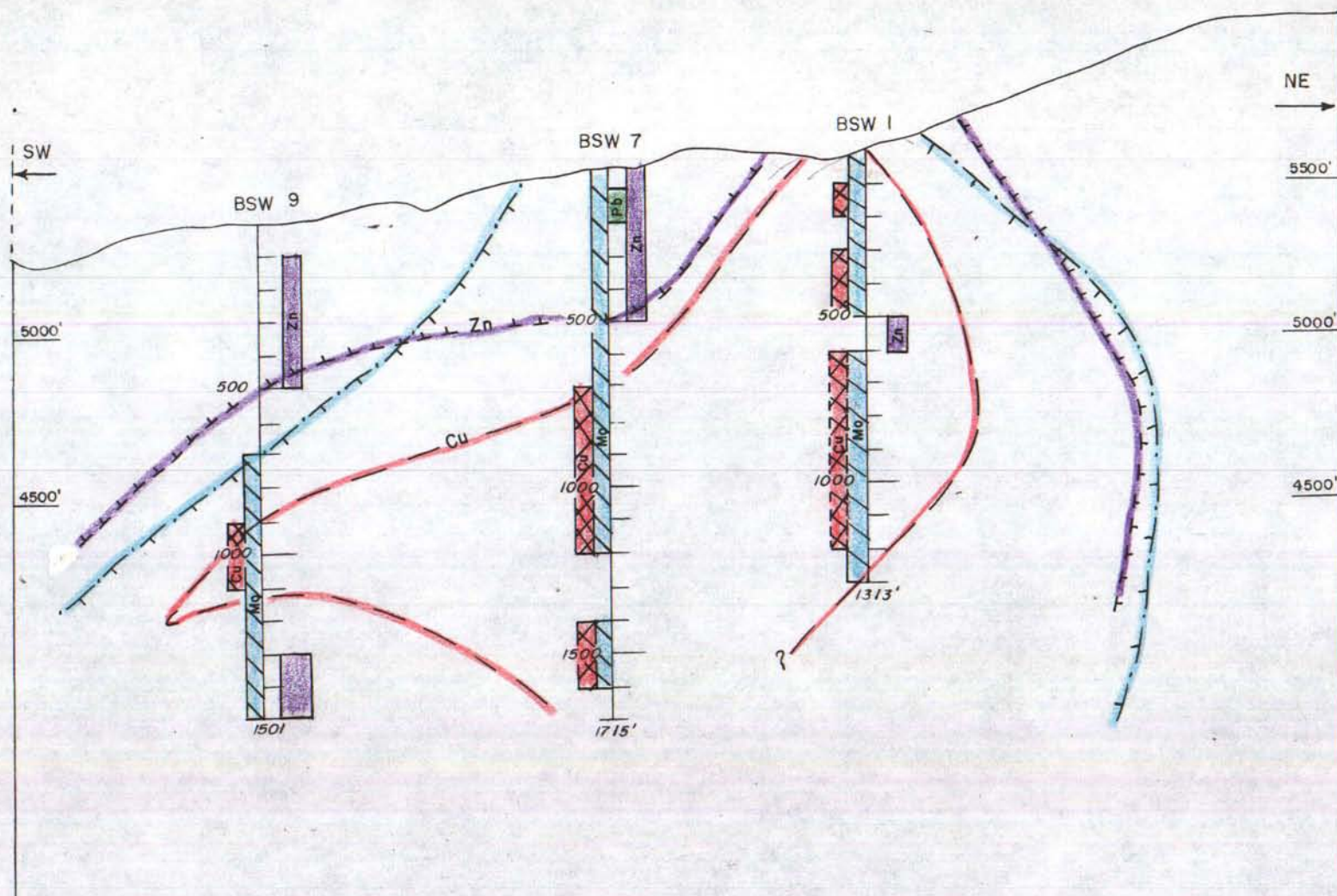












BUCKSKIN PROJECT  
Drill Hole Cross Section  
BSW 1, 7, 9

ANOMALOUS GEOCHEMICAL  
ZONES

— COPPER  
— MOLYBDENUM  
— ZINC



the interior moly anomaly (with weak copper) rising toward the surface. Anticlinal doming of an intraformational volcanic conglomerate (Javb) (Figures 1, 7, and 29; Plate 2) takes place near hole BSW-1. Additional geologic manifestations of this trend and plunge of the system are shown by the presence of quartz.

#### Alteration and Mineralogy

The alteration and mineralogical zoning parallel the geochemical zones in an equally classical pattern. Holes BC-2 and BC-3 show progressive increases in gypsum, anhydrite and secondary biotite toward the bottom of the holes. Secondary orthoclase, albite and anhydrite increase easterly toward and below hole BC-3 (Figure 1, 14, Appendix C). Sericite is predominant throughout both holes, although chlorite with traces of epidote is widespread in both holes, especially in the lead and zinc rich rocks and in some late dikes.

Secondary biotite and orthoclase occurrences in holes BC-2 and BC-3 form consistent downward increasing zones, but should not be mistaken for the classical potassium-rich cores with which most copper ore is associated. Orthoclase and biotite are present in weak to trace amounts, usually recognizable in hand specimen after thin section identification. The true potassium-rich zone expected in such a system has not yet been drilled but forms the intriguing target for 1978 (Figure 1).

All Conoco data coincide to indicate that holes BC-2 and BC-3 were drilled into the top half of a porphyry copper system, dipping 15 to 35 degrees easterly. The copper-rich zone was not entered but only approached by holes BC-2 and BC-3.



BC-1 - 1994' - Conoco - 1976

Conoco's diamond drill hole BC-1 was completed December 20, 1976, at a vertical depth of 1994 feet (Appendix C). It is in the northerly part of the central geochemical anomaly and in the center of the moderate IP anomaly. It intercepted anomalous but weak copper and molybdenum mineralization, but had a very strong zone of lead and zinc from 70 to 1100 feet. Lead and zinc increase again from 1650 feet to the bottom. Sphalerite and galena are finely disseminated in the porphyry dikes, but are almost entirely on fractures in the metavolcanics. Porphyry sills are near strong lead-zinc mineralization in most cases, and are probably the local source. Sericite, clay veins, and weak epidote and chlorite are associated with the mineralization in the upper 600 feet of the hole. Chlorite and epidote are more abundant, and some potassium feldspar veinlets occur in the lower portion of the hole. Several large gouge zones or faults are present high in the hole, but mineralization tends to overlap and generally disregard them. The hole appears to have intersected an outer lead-zinc halo high and off to the side of the core of the "porphyry" system.

Conoco contracted Mr. Robert Bamford to conduct total sulfide analyses of concentrates from 100 foot average and composited pulps (Oriol, 1977; Bamford, 1977). Bamford completed hole BC-1 in early 1977, and the results are discussed in last year's report (Oriol, 1977) and presented in Appendix E. He believed the data indicated that hole BC-1 was drilled toward the inner zoning of a porphyry sulfide system. This conclusion was drawn from the increasing gradients of bismuth, tellurium, tin and possibly arsenic; and the decreasing gradients of lead, zinc and possibly silver. Copper is fairly uniform, but because molybdenum seems



to decrease downward rather than increasing, he believed we were not directly over the top of the target area but off to one side. X-ray diffraction analyses indicate a similar trend as shown by the presence of potassium feldspar near the bottom of the hole. Conoco's data agrees with Bamford's (Appendix D).

The total sulfide analyses of holes BC-2 and BC-3 in Appendix E were compiled from Bamford's data (Bamford, 1977). They are not as well defined as Conoco's data, but show other elemental trends. We are presently awaiting his results for additional holes to complete the comparison of his and Conoco's data. Both sets of data indicate a westward-tilted porphyry system with the copper target eastward, south-eastward, and/or below hole BC-3 (Figure 1).

BSW-1 - 1313' - BCMC - 1970 and 1974

More than 0.10% copper was encountered in hole BSW-1 in two intervals from about 350 feet to 450 feet and 700 feet to 900 feet. The rest of the hole contained about 400 ppm copper. Lead, zinc, and silver were not as abundant in BSW-1 as in other western and central holes. Strong sericitic alteration is present throughout the hole. Rare secondary biotite has been identified in this section.

BSW-2 - 200' - BCMC - 1970

Bear Creek encountered a chalcocite enrichment blanket and holes BSW-8 (1125') and BSW-2 (200') were drilled specifically to test for it. The best enrichment was about 30 feet of 0.6% copper in BSW-2, although holes BSW-1, 7, and 8 also encountered a thin, enriched zone.



BSW-7 - 1715' - BCMC - 1972

By far the best copper assays are in drill hole BSW-7 near the west side of the prospect (Appendix C). An interval of 510 feet averaged 0.16% copper from 750 to 1260 feet; and 88 feet within that interval averaged 0.35% copper. Fine-grained chalcopyrite is disseminated in faulted and broken metavolcanics. One small dike is within the ore zone, but most of the dikes are just below the high copper interval. Weak secondary biotite and strong silica-sericite accompany the sulfide mineralization. A small zinc anomaly (110 feet of 0.145% zinc) was drilled high above the copper interval. Although no core is now available from this specific interval, sphalerite accounts for the zinc elsewhere at the Buckskin prospect.

BSW-8 - 1125' - BCMC - 1972 and 1973

Hole BSW-8 was drilled to test for enrichment and primary copper to the northwest boundary of the small, strong IP anomaly BCMC had outlined. The hole possessed strong moly (up to 45 ppm for 100') and lead and zinc anomalies at the top of the hole. Copper was less than 600 ppm over 100' intervals except for the top of the hole where minor supergene enrichment exists.

BSW-9 - 1501' - BCMC - 1974

Hole BSW-9 was drilled 1000 feet southwest of BSW-7 to follow up on the high grade interval in BSW-7. Bear Creek drilled there because they believed the mineralization in BSW-7 was related to a west dipping "intrusion breccia". Conoco's mapping indicates that this breccia is a volcanoclastic breccia that is conformable to the rest of the volcanics and only fortuitously acted as a good host for mineralization in BSW-7.



BSW-9 did contain a small, 150 foot interval of about 0.1% copper in a fine biotite hornblende porphyry (Figure 29). A zinc anomaly similar to that in hole BSW-7 is also present high above this copper zone.

Other BSW holes #3, 4, 5, and 6 were shallow rotary holes drilled far to the west in alluvium (Hulen, 1972).

Bul-9 - 734' - Phelps Dodge - 1967(?)

Drill log data for hole Bul-9 was attained from Phelps Dodge in 1977. Assays show a lead-zinc anomaly in agreement with Conoco surface and drill data. No core is available and the Phelps Dodge geologic logs are incomplete and completely unreliable.

Bue-1 - 838' - Phelps Dodge - 1967

Phelps Dodge and Anaconda drilled four holes in the eastern portion of the prospect, all less than 1000 feet deep. Hole Bue-1 (Figure 31) was drilled to 838 feet and averaged about 0.05% copper, except for a small supergene-enriched zone of 0.1% copper and a 150 foot interval of chalcopyrite that assayed approximately 0.15% copper. Lead and zinc decreased gradually until the last 28 feet where zinc increased significantly. Molybdenum also appeared to be increasing for the last 78 feet.

Bue-2 - 893' - Phelps Dodge - 1967

Bue-2 was drilled by Phelps Dodge in 1967 to a depth of 893 feet. It also encountered a 10 foot interval of enriched copper but otherwise averaged a little less than 0.1% copper as chalcopyrite(?) for the entire hole. Mesozoic granodiorites are the predominant host rocks. Pyrite increased to about 6% for the last 100 feet with traces of molybdenum and silver.



D-281 - 446' - Anaconda - 1971

Hole D-281 was a short hole (446 feet) that averaged about 0.05% copper in Mesozoic metavolcanics.

D-289 - 958' - Anaconda - 1972

Anaconda drilled holes D-281 and D-289 in 1971 and 1972. Hole D-289 was drilled to 958 feet in Mesozoic granodiorites intruded by numerous porphyry dikes. Anaconda assayed only certain intervals, unfortunately, and these assays do not show any definite trends. The hole appears to average more than 0.05% copper, however.

The holes in the eastern portion are very intriguing. The Mesozoic phaneritic granodiorites are uniformly high in copper and two holes show apparent downward increasing gradients in lead, zinc and total sulfides. Late Tertiary dikes are well mineralized also, and occasionally appear to control locally higher assays. Most of the mineralization in these holes, however, appears to be genetically related to the older granodiorite.

The soil and subsurface geochemistry of the eastern area poses some interesting problems. The strong copper and moly anomalies would indicate that the core of the system is closer in the eastern area. This hypothesis would first necessitate a fault that has upthrown the east flank. There is an old thrust(?) fault nearby, but the upward movement of the block is structurally anomalous in the region and in the Buckskin Range in particular. Additionally, the strong lead and zinc anomalies are coincident with the copper and moly anomalies, giving a different geochemical signature than observed in holes BC-3 and BC-2 and westward. Alteration also appears to be propylitic. Upward movement should have eliminated the propylitic zone through erosion, however, unless movement was extensive enough to erode to the bottom side of the system.



Another possibility is that these geochemical anomalies are related to an older Yerington-aged mineralizing event such as the intrusion of the granodiorites. Conoco does not presently have enough data to define the system on the eastern flank of the prospect, but will pursue the problem in the future.

Detailed written and graphic logs (WMO) 1" = 20' are available for all holes mentioned. Some written logs by BCMC and Phelps Dodge are also available for the appropriate holes, but the geologic interpretation has changed drastically with Conoco's relogging.



APPENDIX A  
PROJECT COST SUMMARY AND PLANNING 1977

PROPERTY: Buckskin - Douglas & Lyon Counties, Nevada  
Township 13 North, Range 23 & 24 East

		1	2	3	4	5	
		BCMC 1977	Smith & Harcourt 1977	Conoco 1977		TOTAL 1977	
1	1,2,28	7776.54	7095.55	264.09		15617.27	1
2	.26	412.78	390.25	110.7		905.16	2
3	5	1203.56	1036.13	1905.28		4237.00	3
4	14 17 18 3 4 7 8 9	418.22	443.45	482.7		975.90	4
5	65 75	1779.598	1770.56	489.77		11770.0	5
6						38298.80	6
OVERHEAD SUB-TOTAL							
7	40 41 42	64237.38	57507.39			162344.52	7
8	31	6338.30	6338.28			12676.58	8
9	16 19 43 49						9
10	25	7654.46	7017.64			14672.10	10
11							11
12	23 50	3080.5	3080.5			978.10	12
13							13
OPERATING SUB-TOTAL						190671.30	14
15	44 45 46		20000.00			20000.00	15
16	21						16
17	27 11 12 13 14 15 16	103.5	176.0	915.50		943.45	17
18	74	1146.7	445.22	617.11		37121.80	18
19						58065.25	19
LAND SUB-TOTAL							20
21		65670.54	98305.12			170449.38	21
22		40599.75	20000.00			60599.75	22
23		106270.29	118305.12	4613.09		231049.13	23
24							24
25		65725.93	53956.71	-		119682.64	25
26		171996.22	172261.83			350731.77	26
27							27
28							28
29							29
30							30
31	Property Obligation	Lease Payments		Work Requirements			31
32	Owner Claim Name No. Acres Type	Date	Amt.	Year	Notes		32
33	BCMC BSW 47 900 Unpat'd			1978 only	1500' drilling		33
34	Smith & Harcourt Nev & Lorraine 43 700 Unpat'd	1976	353.00		A/W only		34
35	Conoco BC 38 554 Unpat'd	1977	200.00		A/W only		35
36							36
37							37
38							38
39							39
40							40



## APPENDIX B

DETAILED ROCK DESCRIPTIONS  
(from Hudson, 1977)

## MESOZOIC ROCKS

Oreana Peak Formation (T op)

The lower unit of the Oreana Peak Formation consists of at least 1200 feet of interfingering beds of volcanic rocks and conglomerate (T opv) and limestone (T opl) in beds a few inches to 25 feet thick. The volcanics consist of felsic pyroclastics, finely-varved water-lain tuffs, tuffaceous sediments, and dense aphanitic basalts with elongate amygdules up to 20 cm long and 2 cm thick. Coarse conglomerates and massive grey limestones are interbedded with the volcanics.

The overlying carbonate unit (T oc) of the Oreana Peak Formation (Upper carbonate member of Noble, 1962) contains about 200 feet of massive-to-medium bedded grey limestone with 1/2 to 3 inch, (often pinkish) siltstone partings. The limestone locally contains abundant fossil fragments. It is occasionally oolitic and locally dolomitic.

Gardnerville Formation (Jg)

The Gardnerville Formation (Noble, 1962) of uppermost Triassic and Lower Jurassic age, conformably overlies the Oreana Peak Formation. One hundred and fifty to 1200 feet of Gardnerville Formation is exposed in the northern Buckskin Range. The sequence is composed primarily of grey, locally pyritic, ashy(?), calcareous siltstone interbedded with thin, green-grey locally pyritic shale (Jgs). The siltstones break parallel to bedding and contain impressions of ammonites. The Gardnerville contains several discontinuous lensoidal or patchy massive grey limestones



that locally contain "fossil hash" and are locally oolitic. The limestones are from 1 to 20 feet thick. Occasionally the limestones contain thin interbeds of crossbedded calcarenite.

Contact metamorphism of the Gardnerville Formation by adjacent Mesozoic and Tertiary intrusions converts the siltstones to a dense, dark grey hornfels with a fine varved "chert" appearance. The hornfels fractures across bedding. The limestones are strongly recrystallized though rarely marbleized. The break between hornfels and apparently unmetamorphosed Gardnerville can be very abrupt.

#### Artesia Sequence (Jav)

The Artesia sequence of probable uppermost Lower Jurassic age is named for Artesia Mountain in the central Buckskin Range. About 2000 feet of the Artesia sequence is exposed south of Churchill Canyon, while 6000 feet or more of the Artesia sequence metavolcanics is exposed in the southern Buckskin Range. These metavolcanics lie in thrust contact with underlying rocks. The majority of the Artesia sequence is composed of dark lavender, red-lavender, and greenish andesitic to dacitic volcanic flows (Jav) with occasional volcanic breccia and conglomerate-sandstone interbeds. The volcanics are typically porphyritic with 10 to 40% phenocrysts up to 3 mm in length; however, they generally are less than 1 mm long. Phenocrysts of pyroxene(?), hornblende, biotite and weakly-to-strongly zoned plagioclase (An 55-35) occur, with rounded quartz phenocrysts up to 0.5 mm in some flows in the upper portion of the sequence. Some flows completely lack phenocrysts.

The aphanitic groundmass contains plagioclase, hornblende, biotite and rare quartz, where the groundmass is not recrystallized by metamorphism or alteration.



Felsic breccias (Jafb) occur strataconformably in about the middle of the Artesia sequence in the southern Buckskin Range. The white-to-buff felsic unit, which is up to 2000 feet thick, is often flow banded and contains rounded felsic fragments up to 20 cm in diameter. The composition of the unit may be rhyolite or quartz latite.

A thick sandstone and conglomerate unit (Jas) occurs in the upper Artesia sequence in the central and southern Buckskin Range. About 500 feet of the clastic unit is present, in NE 1/4, Sec. 3, T13N, R23E, but thins to the north, south, and east. It is apparently absent in the northernmost Buckskin Range. The clastic unit is coarsest in the west central Buckskin Range with clasts up to 15 cm. The clast size becomes progressively finer as the unit thins. The unit contains rounded clasts of Artesia volcanics grading upward into feldspathic sandstone in usually well-defined graded beds 1 cm to 1 meter thick.

A series of andesitic to dacitic volcanic flows up to 1200 feet thick (Fulstone formation of Anaconda terminology) overlies the clastic unit in the central Buckskin Range. These flows thin rapidly to the north and south and are completely absent in the northern Buckskin Range. These flows are essentially the same composition as the rest of the Artesia sequence except that the phenocrysts are up to 5 mm in length.

Thick, discontinuous, nearly strataconformable quartz "reefs" are a distinctive part of the Artesia sequence in the central and southern Buckskin Range. Many of the quartz "reefs" contain breccias or have protobrecciation. Original textures are usually obscured. The quartz "reefs" rarely occur above the thick clastic unit (Jas), and probably represent epigenetic selective silification associated with local hydrothermal activity.



Churchill Canyon Sequence (Jcc)

The Churchill Canyon sequence, conformably overlies the Artesia sequence. Possibly as much as 3000 feet of Churchill Canyon volcanics is exposed. A thin conglomerate unit (Jccs) usually lies along the contact between the Artesia sequence and the Churchill Canyon sequence. The conglomerate is 1 to 3 feet and rarely up to 7 feet thick. It contains clasts of both Artesia and Churchill Canyon lithologies.

Lithologically, the Churchill Canyon sequence consists of dacitic to quartz latitic flows and crystal-rich tuffs (Jccv) with associated porphyritic intrusions (Jcci) of essentially the same composition. The exposed portion of the sequence contains primarily ash-flow tuffs in the northernmost Buckskin Range and dominately flows in the southern and central Buckskin Range. The ash-flow tuffs and flows interfinger in the north-central Buckskin Range. Both rock types are present throughout the range, although structural and erosional complications expose only a small portion of the Churchill Canyon sequence in the southern Buckskin Range. Quartz phenocrysts up to 10 mm in diameter are distinctive of the Churchill Canyon sequence. The quartz phenocrysts are usually strongly resorbed, rounded and rarely bipyramidal. Not all flows and intrusions contain quartz phenocrysts.

The ash-flow tuffs are usually tan to greenish light brown, containing 20 to 60% phenocrysts. Phenocrysts include 0 to 4% hornblende, 1 to 4% biotite, 0 to 10%(?) sanidine, 1/2 to 7% quartz. The remaining phenocrysts are plagioclase (generally An<sub>35-25</sub>) with weak oscillatory zonation. Crystals are usually broken and often bent. Flattened pumice fragments up to 5 cm long are often seen in thin section. Tuff breccias occur in the sequence and are composed of up to 50% deformed fragments of tuff that are strongly welded into a tuff breccia.



The volcanic flows of the Churchill Canyon sequence are grey to grey-green, and generally are thin bedded (5 to 30 feet thick). Occasional conglomerate or breccia interbeds up to 1-1/2 feet thick are present between some flows. The flows contain 5 to 40% subhedral to anhedral, often broken phenocrysts of 1 to 5% biotite, 0 to 5% oxyhornblende(?), 0 to 4% quartz, weakly-zoned plagioclase (generally An30-40) and occasional large, pink sanidine. Phenocrysts up to 1 cm and rarely 3 cm in length with hiatial (sharp break in size between phenocrysts and groundmass) texture are set in an aphanitic groundmass of plagioclase, sanidine(?), and quartz microlites with possible glass.

Intrusions (Jcci) which are very similar in texture and composition to the Churchill Canyon volcanic rocks, intrude the Churchill Canyon sequence. They are difficult to distinguish from the flows except where crosscutting relationships exist. Where intrusives and volcanics are indistinguishable, they are mapped as undifferentiated (Jccu).

Both the Artesia sequence and Churchill Canyon sequence are probably correlative to the Double Springs Formation of Noble (1962). Ammonites collected from the Double Springs Formation near Topaz Lake indicate an uppermost Lower Jurassic age of the Double Springs Formation (H.F. Bonham, oral comm.). Age dates on hornblende obtained by Castor (1972) from Churchill Canyon sequence type rocks in the Pine Nut Range yielded dates of 146 m.y. and 124 m.y. The age dates would indicate an Upper Jurassic to Lower Cretaceous age for the Artesia and Churchill Canyon sequences. However, the dates may have been reset by subsequent intrusions and/or metamorphism. Although the correlation of the Artesia and Churchill Canyon sequences with the Double Springs Formation is tentative, an uppermost Lower Jurassic age for the sequences is used in this report.



## MESOZOIC INTRUSIONS

Granodiorite (Jgd)

Fine to medium grained, equigranular, phaneritic granodiorite is probably the oldest intrusion exposed in the Buckskin Range. It may be equivalent to the Black Mountain Granodiorite in the Singatse and Wassuk Ranges (E. C. Bingler, oral comm.). The granodiorite contains crystals up to 1 mm in length but they are usually 0.7 mm or less. It contains 1 to 3% biotite, trace hornblende, 8 to 12% interstitial quartz, 10 to 15% orthoclase and the remainder unzoned plagioclase locally varying from An35 to An45. South of the Minnesota Mine a small body of coarser-grained granodiorite (Jgdc) intrudes the fine granodiorite. The rock is essentially the same mineralogically except that it contains about 10% biotite and 1 to 2% hornblende with crystals up to 5 mm.

Diorite (Jd)

Several exposures of black, fine-grained diorite that may be older or younger than the granodiorite crop out along Churchill Canyon. The diorite contains 4 to 10% augite, 1 to 2% quartz, trace orthoclase, and plagioclase (An45-55) in very weakly-zoned crystals.

Quartz Monzonite (Jqmp)

A pinkish-brown quartz monzonite porphyry intrudes the granodiorite just east of the Minnesota Mine (Plate 1). Subhedral plagioclase (oligoclase?) and biotite phenocrysts up to 5 mm long occur in seriate texture with groundmass minerals up to 0.4 mm. Overall, the rock contains 1 to 3% biotite, about 15% quartz, 50 to 55% plagioclase, 30 to 35% orthoclase, 1 to 2% magnetite, and trace pyrite. The quartz monzonite porphyry



is cut by abundant quartz-magnetite-pyrite veins. A coarse-grained quartz monzonite breccia (Jqmb) containing clasts of quartz monzonite porphyry in an aphanitic igneous matrix is exposed just south of the Minnesota Mine (Plate 1).

#### Microgranodiorite (KJmgd)

A small body of buff subequigranular microgranodiorite intrudes the Artesia sequence west of Fulstone Spring No. 1 (Plate 1).

Plagioclase is up to 1 mm long, but the crystals are generally less than 0.4 mm. The rock contains 15 to 20% anhedral quartz, 10 to 15% anhedral orthoclase, and subhedral to anhedral unzoned plagioclase, (An35-40). The exact age is unknown except that it is upper Mesozoic.

#### Churchill Canyon Intrusives (Jcci)

The Churchill Canyon intrusives are discussed under Churchill Canyon sequence.

#### Latite Porphyry Dikes (KJlp)

Latite porphyry dikes intrude the older Mesozoic rocks throughout the Buckskin Range. They contain subhedral 1 to 3 cm K-feldspar phenocrysts with small hornblende, biotite, plagioclase, and about 1% quartz phenocrysts in an aphanitic groundmass. Several dikes in the northern Buckskin Range are essentially identical in texture and mineralogy, but contain about 5% quartz phenocrysts (Kqlp). The dikes are similar in composition to Churchill Canyon sequence intrusions, but they are distinctive enough in appearance to map separately. The latite porphyry dikes could be a later phase of Churchill Canyon intrusive activity, based upon intrusive relationships, but may be considerably younger.



## CENOZOIC ROCKS

The Oligocene is represented by a series of ash-flow tuffs and tuffaceous sediments which are thoroughly discussed by Proffett and Proffett (1976) from the Singatse Range. Brief descriptions are given here to note local variations in the tuff units.

Guild Mine Tuff (Tgmt)

The Guild Mine Tuff is a lavender buff, moderately-welded, crystal-rich ash-flow tuff. It contains 20 to 35% phenocrysts, up to 2.5 mm in length, of quartz, sanidine, plagioclase, and 1 to 3% biotite. Rare fragments of petrified wood occur at the base of the Guild Mine Tuff, overlain by 5 to 25 feet of black vitrophyre. The vitrophyre is overlain by 700 to 900 feet of moderately-welded and occasionally densely-welded tuff with moderately abundant small pumice fragments. The top of the Guild Mine Tuff consists of 200 to 300 feet of buff vapor phase tuff with abundant large pumice fragments and frequently large irregular blocks of deep red jasper.

Weed Heights Tuff (Twht)

The Weed Heights Tuff is a lavender to reddish brown, moderately-welded, moderately-crystal rich ash-flow tuff. The tuff contains 5 to 25% phenocrysts including plagioclase, sanidine, quartz, and 1 to 2% biotite. Abundant large white pumice fragments are distinctive of the unit. There appears to have been a very short cooling break between the Guild Mine Tuff and the Weed Heights Tuff in the Buckskin Range as evidenced by the lack of intervening sedimentary units which are present in the Singatse Range (Map Units 4 and 5 of Proffett and Proffett, 1976).



Fifty to 75 feet of rhyolitic sediments overlie the Weed Heights Tuff and are mapped with it. The lower portion is usually bright red, whereas the upper portion is yellowish green. The uppermost sediments contain abundant fragments of petrified wood and sparse leaf impressions.

#### Singatse Tuff (Tst)

The Singatse Tuff is a brown to red-brown, strongly to moderately welded, crystal-rich ash-flow tuff. The tuff contains 30 to 45% phenocrysts of plagioclase, quartz, sanidine, hornblende, and 3 to 6% biotite up to 5mm in length. Hornblende, abundant biotite, large phenocrysts, and its massive nature are distinctive of the Singatse Tuff. The unit contains sparse pumice fragments and 1 to 4% lithic fragments except local zones slightly above the base where the unit contains up to 30% lithic fragments. Five to 10 feet of black vitrophyre occurs locally at the base.

#### Nine Hill Tuff (Tnht)

The Nine Hill Tuff is part of the Bluestone Mine Tuff of Proffett and Proffett (1976). It is a deep lavender to buff, poorly-welded to non-welded, crystal-poor ash-flow tuff and tuff breccia about 100 feet thick. The tuff contains less than 5% phenocrysts of plagioclase, quartz, sanidine, and trace quantities of biotite. White flattened pumice fragments up to 20 cm across are distinctive of the unit. Locally, the Nine Hill Tuff is a tuff breccia.

#### Eureka Canyon Tuff (Tect)

Fifty feet (?) of bright, red-orange, poorly to non-welded, crystal poor tuff of the Eureka Canyon Tuff of E. C. Bingle (oral comm.) conformably overlies the Nine Hill Tuff. The Eureka Canyon Tuff is also part



of the Bluestone Mine Tuff of Proffett and Proffett (1976). The tuff contains sparse pumice fragments and less than 5% phenocrysts of plagioclase, quartz, sanidine, and trace biotite.

#### Andestic Volcanics (Tv)

Miocene andesitic volcanic flows and flow breccias overlie the Oligocene ash-flow tuffs in angular unconformity. More than 1000 feet of andesite is present in north Churchill Canyon, while only small exposures crop out on the east flank of the Buckskin Range. More than 5000 feet of andesite is present north of Lincoln Flat (Proffett and Proffett, 1976). Many of the flows are oxyhornblende and augite-rich andesitic porphyries (Tvhp) and probably were derived from a volcanic intrusive center of the same composition and texture in the northern Buckskin Range. A few flows of possibly younger oxyhornblende biotite andesite porphyry (Tvhp) crop out south of the Minnesota Mine.



## TERTIARY INTRUSIONS

Relative ages of numerous Tertiary dikes and stocks are given where known, but otherwise the Tertiary intrusions are assumed to be Miocene based upon intrusive relationship and age dates. Most intrusives in the southern Buckskin Range are described in the main text.

Biotite Pyroxene Diorite (Tbpd)

A few small stocks of dense, black, fine-grained diorite are present at the north end of the Buckskin Range (Plate 1). The diorite contains 5 to 7% shreddy, greenish-brown biotite less than 0.05 mm in length, usually in magmatic reaction with augite, 7 to 10% subhedral augite up to 2 mm long, less than 2% quartz, weakly-zoned subhedral plagioclase (An<sub>50-55</sub>) up to 2 mm in length, and about 1% magnetite. The rock shows strong foliation in thin section, but it is not evident in hand sample.

Hornblende Pyroxene Diorite Porphyry (Thdpd)

Phaneritic diorite porphyry and locally porphyritic diorite underlies the north end of the Buckskin Range and intrudes the biotite pyroxene diorite. Phases of the diorite vary widely in texture and composition and are generalized here. The rock contains 10 to 40% phenocrysts of hornblende that are commonly 5 to 10 mm in length but range up to 3 cm in length. Plagioclase phenocrysts are up to 5 mm in length and rarely reach 10 mm. The phenocrysts grade downward in size in a seriate texture to a groundmass generally 0.01 to 0.5 mm. The diorite contains 5 to 25% euhedral to subhedral green to green-brown zoned hornblende with numerous inclusions, 3 to 7% subhedral augite, 2 to 3% quartz, oscillatory normally-



zoned plagioclase (An50-40), and trace sphene and apatite. The diorite is often foliated with subparallel plagioclase and hornblende phenocrysts.

Numerous andesitic dikes and small stocks related to the hornblende pyroxene diorite porphyry crop out in the northern portion of the Buckskin Range. They have variable compositions and highly variable textures. These have been generalized into two groups: those with less than 25% phenocrysts (Thppa), and those with 25 to 50% phenocrysts (Thpap). The rocks are similar in composition to the hornblende pyroxene diorite porphyry but have an aphanitic groundmass of plagioclase, augite, and a small amount of quartz, and phenocrysts of plagioclase, hornblende and usually augite. These intrusives are green to greenish-grey and many have weakly developed columnar jointing.

Several intrusive breccias (Timb) crop out in Sec. 22, T14N, R23E. These have a matrix of hornblende pyroxene porphyritic andesite with angular to highly rounded pebbles to boulders of hornblende pyroxene porphyritic andesite, hornblende pyroxene andesite and diorite porphyry, various ash-flow tuffs, and Mesozoic metavolcanic rocks as well as granitic rocks not exposed in the Buckskin Range.

The hornblende-pyroxene andesite and diorite porphyries, hornblende pyroxene porphyritic andesites, and intrusive breccias form a large Miocene intrusive complex in the northern Buckskin Range. This diorite-andesite intrusive complex is Miocene because dikes of the complex intrude Tertiary ash-flow tuffs across Basin and Range faults. Some of the faults have post-intrusive displacement, however. The intrusives may have volcanic equivalents in the Lincoln Flat Andesite (17 to 10 m.y.), Proffett and Proffett (1976) but may be younger, because most of the Basin and Range faults cut by the intrusives do not show later movement.



### Biotite Hornblende Andesite Porphyry (Tbhap)

Several dikes of coarse-grained biotite hornblende andesite porphyry intrude the older rocks in and around the Minnesota Mine. The andesite contains 45 to 50% phenocrysts of euhedral biotite up to 1.2 cm across in books up to 1 cm thick, euhedral hornblende up to 5 mm long, and subhedral plagioclase to 1 cm in length. The phenocrysts are in hial texture with an aphanitic groundmass of glass, plagioclase, microlites, and quartz. The rock contains 7 to 10% biotite, 2 to 3% hornblende, 2 to 4% quartz, 5 to 10% brown glass, and weakly-oscillatory normally-zoned plagioclase (An43-38). The biotite hornblende andesite porphyry is apparently younger than the sheared hornblende pyroxene andesite porphyry exposures in the Minnesota Mine.

### Porphyritic Hornblende Dacite (Tphd)

An intrusive porphyritic hornblende dacite crops out extensively in the southern Buckskin Range. The unit contains 15 to 25% phenocrysts of plagioclase and hornblende up to 1 cm and rarely up to 3 cm in length, although locally as much as 40% phenocrysts may be present. The plagioclase usually has epidotized cores. Locally the dacite contains rounded quartz "eyes" up to 3mm (Tphdp) and locally occurs as an autobreccia (Tphdb). One to 2% biotite occurs locally in the dacite in Sections 2, 3, 10, and 11, T13N, R23E (Plate 1). The rock contains 2 to 8% hornblende, up to 10%(?) K-feldspar, 0 to 7% quartz, and weakly-zoned plagioclase (An30-35?). The groundmass contains K-feldspar and plagioclase microlites with or without quartz and hornblende.

Conflicting evidence indicates possibly both Tertiary and Mesozoic ages for the porphyritic hornblende dacite. In exposures in Sections 11, 12, 13, and 14, T13N, R23E, (Conoco prospect) the dacite lies along low



angle Basin and Range faults with very altered Tertiary ash-flow tuffs above the dacite and highly-altered Mesozoic metavolcanic rocks below. The dacite is only weakly propylitized. The difference in alteration suggests a late Tertiary age for the dacite in these outcrops. In Sections 2 and 3, N1/2 Sec. 15, NW1/4 Sec. 11, and NE1/4 Sec. 10, T13N, R23E (Plate 1), the dacite is locally highly sericitized, intruded by highly-altered Tertiary dikes (N1/2, NE1/4, Sec. 10, T13N, R23E), and apparently may not intrude the Tertiary ash-flow tuffs. This evidence, along with its similar appearance to the intrusives of the Churchill Canyon sequence, indicates a Mesozoic age for these porphyritic hornblende dacite occurrences. Most of the dacite is probably an intrusive portion of the Churchill Canyon sequence. Some of the texturally-identical dacite is Tertiary in the prospect area. The Tertiary and Mesozoic dacites are indistinguishable in the field.

#### Hornblende Biotite Andesite Porphyry (Thbap)

Several intrusions of grey hornblende biotite andesite porphyry crop out in the southern Buckskin Range. Phenocrysts make up 25 to 35% of the rock with 4 to 6% hornblende with thin magnetite rims, 4 to 5% biotite, and oscillatory-zoned subhedral to anhedral plagioclase (An42-48). The phenocrysts are up to 1 cm in length and occur in an aphanitic groundmass of plagioclase, 1 to 4% mafics, up to 4% quartz, 1 to 2% magnetite and trace apatite. The rock generally is propylitized and rarely is sericitized. A large mass of hornblende biotite andesite porphyry in Sec. 2, T13N, R23E (Plate 1) may be either extrusive or intrusive, although contact relationships, where seen, seem to indicate that it is extrusive.



Biotite Hornblende Sanidine Dacite Porphyry (Tbhsdp)

Several biotite hornblende sanidine dacite porphyry dikes crop out on the southern tip of the Buckskin Range. Phenocrysts make up to 60% of the rock, in seriate texture. They consist of subhedral oscillatory normally-zoned plagioclase (An52-39) up to 10 mm long, rounded sanidine up to 5 mm making up 2 to 3% of the phenocrysts, 4 to 6% biotite up to 5 mm across, and 4 to 6% strongly-zoned oxyhornblende up to 4 mm long with magnetite rims. The aphanitic groundmass contains 1 to 2% sanidine, plagioclase microlites, and minor amounts of glass.

Hornblende Andesite Porphyry (Thap)

Dikes of hornblende andesite porphyry intrude along faults in the southern Buckskin Range. The dark green andesites contain 30 to 40% phenocrysts, in seriate texture, composed of 30 to 35% hornblende up to 10 mm long and 10 to 20% subhedral plagioclase up to 15 mm in length.

Basalt (Tb)

A few black, dense, aphanitic basalt dikes crop out in the southern Buckskin Range. These are probably the youngest intrusives in the range, but their exact age is uncertain.



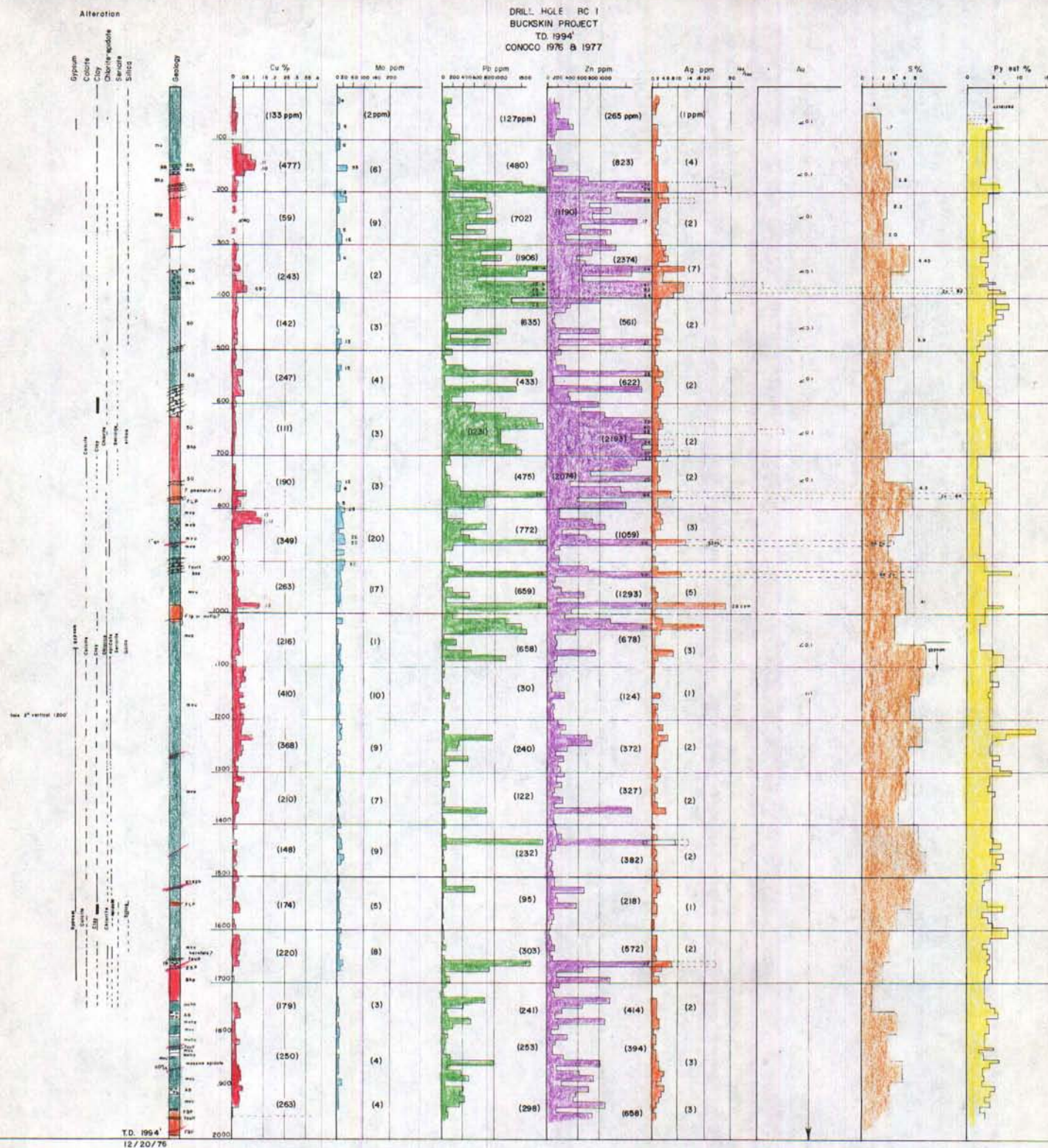


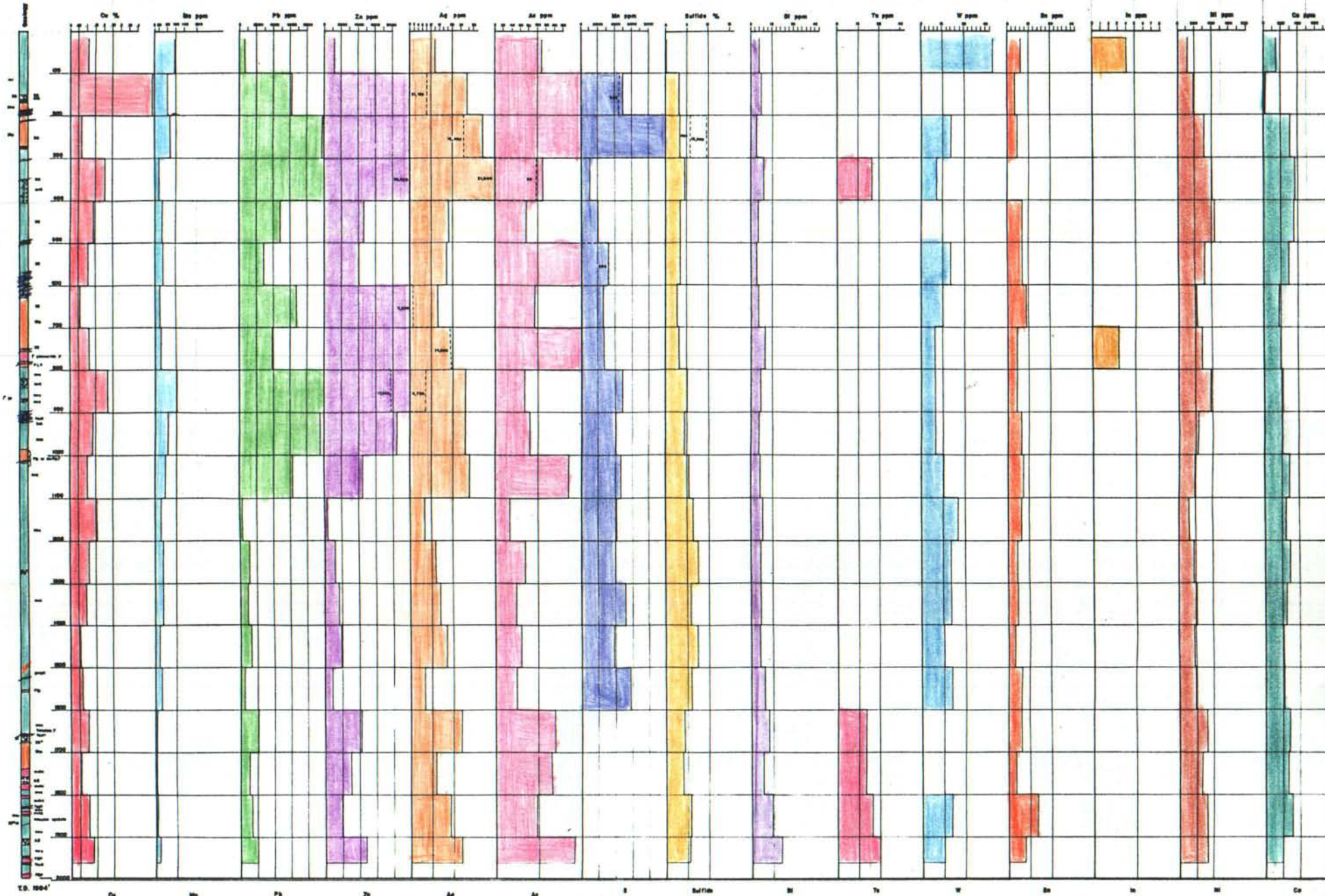
FIGURE C-1



DRILL HOLE BC 1  
SULFIDE CONCENTRATE ANALYSES  
DATA BY BAMFORD MARCH 1977  
BUCKSKIN PROJECT  
T.D. 1994'

Qualitative XRD  
(100' composite)

Pyrite  
Sphalerite  
Galena  
Copper  
Zinc  
Lead  
Silver  
Bismuth



C-2



Alteration

Oxym  
Calcite  
Chalcopyrite  
Sericite  
Biotite

100' Averages in ( ) ppm

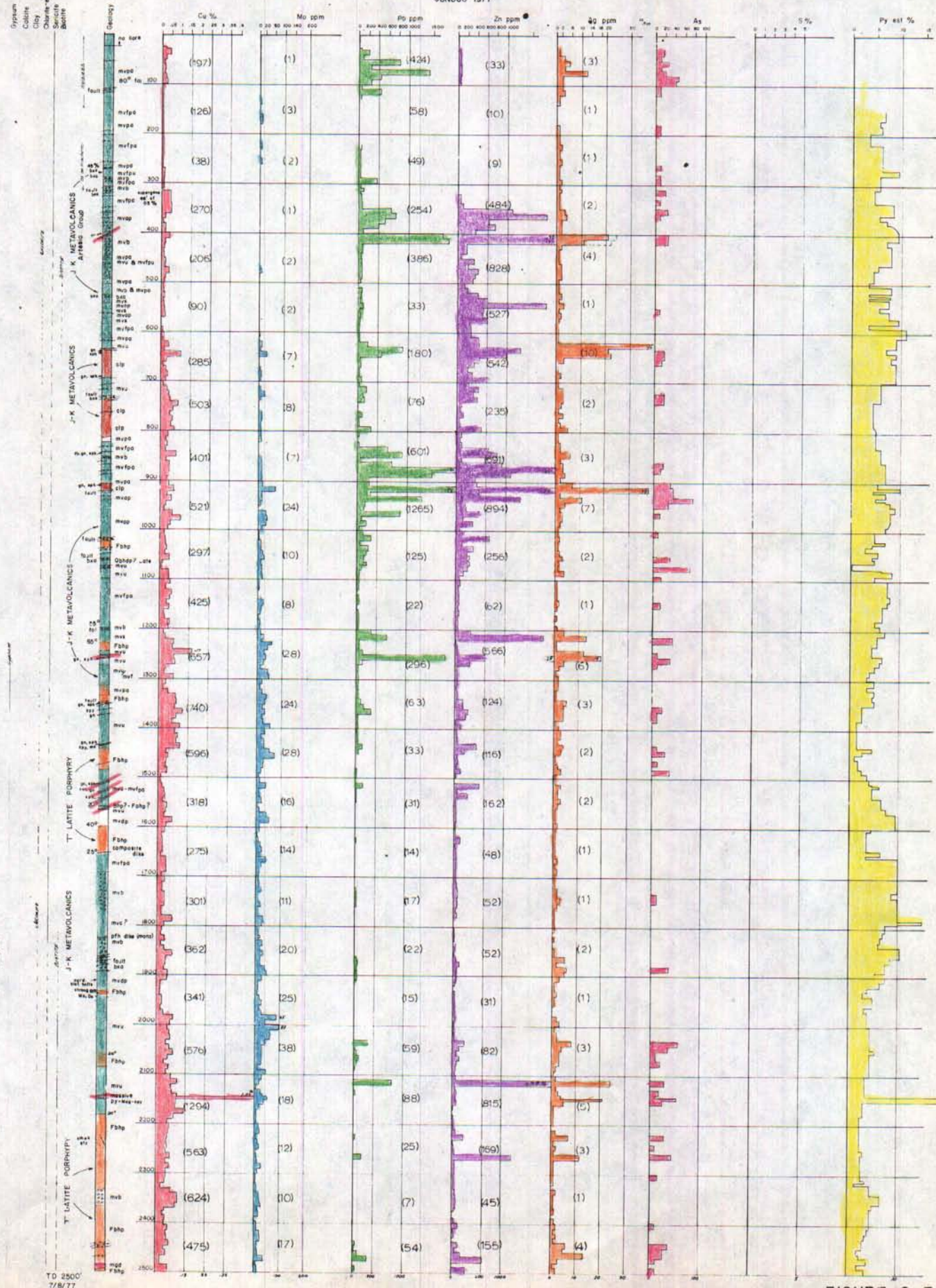
DRILL HOLE BC 2  
BUCKSKIN PROJECT  
TO 2500' 7/8/77  
CONOCO 1977

FIGURE C-3



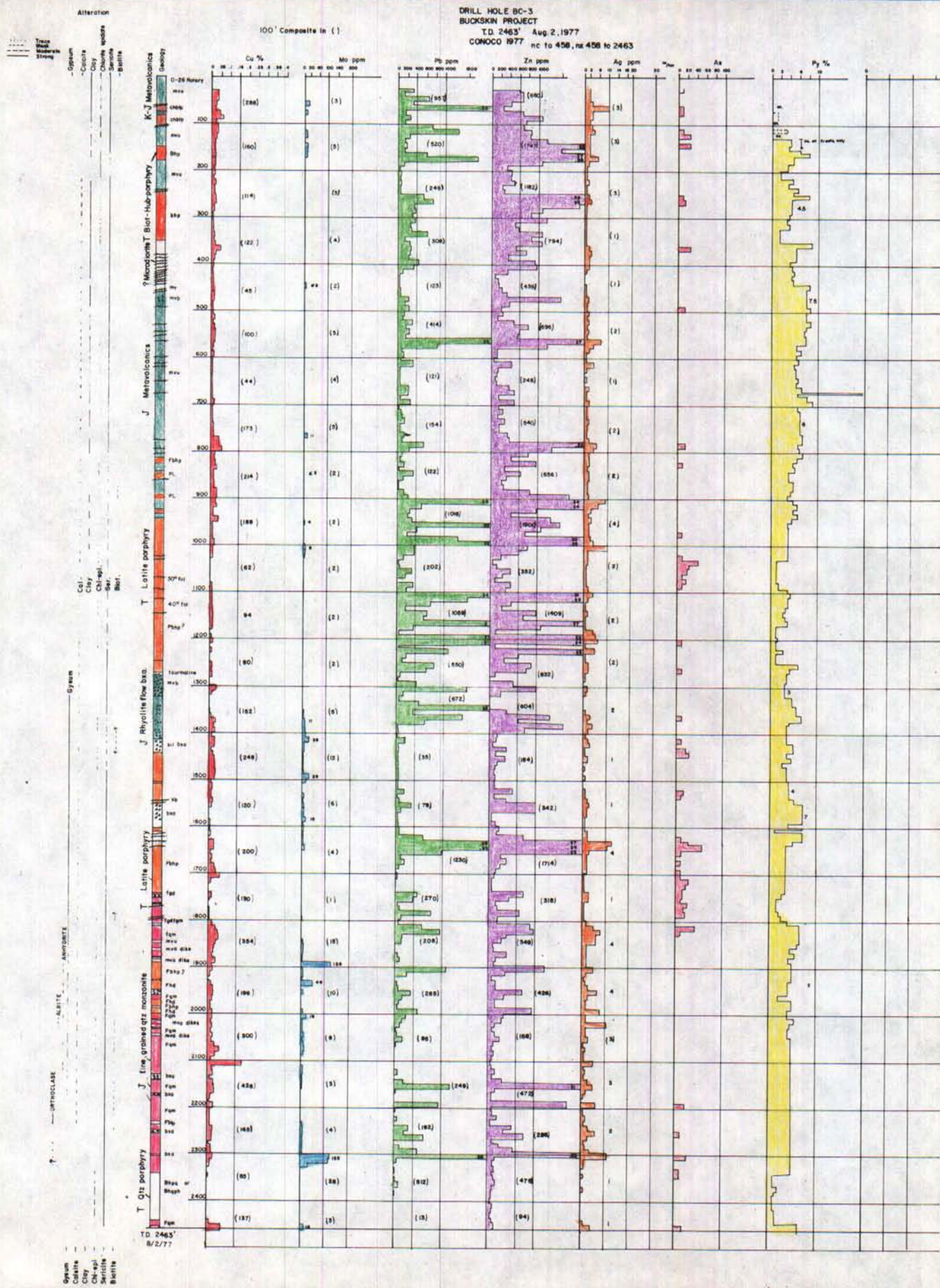


FIGURE C-4



DRILL HOLE BSW 1  
BUCKSKIN PROJECT  
T.D. 1313'  
BCMC 1974

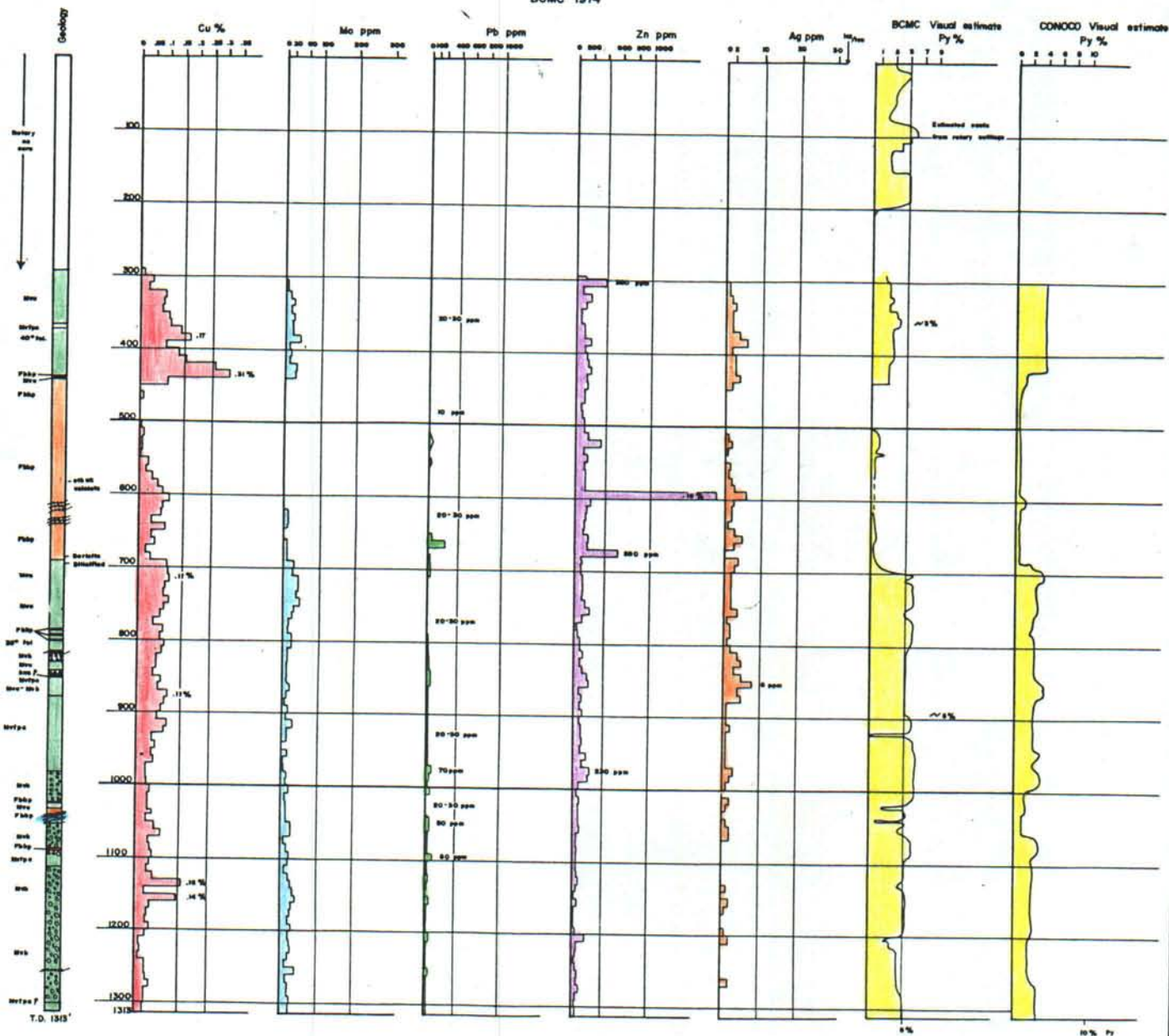


FIGURE C-5



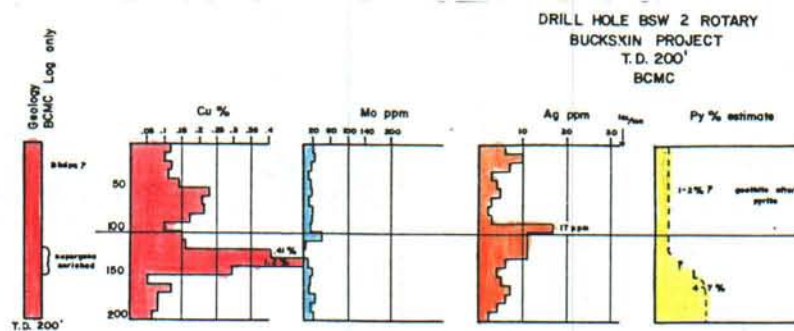


Figure c-6



DRILL HOLE BSW 7  
BUCKSKIN PROJECT  
T.D. 1715'  
BCMC 1972

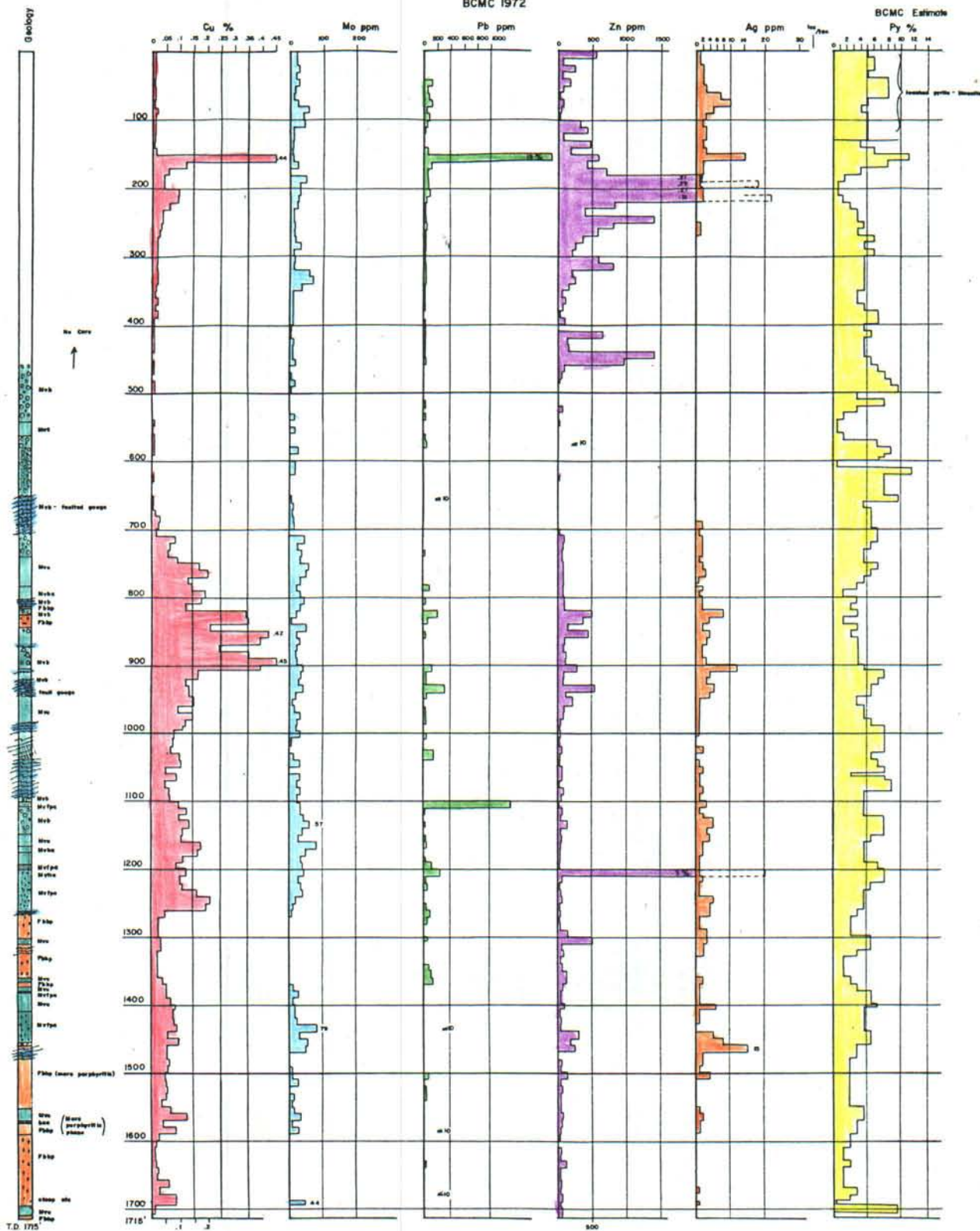


Figure C-7



DRILL HOLE BSW 8  
BUCKSKIN PROJECT  
TD 1125'  
BCMC 1972 & 1973

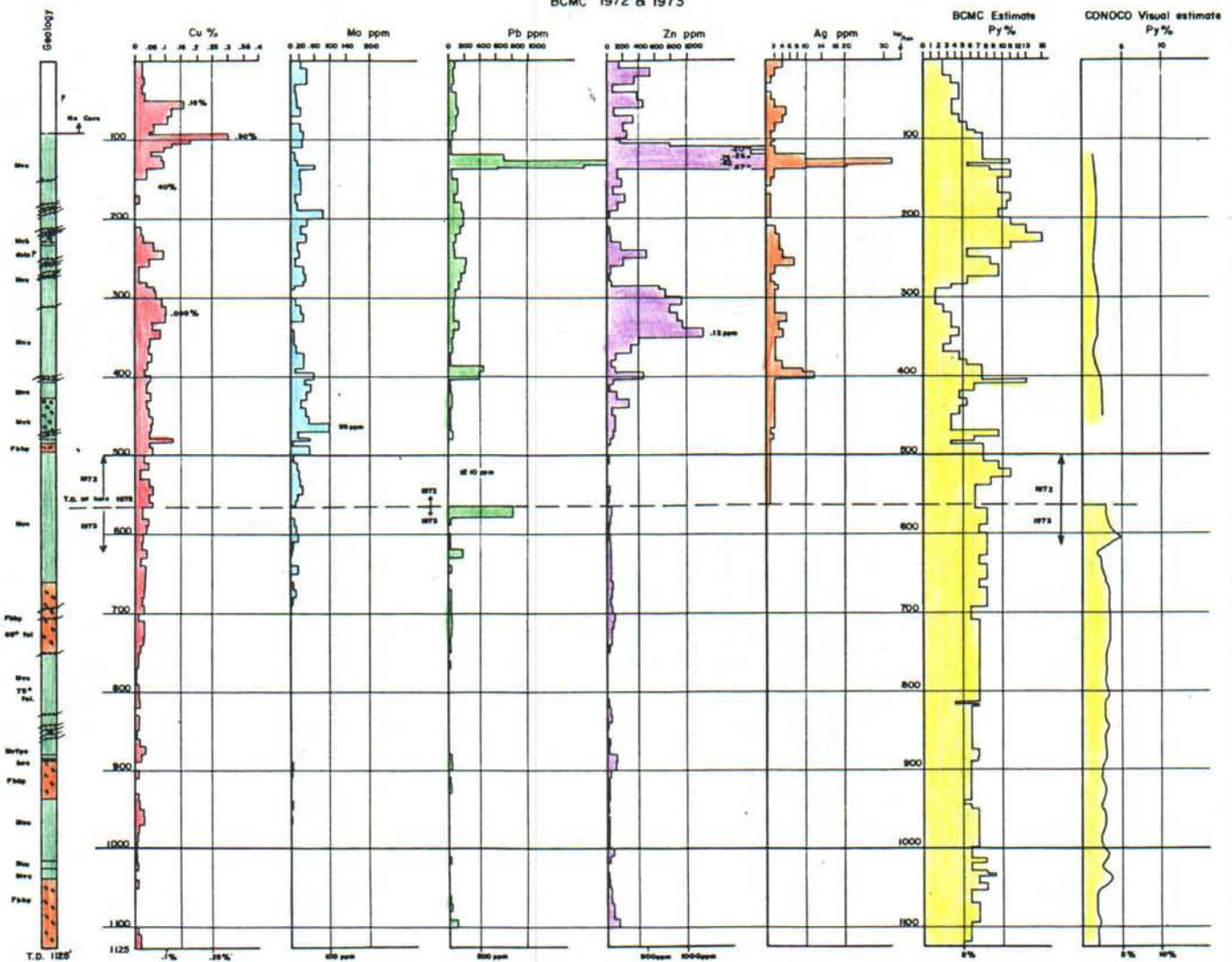


Figure C-8



DRILL HOLE BSW 9  
BUCKSKIN PROJECT  
T.D. 1501'  
BCMC 1974

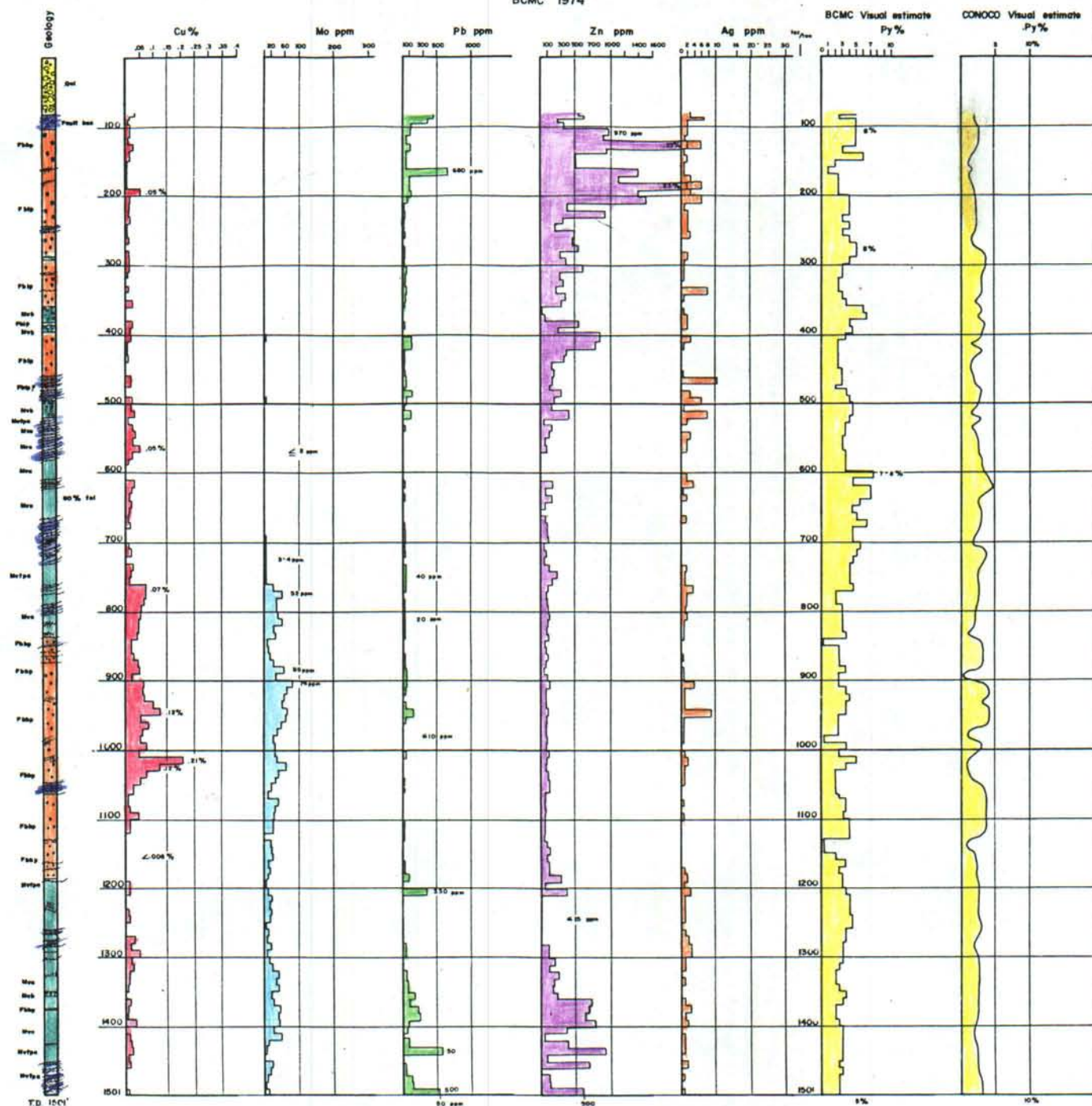
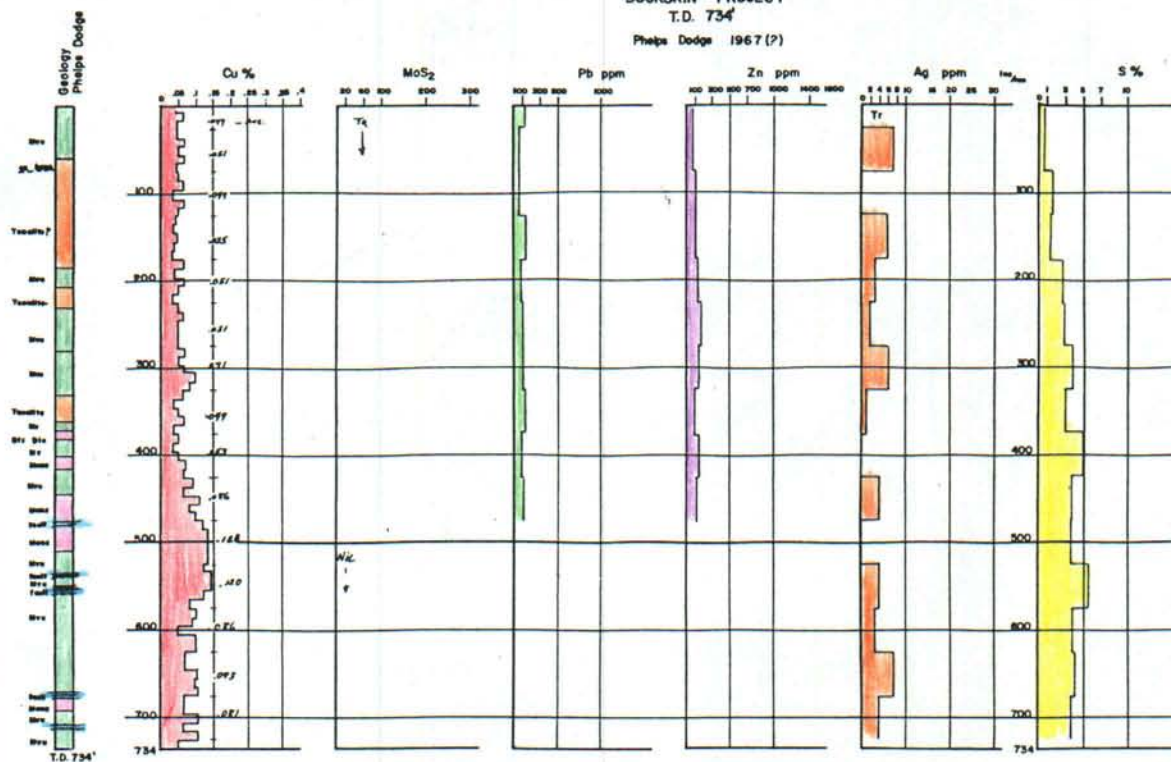


Figure c-9



DRILL HOLE BUL 9  
BUCKSKIN PROJECT  
T.D. 734'  
Phelps Dodge 1967(?)



NOTE - No core available for inspection.  
Hence no correlation can be made  
with other holes. Rock names  
remain as originally described by  
P.D. in 1967.

Figure c-10



DRILL HOLE BUE I  
BUCKSKIN PROJECT  
T.D. 828'  
1967 P.D.

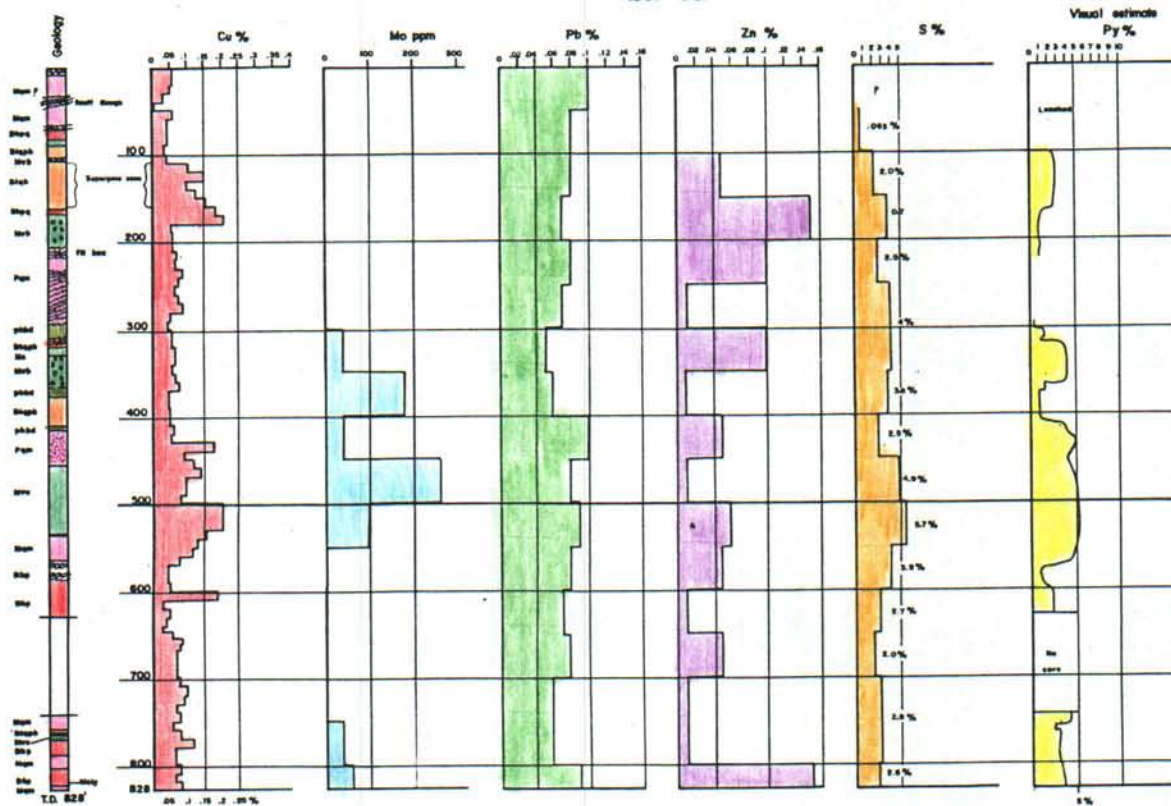


Figure c-11



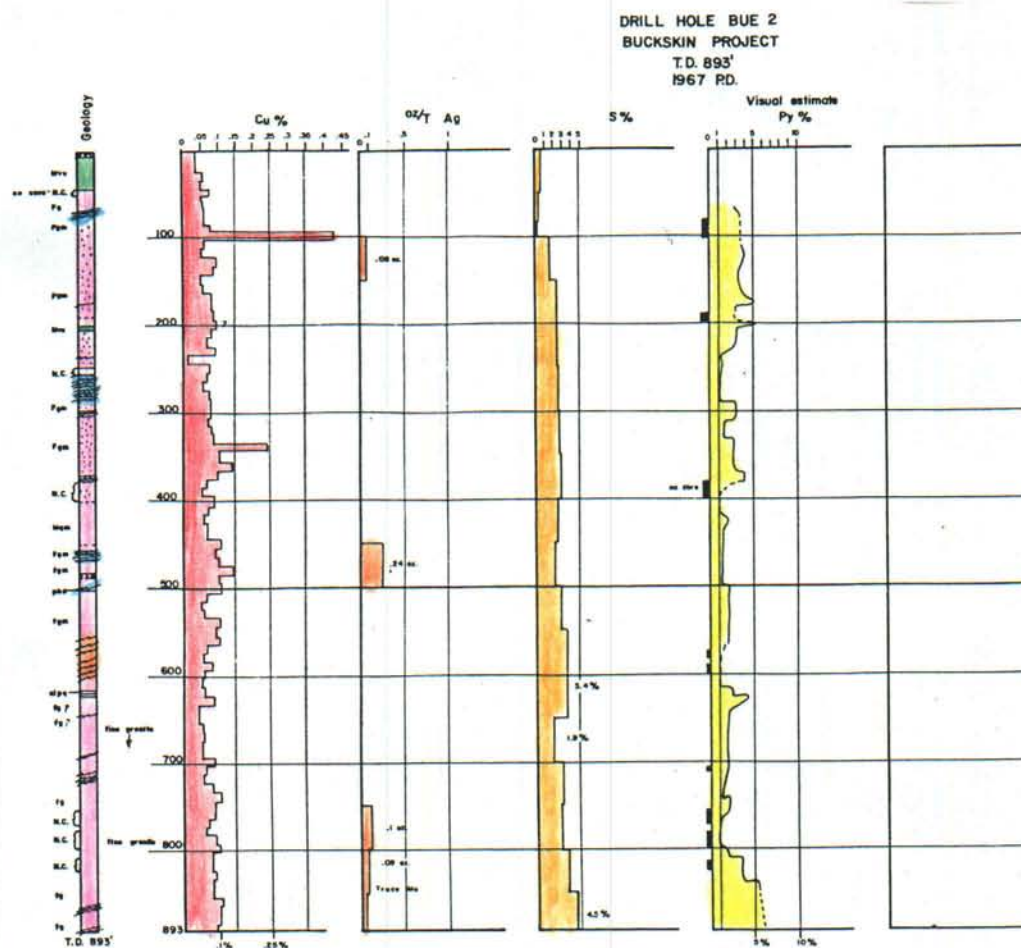


Figure c-12



DRILL HOLE D-281  
BUCKSKIN PROJECT  
T.D. 446'  
1971

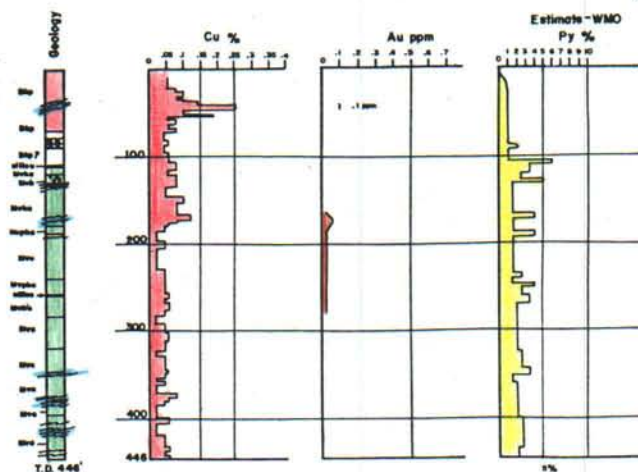


Figure c-13



DRILL HOLE D-289  
BUCKSKIN PROJECT  
T.D. 958'  
1972 ACM

Estimate - WMO  
Py %

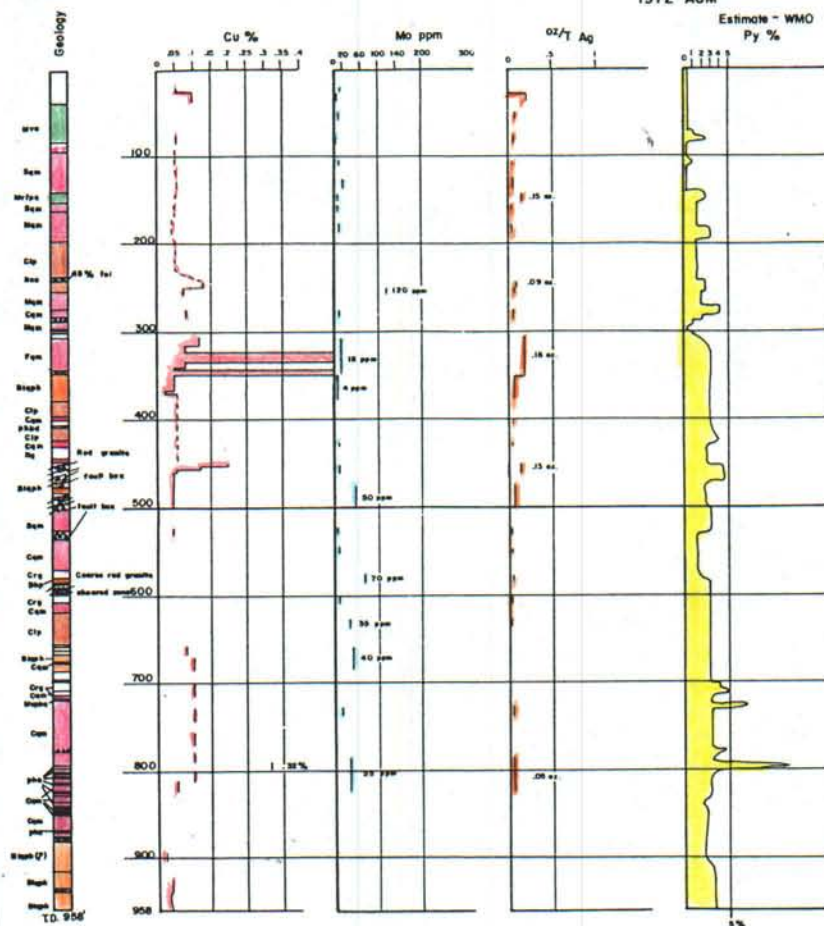
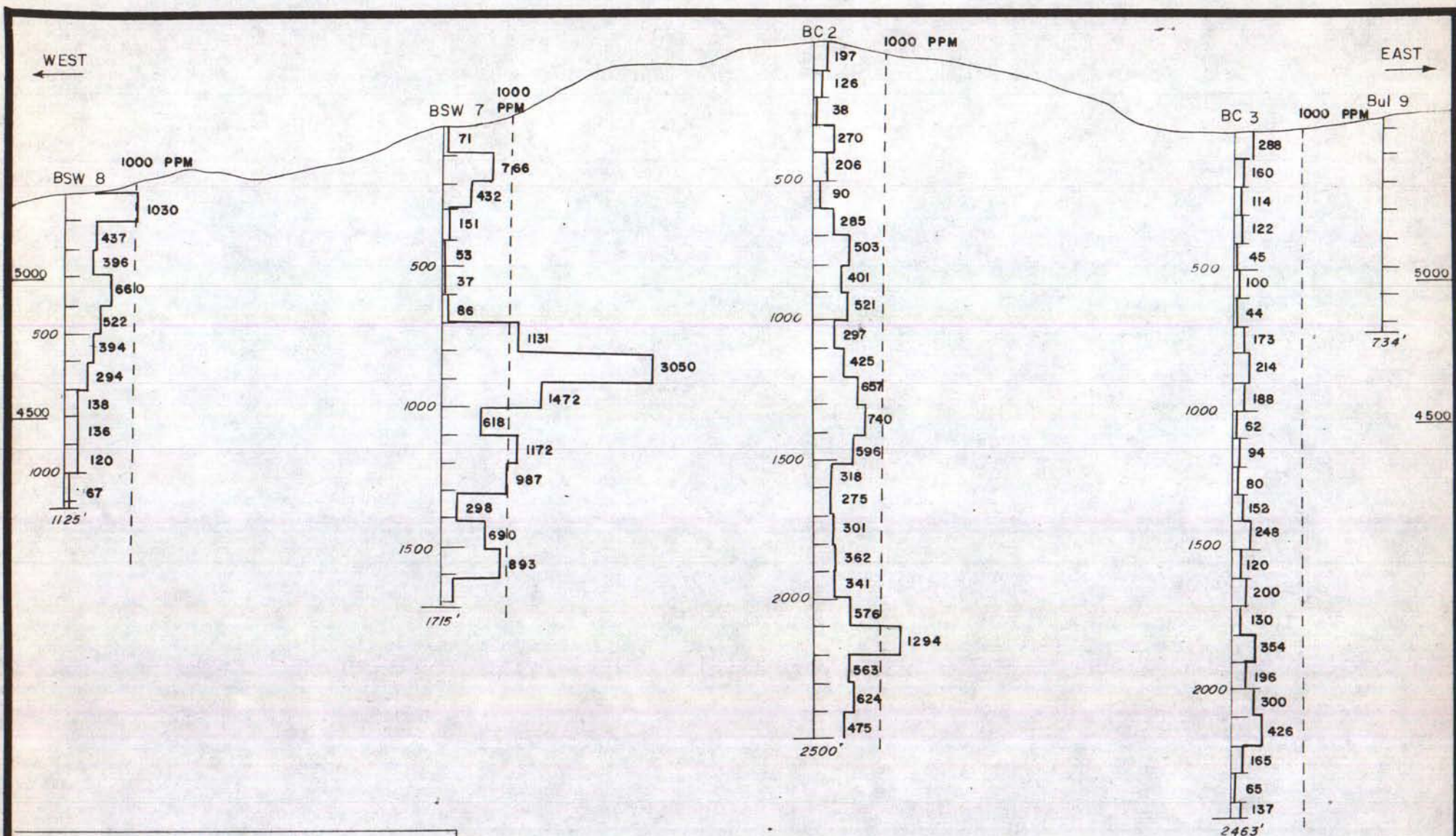


Figure c-14





BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 8, 7, BC-2, 3  
 CONOCO COMPOSITE ASSAYS-100'

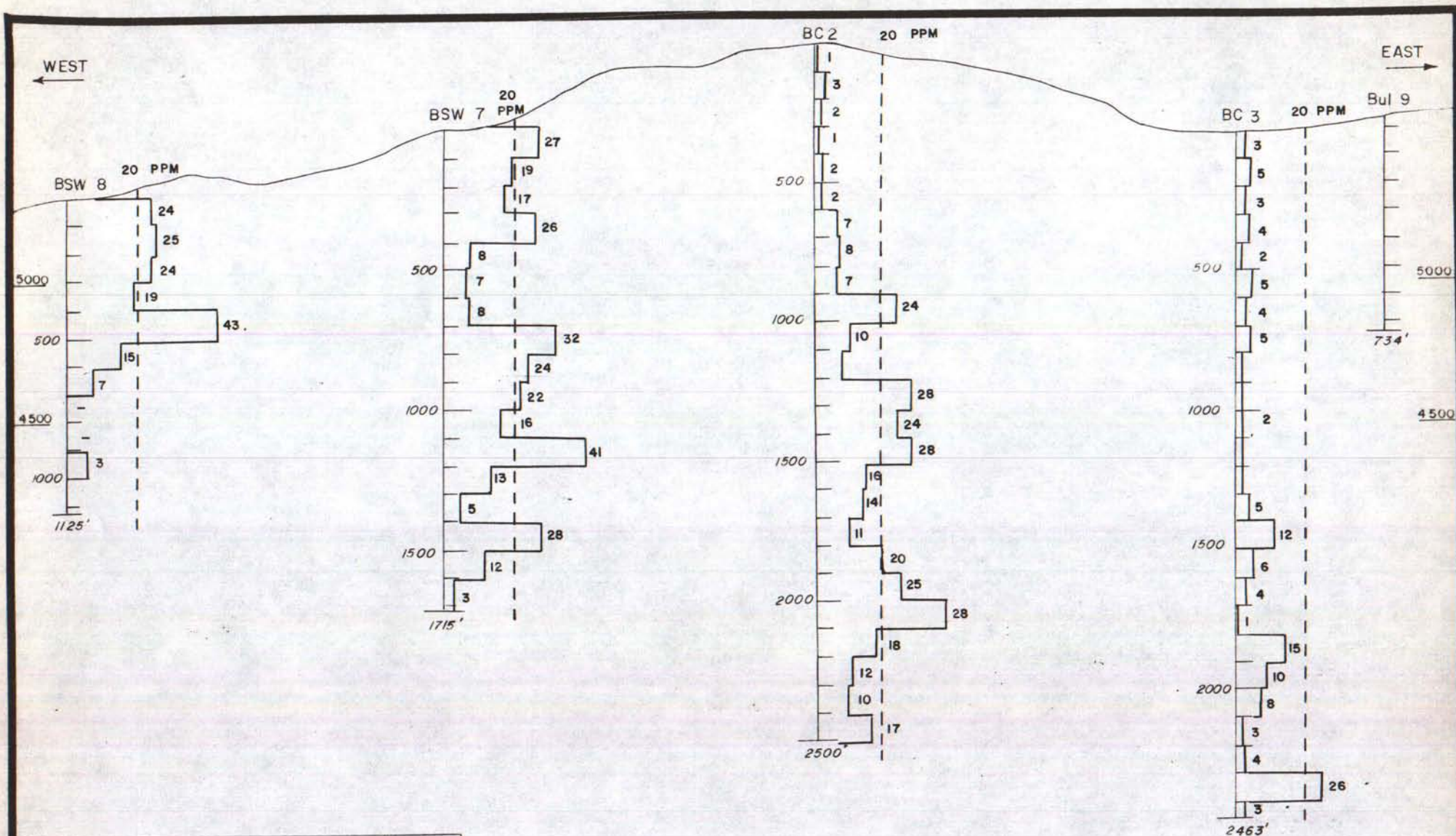
**Cu**

APPENDIX D  
 FIGURE D-1

0 500'

WMO 3/78





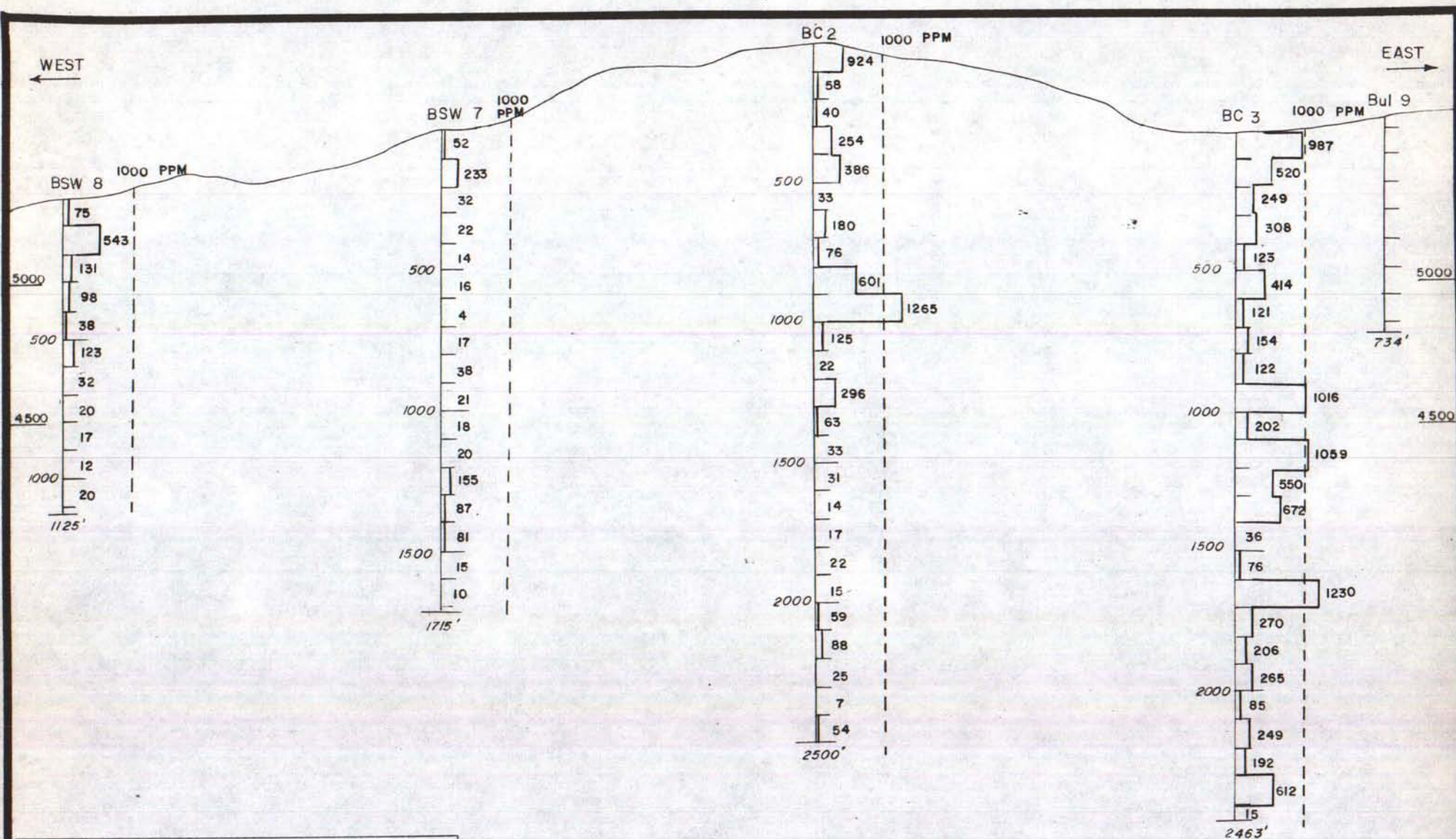
BUCKSKIN PROJECT  
Drill Hole Cross Section  
BSW 8, 7, BC-2, 3  
 CONOCO COMPOSITE ASSAYS-100'

APPENDIX D  
 FIGURE D-2

**Mo**  
 0 500'

WMO 3/78





BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 8, 7, BC-2, 3  
 CONOCO COMPOSITE ASSAYS - 100'

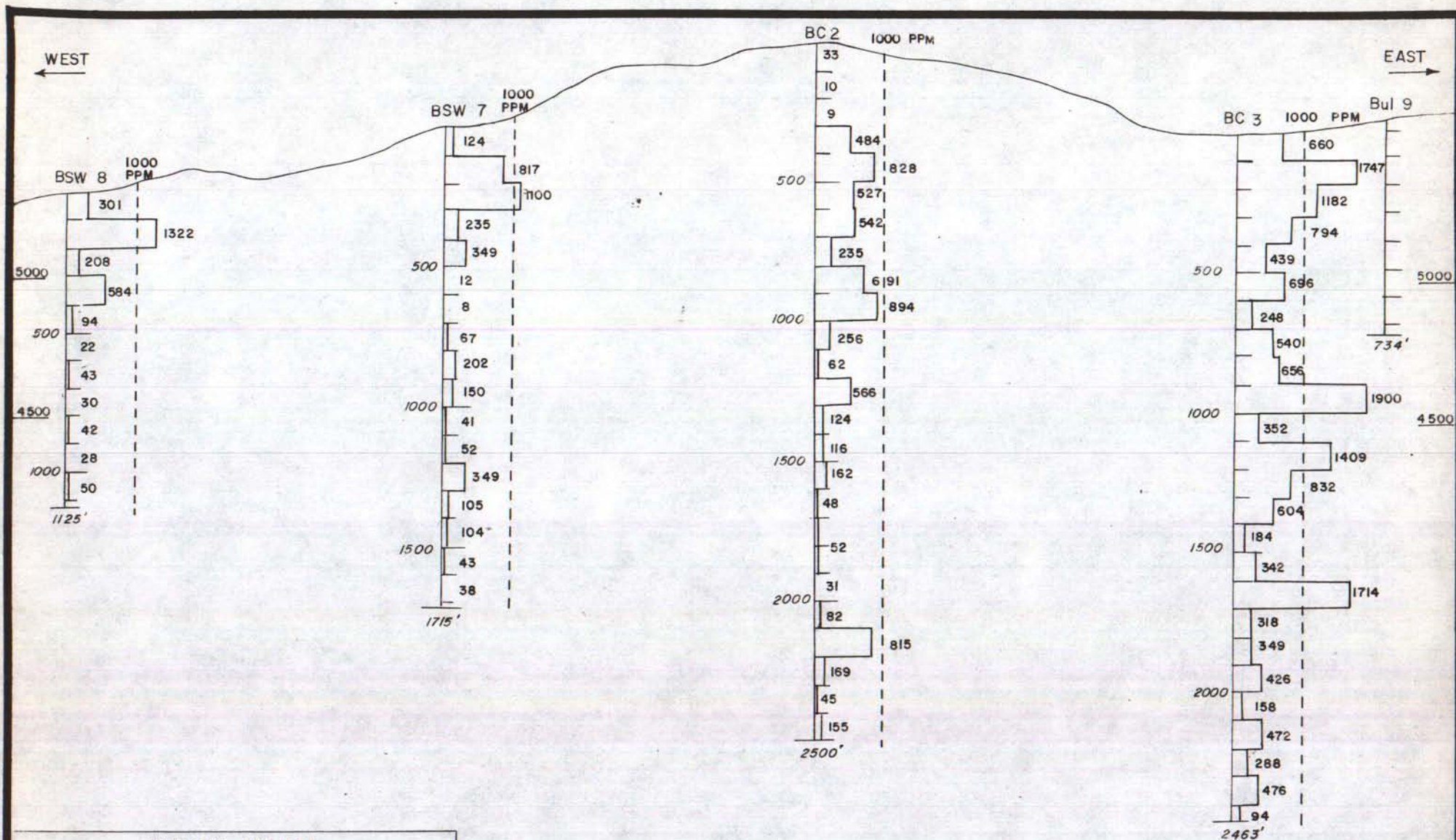
**Pb**

APPENDIX D  
 FIGURE D-3

0 500'

WMO 3/78





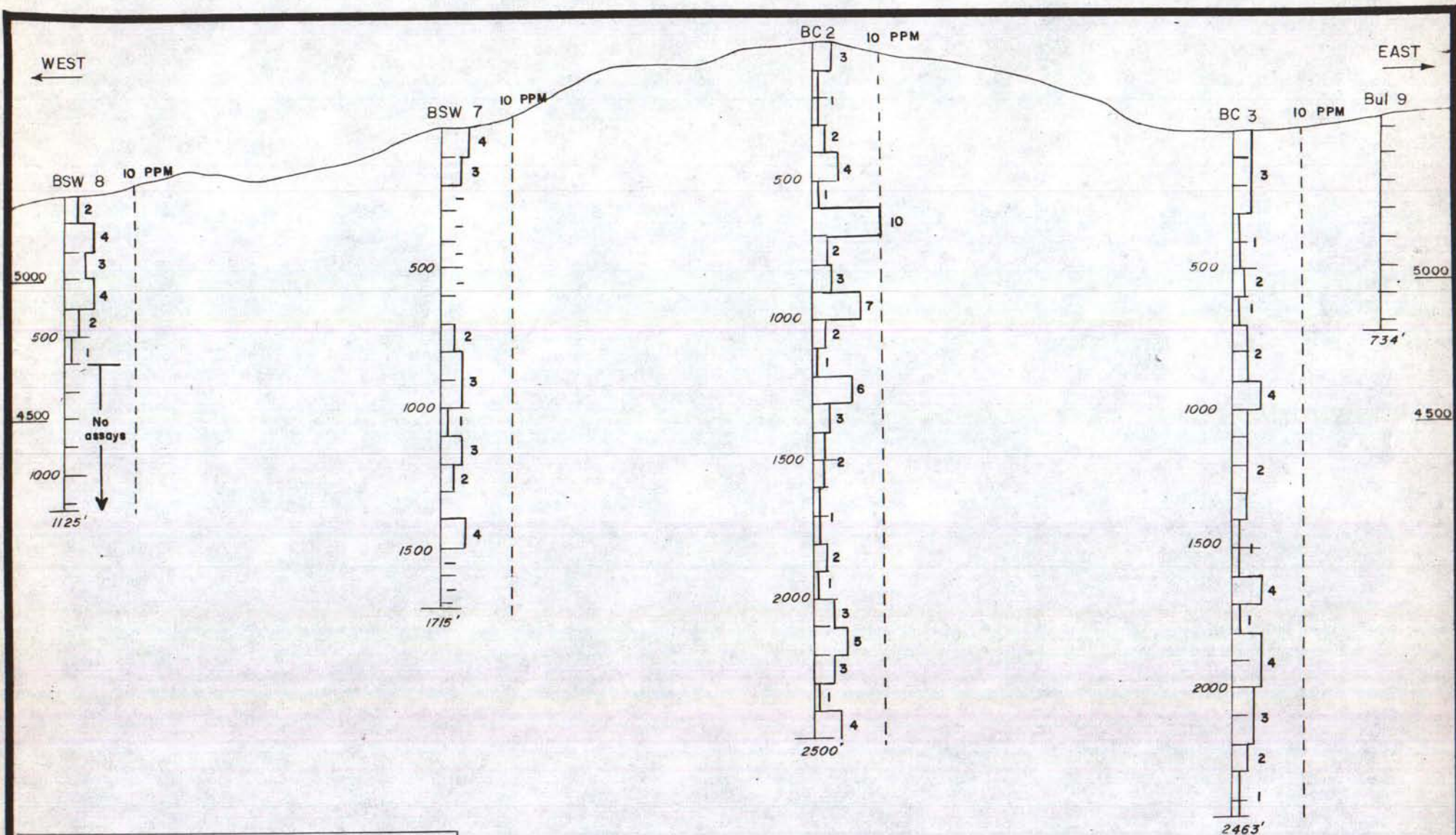
BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 8, 7, BC-2, 3  
 CONOCO COMPOSITE ASSAYS-100'

APPENDIX D  
 FIGURE D-4

**Zn**  
 0 500'

WMO 3/78





BUCKSKIN PROJECT

Drill Hole Cross Section

BSW 8, 7, BC-2, 3

CONOCO COMPOSITE ASSAYS-100'

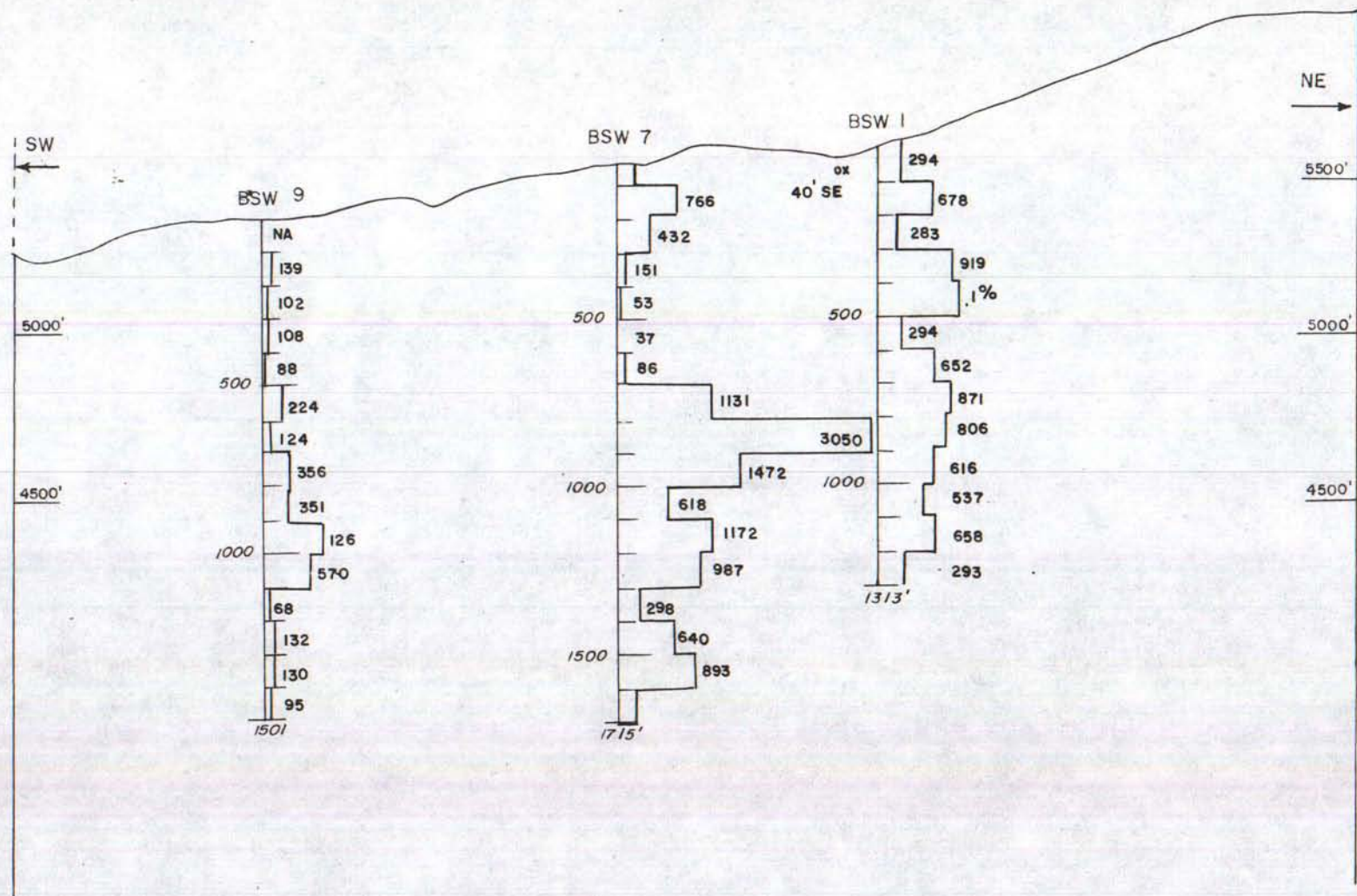
**Ag**

APPENDIX D  
FIGURE D-5

0 500'

WMO 3/78





BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 1, 7, 9  
 CONOCO COMPOSITE ASSAYS-100'

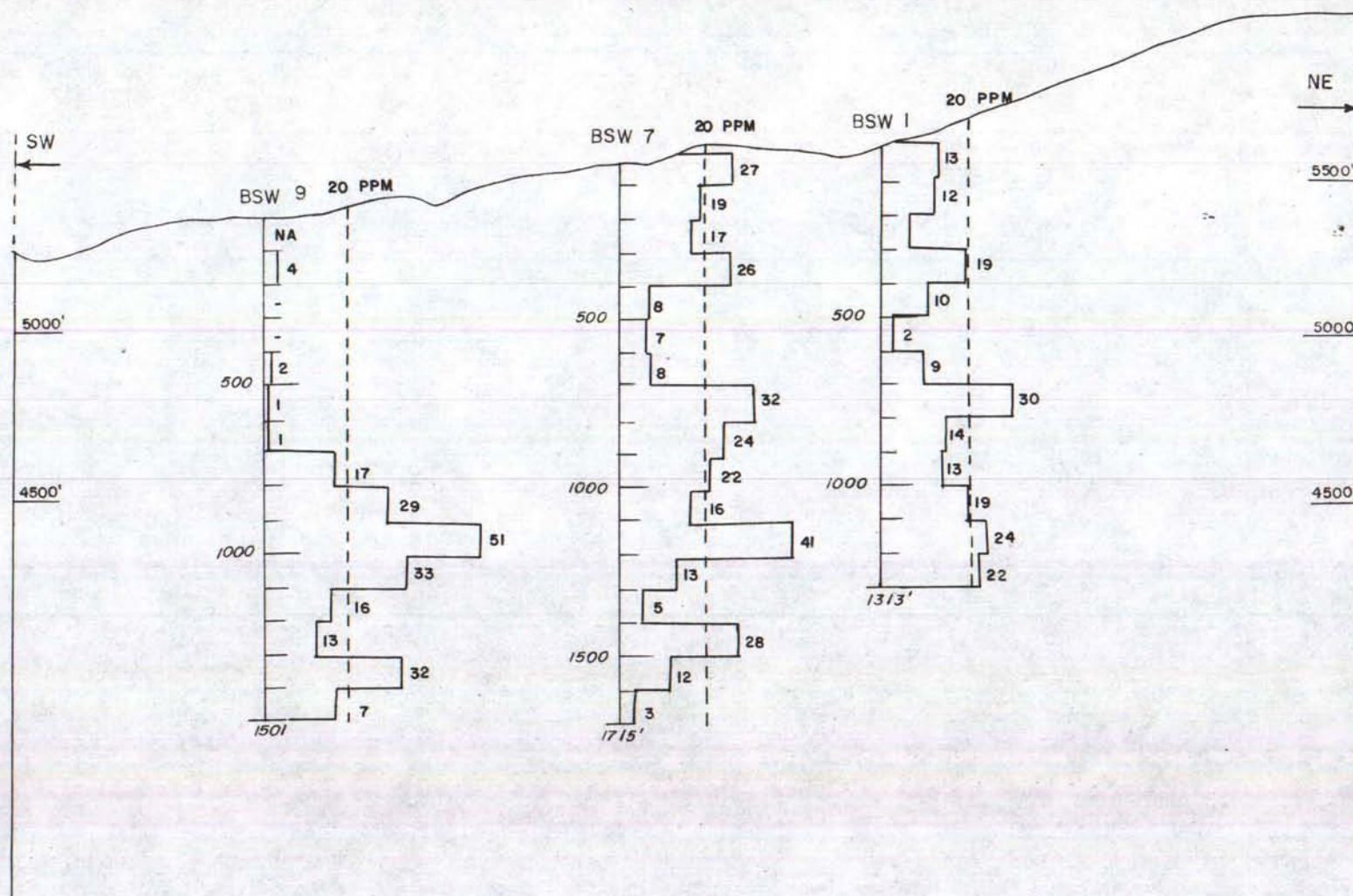
**Cu**

APPENDIX D  
 FIGURE D-6

0 500'

WMO 3/78





BUCKSKIN PROJECT

Drill Hole Cross Section

BSW 1, 7, 9

CONOCO COMPOSITE ASSAYS-100'

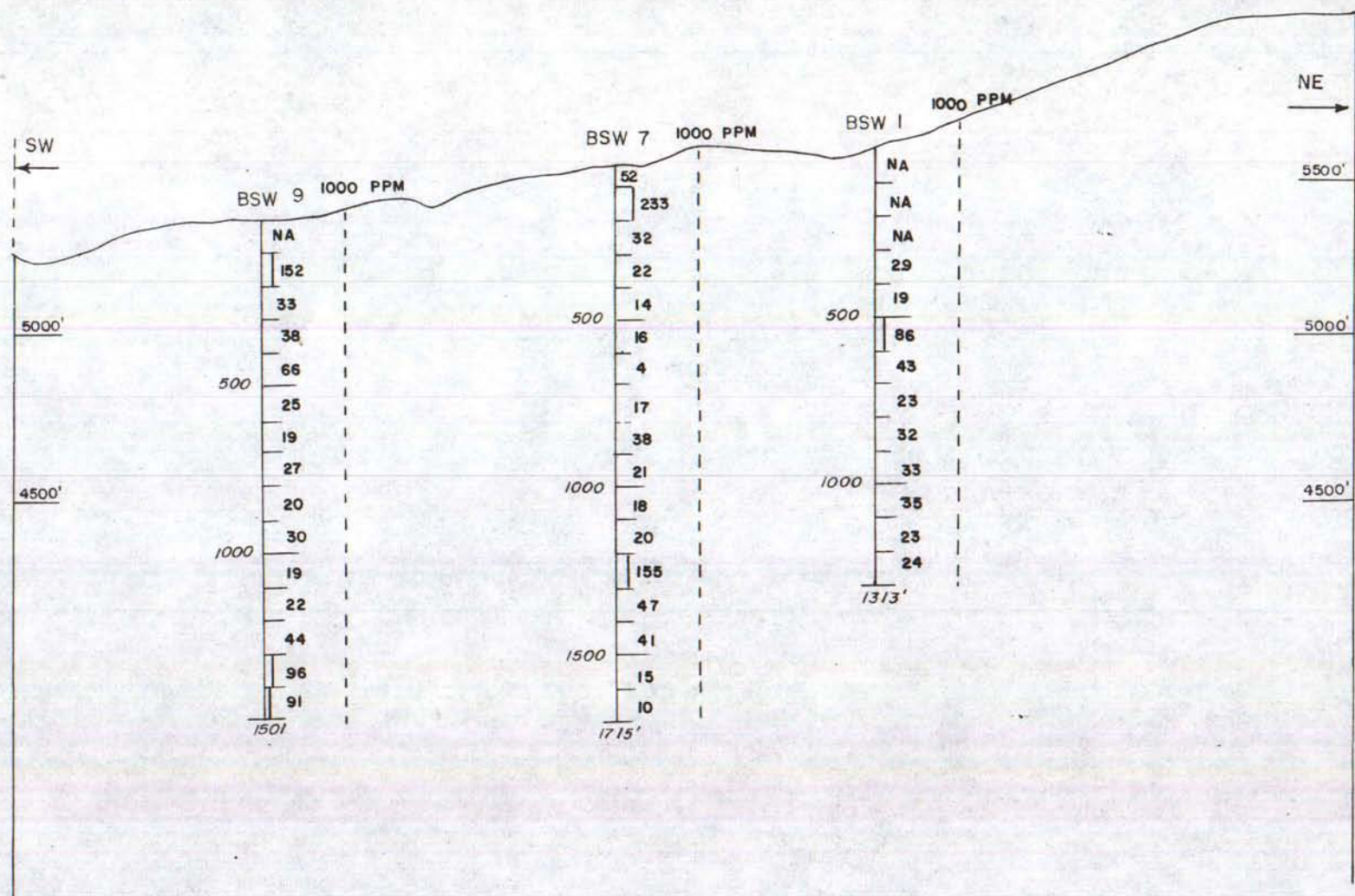
Mo

APPENDIX D  
FIGURE D-7

0 500'

WMO 3/78





BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 1, 7, 9  
 CONOCO COMPOSITE ASSAYS - 100'

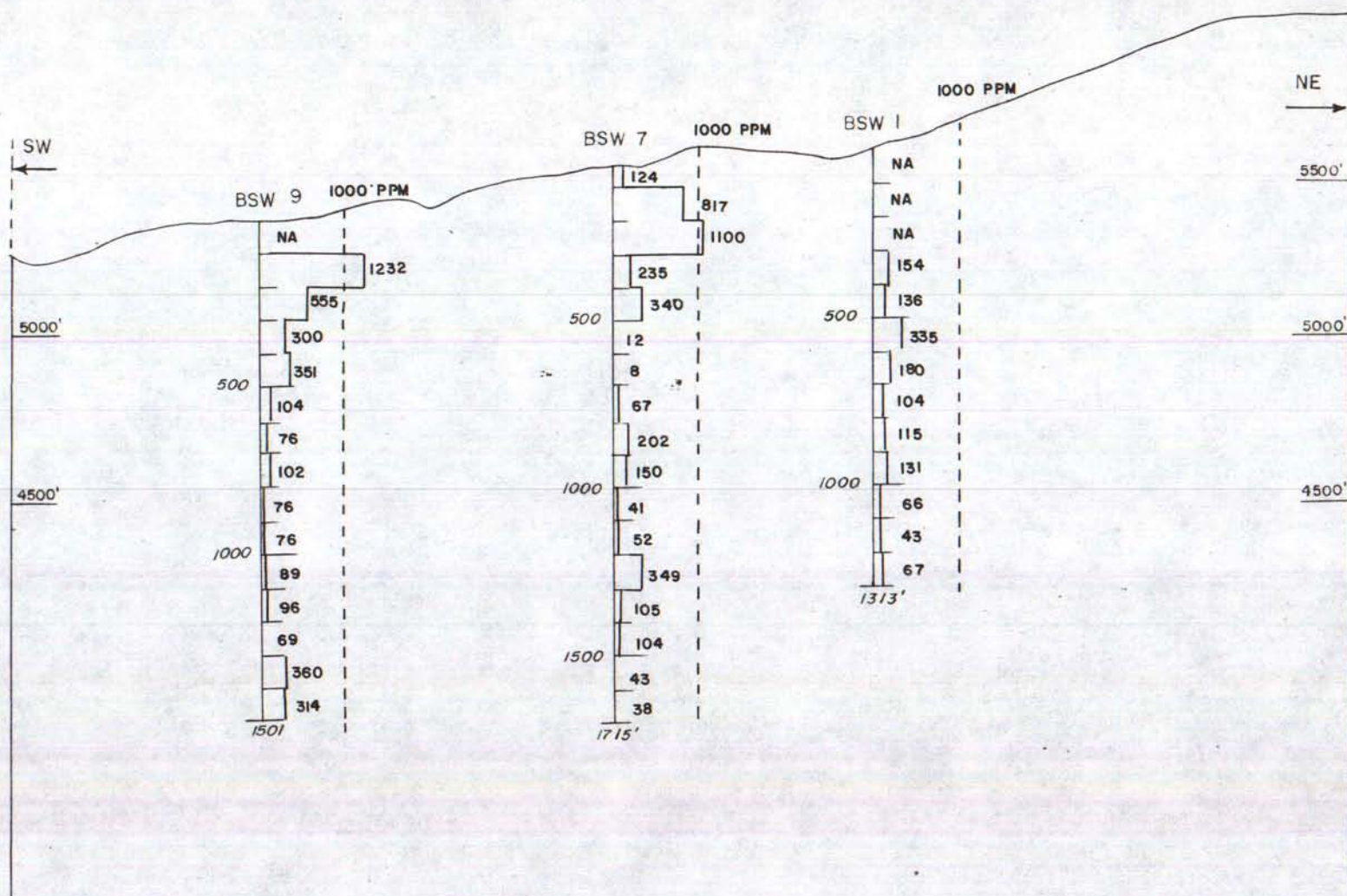
**Pb**

APPENDIX D  
 FIGURE D-8

0 500'

WMO 3/78





BUCKSKIN PROJECT

Drill Hole Cross Section

BSW 1, 7, 9

CONOCO COMPOSITE ASSAYS-100'

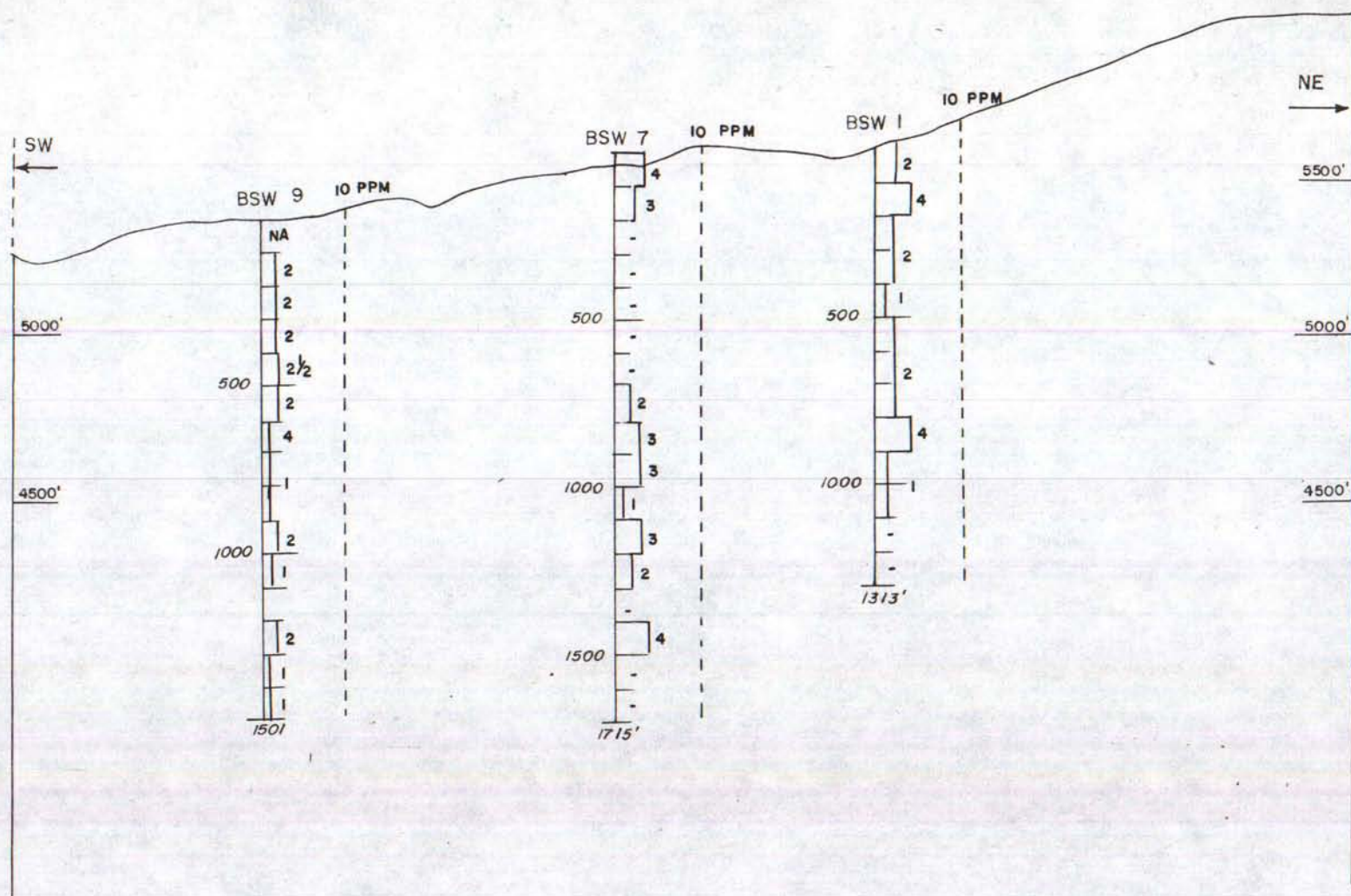
**Zn**

APPENDIX D  
FIGURE D-9

0 500'

WMO 3/78





# BUCKSKIN PROJECT

Drill Hole Cross Section

BSW 1, 7, 9

CONOCO COMPOSITE ASSAYS - 100'

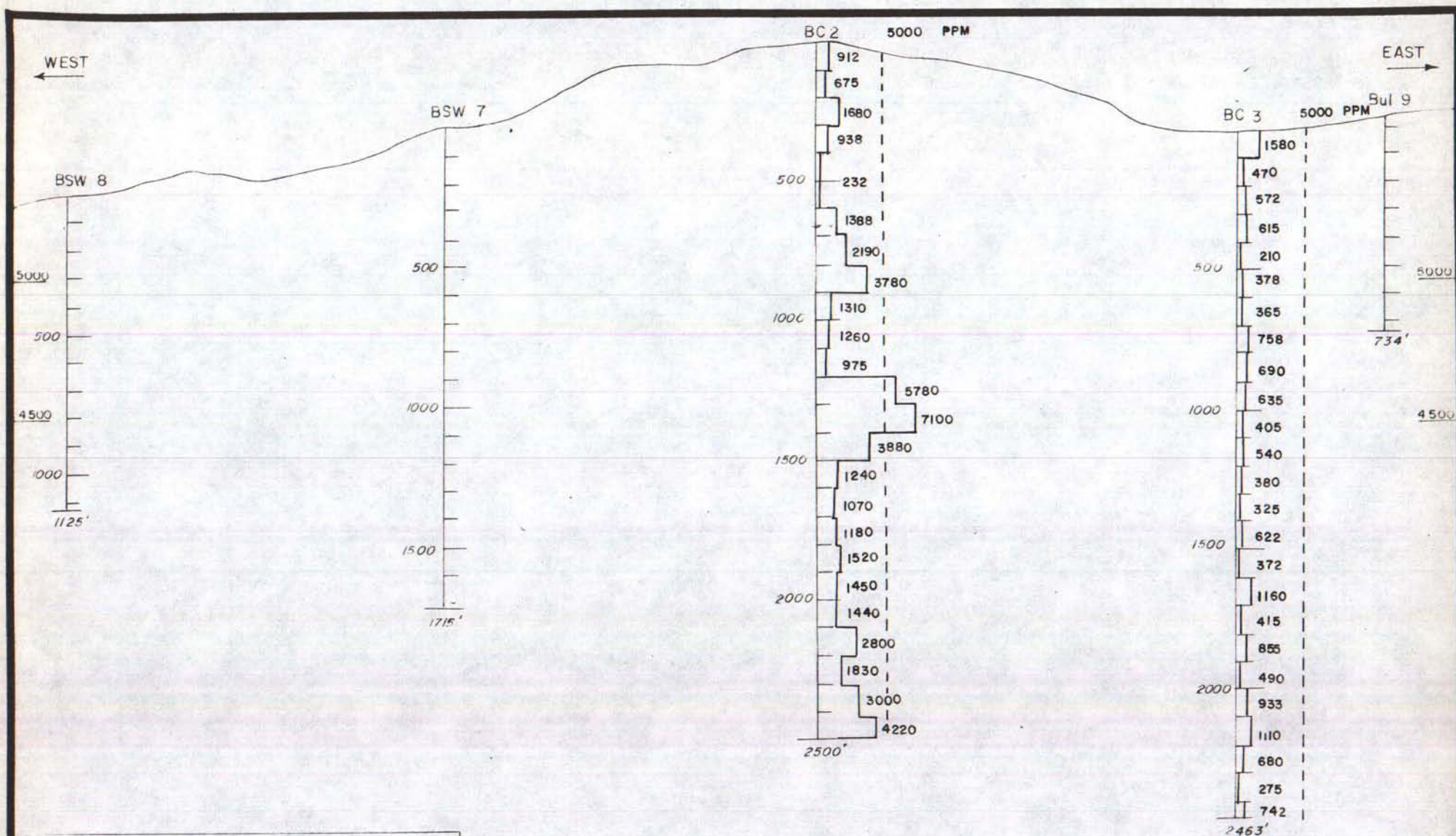
Ag

APPENDIX D  
FIGURE D-10

0 500'

WMO 3/78





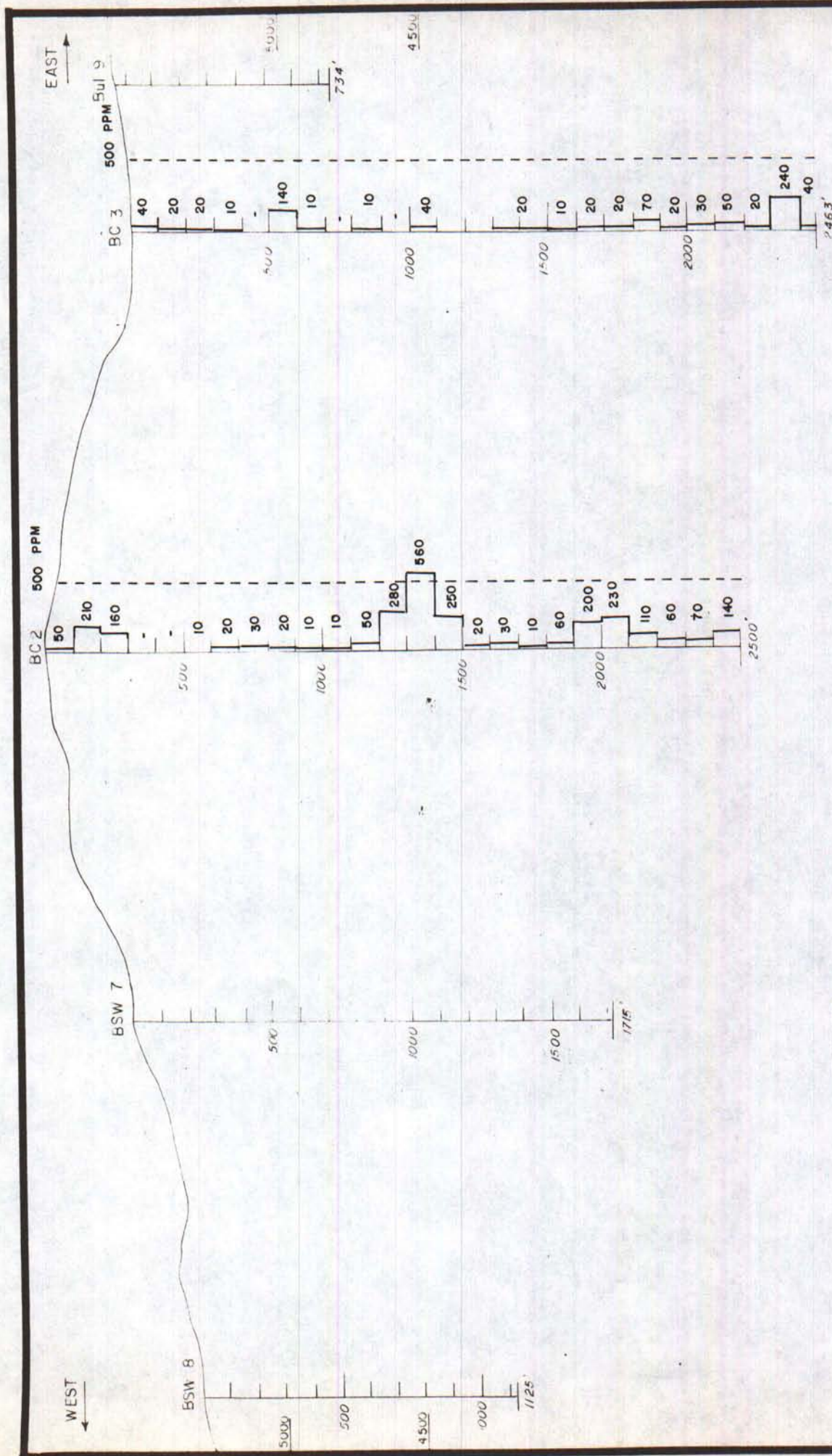
BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 8, 7, BC 2, 3  
 TOTAL SULFIDE CONCENTRATES  
 (Bamfield)  
 COPPER

APPENDIX E  
 FIGURE E-1

0 500'

WMU 3/78





BUCKSKIN PROJECT

Drill Hole Cross Section

BSW 8, 7, BC-2, 3

(Bamfield)

TOTAL SULFIDE CONCENTRATES

MOLYBDENUM

APPENDIX E

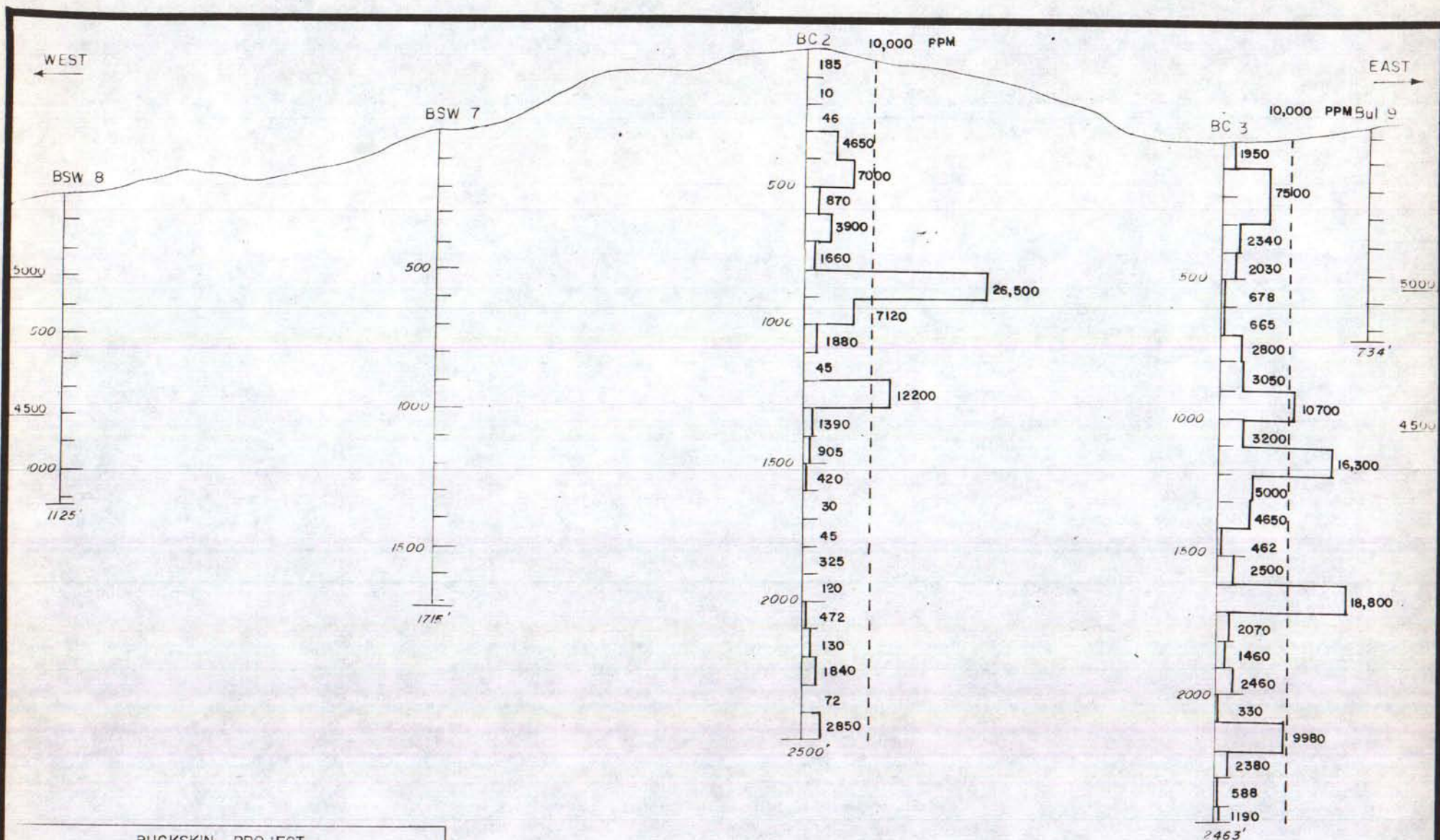
FIGURE E-2

WMU 3/16



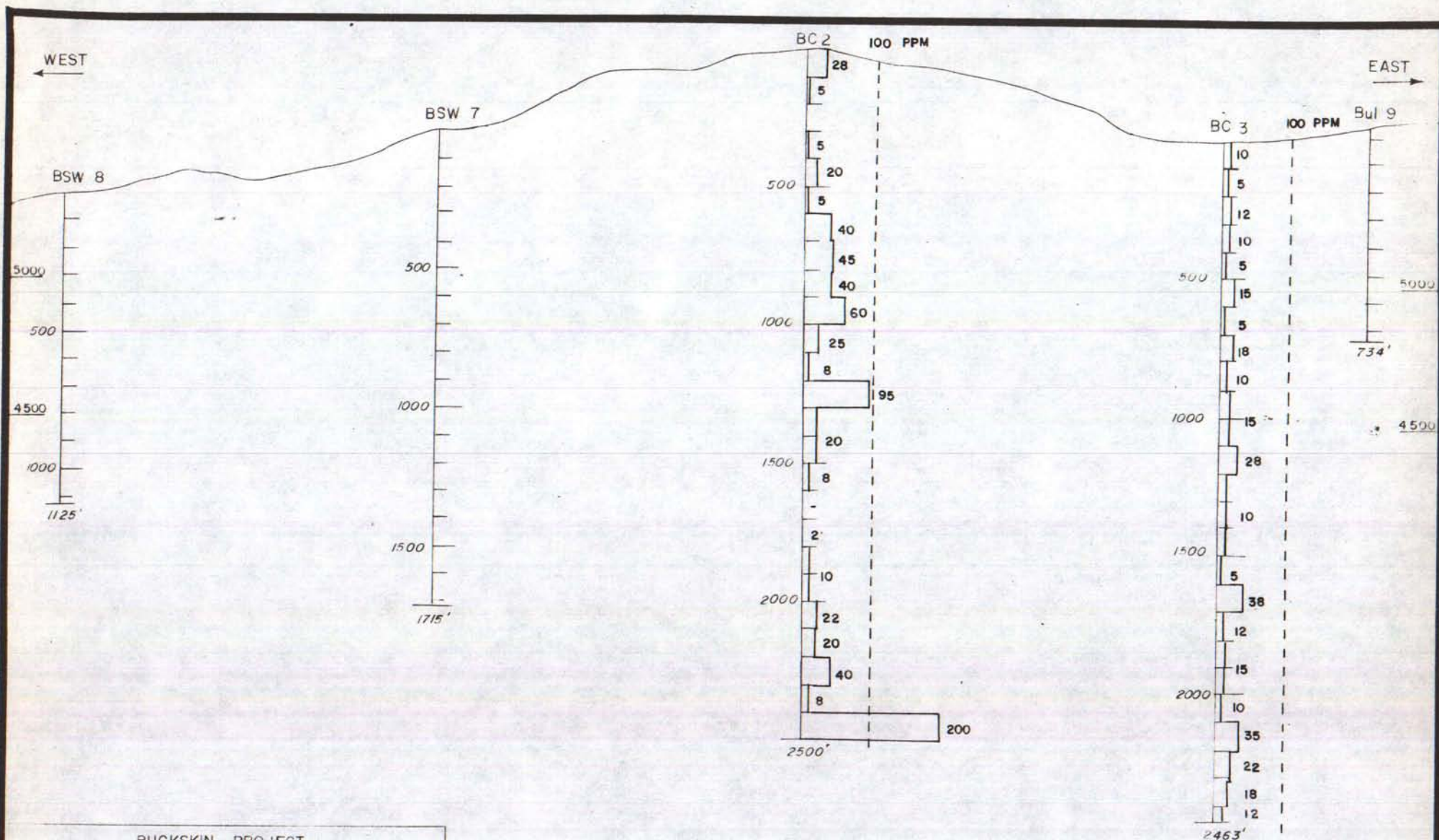






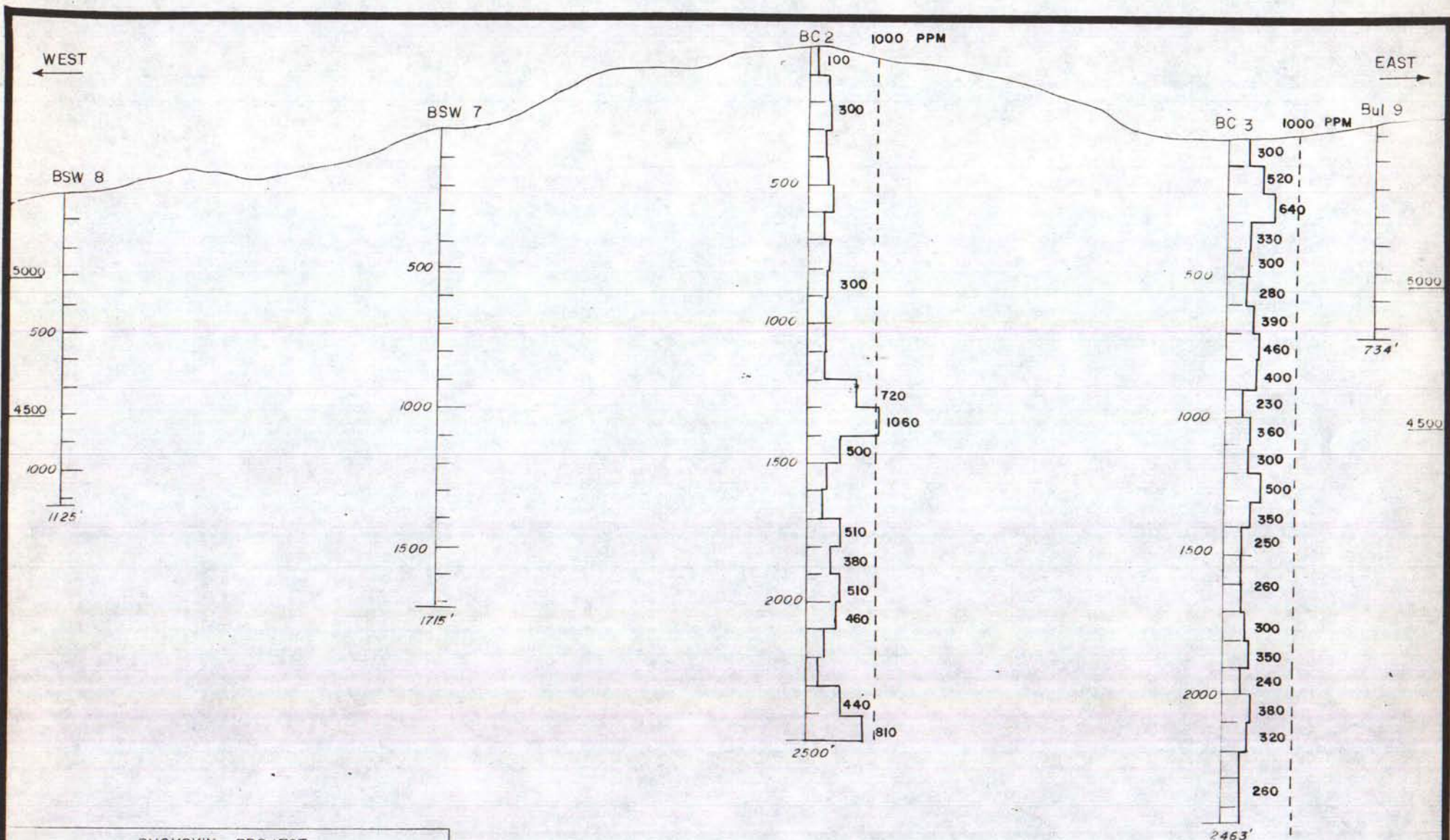
BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 8, 7, BC 2, 3  
 TOTAL SULFIDE CONCENTRATES  
 (Bamford)  
 ZINC  
 APPENDIX E  
 FIGURE E 4  
 0 500'  
 WMU 3/78





BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 8, 7, BC 2, 3  
 TOTAL SULFIDE CONCENTRATES  
 (Bamford)  
 SILVER  
 APPENDIX E  
 FIGURE E-5  
 0 500'  
 WMO 3/78





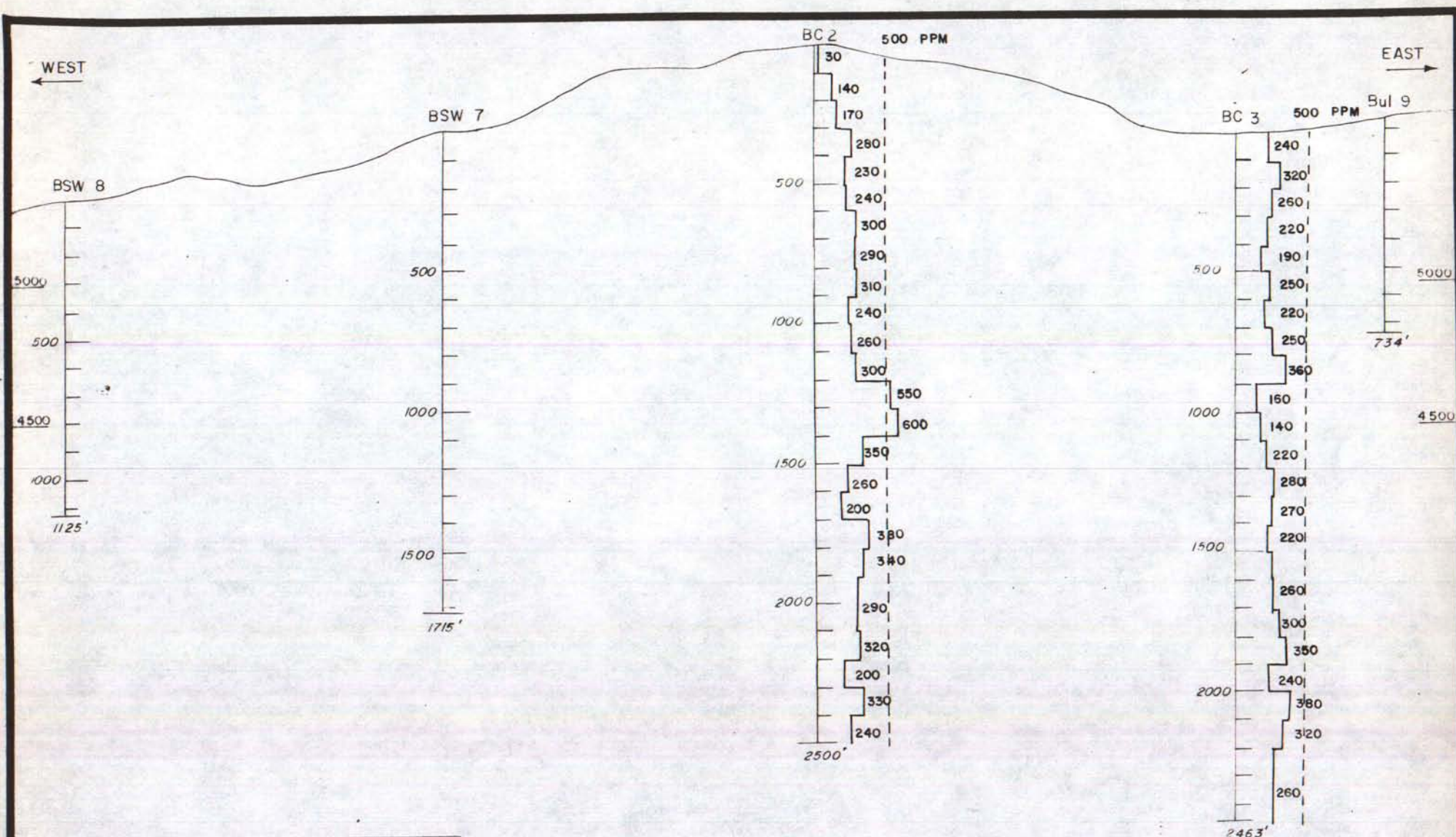
BUCKSKIN PROJECT  
 Drill Hole Cross Section  
 BSW 8, 7, BC-2, 3  
 TOTAL SULFIDE CONCENTRATES  
 (Bamfield)  
 NICKEL

APPENDIX E  
 FIGURE E6

0 500'

WMO 3/78





BUCKSKIN PROJECT

Drill Hole Cross Section

BSW 8, 7, BC-2, 3

TOTAL SULFIDE CONCENTRATES

(Bamford)

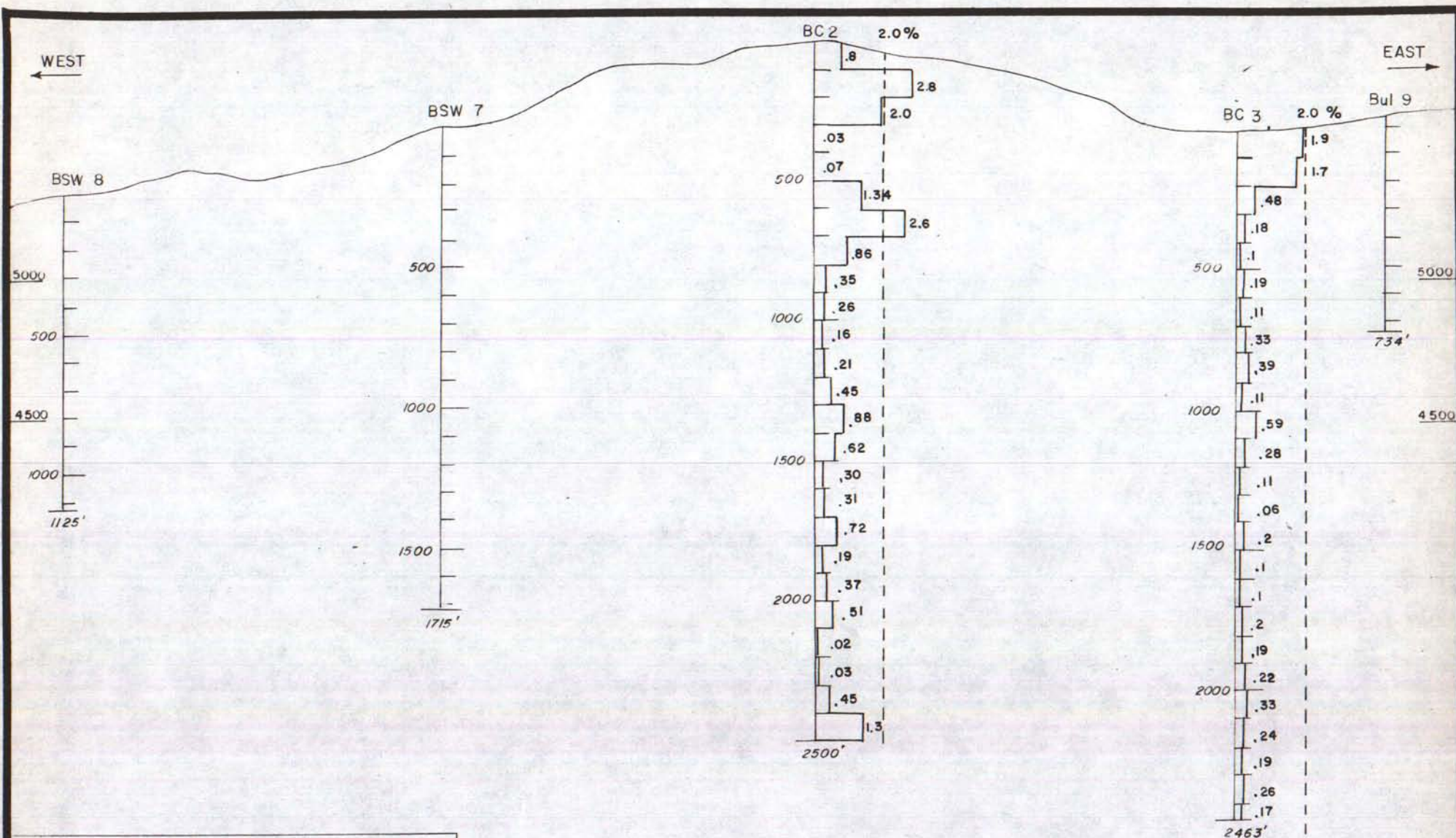
COBALT

APPENDIX E  
FIGURE E 7

0 500'

WMO 3/78





BUCKSKIN PROJECT

Drill Hole Cross Section

BSW 8, 7, BC-2, 3

TOTAL SULFIDE CONCENTRATES

(Bamford)

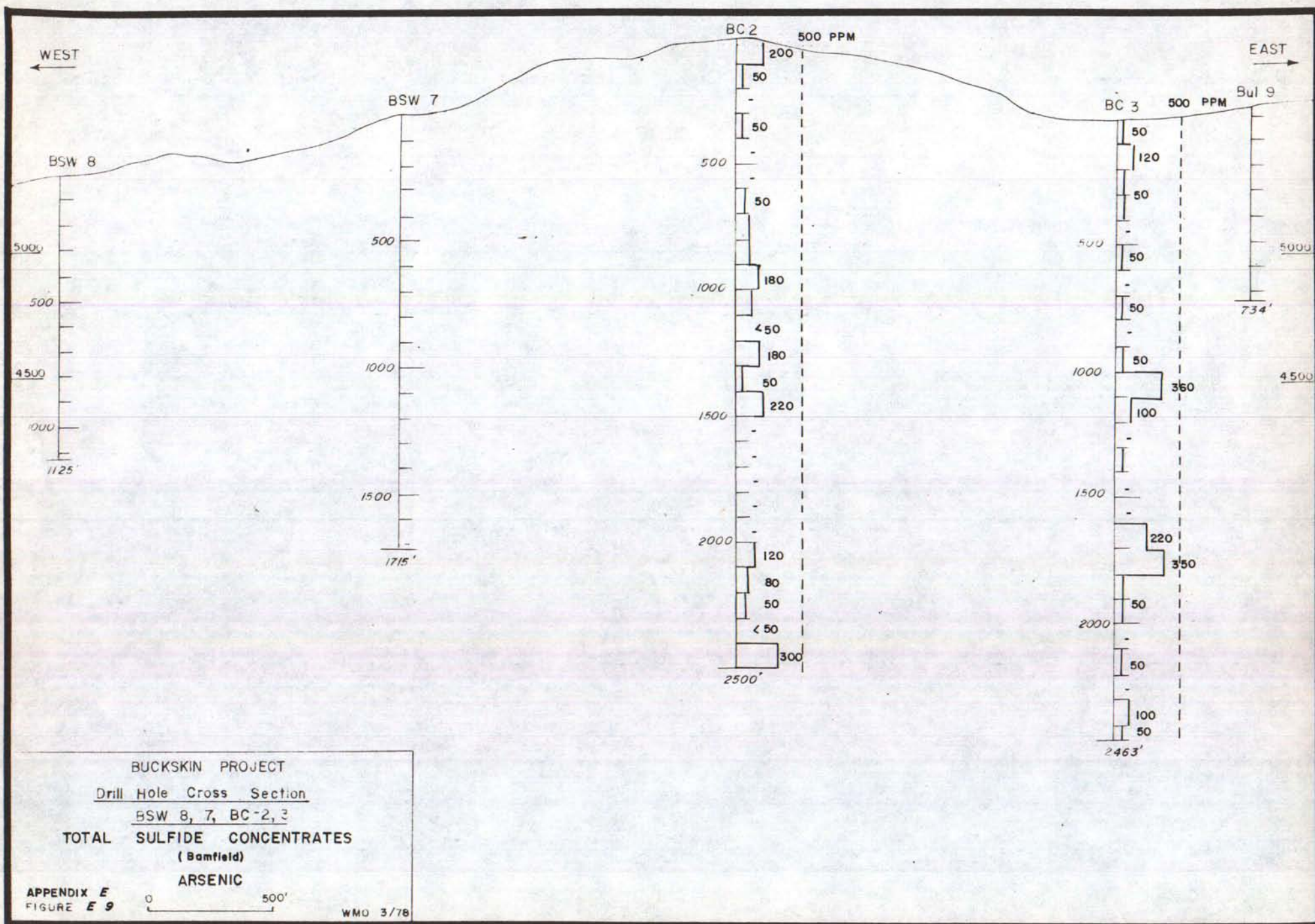
MANGANESE

APPENDIX E  
FIGURE E-8

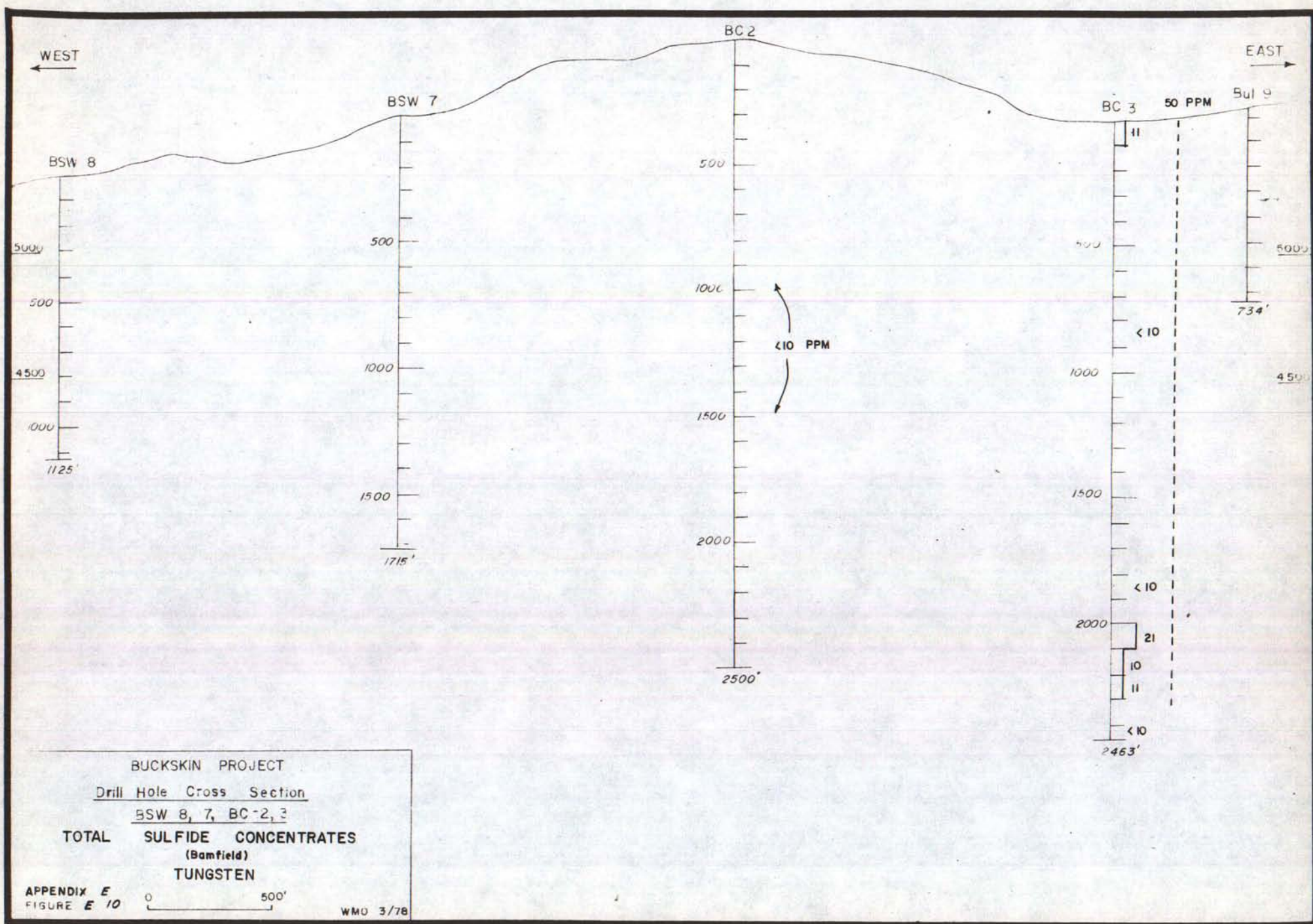
0 500'

WMO 3/78































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